Biphasic Behavior of Anammox Regulated by Nitrite and Nitrate in an Estuarine Sediment

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The production of N₂ gas via anammox was investigated in sediment slurries at in situ NO₂⁻ concentrations in the presence and absence of NO₃⁻. With single enrichment above 10 μ M ¹⁴NO₂⁻ or ¹⁴NO₃⁻ and ¹⁵NH₄⁺, anammox activity was always linear (P < 0.05), in agreement with previous findings. In contrast, anammox exhibited a range of activity below 10 μ M NO₂⁻ or NO₃⁻, including an elevated response at lower concentrations. With 100 μ M NO₃⁻, no significant transient accumulation of NO₂⁻ could be measured, and the starting concentration of NO₂⁻ could therefore be regulated. With dual enrichment (1 to 20 μ M NO₂⁻ plus 100 μ M NO₃⁻), there was a pronounced nonlinear response in anammox activity. Maximal activity occurred between 2 and 5 μ M NO₂⁻, but the amplitude of this peak varied across the study (November 2003 to June 2004). Anammox accounted for as much as 82% of the NO₂⁻ added at 1 μ M in November 2003 but only for 15% in May 2004 and for 26 and 5% of the NO₂⁻ at 5 μ M decreased the significance of anammox as a sink for NO₂⁻. The behavior of anammox was explored by use of a simple anammox-denitrification model, and the concept of a biphasic system for anammox in estuarine sediments is proposed. Overall, anammox is likely to be regulated by the availability of NO₃⁻ and NO₂⁻ and the relative size or activity of the anammox is population.

Recently, it has been demonstrated that in both aquatic sediments and anoxic water basins denitrification is not the only pathway for N_2 formation (3, 8, 19, 22). Anaerobic ammonium oxidation (anammox) can couple the reduction of NO_2^- to the oxidation of NH_4^+ , producing N_2 gas, and hence act as an additional pathway for the removal of N from aquatic ecosystems, yet what regulates the significance of anammox for N removal in aquatic systems is largely unknown (4). Given that N is often limiting for primary production in aquatic ecosystems, further understanding of this novel shunt in the N cycle is essential.

Research using microelectrodes and measurements of solute exchange, for example, O_2 , NO_3^- , and NH_4^+ , has established that the concentrations of solutes in sediment pore water and overlying water are at a steady state (11, 13, 14). Nitrate concentrations can reach 1 mM in the overlying water of hypernutrified estuaries and can still be present in the top 0 to 1 cm of sediment at 400 µM; however, the magnitude of both of these pools is seasonal (6, 9). In contrast, the concentrations of NO₂⁻ in both the overlying water and sediment are usually 2 orders of magnitude less than those of NO₃⁻ and seldom exceed 5 μ M, though in comparison to NO₃⁻, data for NO₂⁻ are scarce. Studies using NO2⁻ microsensors, biosensors, or fine-scale porewater extraction methods have indicated that net NO₂⁻ production occurs in anaerobic sediment as a result of NO₃⁻ reduction (15, 17). In addition, the first stage of nitrification at the aerobic-anaerobic interface produces NO₂⁻, which, depending on the conditions for nitrification and

the overall demand for NO_2^- , may diffuse into the underlying anaerobic layers (5). The critical point is that at a steady state, NO_2^- and NO_3^- can coexist in anaerobic sediments at concentrations separated by orders of magnitude.

To date, measurements of anammox in sediments have relied on anaerobic slurries or anaerobic bag incubations enriched with NO3⁻, NO2⁻, or both at concentrations higher than those in the environment, especially for NO_2^{-} (19, 22). In anaerobic slurries, in the direct presence of NO₃⁻, anammox is reliant on the initial reduction of NO_3^- to NO_2^- by the total NO_3^- -reducing community. Nitrate respiration ($NO_3^- \rightarrow$ NO_2^{-}) results in NO_2^{-} as an actual end product of metabolism which is exported from the cell. In contrast, NO_2^- is an intermediary of the denitrification $(NO_3^- \rightarrow NO_2^- \rightarrow NO \rightarrow N_2O \rightarrow N$ N_2) and dissimilatory reduction of nitrate to ammonium $(NO_3^- \rightarrow NO_2^- \rightarrow NH_4^+)$ pathways. All of this intermediary $\mathrm{NO_2}^-$ may "leak" from the cell into the surrounding environment. The loss of NO_2^- is in part governed by the kinetics of each respective enzyme, the species of NO₃⁻-reducing bacteria, and the quality of the organic substrate (1, 2). The linear formation of $^{29}\mathrm{N}_2$ from $^{15}\mathrm{NH_4^+}$ and $^{14}\mathrm{NO_2^-}$ in anaerobic slurries suggests that NO₂⁻ is nonlimiting for anammox, and hence the pathway of NO_2^{-} formation would be irrelevant to the actual significance of anammox, but these observations (19, 22) come from experiments that were not specifically designed to examine the behavior of anammox at representative in situ NO_2^- concentrations (<10 μ M). Even in the presence of NO_2^- and the absence of NO_3^- , the anammox community will have to compete against the largely heterotrophic NO3⁻ and NO₂⁻-reducing community, which will use NO₂⁻ just as efficiently in the absence of NO_3^{-} (22).

Although we cannot reproduce a steady state in batch sediment slurry experiments, unlike intact sediment cores, the

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purpose of this study was to investigate the fate of NO_2^- and the behavior of anammox at representative sediment $NO_2^$ concentrations in the presence and absence of NO_3^- . Here we show that anammox exhibits pronounced nonlinear behavior at concentrations below 10 μ M NO_2^- in the presence of 100 μ M NO_3^- . In addition, a simple anammox-denitrification model and the concept of a biphasic system for anammox in estuarine sediments are proposed.

MATERIALS AND METHODS

Sample sites and sediment slurry preparation. Sediment samples (oxic and suboxic layers from 0 to 2 cm) were collected from the intertidal flats at low tide from an area in the Thames estuary where significant anammox had previously been measured (site 2 [22]), stored in plastic bags, and returned to the laboratory within 1 h. Anaerobic ammonium oxidation was assayed by measuring the production of $^{29}N_2$ in anaerobic slurries as previously described (22). Sediment slurries were enriched by injection through the septa, using microliter syringes (Hamilton, Bonaduz, Switzerland), with concentrated stocks of ¹⁵NH₄Cl (120 mM ¹⁵NH₄Cl [99.3 ¹⁵N atom%]) (Sigma-Aldrich, Poole, United Kingdom) to give a final concentration of 500 µM $^{15}\mathrm{NH_4^+}$ (approximately 60% labeling of the $\mathrm{NH_4^+}$ pool) and either $\mathrm{Na^{14}NO_2}$ or Na¹⁴NO₃ (stock concentrations of 4 and 93 mM, respectively) (VWR International Ltd., Lutterworth, United Kingdom) to give a range of concentrations (see below). Sediments from this site are capable of reducing NO₃⁻ at approximately 100 μ M h⁻¹, and an overnight preincubation ensured that all traces of potential oxidants were removed before experimental manipulations commenced (22). All incubations were carried out on rollers in the dark in a constant-temperature room at 15°C. Following incubation, gas samples were collected (1 ml) using a gas-tight syringe (SGE Gastight Luer Lock syringe; Alltech Associates Ltd., Carnforth, Lancashire, United Kingdom) and stored in an inverted water-filled gas-tight vial (12-ml Exetainer; Labco Ltd., High Wycombe, United Kingdom). For determinations of the labeling of the NH4⁺ pool, pore waters were recovered from both enriched and nonenriched reference slurries by centrifugation of the slurry, filtered (0.2-µm-pore-size Minisart Plus filter; Sartorius UK Ltd.), and frozen (-20°C) until later analysis.

Anammox activity with single NO₂⁻ or NO₃⁻ enrichment. For the first set of experiments, slurries were enriched with ¹⁵NH₄⁺ and then independently with either ¹⁴NO₃⁻ or ¹⁴NO₂⁻ to give final concentrations of 1, 2.5, 5, 8, 10, 12, 16, and 20 μ M, incubated overnight, and treated as described above.

Nitrate reduction and nitrite accumulation. In order to determine the effect of the NO₃⁻ enrichment concentration on NO₂⁻ accumulation, we enriched prepared sediment slurries with ¹⁴NO₃⁻ at 100, 200, 400, and 600 μ M and sacrificed independent slurries every hour for 4 h. The microbial activity was inhibited by injecting ZnCl₂ (500 μ l of a 50% [wt/vol] solution) through the septum of the serum bottle, and the pore waters were recovered and treated as described above for later analyses of NO₃⁻ and NO₂⁻ (see below).

Anammox activity with dual NO₂⁻ and NO₃⁻ enrichment. Having established the concentration of NO₃⁻ at which NO₂⁻ did not transiently accumulate in the slurries (100 μ M NO₃⁻ [see Results]), we prepared dual-enrichment experiments as outlined above and added isotopes from concentrated stocks to give the following final concentrations, in this order: ¹⁵NH₄⁺ at 500 μ M; ¹⁴NO₃⁻ at 100 μ M followed by mixing for 5 min on rollers (during which time approximately 8% of the NO₃⁻ pool would have turned over); and finally, ¹⁴NO₂⁻ at 1, 2.5, 5, 8, 10, 12, 16, or 20 μ M. Subsequent dual-enrichment experiments were performed with 25, 50, 75, and 100 μ M ¹⁴NO₃⁻ and with ¹⁴NO₂⁻ at just 5 μ M. The slurries were incubated overnight and treated as described above.

Analysis. All nutrient analyses (NO₃⁻, NO₂⁻, and NH₄⁺) were performed by use of a continuous-flow autoanalyzer (San⁺⁺; Skalar, Breda, The Netherlands) and standard colorimetric techniques (7). Salinities were measured using a handheld refractometer. Water contents, specific gravities, and porosities were determined from the dry and wet weights of known volumes of sediment. Samples of the headspace (50 μ l) from the gas-tight vials were injected using an autosampler into an elemental analyzer interfaced with a continuous-flow isotope ratio mass spectrometer calibrated with air, and the mass charge ratios for *m*/*z* 28, 29, and 30 nitrogen (²⁸N₂, ²⁹N₂, and ³⁰N₂) were measured (Delta Matt Plus; Thermo-Finnigan, Bremen, Germany).

Calculation of total anammox activity. Assuming that the ¹⁵NH₄⁺ pool turns over at the same rate as the ambient ¹⁴NH₄⁺ pool, the total anammox N₂ production can be calculated from the production of ²⁹N₂ and the proportionate

 ^{15}N labeling of the $\rm NH_4^+$ pool, determined by accounting for the difference from nonenriched reference slurries (19, 22). Anammox can be measured by labeling either the electron donor ($^{15}\rm NH_4^+$) or acceptor ($^{15}\rm NO_3^-$ and $^{15}\rm NO_2^-$) pool (19), and we have previously demonstrated a high level of agreement between these two techniques (22). For this study, however, only the $\rm NH_4^+$ pool was labeled, and our measures of anammox are therefore expressed relative to total $\rm NO_3^-$ and $\rm NO_2^-$ reduction and not total gas formation.

Modeling the behavior of anammox. A simple model was developed in EcoS (version 3.37, N.E.R.C.; Plymouth Marine Laboratory, Plymouth, United Kingdom) to examine the behavior (shape of response) of anammox in the dual presence of NO_2^- and NO_3^- in a slurry system with a mixed population of anammox and denitrifying bacteria. The model is based around the assumption that anammox, as an autotroph, makes up a relatively small proportion (θ) of the total NO_3^- and NO_2^- reducer population (<5%) and can only use NO_2^- as an oxidant, whereas the denitrifying population will use NO_3^- preferentially but will swap to NO_2^- when NO_3^- falls below a certain threshold (22). In addition, for the sake of simplicity, the additional potential source of NO_2^- for anammox resulting from denitrification was not included (see Results for rationale); hence, as in the above experiments, the mixed population was started with only external sources for both NO_2^- and NO_3^- .

The changes over time of NO_2^- and NO_3^- concentrations in the pore water, subject to consumption by a bacterial population with the proportion θ of anammox and $1 - \theta$ of denitrifiers, were assumed to be governed by the following equations:

$$\frac{d}{dt}[\mathrm{NO}_2^-] = -\theta M f_a - (1-\theta) M f_d \tag{1}$$

$$\frac{d}{dt}[\mathrm{NO}_3^{-}] = -(1-\theta)Mg_d \tag{2}$$

where *M* is the number of bacteria per liter. The rates of consumption of NO_2^- by denitrifiers as a function of $[NO_2^-]$ were assumed to be given by the following equations:

$$f_d = \beta V_{M2D}[NO_2^{-}]/([NO_2^{-}] + k_{NO2})$$
(3)

$$g_d = V_{M3D}[NO_2^{-}]/([NO_2^{-}] + k_{NO3})$$
(4)

In these equations, f_d and g_d are classic Michaelis forms with half-saturation constants k_{NOx} and maximum rates V_{MxD} (µmol N bacterium⁻¹ day⁻¹). The following function acts to block the denitrifier reduction of NO₂⁻ when NO₃⁻ is available:

$$\beta = MAX(1 - [NO_3^{-}]/[NO_3^{-}]_{crit}, 0)$$
(5)

It gives a linear decrease in rate for increasing [NO₃⁻], with no uptake of NO₂⁻ occurring when [NO₃⁻] > [NO₃⁻]_{crit}, in this case, 100 μ M NO₃⁻. The rate of consumption of NO₂⁻ by anammox was modeled as a biphasic process, determined by the following equation:

$$f_a = V_{MA}(f_p + f_m) \tag{6}$$

where V_{MA} (µmol N bacterium⁻¹ day⁻¹) is a maximal processing rate, and the modifying functions f_p and f_m , with $0 < f_p$ and $f_m < 1$, are given by the following equations:

$$f_p = \text{Exp}[1.0] \times ([\text{NO}_2^{-}]/l_N) \text{Exp}(-[\text{NO}_2^{-}]/l_N)$$
 (7)

$$f_m = \varepsilon[\mathrm{NO}_2^-]/([\mathrm{NO}_2^-] + K_m) \tag{8}$$

The function f_p yields an optimum rate when $[NO_2^{-1}] = l_N$, with inhibition for $[NO_2^{-1}]$ above and below this rate (see Fig. 3). The function f_m , which is Michaelis in form, yields a constant reduction rate for NO_2^{-1} at concentrations more than six times K_m . The combination yields a peak of $\sim V_{MA}$ when $[NO_2^{-1}] = l_N$ and smoothly decreases to a proportion, ε , of the maximal rate at high NO_2^{-1} concentrations. We emphasize that it was not the purpose of this modeling exercise to precisely replicate the conditions and dimensions of our experimental slurries, but rather just to explore the overall shape of the anamnox activity. While the half-saturation constants for anamnox and denitrification are realistic, at 2.5 μ M (reference 4 and data reported here) and 15 μ M (1), respectively, the total number of bacteria was set to 10^4 liter⁻¹ and the maximum rate of reaction for anamnox and denitrification was set to $10^{-3} \mu$ mol N bacterium⁻¹ day⁻¹.



FIG. 1. (a and c) Production of ${}^{29}N_2$ from labeled ${}^{15}NH_4^+$ oxidation in the direct presence of ${}^{14}NO_2^-$ or ${}^{14}NO_3^-$ in November and December 2003. (b and d) Total anammox N_2 production (calculated from ${}^{29}N_2$ production) as a percentage of the respective NO_2^- or NO_3^- spike. (e and f) Composite of the above NO_2^- data plus additional NO_2^- data for October 2003. Values are means ± 1 standard error of the means (SEM) (n = 4).

RESULTS

Anammox activity with single NO₂⁻ or NO₃⁻ enrichment. Above 10 μ M NO₂⁻ or NO₃⁻, anammox activity was always linear (P < 0.05) (Fig. 1). In contrast, anammox exhibited a range of activity below 10 μ M NO₂⁻ or NO₃⁻ (Fig. 1). For example, at concentrations below 10 μ M for either NO₂⁻ or NO₃⁻, ²⁹N₂ production remained constant (Fig. 1a) and, in turn, the percentages of the electron acceptors accounted for by anammox increased to 7 and 4%, respectively (Fig. 1b). Alternatively, the production of $^{29}N_2$ increased linearly from 2.5 to 20 μ M NO₂⁻ or NO₃⁻, and the percentage accounted for by anammox remained constant at <2% (Fig. 1c and d). Overall, the single-enrichment activities were highly variable below 10 μ M, and on occasion (October and November 2003), several data points supported elevated activities below 5 μ M (Fig. 1e and f). In contrast, the heterotrophic NO₃⁻- and



FIG. 2. (a) Reduction of NO₃⁻ over time in anaerobic sediment slurries at four NO₃⁻ enrichment concentrations. (b) Subsequent accumulation and reduction of NO₂⁻ over time for each respective NO₃⁻ enrichment. Values are means ± 1 SEM (n = 4). The same symbols are used in each panel.

 NO_2^- -reducing community accounted for >93 and 96% of the NO_2^- or NO_3^- , respectively.

Nitrate reduction and nitrite accumulation. In slurries enriched with 100 μ M NO₃⁻, all of the NO₃⁻ had been consumed within the first half-hour and no significant transient accumulation of NO₂⁻ could be measured (Fig. 2). At 200, 400, and 600 μ M NO₃⁻, the slurries exhibited similar rates of NO₃⁻ reduction, i.e., 103, 105, and 97 μ M h⁻¹, respectively. These data agree with the rate of 85 μ M h⁻¹ measured at the same site in the previous year (22). In addition, those samples enriched with 200 μ M NO₃⁻ and above exhibited a transient accumulation of NO₂⁻ of 19 μ M after 1 h. Therefore, 100 μ M NO₃⁻ could, at least during our short dual incubations, satisfy total NO₃⁻ reduction and enable the initial availability of NO₂⁻ to be regulated.

Anammox activity with dual NO₂⁻ and NO₃⁻ enrichment. In the dual presence of NO₂⁻ (1 to 10 μ M) and NO₃⁻ (100 μ M) there was an even more pronounced nonlinear response in anammox activity, relative to those in the single enrichments (Fig. 3a). In addition, there was a marked increase in overall anammox activity in the dual-enrichment experiments. For example, the anammox activity increased from 0.01 nmol N₂ ml⁻¹ wet sediment at 5 μ M NO₂⁻ without NO₃⁻ to 0.6 nmol N₂ ml⁻¹ wet sediment with NO₃⁻. In turn, the presence of NO₃⁻ increased the percentage of NO₂⁻ accounted for by anammox from 0.15 to 8% at 5 μ M NO₂⁻. As with the previous single-enrichment experiments, anammox exhibited a range of activities, although they all had a similar shape (Fig. 3c). The maximum anammox activity in the presence of both NO_2^- and NO_3^- was measured in November 2003 (2.6 nmol $N_2 \text{ ml}^{-1}$ wet sediment at 5 μ M NO₂⁻), and the lowest, to date, was measured in May 2004 (0.3 nmol $N_2 \text{ ml}^{-1}$ wet sediment at 5 μ M NO₂⁻). In November, anammox accounted for as much as 82% of the NO₂⁻ at 1 μ M, but it only accounted for 15% in May and between 26 and 5% at 5 μ M NO₂⁻, respectively, for these two months (Fig. 3d). Using 5 μ M NO₂⁻ and 100 μ M NO₃⁻ as a reference measure for anammox, we show the data collected to date in Table 1, along with some site characteristics.

In subsequent experiments, dual enrichment was repeated with 5 μ M NO₂⁻ and with NO₃⁻ from 0 to 100 μ M (Fig. 3e and f). Clearly, as the concentration of NO₃⁻ was reduced, the anammox activity at 5 μ M NO₂⁻ declined. In addition, as with all the previous experiments, there was a change in the relative significance of anammox to NO₂⁻ reduction on each occasion.

Modeling the behavior of anammox. The model was run for the component forms (f_p [equation 7] and f_m [equation 8]) and the combined form of the biphasic anammox expression (f_a [equation 6]), with the following parameters: $\varepsilon = 0.15$, $l_N = 2.5$ μ M, and $K_m = 2.5 \mu$ M. In addition, for each component and combined run, the proportion of anammox bacteria (θ) within the model was varied from 0.3 to 4.5%, and the results for the various scenarios are given in Fig. 4.

TABLE 1. Summary of anammox activities measured under dual enrichment with 5 μ M NO₂⁻, 100 μ M NO₃⁻, and 500 μ M ¹⁵NH₄⁻ at 15°C from November 2003 to June 2004, with some site water characteristics

Date	Anammox activity (nmol of N_2 ml of wet sediment ⁻¹) ^a	NO_3^- (μM)	NO_2^- (μM)	Temperature (°C)	Salinity (psu)
November 2003	2.6 ± 0.17	167.14	0.86	12	20
December 2003	1.3 ± 0.13	562.86	2.07	10	12
February 2004	0.5 ± 0.15	622.14	10.64	8	4
May 2004	0.3 ± 0.04	427.29	1.23	13	17
2 June 2004	0.6 ± 0.07	409.59	0.95	17	17
16 June 2004	0.6 ± 0.08	387.63	0.83	22	20

^{*a*} Data are means \pm standard errors (n = 4 for each date).



FIG. 3. (a) Total anammox (AAO) N₂ production in the direct presence of 1 to 10 μ M ¹⁴NO₂⁻, 100 μ M ¹⁴NO₃⁻, and 500 μ M ¹⁵NH₄⁺. (b) AAO as a percentage of each NO₂⁻ spike. (c and d) Representative ranges of patterns for AAO with NO₂⁻ from 1 to 20 μ M and 100 μ M ¹⁴NO₃⁻. (e and f) AAO with 5 μ M ¹⁴NO₂⁻ and decreasing concentrations of ¹⁴NO₃⁻. Note that a measurement with 0 μ M ¹⁴NO₃⁻ was not determined in November 2003. Values are means ± 1 SEM (*n* = 4). The same symbols are used for each respective left and right panel.

DISCUSSION

In our first study of anammox, we had demonstrated a linear response for anammox (production of $^{29}N_2$) between 200 and 800 μ M NO₃⁻ or NO₂⁻, with a constant yield of $^{29}N_2$ per molecule of NO₃⁻ or NO₂⁻ reduced (22). Here we have shown that this relationship holds for either acceptor down to a concentration of 10 μ M but that below this level, anammox exhibits nonlinear behavior, with an optimum at 2 to 5 μ M NO₂⁻. Previous work with sediments had suggested that ana-

mmox has a K_m for NO₂⁻ of 3 μ M (4), and the results presented here confirm this observation. In single-enrichment trials with concentrations above 10 μ M NO₃⁻, anammox could account for approximately 1% of the reduced NO₃⁻. This suggests that as more NO₃⁻ was added, a fixed proportion was lost as NO₂⁻ and, in turn, reduced via anammox, and hence the linear response (Fig. 1a and c). The flat response below 10 μ M NO₃⁻ may have simply been due to the preferential affinity by anammox for the lower concentrations of NO₂⁻ re-



FIG. 4. Model runs for anammox, with θ varying between 0.003 and 0.045 in each case. (a) Anammox as non-Michaelis f_p (equation 7); (c) anammox as Michaelis f_m (equation 8); (e) anammox as a function of the combined components $f_p + f_m$ (equation 6). (b, d, and f) NO₂⁻ accounted for by anammox for each respective scenario (a, c, and e).

sulting from NO₃⁻ reduction or, in part, to a disproportionately greater loss of NO₂⁻ from NO₃⁻ reduction below 10 μ M NO₃⁻ (Fig. 1a). It has been suggested (12) that periplasmic NO₃⁻ reductase has a relatively high affinity (Nap; K_m , 15 μ M) and may enable bacteria to scavenge NO₃⁻ at low concentrations, while membrane-bound NO₃⁻ reductase (Nar; K_m , 50 μ M) is more suited to higher concentrations. Hence, a changing proportion of NO₂⁻ loss may reflect different enzyme systems, e.g., the loss of NO₂⁻ may be greater under Nap rather than Nar activity. The flat response below 10 μ M NO₂⁻ indicates that the anammox community takes proportionately more of the available NO₂⁻ at lower concentrations, again reflecting the high affinity for NO₂⁻. However, it is difficult to rationalize why the yield of anammox remained constant in the direct single presence of either NO₂⁻ or NO₃⁻ (Fig. 1a and c). In the single presence of NO₂⁻ the NO₂⁻ is obviously extra-

cellular and thus directly available to the anammox community, yet the yield of anammox was the same as that for NO₃⁻ when the anammox community was reliant on a supply of NO₂⁻ originating from the NO₃⁻-reducing community. For this to be the case, the availability of NO₂⁻ for anammox would have to be the same under both scenarios, either direct NO₂⁻ or NO₃⁻ reduced to NO₂⁻. This suggests either that NO₂⁻ is never limiting in single-enrichment experiments or that some anammox bacteria may actually be coupled (both physically and metabolically) to members of the NO₃⁻-reducing community and cannot proceed faster than the rate governed by the metabolism of the heterotrophic NO₃⁻ and NO₂⁻ reducers.

At a steady state in sediments, NO₃⁻ and NO₂⁻ coexist at concentrations separated by orders of magnitude, and it was only in the dual-enrichment experiments that the nonlinear behavior became fully apparent (Fig. 3). Clearly, the anammox activity was regulated by the presence of NO_3^{-} , and as the availability of NO3⁻ was reduced, the anammox activity and its significance as a sink for NO₂⁻ declined (Fig. 3e and f). Hence, although anammox bacteria have a high affinity for NO₂⁻, they are most likely vastly outnumbered by the heterotrophic NO₃⁻⁻ and NO₂⁻⁻ reducing community, which will reduce NO_2^- when NO_3^- is limiting (22). The final path of $NO_2^$ reduction is therefore modulated by the dual availability of NO_3^- and NO_2^- and the overall competition for electron acceptors. The data presented in Fig. 2 suggest that if the availability of NO_3^{-1} is greater than that required for NO_3^{-1} reductase to operate at $V_{\rm max}$, then NO₂⁻ is exported to the sediment, while below this it is not, explaining why the nonlinear behavior only became clearly visible when the loss of NO₂⁻ from NO₃⁻ reduction was controlled and NO₂⁻ was made available only under these controlled conditions. Despite the lack of NO_2^- leakage from NO_3^- reduction below 100 μM NO₃⁻, anammox was always measurable in the single-enrichment experiments (Fig. 1), which again suggests that some of the anammox bacteria are tightly coupled to the turnover of NO₃⁻. It is interesting that NO₃⁻ was reduced in these estuarine sediments (Thames) at 20 times the rate ($\sim 100 \ \mu M \ h^{-1}$ compared to 5 μ M h⁻¹) of those in the Skagerrak, where anammox has also been measured (4, 19) and where its relative significance to N₂ formation is much greater than that for the Thames. Assuming that NO_3^- reduction was operating at V_{max} in the Skagerrak (4, 19), as shown for the Thames in Fig. 2b, then of the reduced NO_3^- , 60% accumulated as NO_2^- in the Skagerrak but only 20% accumulated as NO_2^- in the Thames. In the Skagerrak, therefore, the greater availability of NO₂⁻ from NO₃⁻ reduction may maintain a relatively large anammox population, and although the volume-specific rates of anammox are lower in the Skagerrak, its proportionate contribution to N₂ production is greater than that in the Thames, where the anammox population is probably comparatively smaller. Whether this means that NO₃⁻ reduction leaks proportionately more NO₂⁻ in the Skagerrak than in the Thames or that NO3⁻ respiration to NO2⁻, before anammox or denitrification, dominates the initial step of NO₃⁻ reduction is unclear.

In addition to the overall availability of NO_2^{-} - and NO_3^{-} regulating anammox, there were marked fluctuations in ana-

mmox activity (Table 1). Seasonal effects have been clearly documented for the common forms of sediment metabolism, which usually follow smooth curves reflecting seasonal temperatures (6, 21). Short-term changes in sulfate reduction have been reported for upper estuarine sediments which corresponded with peaks in *Desulfovibrio* species activity (20). Although all of the experiments were carried out at 15° C and, hence, do not reflect in situ temperatures, the differences may reflect variations in the in situ abundance and/or activity of the anammox community sampled on each occasion; this is supported by the overall pattern of the data set and model output.

In our anammox model, NO_2^- can be reduced via either anammox or denitrification, with denitrification switching between NO_2^- or NO_3^- depending on the availability of $NO_3^ ([NO_3^-]_{crit})$. Examples of the model output for anammox with the non-Michaelis (f_p [equation 7]) (Fig. 4a and b), Michaelis $(f_m \text{ [equation 8]})$ (Fig. 4c and d), and combined $(f_p + f_m)$ [equation 6]) (Fig. 4e and f) components, with θ varying from 0.003 to 0.045, are given in Fig. 4. When f_a includes only the function f_p , anammox can never respond in a linear fashion at concentrations higher than 10 μ M, since f_p is modeled with a peak in activity when $[NO_2^{-}] = l_N$ and declines exponentially after this. Alternatively, a Michaelis function, such as f_m , can never explain the nonlinear pattern shown at concentrations below 10 µM (Fig. 4c and d). It is only when the two functions are combined (as in equation 6) to give a biphasic behavior for anammox that the model begins to resemble the patterns observed in the data (compare Fig. 4e and f to Fig. 3c and d), in which the amount of anammox represents the sum of gas produced from both the Michaelis and non-Michaelis forms of the model. Whether or not there is a single anammox process with two enzyme systems or, alternatively, two independent anammox systems operating in these estuarine sediments is not known. The data suggest that in studies of anammox to date (19, 22), only the Michaelis form has been measured, which largely explains the previously reported linear production of $^{29}\mathrm{N}_2$ regardless of the concentration of $\mathrm{NO_2}^-$ or $\mathrm{NO_3}^-$ above 10 µM. The non-Michaelis form was certainly masked in our previous measurements of anammox (22), but it is not known whether this behavior would apply to off-shore continental shelf sediments (19). A Michaelis system may apply to anammox bacteria coupled tightly to the reduction of NO₃⁻, and as such, NO₂⁻ will never be limiting and the operational range for NO_2^{-} is less significant. The non-Michaelis system could be optimized to function at actual pore-water NO2⁻ concentrations, although the resolution of NO₂⁻ profiles and pore-water NO_2^{-} data to date are limited.

In addition to the biphasic expression, changing θ from 0.3 to 4.5% reproduced the range of anammox (amplitudes) measured in the slurry experiments (Fig. 4 e and f). Whether the effect of changing θ truly reflects a change in the in situ abundance (which may itself be a mixed population) or, alternatively, a change in specific activity is not known, as changing either in the model produces the same effect. Alternatively, the dampening down of anammox activity can be driven by increasing the rate of denitrification, yet this would seem least likely, as we consistently measured NO₃⁻ reduction at about 100 μ M h⁻¹ for these sediments, while the model would require an increase by a factor of about 8 to reproduce the minimum and maximum amplitudes shown in Fig. 4.

The production of ²⁹N₂ from labeled ¹⁵NH₄⁺ and ¹⁴NO₂⁻ agrees with the 1:1 catabolism of ¹⁴NO₂⁻ + ¹⁵NH₄⁺→²⁹N₂ + 2H₂O. Anammox bacteria (in bioreactors) have, however, been shown to consume NH₃ and NO₂⁻ with an overall stoichiometry of 1:1.3, with the excess NO₂⁻ (0.3 mol) being anaerobically oxidized to NO₃⁻, though this ratio can change with the concentration of NO₂⁻ (16, 18). Hence, 1 mol of ²⁹N₂ accounts for 1 mol of ¹⁵NH₄⁺ oxidized but potentially (depending on the actual sediment reaction) 1.3 mol of NO₂⁻ reduced, and therefore, the total anammox activity potentially needs to be multiplied by 1.3 to fully account for NO₂⁻ reduction via anammox. Allowing for this change in ratio, and depending on in situ conditions, anammox may account for the majority of NO₂⁻ reduction at 2 to 5 μ M NO₂⁻ in these sediments.

The presence of anammox violates the central tenets of the isotope pairing technique (10), and the implications of this where anammox and denitrification coexist in sediments have been explored (15). Their discussion pivots around four assumptions, one of which is that anammox and denitrification are both limited by the supply of $\mathrm{NO_3}^-$ and that the uptake kinetics for its reduction product, NO₂⁻, by denitrifying and anammox bacteria are similar. It is acknowledged that while the kinetics of denitrification as a function of NO_3^- are well characterized for sediments, little is known about the in situ regulation of anammox. The nonlinear response reported here for anammox below 10 μ M NO₂⁻, especially in the presence of NO₃⁻, suggests that the kinetics of the two processes are different and challenges this assumption for anammox in estuarine sediments. In the presence of representative in situ NO₂⁻ and NO₃⁻ concentrations, the production of N₂ via anammox increased. What regulates the proportionate significance of either anammox or denitrification to N_2 formation in intact aquatic sediments is likely to be a combination of the respective availability of both NO₃⁻ and NO₂⁻ and the relative size or activity of the anammox population.

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