

# Abundance, Activity, and Community Structure of Pelagic Methane-Oxidizing Bacteria in Temperate Lakes

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**The abundance and activity of methane-oxidizing bacteria (MOB) in the water column were investigated in three lakes with different contents of nutrients and humic substances. The abundance of MOB was determined by analysis of group-specific phospholipid fatty acids from type I and type II MOB, and in situ activity was measured with a <sup>14</sup>CH<sub>4</sub> transformation method. The fatty acid analyses indicated that type I MOB most similar to species of *Methylomonas*, *Methylomicrobium*, and *Methylosarcina* made a substantial contribution (up to 41%) to the total bacterial biomass, whereas fatty acids from type II MOB generally had very low concentrations. The MOB biomass and oxidation activity were positively correlated and were highest in the hypo- and metalimnion during summer stratification, whereas under ice during winter, maxima occurred close to the sediments. The methanotroph biomass-specific oxidation rate (*V*) ranged from 0.001 to 2.77 mg CH<sub>4</sub>-C mg<sup>-1</sup> C day<sup>-1</sup> and was positively correlated with methane concentration, suggesting that methane supply largely determined the activity and biomass distribution of MOB. Our results demonstrate that type I MOB often are a large component of pelagic bacterial communities in temperate lakes. They represent a potentially important pathway for reentry of carbon and energy into pelagic food webs that would otherwise be lost as evasion of CH<sub>4</sub>.**

Microbiological oxidation of methane takes place in many kinds of aquatic systems and soils (28, 32, 33, 58) and plays an important role in the global budget of this greenhouse gas (15, 43). It can proceed both under aerobic and anaerobic conditions, but the latter seems to be extensive primarily in waters with high ionic strength, e.g., saline lakes and marine systems (29, 31, 55). Many strains capable of aerobic methane oxidation have been isolated and characterized, and phylogenetically all belong to either the  $\gamma$ -proteobacteria (commonly referred to as “type I” methanotrophs) or the  $\alpha$ -proteobacteria (“type II” methanotrophs) (10, 12, 24).

Aerobic methane-oxidizing bacteria (MOB) in the water columns of lakes consume a significant part of the methane produced in the sediment or in anaerobic layers of water (46, 54). Besides metabolizing methane that would otherwise be emitted to the atmosphere, they represent a route for reentry of carbon back into the food-webs (2, 26, 30, 46). In spite of this, direct studies of MOB populations in the pelagic systems of lakes are rare and, with the exception of a small oxbow lake (45) and a tropical hydropower reservoir (20, 21), very little is known about the abundance and population structure of pelagic methanotrophs in lakes.

The aerobic methanotrophic bacteria contain some phospholipid fatty acids (PLFAs) that seem to be unique for this group of organisms (16:1 $\omega$ 8c in type I MOB and 18:1 $\omega$ 8c in type II MOB) (10, 40). Analysis of these PLFAs has been shown to be a very sensitive tool for the direct study of populations of

type I and type II MOB in samples from several different kinds of ecosystems (7, 9, 16, 39, 51).

In the present study we use PLFA analysis to characterize the populations of MOB in the water columns of lakes and estimate the relative contribution of MOB to total water column bacteria. The efficiency of the PLFA method to determine bacterial abundance was evaluated by comparison with direct cell counts using flow cytometry. By combining the measurements of abundance of MOB with their activity estimated with a <sup>14</sup>CH<sub>4</sub> transformation method, we were able to estimate the in situ biomass-specific activity of this particular group of bacteria. Samples were collected during summer and winter conditions in three temperate lakes differing in nutrient status and concentrations of dissolved organic matter.

## MATERIALS AND METHODS

**Study lakes.** Three lakes in south central Sweden were sampled: Mårn (64°95'96N, 15°04'15W), eutrophic with intermediate water color; Illersjön (64°94'96N, 14°52'66W), eutrophic with low water color; and Lillsjön (65°04'01N, 15°19'83W), oligotrophic with high water color (for details about water chemistry, see reference 4). All lakes were sampled twice in 1999 and 2000: once in summer during stratified conditions and once from ice in winter.

**Temperature, oxygen, and water sampling.** Profiles of temperature and oxygen concentrations were obtained with an oxygen probe (Orion model 35). Four to seven depths were sampled, ranging from 0.5 m below the surface to 0.5 m above the sediments. Water samples were pumped from each depth with a submersible pump (Amazon 10; Awimex International AB, Sweden).

**Analysis of PLFAs.** Duplicate subsamples of 40 to 250 ml were filtered using 47-mm 0.2- $\mu$ m-pore-size filters (polyvinylidene difluoride; Millipore) and frozen awaiting further analysis. The methods for extraction of total lipids, solid-phase extraction of the polar fraction, transformation of the fatty acids into fatty acid methyl esters (FAMES) by alkaline methanolysis, dimethyl disulfide derivatization of monounsaturated PLFAs, and identification and quantification of FAMES with GC-FID and GC-MS have been described previously (38, 52, 57). It was not possible to analyze all replicate filters for PLFAs. For three water

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samples where the duplicate filters were analyzed, the coefficient of variation for the total PLFA concentration varied from 11 to 23%.

**Biomass and abundance.** The total bacterial biomass and carbon content were determined from the concentrations of the PLFAs 14:0, i15:0, a15:0, 15:0, all 16:1, 16:0, i17:0, 17:0, 18:1 $\omega$ 7, 18:1 $\omega$ 8, and cy19:0, which are all of general bacterial origin (34, 60), assuming a total bacterial content of PLFA of 100  $\mu\text{mol g}^{-1}$  dry weight (dw) of cells and that carbon constituted 50% of the dw (59). The bacterial abundance (i.e., the cell numbers) could then be calculated from bacterial cell volumes in the lakes and the relation between the dw and cell volume for bacterial cells according to the method of Loferer-Krössbacher et al. (36) ( $\text{dw} = 435 \times V^{0.86}$  fg). The average bacterial cell volume in the water samples, estimated by image analyzed fluorescence microscopy (5), was 0.12  $\mu\text{m}^3$  for epi- and hypolimnion in Illersjön and Mårn, ranging from 0.08 to 0.13  $\mu\text{m}^3$ .

The biomass and cell numbers of MOB were calculated from the concentrations of 16:1 $\omega$ 8c and 18:1 $\omega$ 8c, considered to occur almost exclusively in type I and II MOB, respectively. The same conversion factors were used, and we assumed that 16:1 $\omega$ 8c and 18:1 $\omega$ 8c constituted 25 and 50% of the PLFAs in type I and type II MOB, respectively (10, 40, 61).

The total bacterial abundance based on the PLFA analysis was compared to an independent estimate of bacterial abundance, obtained by using flow cytometry (Becton Dickinson FACSCalibur; CellQuest 3.1 software) after staining with Syto 13 (Molecular Probes) according to the method of del Giorgio et al. (18).

**Methane concentrations.** For each depth sampled, four replicate 330-ml infusion bottles were completely filled with lake water, followed by preservation with NaOH raising pH to above 11 (47). Methane concentrations were analyzed by a headspace equilibration method (3).

**In situ methane oxidation.** The rates of methane oxidation in lake water were measured with a  $^{14}\text{CH}_4$  incorporation method as described in detail previously (3). In brief, water samples were amended with  $^{14}\text{CH}_4$  and incubated at actual depths for 4 to 8 h. The incubations were terminated by NaOH addition. In the lab, subsamples of 20 to 30 ml were filtered through 0.22- $\mu\text{m}$ -pore-size mixed cellulose ester filters (Millipore GSWP). The filtrates were purged with air to remove labeled methane, and the radioactivity of filters and filtrates was measured by liquid scintillation counting. The total radioactivity in these fractions represents the total oxidation of methane as the sum of incorporation into cell biomass, substances released as dissolved organic carbon, and dissolved inorganic carbon from respiration.

**Statistical analyses.** The relationships between bacterial cell numbers determined with the two independent methods (flow cytometry versus calculations based on PLFA concentrations) were investigated with single linear regression analysis. Both single and multiple linear regression analyses were used to evaluate relationships between methane oxidation and environmental parameters and were performed by using Kaleidagraph 3.6.2 (Synergy Software) and JMP 3.1.5 (SAS Institute, Inc.) software, respectively.

## RESULTS

In all, 23 different PLFAs could be identified and quantified, including *trans* isomers of monounsaturated fatty acids. The total PLFA concentration ranged from 3.8 to 56  $\text{pmol ml}^{-1}$ , with the highest concentrations toward the lake surface and the sediment, in most profiles.

The two independent methods to determine total bacterial abundance showed the same trends within the lakes, but the absolute abundances differed among the lakes (Fig. 1). Treatment of all data points in one single regression resulted in a strong linear relationship (Fig. 2a;  $r^2 = 0.53$ ,  $P < 0.0001$ ). Some individual PLFAs were also positively related to the bacterial abundance measured with flow cytometry, most strongly for 18:1 $\omega$ 7c, which gave an even better indication of bacterial abundance than the total concentration of bacterial PLFAs (Fig. 1 and 2b;  $r^2 = 0.57$ ,  $P < 0.0001$ ).

The PLFA 16:1 $\omega$ 8c, considered to be unique for type I MOB, occurred in most samples. The concentrations correlated very closely to those of 16:1 $\omega$ 6c and 16:1 $\omega$ 5t but were more weakly related to other general bacterial PLFAs (Table 1). This is strong evidence that the two latter fatty acids also come from type I methanotrophs in these lakes. In contrast,

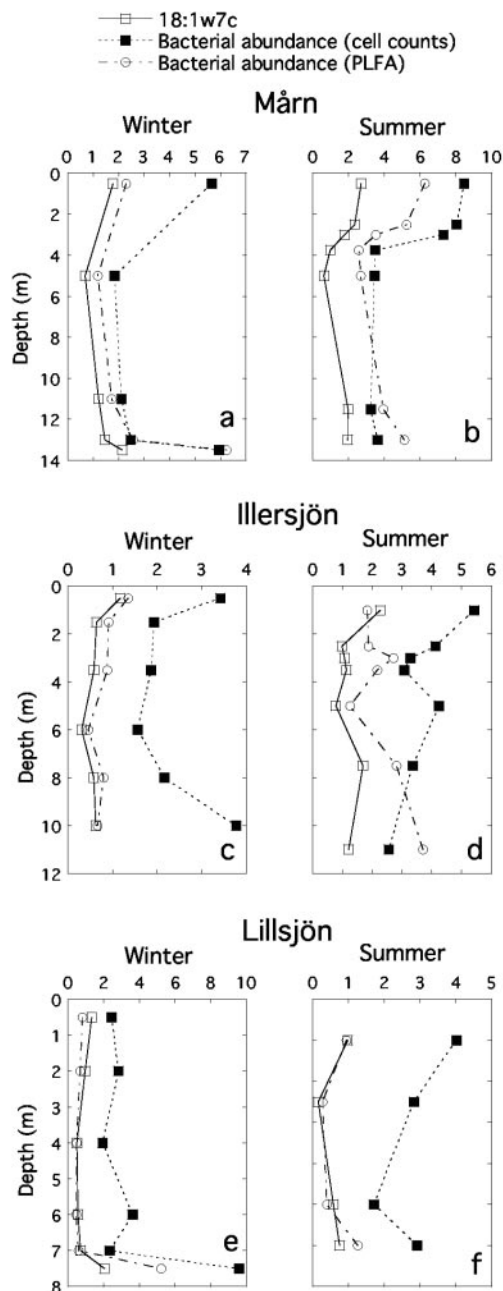


FIG. 1. Total bacterial abundance in lake profiles (millions of cells  $\text{ml}^{-1}$ ), as estimated by direct cell counts and from concentrations of bacterial phospholipid fatty acids. For comparison, concentrations of the common bacterial fatty acid 18:1 $\omega$ 7c ( $\text{pmol ml}^{-1}$ ) are also shown. The profiles are from lakes Mårn (a and b), Illersjön (c and d), and Lillsjön (e and f).

the type II MOB signature fatty acid 18:1 $\omega$ 8c was only detected in roughly half of the samples, with no obvious distributional pattern and always in low concentrations. In the seven samples showing highest oxidation (2.6 to 37.5  $\text{mg CH}_4\text{-C m}^{-3} \text{ day}^{-1}$ ), the biomass of type II MOB was only 0.0 to 4.8% (mean 2.2%) of that of type I. In 11 of the 20 samples where the type II signature 18:1 $\omega$ 8c was detected, the type II biomass was above

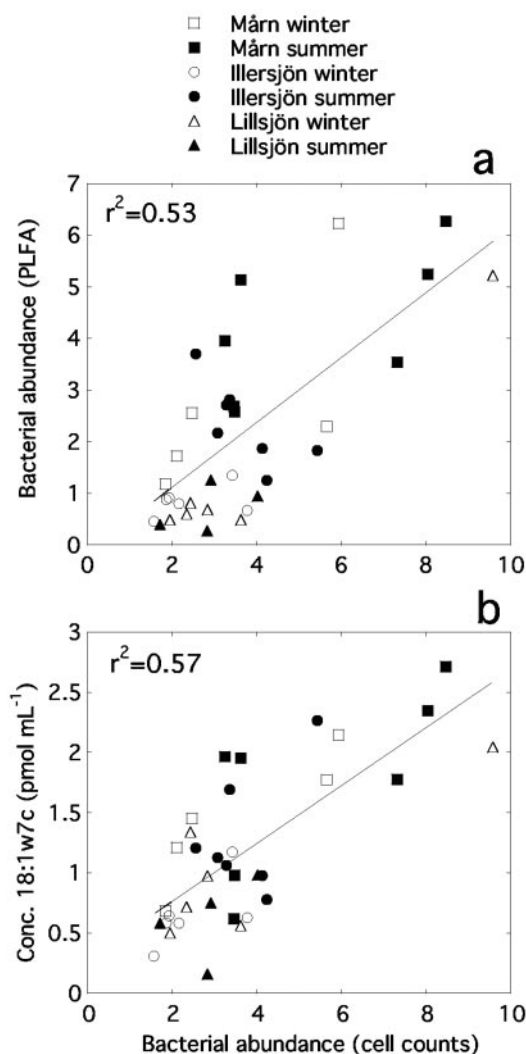


FIG. 2. Relationships between bacterial abundance estimated by direct cell counts and concentrations of phospholipid fatty acids in lake water. Units for bacterial abundance are in millions of cells per milliliter.

10% of the type I, but in all of these samples the oxidation rate was very low (below  $0.15 \text{ mg CH}_4\text{-C m}^{-3} \text{ day}^{-1}$ ).

Because the fatty acid composition implied that type I MOB strongly dominated the MOB communities, we estimated the total biomass of methanotrophs from the concentration of 16:1 $\omega$ 8c. In all individual samples, the biomass ranged from 0.0 to  $178 \text{ ng dw ml}^{-1}$ . By comparing the contribution of 16:1 $\omega$ 8c to the total concentration of bacterial fatty acids, we estimate that type I MOB account for up to 41% of the total bacterial biomass. During summer conditions the MOB biomass generally paralleled the  $\text{CH}_4$  consumption rate (exemplified with Illersjön in Fig. 3), being highest in the hypolimnion. During winter, the MOB abundance gradually increased toward the sediment, while the oxidation rate was very low, with the exception of the samples taken closest to the sediment (Fig. 3). Generally for all three lakes, the oxidation rate ranged from  $0.0004$  to  $37.5 \text{ mg CH}_4\text{-C m}^{-3} \text{ day}^{-1}$ . Six samples showing an especially high oxidation rate, all had low oxygen concentra-

TABLE 1. Correlation coefficients ( $r$ ) between concentrations of typical methanotrophic (16:1 $\omega$ 8c, 16:1 $\omega$ 6c, and 16:1 $\omega$ 5t) and general bacterial (16:0, 18:1 $\omega$ 7c, 16:1 $\omega$ 7c, and total bacterial) phospholipid fatty acids in lake water

Fatty acid	$r$ for fatty acid:					Total bacterial
	16:1 $\omega$ 6c	16:1 $\omega$ 5t	16:0	18:1 $\omega$ 7c	16:1 $\omega$ 7c	
16:1 $\omega$ 8c	0.98	0.98	0.24	0.38	0.82	0.55
16:1 $\omega$ 6c		0.99	0.19	0.35	0.80	0.52
16:1 $\omega$ 5t			0.20	0.34	0.81	0.52
16:0				0.79	0.66	0.92
18:1 $\omega$ 7c					0.74	0.85
16:1 $\omega$ 7c						0.88

tions ( $<25 \mu\text{M O}_2$ ), high methane concentrations (all except one above  $33 \mu\text{M CH}_4$ ), and low temperature ( $3.0$  to  $6.1^\circ\text{C}$ ). The general positive relationship between MOB abundance and total oxidation, estimated with independent methods, is illustrated by a single regression with all data (Fig. 4,  $y = 2.0 + 0.18x$ ;  $r^2 = 0.37$ ,  $P = 0.0003$ ;  $n = 34$ ).

By dividing the in situ methane oxidation with the MOB biomass we obtained estimates of the biomass-specific activity ( $V$ ) of this group of bacteria in lake water. The total range was  $0.001$  to  $2.77 \text{ mg CH}_4\text{-C mg}^{-1} \text{ C day}^{-1}$  (corresponding to  $0.03$  to  $80 \text{ nmol mg}^{-1} \text{ dw min}^{-1}$ ), with the highest value representing a daily consumption of methane carbon corresponding to more than twice the standing stock of methanotroph carbon biomass.  $V$  was positively correlated with  $\text{CH}_4$ -concentration (Fig. 5a,  $y = 0.14 + 0.017x$ ,  $r^2 = 0.47$ ,  $P < 0.0001$ ;  $n = 30$ ) but also with the total methane oxidation (Fig. 5b). There was no clear relationship between  $V$  and temperature, except that the four samples with outstandingly high  $V$  all had low in situ temperatures of  $4.0$  to  $6.1^\circ\text{C}$ . Similarly,  $V$  was not related to oxygen concentration, except that it was quite low in the high-activity samples. These relationships were confirmed by a multiple linear regression with  $V$  as a dependent variable and concentrations of methane and oxygen and water temperature as independent variables. This regression once again demonstrated that methane concentration accounts for a substantial share of the variance ( $r^2 = 0.49$ ,  $P = 0.0007$ ), whereas the addition of either oxygen concentration, temperature, or both as dependent variables increased the explained variance marginally but was not statistically significant ( $P \geq 0.27$ ).

## DISCUSSION

**Total bacterial abundance.** Estimates from direct counts of cells and from concentrations of PLFAs correlated strongly (Fig. 1 and 2). Our data thus demonstrate that PLFAs can be used to estimate bacterial standing stock in lake waters. In all three lakes, the two methods yielded similar patterns of depth distribution, although the ratio of cell numbers from PLFAs to flow cytometry differed slightly among lakes. However, it should be kept in mind that the absolute levels of bacterial biomass derived from the two methods depend on choice of conversion factors. These factors are rough averages that surely differ among bacterial groups and are dependent on environmental conditions. It should also be noted that, since data on bacterial cell sizes were not available for all individual

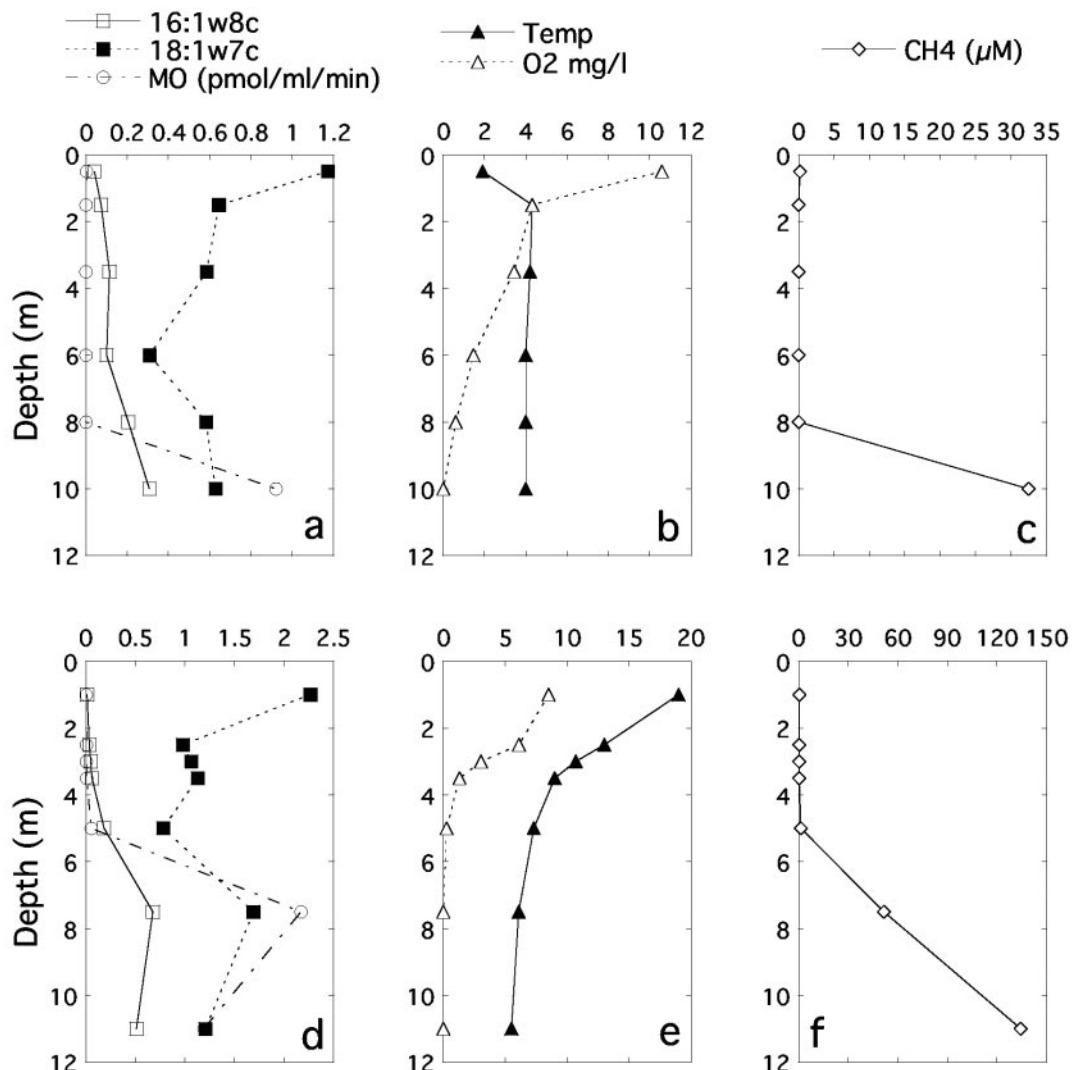


FIG. 3. Depth profiles of methanotrophic biomass (expressed as concentration of the signature PLFA 16:1 $\omega$ 8c, to facilitate comparison with 18:1 $\omega$ 7c), in situ methane oxidation rate, the concentration of the general bacterial fatty acid 18:1 $\omega$ 7c, and abiotic factors in Illersjön. (a to c) Winter conditions under ice; (d to f) stratified summer conditions. Note the different scales for winter and summer.

samples, we used a single average for all samples. Bacterial cell sizes may differ substantially depending on lake, depth, and sampling occasion (14). Such differences in cell size would affect the general ratio of abundances estimated with the fluorescence microscopy and PLFA methods and influence individual samples to a different degree as well. However, given the uncertainties inherent in the conversion factors, the general correlation between total bacterial abundance estimates from flow cytometry and PLFA concentrations still appears robust (Fig. 2).

**Biomass and activity of methanotrophic bacteria.** The strong covariation of 16:1 $\omega$ 8c with both 16:1 $\omega$ 5t and 16:1 $\omega$ 6c supports that all three come from the same group of bacteria. Whereas the former is considered specific to MOB, the latter two are often present simultaneously with 16:1 $\omega$ 8c in methanotrophs. Therefore, they can be used as “qualifying” fatty acids (10, 40, 61) for type I MOB, but they do occur in some other bacteria as well. The

simultaneous presence and the similarities in concentration ratios for these three PLFAs between the lake waters and known isolates of MOB suggest that the methanotrophs in these lakes are most similar to strains of *Methylomonas*, *Methylomicrobium*, and *Methylosarcina* (10, 12, 61).

In all samples with significant methane oxidation activity, the ratio between the concentrations of the  $\omega$ 8c fatty acids from type I and type II methanotrophs suggest that the biomasses of type I MOB were substantially higher than those of type II MOB. Moreover, across all samples the biomass of type II MOB seemed to be unrelated to the oxidation activity. Still, the relevance of type II methanotrophs for methane oxidation cannot be completely ruled out, since it was recently shown that acidophilic type II MOB (genera *Methylocella* and *Methylocapsa*) from peatlands lack 18:1 $\omega$ 8c (17).

There are few other studies with which the dominance of type I methanotrophs in the water columns of the lakes that we



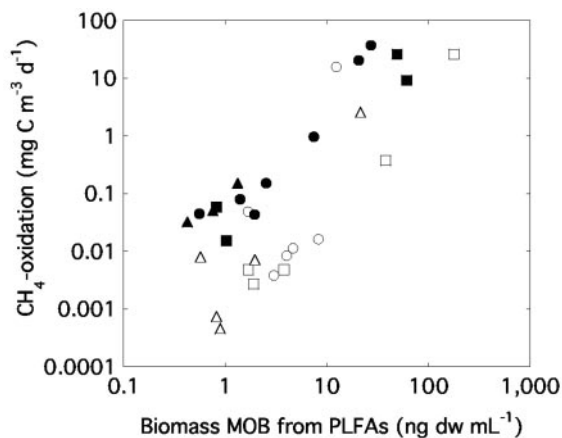


FIG. 4. Relationship between the biomass of methanotrophic bacteria estimated with phospholipid fatty acid analysis and the in situ methane oxidation activity measured with a  $^{14}\text{CH}_4$  transformation method (logarithmic scales). Symbols are as described in Fig. 2.

studied can be compared. However, in line with our results, two studies of MOB populations in pelagic freshwater systems (23, 45) and two in sediments (16, 41) all suggest predominance of type I over type II methanotrophs in lakes, although they represent very different lake ecosystems. However, it cannot be concluded that type I dominance is universal in freshwater bodies, since Dumestre et al. (20, 21), using PLFA analyses and a PCR/DGGE approach, found that type II MOB completely dominated methanotroph communities in a tropical water reservoir. Nevertheless, the available information from freshwater systems suggest that either main group of MOB firmly dominates. This seems to be in contrast to wetland soils, where it has been repeatedly found that the two groups often have similar population sizes (27, 37, 51, 56).

The total range of the biomass-specific methane oxidation rates in these lakes,  $V$ , was 0.03 to 80  $\text{nmol mg}^{-1} \text{dw min}^{-1}$ . These values fall below or in the lower range of published data for  $V_{\text{max}}$  in laboratory cultures of type I MOB, with a reported range of roughly 50 up to 600  $\text{nmol mg}^{-1} \text{dw min}^{-1}$  (13, 25, 49, 50). This literature range is based on MOB growing under nonlimiting nutrient conditions in the laboratory. However, taking, for example, the low temperatures of the samples exhibiting the highest  $V$  in our study lakes into account, our estimates of  $V$  generally appear as quite realistic for type I methanotrophic bacteria. We are not aware of any previous estimates of the biomass-specific  $V$  for methane oxidation under natural conditions.

In terms of carbon transformation,  $V$  ranged from 0.001 to 2.77  $\text{mg CH}_4\text{-C mg}^{-1} \text{C day}^{-1}$ . It was lowest in the epilimnion in summer and at intermediate depths under ice in winter, at methane concentrations below 0.1  $\mu\text{M}$ . The strong positive correlation of  $V$  with methane concentration, but much weaker with temperature and  $\text{O}_2$ -concentration (Fig. 5 and the multiple regression studies), suggests that methane concentration has the overriding influence on the activity and biomass distribution of methanotrophic bacteria in these three lakes, similar to the situation reported by Liikanen et al. (35).

The peaks in activity and biomass of MOB occurred at very low  $\text{O}_2$  concentrations, for several deep water samples below

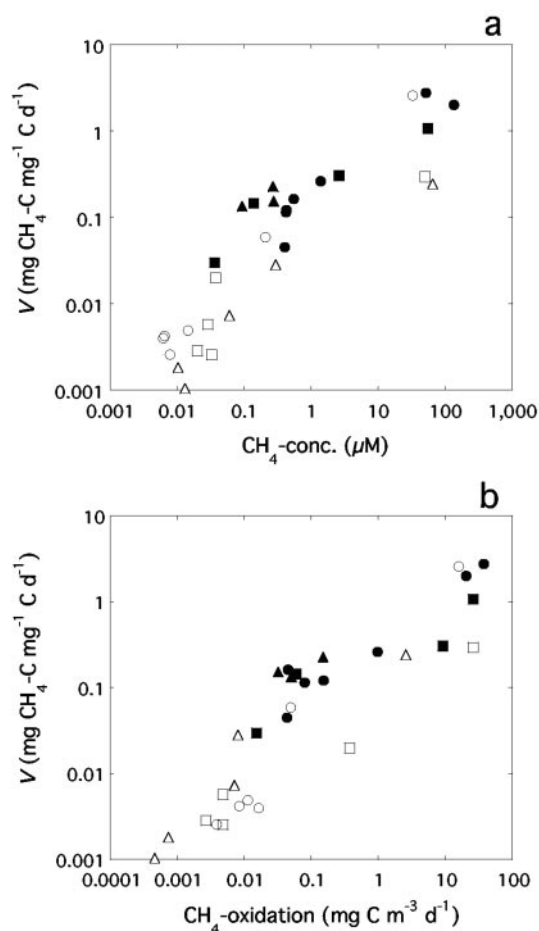


FIG. 5. Relationships of methane concentration (a) and absolute methane oxidation rate (b) in lake water with  $V$ , the methanotrophic biomass-specific methane oxidation activity (logarithmic scales). Symbols are as described in Fig. 2.

5.7  $\mu\text{M}$ , representing the detection limit. Some other studies have also reported methane oxidation in lake water or sediments in the absence of  $\text{O}_2$ , or at least at very low concentrations (1, 35, 42, 54, 62). Anaerobic oxidation of methane coupled to reduction of sulfate is well-known from various marine sediments (6, 19, 55) and saline lakes (29, 31). In order to get a rough idea whether the low  $\text{O}_2$  concentrations were sufficient to support the methane oxidation rates, we calculated turnover times for  $\text{O}_2$  based on the oxygen demand indicated by the measured methane oxidation rates. In the cases where  $\text{O}_2$  was at or below the detection limit, 5.7  $\mu\text{M}$  was used. The standing stock of oxygen had to be replaced within 1.9 to 4.8 days (with one extreme of 34 days) in order to support aerobic methane oxidation. These turnover times are substantially longer than the incubation times in our  $^{14}\text{CH}_4$  uptake assay (4 to 8 h), supporting that methane oxidation was aerobic, provided that the affinity to  $\text{O}_2$  of the MOB was at least comparable to the affinities in the nonmethanotrophic heterotrophic bacteria, which has been demonstrated previously (44).

The largest MOB biomasses and activity in the lakes always occurred at a temperature below 7°C, suggesting that psychrotrophic or psychrophilic methanotrophs may have been in-

volved. Several type I methanotrophs with a growth optimum at 15°C or lower have been isolated (11, 32, 53). In addition, in studies in a soil and a biofilter on landfills (8, 22) low temperatures favored the development of type I methanotroph populations over type II populations, whereas at 20°C or higher, either both groups or preferably type II populations grew. Together with our data, these studies suggest that low temperature selects for dominance of type I methanotrophic bacteria under several different kinds of environmental conditions.

In conclusion, our results demonstrate that MOB may account for a substantial fraction of water column bacterial biomass and carbon transformations. Thus, through grazing on bacterial communities by zooplankton that are able to tolerate O<sub>2</sub> concentrations close to anoxia (48), they may constitute an important link for reentry of carbon into aquatic food webs.

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