

Lack of Correlation between Turnover of Low-Molecular-Weight Dissolved Organic Carbon and Differences in Microbial Community Composition or Growth across a Soil pH Gradient^{∇†}

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We studied how soil pH (pHs 4 to 8) influenced the mineralization of low-molecular-weight (LMW)-dissolved organic carbon (DOC) compounds, and how this compared with differences in microbial community structure. The mineralization of LMW-DOC compounds was not systematically connected to differences in soil pH, consistent with soil respiration. In contrast, the microbial community compositions differed dramatically. This suggests that microbial community composition data will be of limited use in improving the predictive power of soil C models.

The primary connection between soil microbial activity and carbon (C) cycling is well established (29, 46, 51); however, there has recently been a call for a more detailed understanding and integration of the microbial community into models for soil organic C (SOC) cycling to improve their predictive power (5, 6, 10, 11, 20, 29, 35, 44, 47). To accomplish this, we need to study the drivers of respiration and composition of microbial communities together (36).

Low-molecular-weight (LMW)-C compounds dominate the C used in soil respiration (13, 18, 22, 31, 48, 49, 50). Although observations regarding factors that affect LMW-DOC cycling have started to emerge (13, 23, 31), the canonical environmental factors influencing their mineralization remain unclear. This study's aims were to test how one of the most influential factors for the composition of the soil microbial community, soil pH (4, 8, 26, 32, 39, 40, 41, 42), influenced the mineralization of LMW-C compounds across a wide pH gradient and to compare our findings with microbial community composition.

Soil was obtained from the Hoosfield pH gradient at Rothamsted Research, United Kingdom (1, 2, 3, 7, 39). Thirty topsoil samples (0- to 23-cm depth) were tested along the gradient in March 2010, sieved (<2 mm), and characterized (39). The 30 individual soil samples were used for microbial and chemical analyses (e.g., organic C, total N, pH, phospholipid fatty acid [PLFA] composition, and bacterial growth), while the gradient was split into four pH levels for the soil solution analysis and C substrate mineralization assays. For these, independent replicates ($n = 3$) were used for each of the four pH levels (pH 4.1 ± 0.04 , pH 5.0 ± 0.07 , pH 6.0 ± 0.05 ,

and pH 7.1 ± 0.08). Soil solution was extracted by centrifugal drainage (19, 43; see also Supplement S1 in the supplemental material), and the free amino acid and sugar concentrations of the solutions were determined (21, 30).

Soil (5 g) from each pH level was weighed into polypropylene tubes. Soil solution from each replicate (450 μ l) was then individually spiked with one of eight different ¹⁴C-labeled substrates (50 μ l) at a trace level and added to the soil and the mineralization monitored using 1 M NaOH CO₂ traps at 22°C for 7 days (see Supplement S1 in the supplemental material). The ¹⁴CO₂ level in the NaOH traps was determined by liquid scintillation. Sorption of the added ¹⁴C LMW-DOC compounds across the pH gradient was determined in sterilized soil samples (see Supplement S1). Bacterial growth was estimated using leucine incorporation (9, 25), and microbial community structure was determined from PLFA patterns (16, 17, 32).

¹⁴C substrate mineralization and half-lives ($t_{1/2}$) were modeled by fitting a double-first-order decay equation to the experimental results (13, 23, 33, 48, 49; see also Supplement S1 in the supplemental material). The PLFA composition (mol% of the 26 most abundant PLFAs) was analyzed by principal-component analysis (PCA) after values were standardized to unit variance. Analysis of variance (ANOVA) with Tukey's honestly significant difference test *post hoc* comparisons were used to determine differences in variables with soil pH, while regression analyses were used to describe relationships with pH across all 30 samples. Results were compared to those for an identical sampling of the pH gradient from 2008 (40). In addition, a 454-pyrosequencing-based analysis of the bacterial community composition was also previously performed on the same 2008 samples (42; based on analyses from references 14, 27, 28, and 34). A type II major-axis regression analysis was used to investigate the connection between the sequence composition (the principal coordinate of variation of the sequence composition [see Supplement S1]) of the bacterial community with microbial PLFA composition.

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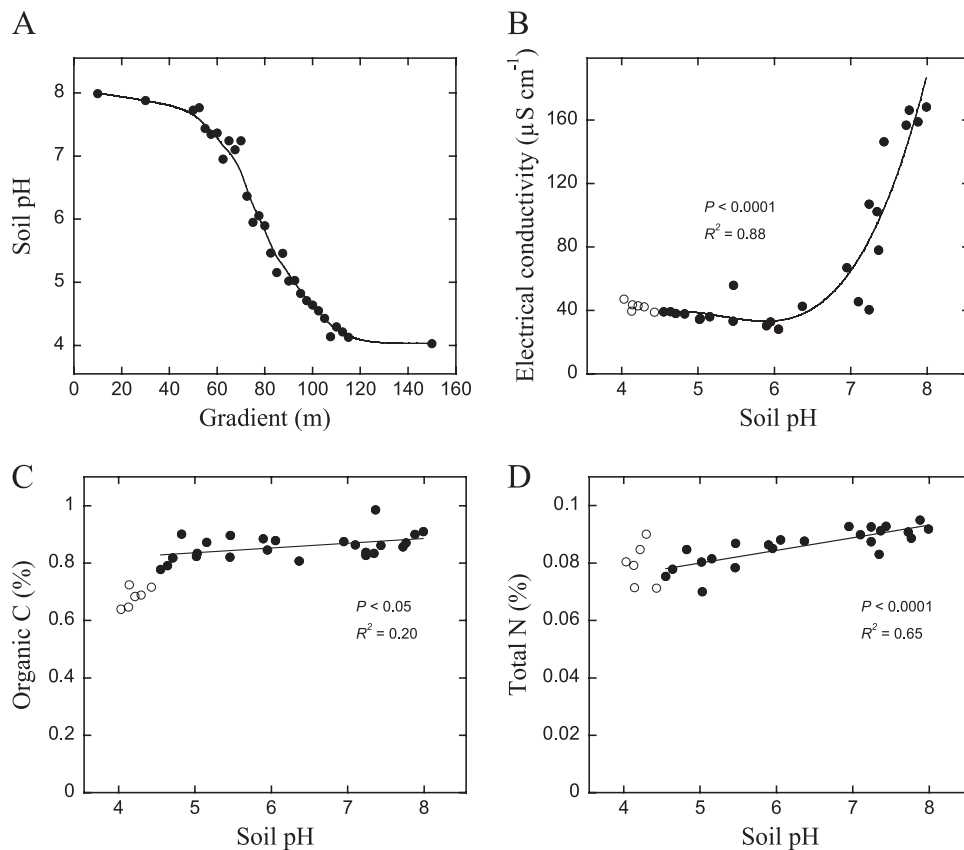


FIG. 1. Soil pH along the Hoosfield acid strip (A), the effect of soil pH on electrical conductivity (B), the organic C (C), and total N (D). Data points below pH 4.5 (open symbols) were not used in the regression analyses (see Materials and Methods).

Soil measurements showed a smooth pH gradient between pH 4 and 8 with only small differences in most other chemical variables (39) (Fig. 1). DOC and dissolved organic nitrogen (DON) concentrations did not systematically change over the pH gradient (see Table S1 in the supplemental material). The DOC and DON concentrations derived from both sugars and amino acids were similar over the gradient ($P > 0.05$) (Table S1), averaging 53.5 ± 7.1 nmol of C derived from sugar (sugar C) g^{-1} soil and 11.4 ± 2.1 nmol amino acid C g^{-1} soil. The mineralization patterns conformed well to a double-first-order kinetic model ($R^2 > 0.99$ for all substrates) (Fig. 2 and 3). The individual amino acids and sugars degraded at different rates (Fig. 2); however, this did not systematically change with pH (see Table S2 in the supplemental material).

Combining our estimates of free sugar and amino acid concentrations with estimates of their turnover rate enabled estimation of the contribution of LMW-C mineralization to soil respiration. We estimate that approximately 0.32 and 0.09 μg of C derived from CO_2 (CO_2 C) $\text{g}^{-1} \text{h}^{-1}$ of soil respiration were derived from sugar and amino acid degradation, respectively. Taken together, this is within the range of previous assessments of the basal respiration rate at the site (0.20 to 0.50 $\mu\text{g CO}_2 \text{ C h}^{-1} \text{ g}^{-1}$) (3, 39, 41).

Bacterial growth increased about 6-fold between pH 4.5 and 8.0 ($P < 0.0001$) (see Fig. S1 in the supplemental material). The first component from a PCA of the microbial PLFA composition explained 40% of the variation in the data and was

closely related to soil pH ($P < 0.0001$, $R^2 = 0.96$) (see Fig. S2A in the supplemental material). The variation of microbial PLFA composition across the Hoosfield site was highly reproducible, with the major component of variation from a PCA aligning identically with soil pH (Fig. 4A). There was also a very close relationship between the variation in the sequence composition of the bacterial community (42) (Fig. 4B).

The Hoosfield site effectively isolated soil pH from most other variables associated with microbial community differences. In addition, it was previously shown that excess Al (or other unidentified inhibitors) or a lack of P or N did not limit bacterial and fungal growth across the pH gradient (41). Soluble LMW-C compound concentrations were also not systematically affected by pH. Similarly, rates of turnover of individual LMW-C compounds did not change consistently with pH. In contrast, the microbial community composition radically differed along the pH gradient. Direct comparison between the microbial PLFA composition and its sequence composition (obtained by 454 pyrosequencing [42]) (Fig. 4B) indicated that the difference in microbial PLFA composition across the gradient was a consequence of a difference in bacterial species composition. The enormous pH-related differences in microbial community composition did not affect the mineralization of LMW C, the dominant source of soil respiration, across the pH gradient. Therefore, we could not find support for a connection between microbial community structure and function (the turnover of the LMW-DOC compounds that govern soil

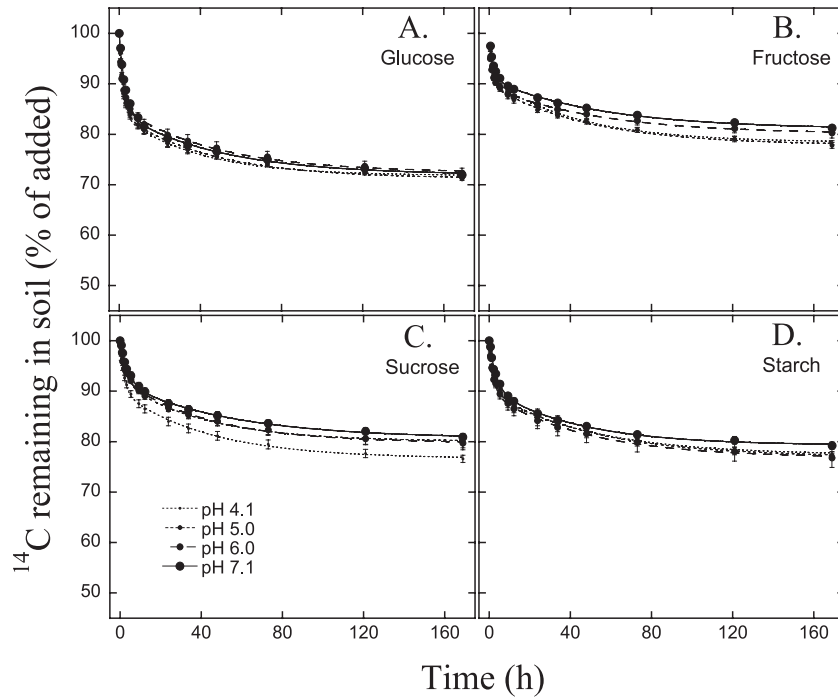


FIG. 2. Mineralization of sugars. Amounts of ¹⁴C-labeled glucose (A), fructose (B), sucrose (C), and starch (D) remaining in soils of the four different pH levels after the injection of the isotopically labeled soil solution into the soil. Values represent means ± 1 SE (*n* = 3). The lines represent the best fits of double-first-order exponential-decay functions ($R^2 > 0.99$ for all curve fits).

respiration). This is consistent with conjectures derived from theoretical models of soil organic matter (SOM) turnover that have suggested high redundancy in the microbial processing of LMW C (12, 24, 45). We cannot discount the possibility that

processes transforming SOM to LMW DOC are closely related to microbial community structure. However, there was only a minor discrepancy between basal respiration (30% difference across the gradient) and LMW-DOC mineralization across the

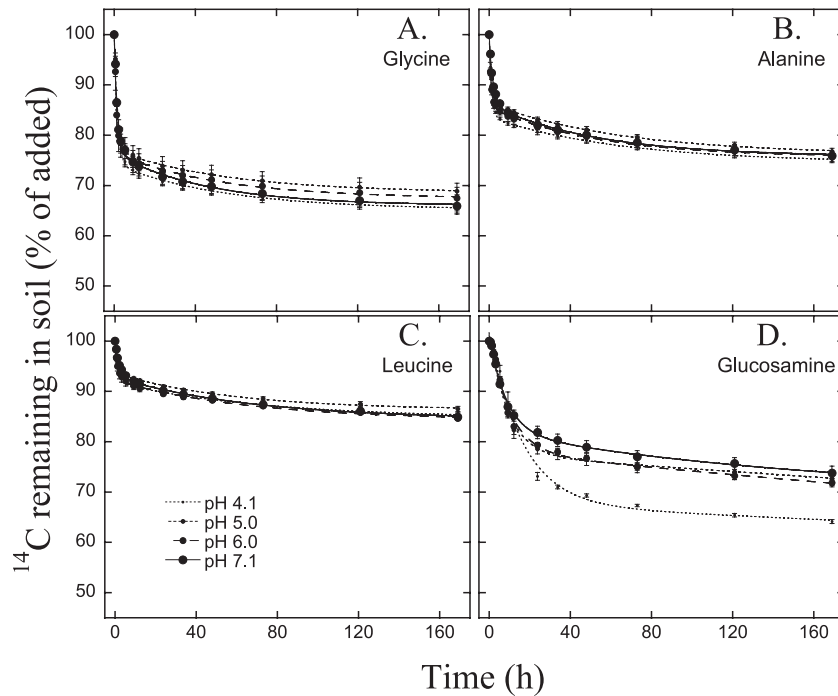


FIG. 3. Mineralization of amino acids and amino sugars. Amounts of ¹⁴C-labeled glycine (A), alanine (B), leucine (C), and glucosamine (D) remaining in soils of the four different pH levels after the injection of the isotopically labeled soil solution into the soil. Values represent means ± 1 SE (*n* = 3). The lines represent the best fits of double-first-order exponential-decay functions ($R^2 > 0.99$ for all curve fits).

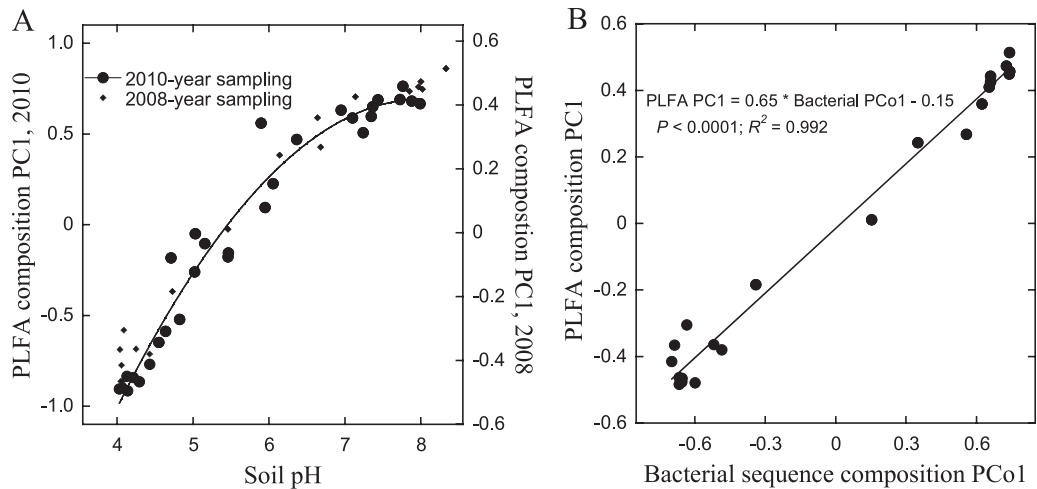


FIG. 4. (A) Comparison of the levels of pH dependence of the microbial PLFA compositions across the Hoosfield acid strips in the soil samples of the present study compared to those sampled in 2008 (data are from reference 40); (B) first principal component (PC1) of the variation in the PLFA compositions (explaining 39.6% of the variation in PLFA composition) across the Hoosfield acid strip from the 2008 sampling (data are from reference 40) compared with the first component of a principal-coordinate (PCo1) analysis (explaining 29.1% of the total sequence variation) of the compositions of bacterial sequences obtained with a 454-pyrosequencing assessment (data are from reference 42). A type II major-axis regression was used to fit the line.

gradient (no difference across the gradient), setting the upper limit for the functional relevance of the microbial community difference. This does not lessen the need to identify and incorporate the rate-limiting mechanisms in models for SOC turnover (6, 29), but it does suggest that adding information about the microbial community composition may be of limited use.

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