

# Denitrification and Nitrogen Fixation Dynamics in the Area Surrounding an Individual Ghost Shrimp (*Neotrypaea californiensis*) Burrow System

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**Bioturbated sediments are thought of as areas of increased denitrification or fixed-nitrogen (N) loss; however, recent studies have suggested that not all N may be lost from these environments, with some N returning to the system via microbial dinitrogen (N<sub>2</sub>) fixation. We investigated denitrification and N<sub>2</sub> fixation in an intertidal lagoon (Catalina Harbor, CA), an environment characterized by bioturbation by thalassinidean shrimp (*Neotrypaea californiensis*). Field studies were combined with detailed measurements of denitrification and N<sub>2</sub> fixation surrounding a single ghost shrimp burrow system in a narrow aquarium (15 cm by 20 cm by 5 cm). Simultaneous measurements of both activities were performed on samples taken within a 1.5-cm grid for a two-dimensional illustration of their intensity and distribution. These findings were then compared with rate measurements performed on bulk environmental sediment samples collected from the lagoon. Results for the aquarium indicated that both denitrification and N<sub>2</sub> fixation have a patchy distribution surrounding the burrow, with no clear correlation to each other, sediment depth, or distance from the burrow. Field denitrification rates were, on average, lower in a bioturbated region than in a seemingly nonbioturbated region; however, replicates showed very high variability. A comparison of denitrification field results with previously reported N<sub>2</sub> fixation rates from the same lagoon showed that in the nonbioturbated region, depth-integrated (10 cm) denitrification rates were higher than integrated N<sub>2</sub> fixation rates (~9 to 50 times). In contrast, in the bioturbated sediments, depending on the year and bioturbation intensity, some (~6.2%) to all of the N lost via denitrification might be accounted for via N<sub>2</sub> fixation.**

Nitrogen (N) is an essential element contributing to the biological processes of all organisms (16), yet biologically available (fixed) N is often limiting for marine productivity (17, 78). Sediments play a key role in the global N cycle, yet our understanding of the benthic microbial processes involved in the N cycle, especially in regard to the sinks and sources of fixed N (16, 17), is incomplete and often debated. In coastal marine sediments, N is lost through the production of dinitrogen gas (N<sub>2</sub>), typically through the microbial processes of denitrification or through the conversion of ammonium and nitrite to N<sub>2</sub> (anaerobic ammonium oxidation [anammox]) (26, 36). In sediments impacted by bioturbation (the biological reworking and ventilation of sediments by infauna [see reference 49 for a comprehensive overview of the current definition of bioturbation]), the rates of coupled nitrification-denitrification are believed to be increased (28), resulting in an even greater loss of N from sediments. It has been known that N<sub>2</sub> fixation, the input of N into the system, plays a role in specific benthic environments, primarily within the photic zone (photosynthetic microbial mats, coral reef sediments, and sediments vegetated by sea grasses and marsh plants) (13, 14, 17, 34, 35). Only recently was it shown that subsurface sediments in bioturbated areas can represent habitats of significant N<sub>2</sub> fixation by sulfate-reducing bacteria (SRB) (5, 6), further demonstrating that our current understanding of N cycling in benthic environments is limited.

Animal-sediment-microbe interactions and the environmental factors controlling these interactions are of great interest to benthic ecologists. In particular, the study of bioturbation has gained momentum in many areas of science (e.g., geology, sedimentology, and paleontology) because burrowing organisms may affect the majority of all surfaces on Earth (54), potentially >20,700 km<sup>3</sup> of sediments worldwide (76), and over geological

time scales (10, 65, 77). Bioturbating organisms can be regarded as “ecosystem engineers,” organisms that significantly modify the environment in such a way as to greatly change the availability of resources for other organisms (39, 54). Macrofaunal burrows are characterized by a species-specific architecture (32, 48) and vary substantially in size and permanence (4, 24, 32, 58, 84), making it challenging to evaluate the impact of bioturbation (the movement of particles) or bioirrigation (the movement of fluids) on the geochemistry and microbial ecology of inhabited sediments. The construction and maintenance of burrows increase the oxic-anoxic interface, which may enhance microbial activity and result in increased cell abundances (1, 58, 59). Processes within the N cycle are particularly linked to the oxic-anoxic interface. Ammonium, which is released through the remineralization of organic matter, is reoxidized to nitrate in the oxic zone (nitrification). The denitrification of nitrate, in the absence of oxygen, to nitrite and further to N<sub>2</sub> is considered to be the major loss of N from the benthic system. Bioturbation, by enhancing coupled nitrification-denitrification activities, is thought to further increase the loss of N<sub>2</sub>. In a previous study (5), we showed that subsurface N<sub>2</sub> fixation by SRB can be an important source of N in bioturbated sediments. The

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question is whether the input of fixed N balances the loss of N via denitrification in bioturbated systems.

In this study, we used field and laboratory measurements to examine the influence of bioturbation on denitrification and  $N_2$  fixation in coastal sediments bioturbated by the bay ghost shrimp *Neotrypaea californiensis* Dana 1854 (Crustacea: Decapoda: Thalassinidea). *N. californiensis* inhabits intertidal areas along the west coast of North America from Alaska to Baja California (51) and belongs to a group of decapods, *Thalassinidea*, of which there are more than 550 known taxa worldwide (25). Thalassinidean burrow morphology is species specific (23, 24, 31, 73), with some thalassinids being capable of creating 3-m-deep burrows (60). The burrow structure relates to the feeding mode (i.e., deposit feeders, drift catchers, or filter or suspension feeders) of the thalassinid (32, 73), but environmental and biological conditions, such as grain size or population density, can also influence burrow morphology (31). *N. californiensis* is known to construct a highly branching and deep-reaching (~80-cm) burrow system (51, 75) and spends the majority of its life subsurface, constantly reworking and maintaining its burrow while deposit feeding (12, 51). We investigated the denitrification and  $N_2$  fixation activity associated with burrow systems created by *N. californiensis* in an intertidal lagoon on Santa Catalina Island off the coast of Los Angeles, CA. The overall goal of these detailed and high-resolution measurements was to quantify N inputs and losses in a bioturbated system to evaluate the role of coastal sediments as sinks or sources of fixed N.

## MATERIALS AND METHODS

**Study site.** Investigations were carried out in an intertidal lagoon located in Catalina Harbor, Catalina Island, CA (33°25.23'N, 118°19.42'W), in July and August 2009. The lagoon is a shallow (<2-m) area consisting of muddy sand (with the majority of grains being <500  $\mu\text{m}$ ). Tides at this location are mixed, with the higher high water preceding the lower low water and a range of ~1.7 m (19). During sampling, the water temperature varied little and was typically in the range of 18°C to 20°C, and salinity was 34.5‰.

Two intertidal sampling areas, one with *N. californiensis* bioturbation (~500 burrow openings  $\text{m}^{-2}$  seafloor) and one without (0 burrow openings  $\text{m}^{-2}$  seafloor), were chosen for detailed investigations. The same areas (within a few meters) had previously been investigated to determine  $N_2$  fixation rates in June 2007 and May 2008 (5). The burrow density at each sampling location was determined by counting the number of burrow openings within a 25-cm by 25-cm frame, with 10 replicates counted. Note that the nonbioturbated area received higher levels of subsurface organic carbon input from the root systems of a surrounding marsh area and was characterized by slightly coarser sediment. At the bioturbated location, *N. californiensis* burrows reached ~20 cm deep into the sediment (4). Typically, each burrow system has multiple branches and 3 to 4 openings to the sediment surface. The burrows consist of shafts (~1-cm diameter) and chambers (~2-cm diameter) that the shrimp maintains and frequently flushes with oxygen-rich water.

**Sampling and sediment characteristics.** A set of three parallel sediment push cores (diameter of 5.4 cm; 30 cm long) were collected during high tide from each sampling location. Each core was collected randomly, with no specific orientation toward *N. californiensis* burrow openings. Two cores from each set were sliced in 2-cm intervals down to a depth of 20 cm under an  $N_2$  atmosphere, and each section was subsampled for further geochemical processing, as follows. Pore water was collected from each interval of one core by centrifugation (10 min at 3,500 rpm) using 50-ml Macrosep centrifugal cell concentrators (Pall Corporation, Life Sciences). Pore water samples (~3 ml) were immediately frozen at -20°C for later determinations of fixed N concentrations: ammonium by flow

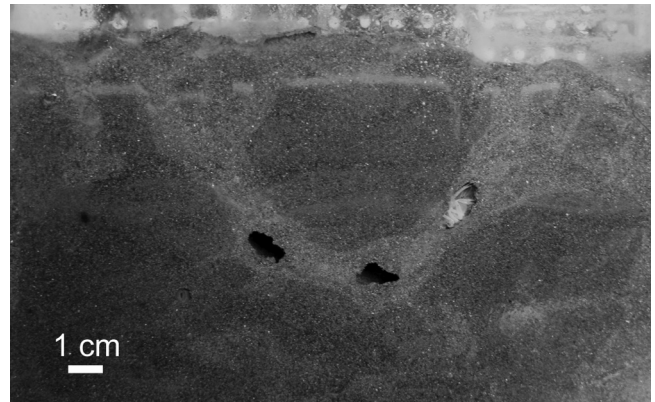


FIG 1 Photograph of the narrow aquarium and burrow system, with the bioturbating ghost shrimp visible.

injection analysis modified for small sample volumes (33) and nitrate by reduction to nitrite with spongy cadmium followed by spectrophotometry (40). On the second core, the total organic carbon (TOC), calculated as the loss on ignition (LOI), was determined for each 2-cm section by drying a known volume of sediment at 65°C for 24 h and then combusting the sample at 450°C for 24 h.

**Catalina Harbor denitrification.** The third sediment core from each location was sliced in 2-cm intervals and analyzed for denitrification rates by use of an acetylene ( $C_2H_2$ ) inhibition method in which  $C_2H_2$  blocks the transformation of  $N_2O$  to  $N_2$  in the denitrification pathway (70), causing an accumulation of  $N_2O$ , which can be measured by gas chromatography (20) or by using  $N_2O$  microsensors (7, 63). Two known potential drawbacks of the inhibition method are that  $C_2H_2$  inhibition may be incomplete (66), especially when hydrogen sulfide is present (47), and that  $C_2H_2$  may inhibit nitrification, causing a decrease in levels of  $NO_3^-$  over time (20, 37). To help alleviate this second problem, it was suggested previously that  $NO_3^-$  be added to incubation mixtures (44). Both potential drawbacks would lead to an underestimation rather than an overestimation.

In this study, triplicate 5-cm<sup>3</sup> samples from each location and each depth were placed into 9-ml serum vials that were flushed with  $N_2$  and contained 400  $\mu\text{l}$  of 110  $\mu\text{M}$  potassium nitrate. This addition of potassium nitrate led to a final concentration of ~20  $\mu\text{M}$   $NO_3^-$  in the samples, which is half the highest  $NO_3^-$  concentration seen previously at this study site (4). Prior to the start of incubation, initial  $N_2O$  concentrations in each vial were determined by using a  $N_2O$  microsensor that was inserted into the sediment (Unisense, Aarhus, Denmark). Microsensor signals were amplified and transformed into millivolts by a 2-channel picoammeter (PA 2000; Unisense) and directly recorded on a computer by using Profix software (Unisense). Each vial was then closed with a butyl stopper, crimp sealed, and injected with 1 ml of  $C_2H_2$ . Vials were kept in the dark and at the *in situ* temperature (20°C). After approximately 6 h, an analysis of  $N_2O$  with microsensors was performed again as an endpoint measurement. Additionally, over the course of the experiment, another set of vials was prepared similarly, using sediment from the 3-cm-depth horizon of the bioturbated sampling location. The sediment was thoroughly mixed, and 36 5-cm<sup>3</sup> samples were placed into 9-ml serum vials that were flushed with  $N_2$  and contained 400  $\mu\text{l}$  of 110  $\mu\text{M}$  potassium nitrate. Each vial was closed with a butyl stopper, crimp sealed, injected with 1 ml of  $C_2H_2$ , and kept in the dark and at the *in situ* temperature. Every hour, over the course of 36 h, the  $N_2O$  in one vial was measured by using the  $N_2O$  microsensor to ensure that  $N_2O$  production remained linear over this time frame.

**Laboratory mesocosm studies.** A detailed examination of an *N. californiensis* burrow was performed by using a narrow aquarium (Fig. 1) that was kept in a wet laboratory at the Wrigley Institute for Environmental Studies, located on Santa Catalina Island. The aquarium was 15 cm tall by 20 cm wide by 5 cm deep, providing enough space for the ghost shrimp to

establish a burrow system. The aquarium was filled with sediment collected from the study site that was sieved through a 500- $\mu\text{m}$  sieve to remove large sediment particles and sediment-dwelling macrofauna. One shrimp was added to the aquarium after 24 h, when the sediment had sufficiently settled. The narrow aquarium was placed into a larger container so that the sediment surface was continually flushed with seawater pumped to the laboratory from an offshore location. Before it reached the laboratory, the seawater passed through a rough sieve to remove larger particles. The shrimp was not fed in order to keep the laboratory setup as similar to natural conditions as possible (i.e., to not alter the amount of organic carbon available). Because the shrimp feeds on suspended matter that is supplied in the ventilation current, as well as on sediment microbiota, the shrimp had access to its natural nutrition sources. The entire setup was left for 1 month, allowing enough time to establish a burrow system. Once a burrow was established, transparencies were placed against the aquarium so that the visible burrow structures could be traced and exact sampling locations could be selected and recorded. Photographs of the aquarium were also taken with a Nikon model D40 digital camera (Fig. 1). The front wall of the aquarium was perforated with holes in a 1-cm grid that were filled with aquarium silicone. The silicone-filled holes served as ports for pore water sampling, especially for targeting specific burrow structures. This front wall was also easily removable so that samples could be collected from the exposed sediment in selected areas.

Pore water (3 ml) was drawn from 6 different burrow compartments for the determination of ammonium (33) and nitrate (40) concentrations within the burrow system. Vertical microprofiles of the redox potential (Unisense, Aarhus, Denmark) were performed at 1-mm increments from the overlying water into the sediment and in 2-cm horizontal intervals across the width of the narrow aquarium. The sensor was attached to a motorized micromanipulator (Märzhäuser, Wetzlar, Germany) controlled by a computer and connected to a high-impedance mV meter (WTW-pH 340; WTW GmbH, Weilheim, Germany).

**Mesocosm microbial rate measurements.** To map variations in denitrification and  $\text{N}_2$  fixation (measured as nitrogenase activity [NA]) around an individual shrimp burrow, sediment samples were taken from the aquarium in a grid with a 1.5-cm spatial resolution. Because denitrification and NA measurements both use acetylene ( $\text{C}_2\text{H}_2$ ), we were able to perform both analyses on the same individual samples, as was done in previous studies (e.g., see reference 44). For sediment sampling, the aquarium was laid on a slight incline so that all sediment and pore water would remain undisturbed while the front wall of the aquarium was removed. A center portion of the aquarium containing visible burrow structures was selected for sampling (13.5 cm wide and 9.5 cm high), avoiding edge effects. The sediment was cored by using 5-ml plastic syringes with the tip cut off following a 1.5-cm resolution grid. Single 5-cm $^{-3}$  samples were placed into 9-ml serum vials that were flushed with  $\text{N}_2$  and contained 200  $\mu\text{l}$  of 220  $\mu\text{M}$  potassium nitrate, again resulting in a final concentration of  $\sim 20 \mu\text{M}$   $\text{NO}_3^-$  in the samples. A higher concentration of potassium nitrate was used to ensure that denitrification measurements could run longer. Prior to the start of incubation, initial measurements of the  $\text{N}_2\text{O}$  level in each vial were made by using the  $\text{N}_2\text{O}$  microsensor (Unisense, Aarhus, Denmark). Each vial was then crimp sealed and injected with 1 ml of  $\text{C}_2\text{H}_2$ . Vials were kept in the dark, at an *in situ* temperature, and were incubated over a 36-h time period.

NA measurements of the aquarium samples were obtained by using a  $\text{C}_2\text{H}_2$  reduction assay (15).  $\text{C}_2\text{H}_2$  is a substrate analog of  $\text{N}_2$  gas and is preferentially reduced by the conventional (molybdenum-based) nitrogenase enzyme to ethylene ( $\text{C}_2\text{H}_4$ ), which is easily quantified by gas chromatography using a flame ionization detector (FID). It is important to note that  $\text{C}_2\text{H}_2$  has been found to inhibit a number of physiological groups, including some nitrifiers, methylotrophs, and methanogens, and caution is advised for interpreting data for environments where those groups are important (57).

In this experiment,  $\text{C}_2\text{H}_4$  was assayed over a 36-h incubation period (while the NA was linear) with a Shimadzu Mini-2 gas chromatograph

TABLE 1 Comparison of NAs between slurry and non-slurry incubation mixtures

Amt of water added (ml)	Mean NA (nmol $\text{C}_2\text{H}_4$ $\text{cm}^{-3} \text{h}^{-1}$ ) $\pm$ SE	% decrease in NA
0	0.48 $\pm$ 0.15	
0.5	0.17 $\pm$ 0.07	64
2	0.16 $\pm$ 0.03	66
4	0.20 $\pm$ 0.07	58

(GC) fitted with an FID and a 6-ft Alltech column packed with Hayesep A. Data were taken from a total of 5 time points, including sampling at time zero. At each time point, 100  $\mu\text{l}$  of headspace gas from each vial was sampled and directly injected into the GC.  $\text{N}_2$  fixation rates were calculated from the NA results by using a conversion factor of 3 moles of  $\text{C}_2\text{H}_4$  to 1 mole of  $\text{N}_2$ . After the  $\text{C}_2\text{H}_4$  measurements were finished, analysis of  $\text{N}_2\text{O}$  with microsensors was performed again as an endpoint for denitrification. Both results of the determination of rates were then plotted on a two-dimensional (2-D) contour graph. The location of the burrow (as traced on transparencies) was overlaid onto these graphs so that direct comparisons between the measured rates in relation to where the sample was taken could be made.

**Testing of the effect of the addition of seawater and potassium nitrate on  $\text{C}_2\text{H}_2$  reduction.** Perturbation experiments are generally performed by the addition of the test compound diluted in seawater (e.g., in this experiment,  $\text{NO}_3^-$  was added). To test the effect of the addition of the seawater itself on the reduction of the  $\text{C}_2\text{H}_2$  concentration, a set of vials was prepared by using sediment from the 0- to 2-cm-depth horizon of the selected bioturbated location from our field studies. The sediment was thoroughly mixed, and 4 sets of triplicate 4-cm $^3$  samples were placed into 9-ml serum vials that were flushed with  $\text{N}_2$ . Each set was then subjected to a different volume of filtered anoxic seawater: 0, 0.5, 2, and 4 ml. Each vial was crimp sealed and injected with 1 ml of  $\text{C}_2\text{H}_2$ , and the increase in the level of  $\text{C}_2\text{H}_4$  was assayed over a 16-h period with a GC. Over this time period, data from a total of 4 time points were taken, including sampling at time zero. Incubation mixtures were kept in the dark and at the *in situ* temperature (20°C). Afterwards, the percent decrease due to the addition of water was calculated by using the 0-ml addition as the starting point. These calculations accounted for the differences in volume in each vial as well as the partitioning of  $\text{C}_2\text{H}_4$  between the water and gas phases. Similarly, to test the effect of the addition of potassium nitrate on the measurements of the reduction of  $\text{C}_2\text{H}_2$  levels, another set of vials was prepared by using sediment from the 0- to 5-, 5- to 10-, and 10- to 15-cm-depth horizons from the same bioturbated sampling location. Samples were collected down to 15 cm so as to match the depth of our aquarium setup. The sediment from each depth was thoroughly mixed, and 4-cm $^3$  samples were placed into 9-ml serum vials that were flushed with  $\text{N}_2$ . For each depth, 2 sets of 5 replicates were prepared. Four hundred microliters of 110  $\mu\text{M}$  potassium nitrate in filtered anoxic seawater was added to one set, and 400  $\mu\text{l}$  of filtered anoxic seawater was added to the other set. Each vial was then crimp sealed and injected with 1 ml of  $\text{C}_2\text{H}_2$ , and the increase in the  $\text{C}_2\text{H}_4$  level was assayed over a 24-h period by gas chromatography. Over this time period, samples were taken at a total of 3 time points, including the sampling at time zero. Incubation mixtures were kept in the dark and at the *in situ* temperature (20°C). Statistical comparisons of the NA between samples after the addition of potassium nitrate and samples after the addition of seawater from the same sediment depth horizon were performed by using a one-tailed *t* test. The results of these methodological tests showed that the addition of anoxic filtered seawater to the sediment slurry caused the NA to decrease by  $\sim 63\%$  on average, regardless of the amount of seawater added (Table 1). In contrast, the addition of nitrate to slurry incubation mixtures did not significantly increase or decrease NA measurements compared to samples with only seawater added (Table 2).

TABLE 2 Effect of the addition of nitrate on NA

Sampling location depth (cm)	% decrease in NA	P value <sup>a</sup>
0–5	5.7	0.53
5–10	8.4	0.31
10–15	–9.2	0.07

<sup>a</sup> Determined by a two-tailed *t* test.

## RESULTS

**Environmental parameters of Catalina Harbor cores.** Ammonium ( $\text{NH}_4^+$ ) concentrations in the top 2 cm from both locations (the bioturbated area, with  $\sim 500$  burrow openings  $\text{m}^{-2}$ , and the nonbioturbated area, with 0 burrow openings  $\text{m}^{-2}$ ) were characterized by very similar values of  $\sim 10 \mu\text{M}$  (Fig. 2). Below this depth, both locations showed increased  $\text{NH}_4^+$  concentrations down to a 3-cm depth, and below this depth, the concentrations remained stable at  $\sim 40 \mu\text{M}$  for the bioturbated location and  $\sim 50 \mu\text{M}$  for the nonbioturbated location. Nitrate ( $\text{NO}_3^-$ ) concentrations, although varying with depth, showed no clear pattern and were below  $8 \mu\text{M}$  for both locations (Fig. 2). The loss on ignition (LOI) at both sites typically ranged from 1 to 2% sediment dry weight (Fig. 2). The nonbioturbated location showed a large peak in the LOI (3.6% dry weight) at an 11-cm depth, where an oily tar-like substance and an abundance of plant roots were noted during core slicing.

### $\text{N}_2\text{O}$ production in Catalina Harbor sediments over time.

Over the course of 36 h,  $\text{N}_2\text{O}$  production (Fig. 3) remained linear ( $R^2 = 0.96$ ), indicating that our Catalina Harbor and mesocosm denitrification rates were not compromised by the incubation time. These results support data reported previously by Jenkins and Kemp (38), who found that  $^{15}\text{N}_2$  production in intact estuarine sediment cores remained linear over a 50-h time span.

**Catalina Harbor denitrification activity.** Denitrification rates for both locations were highly variable with depth, ranging from  $18.0 \pm 0.5$  to  $139.9 \pm 0.8 \text{ nmol N cm}^{-3} \text{ day}^{-1}$  for the nonbioturbated area and from  $9.1 \pm 0.5$  to  $99.9 \pm 17.9 \text{ nmol N cm}^{-3} \text{ day}^{-1}$  for the bioturbated area (Fig. 4). On average, the bioturbated location showed more variability in denitrification rates at each depth, with the exception of the 3- and 13-cm depths. The two depths that showed the greatest difference were the 11- and 13-cm depths, where the nonbioturbated rates were much higher

than the bioturbated rates (Fig. 4). These depths coincided with the large observed peak in the organic carbon concentration (or LOI).

**Mesocosm ammonium and nitrate concentrations.** The  $\text{NH}_4^+$  and  $\text{NO}_3^-$  concentrations within the aquarium did not vary greatly between locations in the burrow and in the overlying water (Table 3). The deepest sampling location (a burrow compartment at a 9-cm depth) had no detectable  $\text{NO}_3^-$  and the highest  $\text{NH}_4^+$  concentrations ( $8.27 \pm 0.24 \mu\text{M}$ ). However, this  $\text{NH}_4^+$  concentration was lower than what was measured in the field at the same depth ( $\sim 45 \mu\text{M}$ ), suggesting that the shrimp periodically ventilates its burrow, keeping the burrow at far lower  $\text{NH}_4^+$  concentrations than those in the surrounding sediment. Vertical profiles of redox potential (mV) measurements depicted as a contour plot (Fig. 5) showed an oxidized zone (positive redox potential) at the sediment surface. With depth, the redox potential decreased and eventually became negative (indicating a reduced zone) near part of the burrow at around a 5-cm depth.

**Mesocosm microbial rates.** Denitrification rates surrounding the burrow (Fig. 5) showed a patchwork distribution, with no clear trend in relation to distance from the burrow or with sediment depth. Integrated denitrification rates at each sampling row ranged from 71 to  $118 \text{ nmol N cm}^{-2} \text{ day}^{-1}$ .  $\text{N}_2$  fixation rates also showed a patchwork distribution, with no statistically significant relation to distance from the burrow or with sediment depth, but many of the areas displaying the highest  $\text{N}_2$  fixation rates appeared to occur along the underside of the burrow. Integrated  $\text{N}_2$  fixation rates at each sampling row ranged from 3.6 to  $4.6 \text{ nmol C}_2\text{H}_4 \text{ cm}^{-2} \text{ day}^{-1}$  (or 2.4 to  $3.1 \text{ nmol N cm}^{-2} \text{ day}^{-1}$ ). Looking at the individual sampling locations, the ratio of denitrification to  $\text{N}_2$  fixation (Fig. 5) ranged from 0.73 to 240.9, with the highest value seen for the samples from the area located beneath the burrow at a 6-cm sediment depth. Similar to the rate values themselves, the ratio showed no statistically significant distribution in relation to distance from the burrow or with sediment depth. Additionally, a comparison of denitrification versus  $\text{N}_2$  fixation at each location showed no correlation (Fig. 6). For a comparison of the 70 locations where both rates were measured, the denitrification-to- $\text{N}_2$  fixation ratios were  $< 10$  for 5 locations and  $< 50$  for 38 niches, and a ratio of  $> 100$  was measured for only 9 locations. However, both denitrification and  $\text{N}_2$  fixation were measured in the same vials, which contained  $200 \mu\text{l}$  of  $220 \mu\text{M}$  potassium nitrate. While the addition of nitrate did not show an effect on measurements of the

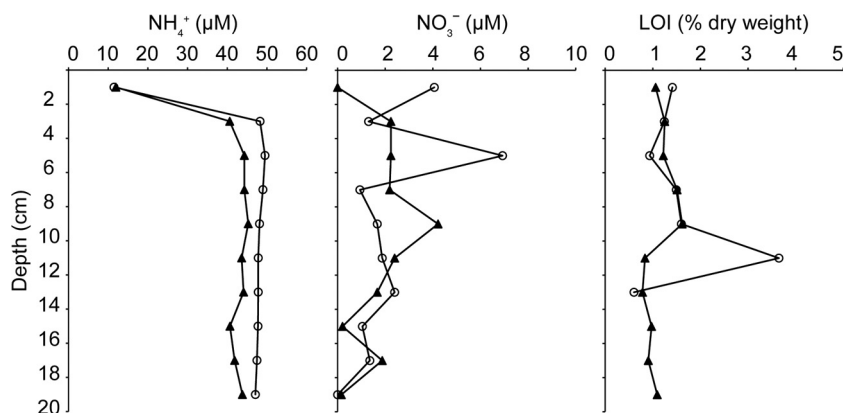


FIG 2 Depth profiles of ammonium ( $\text{NH}_4^+$ ), nitrate ( $\text{NO}_3^-$ ), and total organic carbon calculated by the loss on ignition (LOI) in sediment cores collected from a bioturbated area (O) and a nonbioturbated area (▲).

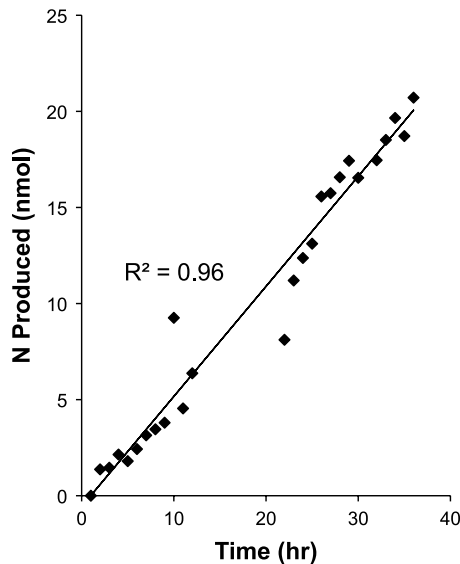


FIG 3 Nitrogen production in serum vials measured by an  $\text{N}_2\text{O}$  microsensor. These measurements were performed to test the linearity of denitrification over a 36-h time period.

NA, the addition of water showed an adverse effect on the NA, averaging a 63% reduction in measured rates. When the  $\text{N}_2$  fixation results were corrected according to this percentage, the ratios of denitrification to  $\text{N}_2$  fixation shifted to 0.46 to 151.7, with 9 locations showing a ratio of  $<10$ , 56 locations showing a ratio of  $<50$ , and only 1 location exhibiting a ratio of  $>100$ .

## DISCUSSION

**N in benthic environments.** Marine sediments play an extremely important, if complicated and not-well-understood, role in oceanic N cycling. It has been estimated that the remineralization of organic matter in benthic environments can supply up to 80% of the N required by phytoplankton in coastal waters (9, 11, 21). When bioturbation or bioirrigation is present, the flux of dissolved N from the sediment into the water column can greatly

TABLE 3 Ammonium and nitrate concentrations of selected burrow compartments

Location	Mean $\text{NH}_4^+$ concn ( $\mu\text{M}$ ) $\pm$ SE	$\text{NO}_3^-$ concn ( $\mu\text{M}$ )
Overlying water	$4.15 \pm 0.13$	1.35
Burrow opening	$3.95 \pm 0.17$	1.41
Burrow opening	$4.41 \pm 0.00$	1.82
4-cm-deep burrow compartment	$4.15 \pm 0.17$	1.20
8-cm-deep burrow compartment	$5.92 \pm 0.35$	2.19
9-cm-deep burrow compartment	$8.27 \pm 0.24$	0.00

increase. A recent study showed that 76% of the benthic efflux of  $\text{NH}_4^+$  within a shallow bay was due to bioirrigation (22). In some areas of the world ocean, according to isotope mass balance, up to 75% of the N lost from the marine system, via the conversion of fixed N into  $\text{N}_2$  by denitrification, occurs in benthic environments (18, 68). It was suggested previously that this loss of N may indirectly regulate primary production in the coastal ecosystem (46, 56). Conversely, climate-induced decreases in primary production may cause heterotrophic marine sediments to switch from being a net sink, via denitrification, to being a net source of N, via  $\text{N}_2$  fixation, which is explained by a decreased export of organic matter to the seafloor under these conditions and, possibly, N limitation (27). However, most estimates of benthic N processes are based on overall fluxes of N species across the sediment-water interface, with few studies directly exploring the N cycling occurring subsurface within different microniches that are important in bioturbated sediments (4).

Marine sediments are generally characterized by geochemical gradients created by the consumption of electron acceptors coupled to the oxidation of organic carbon. Oxygen, the most energetically favorable electron acceptor, is rapidly consumed at the sediment-water interface, followed by other electron acceptors in the order of decreasing energy gain (3, 42). As macrofauna move through the sediment, particles are rapidly transported between oxic and anoxic conditions, causing redox oscillations that disrupt the local electron acceptor succession with depth (2, 45). Those redox oscillations produced by bioturbation may result in a faster

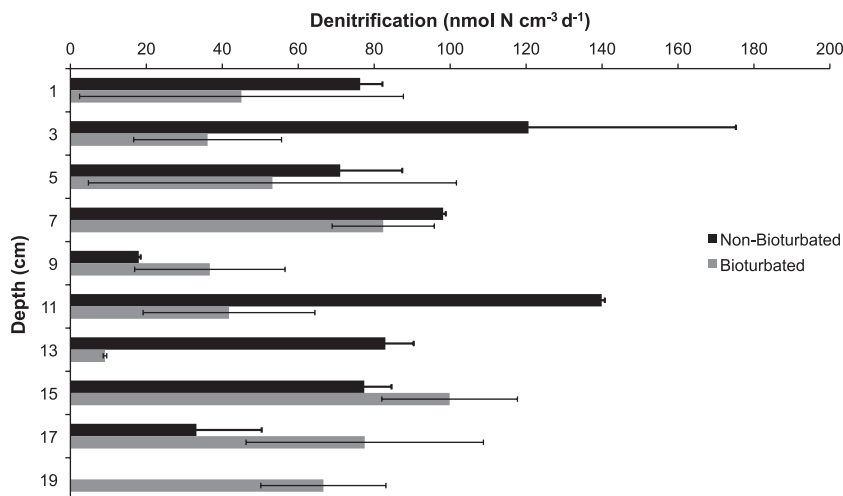


FIG 4 Denitrification measurements in sediment cores collected from a bioturbated area (light) and a nonbioturbated area (dark). Error bars represent standard errors ( $n = 3$ ).

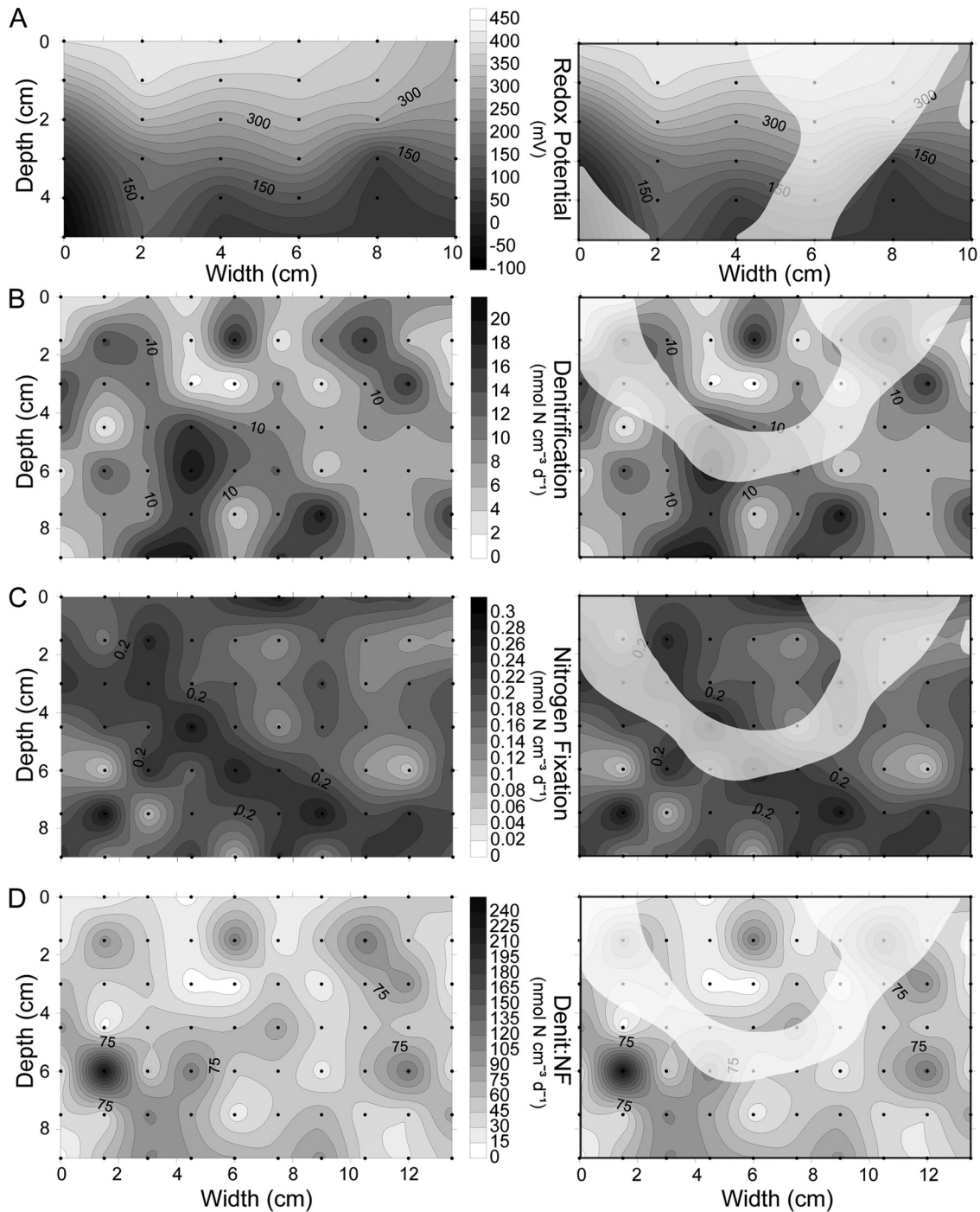


FIG 5 Contour plots of redox potential (A), denitrification (B), N<sub>2</sub> fixation (C), and ratio of denitrification to N<sub>2</sub> fixation (denit:Nf) (D) measured within an aquarium inhabited with a single ghost shrimp. The figures on the right show those portions of the burrow (in white) that were visible at the aquarium wall at the time of sampling, and dots indicate where sampling was performed.

and more complete decomposition of organic matter (74), as well as producing unique microbial communities surrounding the bioturbated area (52, 53, 62). The formation of this unique microbial community structure can have major impacts on the local benthic N cycle through enhanced microbial remineralization processes (see reference 50 and references therein). Previous studies have shown that bioturbation can greatly increase coupled ni-

trification-denitrification rates through the extension of the sediment-water interface, an increase in solute transport, and the oxygenation of sediments (28, 29, 34). Other studies have shown that bioturbation can significantly increase subsurface benthic N<sub>2</sub> fixation rates, linked to sulfate reduction (5, 6), showing that N cycling in bioturbated sediments may be more complex than previously thought.

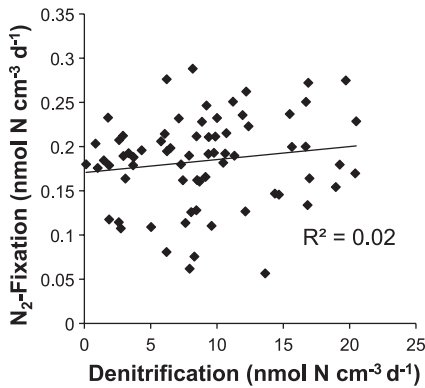


FIG 6 Denitrification versus  $N_2$  fixation from aquarium sediment samples that were incubated in the same serum vials.

**Bioturbation and organic carbon availability.** In Catalina Harbor sediments, denitrification appears to be more closely related to organic carbon availability than to bioturbation activity alone. The highest average denitrification rate occurred in the nonbioturbated location at an 11-cm sediment depth (Fig. 4), where a spike in the organic carbon concentration (LOI) was observed (Fig. 4). This result is not surprising, as organic matter is known to enhance denitrification. Additionally, van Luijn et al. (79) reported previously that denitrification rates in muddy versus sandy sediments were greatly influenced by the level of fresh organic matter. Bioturbation itself can increase the organic matter content of sediments, and this has been shown to increase sulfate reduction activity in the vicinity of burrows (5, 6, 30). The product of sulfate reduction (hydrogen sulfide) can, in turn, have inhibitory effects on nitrification and thus potentially limit the loss of N through coupled nitrification-denitrification (43). However, in iron-rich sediments, such as those in Catalina Harbor, hydrogen sulfide is rapidly precipitated as iron oxides, and although sulfate reduction rates are high, hydrogen sulfide concentrations in the pore water remain very low (at or below the detection limit of hydrogen sulfide microsensors). More importantly, SRB have the genetic capability to fix  $N_2$  (82), which was supported by data from previously reported culture and field studies (e.g., see references 55, 69, and 71), and have been shown to be responsible for the majority of  $N_2$  fixation in the bioturbated sediments of our study site (5). We speculate that the localized increase in the amount of organic matter enhanced  $N_2$  fixation by SRB, while the resulting sulfide inhibited nitrification and a loss of N through subsequent denitrification. In contrast, in the nonbioturbated location, where sulfate reduction rates have been shown to be lower than those found in a bioturbated area, supporting a bioturbation intensity similar to that seen in this study (5), the majority of the organic matter may be supporting increased denitrification rates.

**Impacts of microniche formation on benthic N cycling.** The bioturbated location of Catalina Harbor typically showed extremely variable denitrification rates, indicating high levels of sediment heterogeneity and the possible presence of denitrifying microniches. This high spatial variability was also seen in the measurements of denitrification and  $N_2$  fixation rates in the narrow aquarium (Fig. 5). Interestingly, a correlation between denitrification,  $N_2$  fixation, sediment depth, and distance from the burrow could not be seen (Fig. 6). Because denitrification is an anaerobic process, one would expect that denitrification rates

would increase with distance from the burrow as the sediment becomes more reduced. However, denitrification has been shown to occur under aerobic conditions within intact fecal pellets that form an anaerobic organic-rich microniche within the sediment (64). In theory, the diameter of these fecal pellets would need to be  $>2$  mm to maintain a reduced microniche within an oxygen-saturated environment (41). While thalassinids often deposit their fecal pellets on the sediment surface, pellets can sometimes be pushed into the burrow wall as a substrate (8). Thalassinids are known to produce rod-shaped fecal pellets that can range from 0.75 to 4 mm in diameter and 2 to 10 mm in length (61, 67), meaning that thalassinid fecal pellets could be large enough to produce an anaerobic microniche where denitrification could occur. These reduced microniches in *N. californiensis*-bioturbated sediments can also support increased sulfate reduction rates (6) and, through  $N_2$  fixation by SRB, may cause an increase in  $N_2$  fixation instead of denitrification, which may be inhibited by the increased sulfide concentrations (43) within the microniche.

Beyond fecal pellets, in particular, thalassinidean shrimp are known to secrete a mucus (polysaccharides), which, when mixed with sediment particles, is used to stabilize burrow walls (32), enriching burrow lining in organic carbon ( $C_{org}$ ) and possibly creating organic-rich microniches. Furthermore, microniches can also form around decaying organisms (83), algal aggregates (81), other forms of organic matter (such as marine snow) that are deposited on the sediment surface and transported deeper into the sediment via bioturbation (72), and organic particles in seawater that are carried into the burrow via bioirrigation. As the sediment is bioturbated, these microniche-forming organic aggregates can be shifted, causing a complex and patchy distribution of microniches throughout the sediment column that can be seen by using whole-core resin embedding (80). This patchiness in organic aggregates may be the reason why there appeared to be no correlation between denitrification,  $N_2$  fixation, sediment depth, and distance from the burrow (Fig. 6) within the experimental setup. These findings support the notion that N cycling within bioturbated sediments is extremely complex and most likely driven by the formation of microniches surrounding the burrow system.

**Denitrification versus  $N_2$  fixation in Catalina Harbor sediments.** The above-mentioned  $N_2$  fixation study of these same Catalina Harbor sediments (5) showed that summer (May through June) benthic  $N_2$  fixation rates, integrated down to 10 cm, can vary from 0.25 to 8.05  $mmol\ N\ m^{-2}\ day^{-1}$ , depending on the bioturbation intensity and the presence of a microbial mat at the surface. In contrast, it was shown that an area without bioturbation supported integrated  $N_2$  fixation rates of only 0.15 to 0.80  $mmol\ N\ m^{-2}\ day^{-1}$ . The integration of the denitrification rates found in this study down to a sediment depth of 10 cm revealed areal denitrification rates (denitrification rates per surface area) of 7.68  $mmol\ N\ m^{-2}\ day^{-1}$  for the nonbioturbated location and 5.07  $mmol\ N\ m^{-2}\ day^{-1}$  for the bioturbated location (Fig. 4). Comparing these rates with the rates previously reported for Catalina Harbor  $N_2$  fixation (5), the denitrification rate at the nonbioturbated location was  $\sim 9$  to 50 times higher than  $N_2$  fixation rates, suggesting a possible loss of fixed N. For the bioturbated location, the rate of denitrification was  $\sim 20$  times higher than the lowest rate reported previously (5) for  $N_2$  fixation (0.25  $mmol\ N\ m^{-2}\ day^{-1}$ ), while the highest reported (5)  $N_2$  fixation rates (8.05  $mmol\ N\ m^{-2}\ day^{-1}$ ) were higher than the denitrification rates by

$\sim 2.98 \text{ mmol N m}^{-2} \text{ day}^{-1}$ . Because these denitrification and  $\text{N}_2$  fixation measurements were made in separate years, these results should be taken as an estimate of possible fixed N loss or gain and not as an absolute budget.

**Conclusions.** An understanding of N cycling in coastal environments is critical for an understanding of global element cycling because nitrification and denitrification rates are significantly higher in these shallow-water environments than in the deep sea (43). However, the question of whether these systems represent a sink or source of N is difficult to answer, because bioturbation activity impacts not only nitrification-denitrification activity but also subsurface  $\text{N}_2$  fixation. Although this heterogeneity of redox processes induced by macrofaunal activity represents a challenge in determining a net loss or gain, we can conclude that in bioturbated sediments, it is possible for the loss of N to be balanced by a gain of N via subsurface  $\text{N}_2$  fixation. Thus, shallow-water sediment environments can represent a source of bioavailable N. The processes within the benthic N cycle are impacted greatly by a change in environmental factors (temperature, salinity, oxygen availability, and nutrient and organic carbon loading); therefore, it is critical that more detailed studies are carried out to improve our understanding of the role of coastal environments in global N cycling.

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