

Wheat leaves emit nitrous oxide during nitrate assimilation

David R. Smart* and Arnold J. Bloom

Department of Vegetable Crops, University of California, Davis, CA 95616-8746

Edited by Pamela A. Matson, Stanford University, Stanford, CA, and approved April 16, 2001 (received for review December 4, 2000)

Nitrous oxide (N₂O) is a key atmospheric greenhouse gas that contributes to global climatic change through radiative warming and depletion of stratospheric ozone. In this report, N₂O flux was monitored simultaneously with photosynthetic CO₂ and O₂ exchanges from intact canopies of 12 wheat seedlings. The rates of N₂O-N emitted ranged from <2 pmol·m⁻²·s⁻¹ when NH₄⁺ was the N source, to 25.6 ± 1.7 pmol·m⁻²·s⁻¹ (mean ± SE, n = 13) when the N source was shifted to NO₃⁻. Such fluxes are among the smallest reported for any trace gas emitted by a higher plant. Leaf N₂O emissions were correlated with leaf nitrate assimilation activity, as measured by using the assimilation quotient, the ratio of CO₂ assimilated to O₂ evolved. ¹⁵N isotopic signatures on N₂O emitted from leaves supported direct N₂O production by plant NO₃⁻ assimilation and not N₂O produced by microorganisms on root surfaces and emitted in the transpiration stream. *In vitro* production of N₂O by both intact chloroplasts and nitrite reductase, but not by nitrate reductase, indicated that N₂O produced by leaves occurred during photoassimilation of NO₂⁻ in the chloroplast. Given the large quantities of NO₃⁻ assimilated by plants in the terrestrial biosphere, these observations suggest that formation of N₂O during NO₂⁻ photoassimilation could be an important global biogenic N₂O source.

Plants play a critical role in regulating the chemical and physical state of the atmosphere through the exchange of biogenic greenhouse gases. Most notable are plant-atmosphere exchanges of CO₂, O₂, and H₂O, but leaves also emit a variety of carbon- and nitrogen-based trace gases involved in climate alteration processes (1). One such trace gas is nitrous oxide (N₂O). Plants—either aerenchymous (2) or nonaerenchymous (3, 4)—can serve as conduits for N₂O between the soil and atmosphere. They transpire significant quantities of N₂O when its concentration in the soil solution greatly exceeds the solution equilibrium concentration with ambient N₂O, currently at ≈315 nmol·mol⁻¹ (5).

The global N₂O budget is beset by uncertainty, and sources of N₂O have historically fallen short of the primary sink, photolysis in the upper atmosphere (6–9). The primary biogenic N₂O sources are from soils (70%) and involve the microbial nitrogen transformations brought about by nitrification and denitrification (6). Although nitrification and denitrification are the major N₂O sources, several microbial organisms that do not nitrify or denitrify can also produce N₂O during NO₃⁻ assimilation (10, 11). These observations have led to the general hypothesis that any enzymatic nitrogen transformation through the +2 to +1 oxidation state may generate N₂O (12). One such transformation in higher plants is NO₂⁻ assimilation in chloroplasts. Nitrite assimilation in chloroplasts can generate intermediates capable of reacting to produce N₂O, including NO₂⁻ (as HNO₂) with hydroxylamine (13) or reaction of NO released during NO₂⁻ reduction (14, 15) with ascorbate (16). Nonetheless, early attempts to observe N₂O production by higher plant tissues were not successful (10) and were probably limited by lack of an analytical method capable of detecting plant N₂O emission at the exceptionally slow rates reported here. We developed an analytical approach by using cryogenic trapping (17) and gas chromatography coupled to high-precision isotope ratio mass spec-

trometry (18). This approach resolved leaf N₂O emissions at more than six orders of magnitude lower than photosynthetic gas exchanges of CO₂ and O₂ (Table 1), placing such fluxes among the smallest ever reported for any trace gas emitted by a higher plant (1).

Identifying and quantifying plant N₂O exchange is important. Atmospheric N₂O concentration is increasing at a rate of about 0.27% per year (19), and each mol of N₂O has ≈290 times the radiative forcing potential of CO₂ (20). Consequently, N₂O will account for as much as 7% of projected atmospheric warming (21). In addition to its greenhouse gas properties, photolytic reaction with excited oxygen [O(¹D)] in the upper atmosphere produces nitric oxide (NO), and NO, in turn, consumes stratospheric ozone (22). Biogenic and anthropogenic sources of N₂O are poorly constrained (23) and often do not account for the quantity of N₂O known to undergo photolysis in the upper atmosphere (6). Extreme heterogeneity of soil N₂O emissions largely contributes to such uncertainty, but unidentified hydrologic or biogenic sources may also play a role (7, 8). Kroeze *et al.* (24) have argued for closure of the global N₂O budget, but the theoretical uncertainty range for estimates of N₂O from agricultural soils is extreme [0.6–14.8 Tg N₂O-N y⁻¹ (6)], and new evidence suggests that N₂O emissions from agricultural soils may be in the lower range of recent Intergovernmental Panel on Climate Change (IPCC) estimates (25). To moderate the atmospheric increases of N₂O and to better understand the role of the biosphere in its production, it is critical to identify all major N₂O sources and exchange pathways, including any potential contributions by plants.

Methods

Wheat seeds (*Triticum aestivum* L. cv. Veery 10) were surface sterilized, germinated on rolled moist germination paper, and transferred to opaque hydroponics tanks containing dilute nutrient solutions (26). The hydroponics systems were kept in a controlled environment chamber (PGV36, Conviron, Winnipeg, MB, Canada) with a 25°C, 16-h, 600 μmol·m⁻²·s⁻¹ photosynthetic photon flux density day and a 20°C, 8-h night. After 10 days, when the plants had three true leaves, 12 individuals were transferred into a gas exchange chamber where their shoots produced a canopy with a leaf surface area of about 0.02 m². Roots were sealed into 12 individual gas-tight cuvettes that were connected to the lower surface of a platform over which the canopy chamber was sealed (27).

Net canopy CO₂ and O₂ exchanges were measured under steady-state conditions. A differential infrared gas analyzer (VIA-500R, Horiba, Sunnyvale, CA) monitored net CO₂ exchange. From a second parallel gas stream, a custom oxygen

This paper was submitted directly (Track II) to the PNAS office.

Abbreviation: IPCC, Intergovernmental Panel on Climate Change.

*To whom reprint requests should be sent at the present address: Department of Viticulture and Enology, University of California, One Shields Avenue, Davis, CA 95616-8749. E-mail: drsmart@ucdavis.edu.

The publication costs of this article were defrayed in part by page charge payment. This article must therefore be hereby marked "advertisement" in accordance with 18 U.S.C. §1734 solely to indicate this fact.

Table 1. Shoot and rhizosphere gas exchange rates for 14-day-old wheat seedlings (*T. aestivum* L. cv. Veery 10)

N source	Photosynthesis			N ₂ O-N emissions	
	CO ₂ , μmol·m ⁻² ·s ⁻¹	O ₂ , μmol·m ⁻² ·s ⁻¹	AQ, CO ₂ /O ₂	Shoot, pmol·m ⁻² ·s ⁻¹	Rhizosphere, pmol·g ⁻¹ ·s ⁻¹
50 μM (¹⁵ NH ₄) ₂ SO ₄	14.31 ± 1.32	12.06 ± 0.77	1.21 ± 0.06	ND	40.4 ± 2.1
100 μM K ¹⁵ NO ₃	15.32 ± 1.48	13.63 ± 0.84	1.13 ± 0.05	25.59 ± 1.68	304.1 ± 93.8

Shown for shoots are net photosynthetic CO₂ assimilation rates, net photosynthetic O₂ evolved, AQ (the ratio of CO₂ assimilated to O₂ evolved), and shoot N₂O emissions (mean ± SE, *n* = 11). Shown for the rhizosphere are N₂O production rates (*n* = 7).

analyzer that resolves a 2 μmol·mol⁻¹ O₂ partial pressure difference on a background of 209,460 μmol·mol⁻¹ monitored net O₂ exchange (28). Before passing through the O₂ electrodes, gases that can interfere with O₂ measurements, like water vapor and plant secondary carbon compounds, were cryogenically condensed to constant trace levels by using a liquid argon trap. The CO₂ concentration during the gas exchange measurements was sustained at 330 μmol·mol⁻¹ and the N₂O concentration at 280 nmol·mol⁻¹. A third parallel gas stream exiting the plant chamber flowed at 1.0–1.5 dm³ s⁻¹ through an ascarite filter to remove CO₂, a drierite filter to remove H₂O, and then a second cryogenic trap cooled with liquid argon to condense N₂O (17). This trap concentrated up to 10 μmol·mol⁻¹ N₂O in 15 min when the N₂O concentration in the gas stream was at 280 nmol·mol⁻¹. Gas trapped in the condenser was injected into a mass spectrometer tuned to determine the ratio of ion currents *m/z* at 45/44 and 46/44 (18). The rate of N₂O exchange by shoots (mol·m⁻² s⁻¹) was calculated as $F_{N_2O} = (J_r d_r C_l / A)$ (29). *J_r* is the flow rate of air through the plant chamber (mol·s⁻¹). *d_l* is the mol fraction of N₂O proportional to canopy N₂O emission, calculated according to the isotopic enrichments in masses ⁴⁵N₂O and ⁴⁶N₂O (30). *C_l* is the concentration of N₂O in the gas stream exiting the chamber. *A* is leaf area (m²). The theoretical detection limit for leaf N₂O emission by the mass spectrometer, on the basis of the observed variation in the mass ratios reported for a 1.5 μmol·mol⁻¹ N₂O standard, was ≈2 pmol·m⁻²·s⁻¹.

During the first 24 h of our experiments, roots received a nitrogen source containing 50 μM (¹⁵NH₄)₂SO₄ (99.6 atom % ¹⁵N), and shoot gas fluxes of CO₂, O₂ and N₂O were assessed during the final 6 h. This pretreatment allowed us to purge the xylem stream of NO₃⁻ and establish a baseline value for assimilation quotient (AQ, CO₂/O₂) when little NO₃⁻ was undergoing assimilation in leaves. The nitrogen source was then shifted to 100 μM K¹⁵NO₃ (99.6 atom % ¹⁵N) for 24 h, and shoot gas fluxes were again assessed during the final 6 h.

For rhizosphere N₂O production, the nutrient solution was delivered to roots by using a gas tight, continuous flow system (31). Fluxes of N₂O from the root cuvettes represent the production of N₂O by roots and any microbial organisms on the root surface capable of generating N₂O (32). Before passing through the root cuvettes, the solution concentrations of O₂ and N₂O were brought to their respective saturation concentrations. Dissolved oxygen concentration in the nutrient solution passing through the root cuvettes declined by only about 20% and the solutions were well stirred, so denitrification in the rhizosphere was minimized. During the experiments, two 5-ml samples of nutrient solution were collected into 15-ml septum bottles. One sample was collected before the nutrient solution entered the root cuvette and another after it exited the root cuvette. The head space gases from the septum bottles were injected into the mass spectrometer after equilibration at 22.5°C to determine the concentration of N₂O and the ratio of ion currents *m/z* at 45/44 and 46/44. The rate of N₂O production by the rhizosphere (mol·g⁻¹·s⁻¹) was calculated as $F_{N_2O} = (J_r d_r C_r / W)$. *J_r* is the flow rate of nutrient solution through the root cuvette (liter·s⁻¹). *d_r* is the mol fraction of N₂O in the nutrient solution proportional to rhizosphere N₂O production, and calculated according to the

isotopic enrichments in masses ⁴⁵N₂O and ⁴⁶N₂O (30). *C_r* is the total concentration of N₂O (mol·liter⁻¹) in the nutrient solution (33). *W* is root dry mass.

For chloroplast assays, ≈40 g of fresh leaves from 2-week-old hydroponics-grown wheat plants was blended in a buffer solution containing 0.05 M K-Hepes (pH 7.3), 0.33 M Sorbitol, 1 mM MgCl₂, 1 mM MnCl₂, 2 mM Na₂EDTA, and 0.1% BSA. The extract was centrifuged at 3,000 × *g* for 5 min, resuspended in a 50/50 Percol gradient, and centrifuged at 7,000 × *g* for 10 min. Intact chloroplasts were then collected and washed with 0.05 M K-Tricine (pH 8.0) and 0.33 M Sorbitol. The washed chloroplasts were introduced into 5 ml of an incubation medium consisting of 0.3 mM K¹⁵NO₂ (99.6 atom % ¹⁵N) in 0.05 M K-Tricine (pH 8.0), 0.33 M sorbitol, and 0.3 mM NaHCO₃. The chloroplasts were incubated for 40 min in 15-ml septum bottles at 25°C and a light intensity of 600 μmol·m⁻²·s⁻¹ photosynthetic photon flux density. At the end of the incubation period, the head-space gases were collected and injected into the mass spectrometer. The concentration of N₂O and ratio of ion currents *m/z* at 45/44 and 46/44 were used to calculate the quantity of N₂O produced, as previously described for the dynamic flow measurements.

For nitrate and nitrite reductase assays (34), 2 g of fresh leaves from 2-week-old hydroponics-grown wheat plants was ground in a mortar and pestle with cold-purified N-free sand in 8 ml of a chilled buffer solution. The buffer solution consisted of 0.05 M Tris (pH 8.5), 1 mM EDTA, 1 μM NaMoO₄, 10 μM FAD, 1 mM DTT, 10 μM leupeptin, and 1 μg ml⁻¹ pepstatin. The extract was centrifuged at 30,000 × *g* for 20 min. Then 0.05 ml of the supernatant was added to 0.95 ml of an assay solution containing 62.5 mM potassium phosphate buffer (pH 7.5), 0.7 mM K¹⁵NO₂ (99.6 atom % ¹⁵N), and 0.052 g ml⁻¹ of methyl viologen in a 15-ml septum bottle. The reaction was initiated by injecting 0.2 ml of a solution containing 8.3 mg ml⁻¹ Na₂S₂O₄, and incubated at 30°C for 15 min. For nitrate reductase, 0.05 ml of the supernatant was added to 0.95 ml of assay buffer containing 1.4 mM K¹⁵NO₃ (99.6 atom % ¹⁵N). The reaction was initiated by adding 0.2 ml of a solution containing 2 mg ml⁻¹ NADH and incubated at 30°C for 15 min. The head-space gases from the nitrite and nitrate reductase assays were collected and analyzed on the mass spectrometer as previously described.

Results and Discussion

This is the first study, to our knowledge, to report quantitative leaf N₂O emissions under normal physiological conditions for an intact plant by using steady-state gas exchange methods. Earlier investigations have examined detached leaves or used static chamber methods, where CO₂, H₂O, and temperature change rapidly (3, 4, 10), and nitrogen source (NH₄⁺ and NO₃⁻) cannot be controlled. These preliminary studies were valuable in that they clearly demonstrated N₂O can move from soil to atmosphere via the plant transpiration stream, but they did not resolve the question of N₂O production by plant nitrogen metabolism.

In our investigation, leaves did not emit N₂O at a detectable rate while metabolizing ¹⁵NH₄⁺. The AQ during ¹⁵NH₄⁺ exposure averaged 1.21 ± 0.05 units (Table 1). When the N source was shifted to ¹⁵NO₃⁻, canopy leaves emitted N₂O-N at an average rate of 25.6 ± 1.7 pmol·m⁻²·s⁻¹, and the AQ declined to 1.13 ±

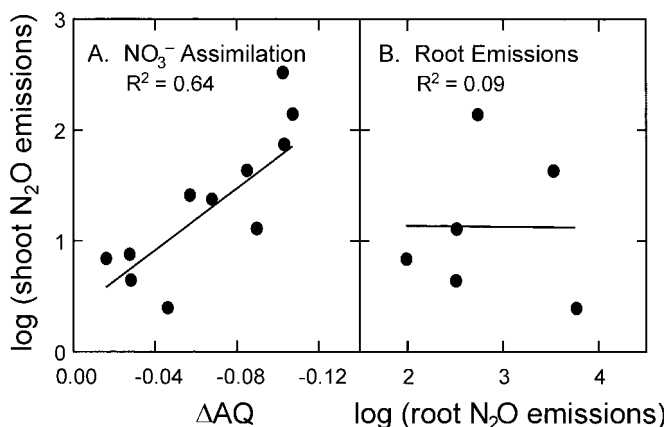


Fig. 1. The relationship between N_2O emission from wheat leaves ($\text{pmol N}_2\text{O-N m}^{-2}\text{s}^{-1}$) and (A) the change in assimilatory quotient (ΔAQ) when the nitrogen source was shifted from $^{15}\text{NH}_4^+$ to $^{15}\text{NO}_3^-$, or (B) N_2O production in the rhizosphere ($\text{pmol N}_2\text{O-N g}^{-1}\text{s}^{-1}$). Shown are the regression lines and the R^2 statistic. The rates of N_2O emission were normalized by using the \log_{10} transformation.

0.06 units (Table 1). Both observations are important. The change in AQ (ΔAQ) gives a nondestructive measure of NO_3^- assimilation under steady-state conditions. Net O_2 exchange provides a measure of photosynthetic electron transport, and either CO_2 or NO_3^- reduction can be coupled to such electron transfer (35). As NO_3^- photoassimilation increases, the AQ declines because reductant produced by photosynthetic electron transport increases to support NO_3^- and NO_2^- reduction in addition to CO_2 fixation (28, 36). Thus, not only did shoot N_2O emission become detectable when the N source was shifted to $^{15}\text{NO}_3^-$, but the strong correlation with ΔAQ (Fig. 1A) suggested that leaf N_2O flux was driven by leaf NO_3^- assimilation.

On the other hand, rhizosphere N_2O production also increased more than 7-fold (Table 1) with the change in N source from $^{15}\text{NH}_4^+$ to $^{15}\text{NO}_3^-$ (Table 1), raising N_2O concentration in the nutrient solution around roots to an average of $89.6 \text{ nmol}\cdot\text{liter}^{-1}$. Such a concentration is an order of magnitude higher than the solution equilibrium concentration when N_2O in the shoot chamber is near ambient, so that movement of N_2O in the transpiration stream could have been significant. At least two experimental observations did not support this scenario. First, shoot N_2O emission during exposure to $^{15}\text{NO}_3^-$ was not correlated with root zone N_2O production (Fig. 1B). Second, leaf N_2O flux fell below detectable limits during exposure to $^{15}\text{NH}_4^+$ despite high rates of N_2O production by the rhizosphere (Table 1). The production of N_2O during $^{15}\text{NH}_4^+$ exposure comes from the activity of nitrifying bacteria on the root surface (37), an activity we were able to completely shut down by using the nitrification inhibitor nitrapyrin. Thus, the absence of a correlation between shoot N_2O emission and rhizosphere N_2O production provides further evidence that photoassimilation of NO_3^- was the major source of N_2O emitted from leaves in this investigation. For investigations where N_2O transpiration was responsible for leaf N_2O flux (6), the concentrations of N_2O in the soil solution were nearly four orders of magnitude higher, at $\approx 326 \text{ }\mu\text{mol}\cdot\text{liter}^{-1}$, than N_2O concentrations observed in our nutrient solutions. This may help to explain why N_2O transpiration was not a factor in this investigation.

The isotopic composition of N_2O emitted from leaves provided further verification that shoot NO_3^- assimilation was largely responsible for the observed N_2O flux. In the course of the experiments, we used nitrogen sources in nutrient solutions highly enriched in ^{15}N ($>99.6 \text{ atom } \% ^{15}\text{N}$ as NH_4^+ or NO_3^-).

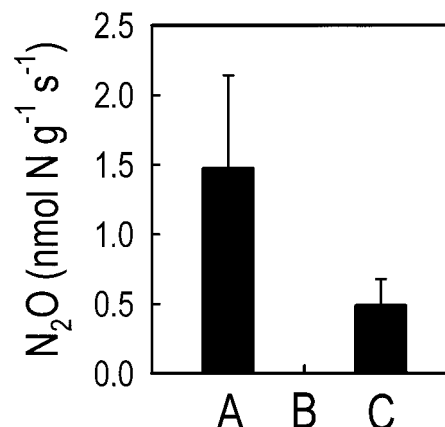


Fig. 2. N_2O production *in vitro* from (A) intact chloroplasts ($\text{nmol N}_2\text{O-N g}^{-1}$ chlorophyll s^{-1}), (B) nitrate reductase ($\text{nmol N}_2\text{O-N g}^{-1}$ protein s^{-1}), and (C) nitrite reductase ($\text{nmol N}_2\text{O-N g}^{-1}$ protein s^{-1}). Intact chloroplasts, nitrate reductase, and nitrite reductase were extracted from fully expanded leaves of 2- to 3-week-old wheat plants (*T. aestivum* L. cv. Veery 10).

Thus, any N_2O produced by microbial metabolism in the rhizosphere, nitrification or denitrification, and subsequently transpired, would be predominantly composed of the $^{15}\text{N}^{15}\text{N}^{16}\text{O}$ or $^{46}\text{N}_2\text{O}$ isoform. We grew our wheat plants in a nutrient solution with $200 \text{ }\mu\text{M NH}_4^+$ plus $200 \text{ }\mu\text{M NO}_3^-$ containing ^{15}N close to natural abundance levels ($0.3674\% ^{15}\text{NH}_4^+$ and $0.3665\% ^{15}\text{NO}_3^-$, or $>99.6 \text{ atom } \% ^{14}\text{N}$). These plants stored relatively large quantities of NO_3^- in their tissues ($8.76 \pm 2.23 \text{ mM}$, mean \pm SE, $n = 6$), of which $2.25 \pm 1.65 \text{ mM}$ was withdrawn and assimilated during the first 24-h exposure to $^{15}\text{NH}_4^+$. N_2O emitted during the assimilation of this internal NO_3^- would not be detected during exposure of roots to $^{15}\text{NH}_4^+$, because the isotopic signature on such stored NO_3^- was very similar to that of N_2O in the background air. However, if stored NO_3^- were being mobilized and assimilated in leaves during the subsequent period of root exposure to $^{15}\text{NO}_3^-$, or chloroplast reactants were participating in N_2O production, then some of the N_2O emitted would contain the $^{45}\text{N}_2\text{O}$ isoforms of $^{15}\text{N}^{14}\text{N}^{16}\text{O}$ and $^{14}\text{N}^{15}\text{N}^{16}\text{O}$. Indeed, $^{45}\text{N}_2\text{O}$ was detected in the N_2O stream emitted from wheat leaves during exposure to $^{15}\text{NO}_3^-$. From the mass ratios of $^{44}\text{N}_2\text{O}$, $^{45}\text{N}_2\text{O}$, and $^{46}\text{N}_2\text{O}$ collected during such experiments, we estimated (30) that $5.19 \pm 0.92 \text{ pmol}\cdot\text{m}^{-2}\text{s}^{-1}$ (mean \pm SE, $n = 13$), or about 20%, of the $\text{N}_2\text{O-N}$ emitted by leaves came from the ^{14}N stored in the plant. It is unlikely that bacterial mineralization of organic-N in the rhizosphere followed by nitrification and denitrification contributed to the $^{45}\text{N}_2\text{O}$ emitted. Roots treated with antibiotic cocktails (38) and nitrapyrin to eliminate rhizosphere microbial nitrogen transformations still emitted detectable quantities of $^{45}\text{N}_2\text{O}$. Nonetheless, the largest fraction of the observed leaf N_2O flux was clearly generated by plant NO_3^- assimilation and not from N_2O produced by microbial processes in the rhizosphere.

The two enzymes responsible for plant NO_3^- assimilation, NO_3^- reductase (NR) and NO_2^- reductase (NiR), are located in the cytoplasm and chloroplasts, respectively. To determine whether NO_3^- or NO_2^- reduction was responsible for the N_2O emitted, we extracted NR, NiR, and intact chloroplasts from wheat leaves and assayed for N_2O production *in vitro* as described. Whereas NR assays did not produce any detectable N_2O , NiR assays and intact chloroplasts did (Fig. 2). These experiments provide *in vitro* evidence that the N_2O emitted from wheat leaves was generated by NO_2^- reduction to NH_3 in the chloroplasts, where NO_2^- undergoes transformation through the +2 to +1 oxidation state, as predicted.

The leaf N₂O emissions we measured were small (Table 1), and it is not currently known how important fluxes of this magnitude are in relation to the global N₂O budget. Some recent micrometeorological N₂O flux measurements over grass swards failed to detect a higher N₂O emission rate than that observed by using static chambers over soil alone (39). However, the error in such measurements was too large, with coefficients of variation for the chamber data ranging from 0.42 to 1.83 (40, 41), to resolve plant N₂O emissions of the size measured in this investigation. In support of a role for plant N₂O emission, Hutchinson and Mosier (42) found that aerodynamic flux measurements over an irrigated corn field were always higher than, although never exceeding twice the mean of, fluxes measured by using soil chambers. Some recent chamber studies found that ryegrass canopies were responsible for 21.1% of the total N₂O emissions (4), and that *Linum perenne* canopies contributed to as much as 50% of N₂O flux to the atmosphere (43). Our results indicated that 0.02–0.2% of the NO₃⁻-N assimilated by wheat was released as N₂O-N. Current estimates are that terrestrial land plants assimilate 1,200 Tg of N annually (44). Nearly half of this N is thought to be absorbed and assimilated as NO₃⁻ (45), of which 25–75% is assimilated in leaves (46). The NR and NiR enzymes involved are highly conserved among higher plants (47). Thus, if

other terrestrial land plants behave as wheat does, we calculate that NO₃⁻ photoassimilation alone could produce from 0.03 to 0.9 Tg N₂O-N yr⁻¹. The uncertainty in the IPCC's estimates of global N₂O source emissions is substantial, with nearly two orders of magnitude in the range of estimated emission of N₂O from soils [0.6–14.8 Tg N y⁻¹ (1)]. Robertson *et al.* (25) tracked N₂O emissions for 9 years from soils of six cropping systems, including a nitrogen-intensive corn rotation and four successional communities, and found emissions to be at the lower end of the IPCC's calculated emission factor, which is based on fertilizer application. Our estimates of plant N₂O emissions represent ≈5–6% of the total amount of N₂O-N thought to be emitted by agricultural plant–soil systems alone (1, 44). These approximations do not include the quantity of N₂O that might be conducted to the atmosphere via the plant transpiration stream. Thus, our results suggest that higher plants could play an intriguing role in N₂O exchange not previously considered in biosphere–atmosphere interactions.

We thank Na Trinh for assistance in many phases of these experiments and Duy Nguyen for assistance with chloroplast and enzyme preparations. This study was supported by the Department of Energy under Grant 95ER62128 TECO and the National Science Foundation under Grant IBN-99-74927.

- Sharkey, T. D., Holland, E. A. & Mooney, H. A. (1991) *Trace Gas Emissions by Plants* (Academic, San Diego).
- Mosier, A. R., Mohanty, S. K., Bhadrachalam, A. & Chakravorti, S. P. (1990) *Biol. Fertil. Soils* **9**, 61–67.
- Chang, C., Janzen, H. H., Cho, C. M. & Nakonechny, E. M. (1998) *Soil Sci. Soc. Am. J.* **62**, 35–38.
- Chen, X., Boeckx, P., Shen, S. & Van Cleemput, O. (1999) *Biol. Fertil. Soils* **28**, 393–396.
- Battle, M., Bender, M., Sowers, T., Tans, P. P., Butler, J. H., Elkins, J. W., Ellis, J. T., Conway, T., Ahang, N., Lang, P. & Clarke, A. D. (1996) *Nature (London)* **383**, 231–235.
- Mosier, A. C., Kroeze, C., Nevison, C., Oenema, O., Seitzinger, S. & van Cleemput, O. (1998) *Nutr. Cycl. Agroecosyst.* **52**, 225–248.
- Naqvi, S. W. A., Yoshinari, T., Jayakumar, D. A., Altabet, M. A., Narvekar, P. V., Devol, A. H., Brandes, J. A. & Codispoti, L. A. (1998) *Nature (London)* **394**, 462–464.
- Bouwman, A. F., Van der Hoek, K. W. & Olivier, J. G. J. (1995) *J. Geophys. Res.* **100**, 2785–2800.
- Houghton, J. T., Callander, B. A. & Varney, S. K., eds. (1992) *Climate Change. The Supplementary Report to the IPCC Scientific Assessment* (Cambridge Univ. Press, Cambridge, U.K.).
- Bleakley, B. H. & Tiedje, J. M. (1982) *Appl. Environ. Microbiol.* **44**, 1342–1348.
- Hutchinson, G. L. & Davidson, E. A. (1993) in *Agricultural Ecosystem Effects on Trace Gases and Global Climate Change*, ASA, Special Publication No. 55, eds. Harper, L. A., Mosier, A. R., Duxbury, J. M. & Rolston, D. E. (ASA/CSSA/SSSA, Madison WI), pp. 79–94.
- Firestone, M. K. & Davidson, E. (1989) in *Exchange of Trace Gases Between the Terrestrial Ecosystems and the Atmosphere: Report of the Dahlem Workshop on Exchange of Trace Gases between Terrestrial Ecosystems and the Atmosphere, Berlin 1989, February 19–24*, eds. Andreae, M. O. & Shimel, D. S. (Wiley, New York), pp. 7–21.
- Bothner-By, A. & Friedman, J. (1952) *J. Chem. Phys.* **20**, 459–462.
- Aparicio, P. J., Knaff, D. B. & Malkin, R. (1975) *Arch. Biochem. Biophys.* **169**, 102–107.
- Dean, J. V. & Harper, J. E. (1986) *Plant Physiol.* **82**, 718–723.
- Brudvig, G. W., Stevens, T. H. & Chan, S. I. (1980) *Biochemistry* **19**, 5275–5285.
- Huff, A. K., Cliff, S. S. & Thiemens, M. H. (1997) *Anal. Chem.* **69**, 4267–4270.
- Brooks, P. D., Herman, D. J., Atkins, G. J., Prosser, S. J. & Barrie, A. (1993) in *Agricultural Ecosystem Effects on Trace Gases and Global Climate Change*, ASA Special Publication No. 55, eds. Harper, L. A., Mosier, A. R., Duxbury, J. M. & Rolston, D. E. (ASA/CSSA/SSSA, Madison, WI), pp. 193–202.
- Ramanathan, V., Cicerone, R. J., Singh, H. B. & Kiehl, J. T. (1985) *J. Geophys. Res.* **90**, 5547–5566.
- IPCC (1996) in *Climate Change 1995, The Science of Climate Change*, eds. Houghton, J. T., Meira Filho, L. G., Callander, B. A., Harris, N., Kattenberg, A. & Maskell, K. (Cambridge Univ. Press, Cambridge, U.K.), pp. 9–50.
- Robertson, G. P. (1993) in *Agricultural Ecosystem Effects on Trace Gases and Global Climate Change*, ASA Special Publication No. 55, eds. Harper, L. E., Mosier, A. R., Duxbury, J. M. & Rolston, D. E. (ASA/CSSA/SSSA, Madison, WI), pp. 96–108.
- Cicerone, R. J. (1987) *Science* **237**, 35–42.
- IPCC (1997) *Intergovernmental Panel on Climate Change Guidelines for National Greenhouse Gas Inventories. Chapter 4. Agriculture: Nitrous Oxide from Agricultural Soils and Manure Management, Paris, France, 1997* (Organization for Economic Cooperative Development, Paris).
- Kroeze, C., Mosier, A. R. & Bouwman, L. (1999) *Global Biogeochem. Cycles* **13**, 1–8.
- Robertson, G. P., Paul, E. A. & Harwood, R. R. (2000) *Science* **289**, 1922–1925.
- Smart, D. R., Ritchie, K., Bloom, A. J. & Bugbee, B. (1998) *Plant Cell Environ.* **21**, 753–763.
- Kosola, K. R. & Bloom, A. J. (1994) *Plant Physiol.* **104**, 435–442.
- Bloom, A. J., Caldwell, R. M., Finazzo, J., Warner, R. L. & Weissbart, J. (1989) *Plant Physiol.* **91**, 352–356.
- Sestak, Z., Catsky, J. & Jarvis, P. G. (1971) *Plant Photosynthetic Production: Manual of Methods* (W. Junk, The Hague, The Netherlands).
- Mulvaney, R. L. & Boast, C. W. (1986) *Soil Sci. Soc. Am. J.* **50**, 360–363.
- Bloom, A. J. (1989) in *Application of Continuous and Steady State Methods to Root Biology*, eds. Torrey, J. G. & Winship, L. J. (Kluwer, Dordrecht, The Netherlands), pp. 147–163.
- Smart, D. R., Ritchie, K., Stark, J. M. & Bugbee, B. (1997) *Appl. Environ. Microbiol.* **63**, 4621–4624.
- Wilhelm, E., Battino, R. & Wilcock, R. J. (1977) *Chem. Rev.* **77**, 219–248.
- Aslam, M. & Huffaker, R. C. (1989) *Plant Physiol.* **91**, 1152–1156.
- Knaff, D. B. (1996) in *Oxygenic Photosynthesis: The Light Reactions*, eds. Ort, D. R. & Yocum, C. F. (Kluwer, Dordrecht, The Netherlands), Vol. 4, pp. 333–361.
- Cramer, M. & Myers, J. (1948) *J. Gen. Physiol.* **32**, 93–102.
- Padgett, P. E. & Leonard, R. T. (1993) *Plant Physiol.* **101**, 141–146.
- Smart, D. R., Ferro, A., Ritchie, K. & Bugbee, B. B. (1995) *Physiol. Plant.* **95**, 533–540.
- Hargreaves, K. J., Skiba, U., Fowler, D., Arah, J., Wienhold, F. G., Klemmedtsen, L. & Galle, B. (1994) *J. Geophys. Res.* **99**, 16569–16574.
- Wienhold, F. G., Frahm, H. & Harris, G. W. (1994) *J. Geophys. Res.* **99**, 16557–16567.
- Clayton, H., Arah, J. R. M. & Smith, K. A. (1994) *J. Geophys. Res.* **99**, 16599–16607.
- Hutchinson, G. L. & Mosier, A. R. (1979) *Science* **205**, 1125–1127.
- Anderson, L. & Hopkins, D. W. (1997) in *Gaseous Nitrogen Emissions from Grasslands*, eds. Jarvis, S. C. & Pain, B. F. (CAB International, New York), pp. 222–224.
- Schlesinger, W. H. (1997) *Biogeochemistry: An Analysis of Global Change* (Academic, San Diego).
- Raven, J. A., Wollenweber, B. & Handley, L. L. (1993) *Physiol. Plant* **89**, 512–518.
- Andrews, M. (1986) *Plant Cell Environ.* **9**, 511–519.
- Caboche, M. & Rouze, P. (1990) *Trends Genet.* **6**, 187–192.