

# Assessing the role of cladogenesis in macroevolution by integrating fossil and molecular evidence

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**Assessing the extent to which population subdivision during cladogenesis is necessary for long-term phenotypic evolution is of fundamental importance in a broad range of biological disciplines. Differentiating cladogenesis from anagenesis, defined as evolution within a species, has generally been hampered by dating precision, insufficient fossil data, and difficulties in establishing a direct link between morphological changes detectable in the fossil record and biological species. Here we quantify the relative frequencies of cladogenesis and anagenesis for macroperforate planktic Foraminifera, which arguably have the most complete fossil record currently available, to address this question. Analyzing this record in light of molecular evidence, while taking into account the precision of fossil dating techniques, we estimate that the fraction of speciation events attributable to anagenesis is <19% during the Cenozoic era (last 65 Myr) and <10% during the Neogene period (last 23 Myr). Our central conclusion—that cladogenesis is the predominant mode by which new planktic Foraminifera taxa become established at macroevolutionary time scales—differs markedly from the conclusion reached in a recent study based solely on fossil data. These disparate findings demonstrate that interpretations of macroevolutionary dynamics in the fossil record can be fundamentally altered in light of genetic evidence.**

punctuated equilibrium | phyletic gradualism | lineage | morphospecies

**B**iological evolution proceeds through two distinct modes: anagenesis, which occurs within a species, and cladogenesis, which occurs during speciation and results in the subdivision of a species into two reproductively isolated, independently evolving taxa (1, 2). Distinguishing between these modes is essential to determining the role of population subdivision in evolutionary dynamics (3, 4). This topic has received considerable attention in paleontology (1, 2), perhaps because the fossil record provides the only direct record of long-term morphological change (5), but its importance extends to other biological disciplines. For example, in population genetics, models only predict that phenotypic change is contingent upon or associated with the evolution of reproductive isolation under specific circumstances (3). Thus, a primary role for cladogenesis may highlight the importance of particular speciation modes (2), metapopulations dynamics (6), or adaptive peaks in fitness landscapes (7, 8). In community ecology, both evolutionary modes may influence biodiversity, but the mechanisms differ: only cladogenesis results in an increase in diversity (9–11), but anagenesis may help maintain diversity if it results in niche differences that facilitate species coexistence (12, 13).

Remarkably few studies have attempted to quantify the relative frequencies of anagenesis and cladogenesis for entire species assemblages (e.g., refs. 5, 14, 15). Addressing this question is challenging due to insufficient fossil data for most taxa and the difficulties in relating morphological change to genetic divergence and the evolution of reproductive isolation (16). Fossil data can only be used to distinguish anagenesis from cladogenesis, at the assemblage level, to the extent that four general assumptions are upheld (5): (i) morphologically distinct taxa (hereafter morphospecies) that coexist temporally are reproductively isolated, and therefore independently evolving; (ii)

ancestor–descendant relationships among morphospecies are inferred accurately; (iii) the vast majority of morphospecies have been identified; and (iv) the temporal range of each morphospecies, as defined by its first appearance datum (FAD) and last appearance datum (LAD), is estimated with known precision. Together these assumptions imply that the mode of evolution may be anagenetic if the FAD of the descendant coincides with the LAD of the ancestor within the bounds of the dating precision. Conversely, the mode is cladogenetic if, after accounting for dating precision, the FAD of the descendant still precedes the LAD of the ancestor, because this implies that the ancestor and descendant existed concurrently for some time period. Overlap in the temporal ranges of the ancestor and descendant has long been considered the “best de facto evidence of true cladogenesis” (17).

The fossil record for planktic Foraminifera—a group of single-celled eukaryotes that occupy marine pelagic zones throughout the world (18)—arguably adheres to assumptions *i–iv* better than that of any other group, providing an unprecedented opportunity to assess the relative frequencies of anagenesis and cladogenesis. Early fossil studies suggested that Foraminifera morphospecies generally arise through gradual, anagenetic change (19, 20). However, more recent studies of selected taxa indicate that cladogenesis may be far more prevalent than previously thought (21, 22). These recent studies highlight the need to integrate fossil data with other forms of evidence, particularly molecular data (23).

Here we assess the relative frequencies of anagenesis and cladogenesis for Cenozoic macroperforate planktic Foraminifera. We do so by analyzing the recently published phylogeny of Aze et al. (24), using a unique approach that incorporates assumptions *i–iii* above while taking into account uncertainties in the precision of the FAD and LAD estimates (assumption *iv*). This phylogeny summarizes current understanding of ancestor–descendant relationships among extinct and extant morphospecies based on the work of multiple researchers. It is anchored by a comprehensive compilation of FAD and LAD estimates (24), most of which were obtained using the most precise dating techniques currently available (25–27). Interpreting this phylogeny in light of molecular evidence for extant members of the Foraminifera (28–31), we calculate upper-bound estimates for the fractions of speciation events that may be the result of anagenesis for the Cenozoic era (last 65 Ma) and for the Neogene period (last 23 Ma), as described in *Methods*. Before presenting the results of our calculations, we consider evidence for and against assumptions *i–iv* for planktic Foraminifera. We conclude by considering the limitations of our approach and the potential implications of our findings.

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## Evaluating Assumptions

Adherence to assumptions *i–iv* above requires a largely complete fossil record. The Cenozoic fossil record for planktic Foraminifera is among the most well resolved for any group in terms of both taxonomic and temporal resolution owing to: high abundance and global ubiquity in marine pelagic zones (18), high preservation potential of their calcium carbonate shells in marine sediments, extensive spatial and temporal coverage of sampling, and a consistent and well-established system of taxonomic classification (24).

Assumptions *i* and *ii* are generally consistent with molecular studies, but exceptions have been observed. Analyses of small subunit ribosomal DNA (SSU rDNA) data for 19 extant morphospecies (of ~50) indicate that morphospecies are generally monophyletic, and that ancestor–descendant relationships among morphospecies are accurately inferred from fossils at lower levels of the phylogeny (33, 34). One noted exception to monophyly among extant morphospecies is *Globigerinoides ruber*, which is now recognized as comprising two distinct clades on the basis of molecular evidence (35). Inconsistencies between molecular and fossil data are largely confined to the placement of higher-order clades, where SSU rDNA divergences are high and relationships are difficult to resolve (36). Although not all ancestor–descendant relationships are yet well established (24), the general concordance of molecular- and fossil-based methods at lower taxonomic levels suggests that the “stratophenetic” method of delineating ancestor–descendant relationship by tracing specimens of intermediate morphology back through time (37) is quite robust. More generally, this concordance is consistent with data indicating that shell traits often have a genetic basis that is sufficiently strong and conservative to allow discrimination among clades at macroevolutionary time scales (Table 1). Given that the shell characteristics used for planktic Foraminifera identification are the same for extinct and extant taxa, the findings of molecular studies are relevant for interpreting fossil assemblages (30). In particular, they suggest that assumptions *i* and *ii* are also upheld for extinct morphospecies.

Assumption *iii* is supported by compiled data (38) presented in Fig. 1, which provides clear evidence of an asymptote in the rate of discovery of new morphospecies between 1826 and 2001 (Fig. 1A). Analyzing these data using the method proposed by Bebbler et al. (39), we estimate that ~96% [95% confidence interval (CI) of 96–100%] of all planktic Foraminifera morphospecies (both extinct and extant) have already been described (Fig. 1B). Consequently, it is unsurprising that recent taxonomic work has focused primarily on clarifying and refining phylogenetic relationships among already described morphospecies (40, 41) and

subdividing morphospecies on the basis of finer-scale morphological differences (28, 31, 42, 43) (Table 1).

Assigning precision estimates (assumption *iv*) to the FAD and LAD data, in part, entails consideration of three sources of dating error: disturbance of the sediments in which the taxon occurs after initial deposition or during collection, resulting in a mixing of sediments of varying age; regional diachroneity of taxon occurrence due to expansion and contraction of a taxon’s geographic range; and resolution of the chronostratigraphic techniques used to date the sediments (44). The first two sources affect dating at particular locations, but are largely controlled for when estimating FAD and LAD because consensus estimates are obtained by selecting the data deemed to be most reliable from among multiple locations (24). Chronostratigraphic resolution is therefore the primary source of error and, importantly, varies with sediment age. For the Neogene period (last 23 Ma), changes in sediment properties through time can be detected and correlated with Milankovitch glacial–interglacial cycles, resulting in a dating resolution of  $\pm 0.005$  Ma to  $\pm 0.046$  Ma (45). By contrast, for the Paleogene period (65–23 Ma ago), the most accurate dating method makes use of cycles in geomagnetic polarity, resulting in a dating resolution of  $\pm 0.380$  Ma (46). As these dating methods have known levels of precision, they can be used to assign maximum precision estimates to the FAD and LAD data, following assumption *iv*.

Despite evidence in support of assumptions *i–iv*, there are two limitations in using morphospecies-level FAD and LAD data to distinguishing anagenesis from cladogenesis. First, Foraminifera specimens are generally assigned to morphospecies based on their similarity to designated holotypes (47). Whereas the holotype concept is inherently static and typological, a population is a dynamic entity composed of individuals of varying morphology (37). If individuals at the extremes of the morphotype distribution of an evolving population, rather than the mean morphotype, are used to estimate FAD and LAD, overlap in the temporal ranges of the “ancestor” and “descendant” of a single population undergoing anagenesis may occur (Fig. 2A). This phenomenon, which we refer to as “typological error,” can be avoided by defining FAD and LAD on the basis of frequency distributions of morphotypes, but detailed morphometric studies of this type are rare (e.g., ref. 48). In the absence of data to directly estimate the magnitude of typological error, we consider a range of values for the total error, defined as the sum of the dating error and the typological error, as well as other forms of evidence (*Methods*).

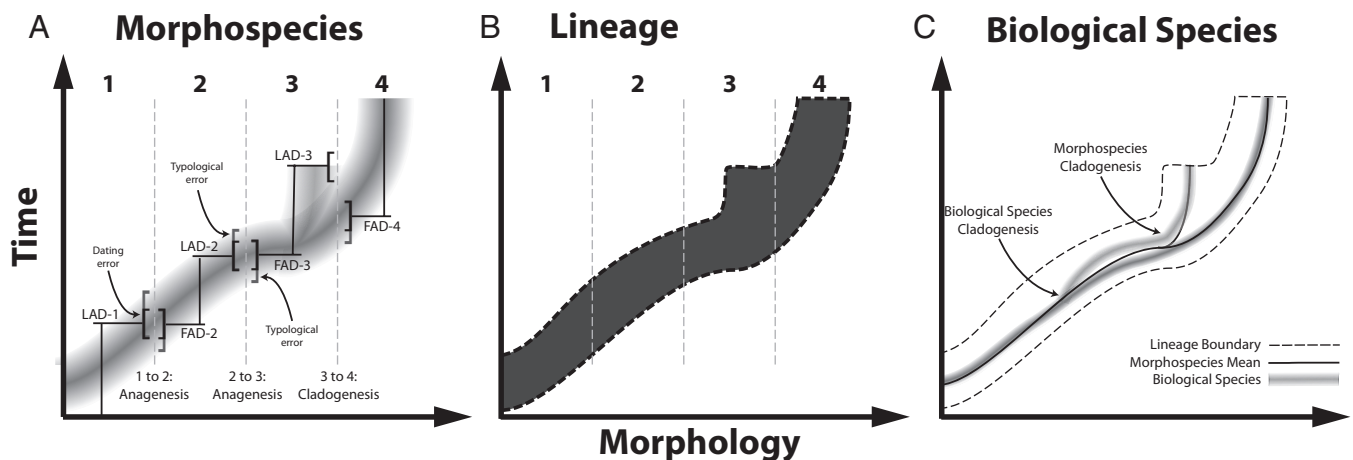
The second challenge is that a number of studies have presented molecular evidence that morphospecies generally contain multiple “cryptic” genotypes that correspond to distinct biological species (30, 35, 49). Whereas these findings lend further

**Table 1. Foraminiferal morphological characters that have been demonstrated to reflect genetically distinct populations**

Morphological character	Independent evidence
Test outline	Test outlines differ among genetically, ecologically, and biogeographically distinct populations within <i>Globigerinella siphonifera</i> (42) and <i>Truncorotalia truncatulinoides</i> (31, 54).
Chamber morphology	Compression of the final two chambers can be used to discriminate among genetically distinct populations within <i>Globigerinoides ruber</i> (49).
Coiling direction	Sinistral and dextral forms of the morphospecies <i>Neogloboquadrina pachyderma</i> correspond to ecologically and genetically distinct species with different temperature tolerances and biogeographic distributions (68, 69). Relative frequencies of sinistral and dextral forms vary among genetically, ecologically, and biogeographically distinct populations of <i>Truncorotalia truncatulinoides</i> (31).
Porosity	Populations of <i>Orbulina universa</i> that vary in porosity are genetically distinct and have different biogeographic distributions (28, 42, 43).

Test outline and chamber morphology have been used historically to differentiate morphospecies. Coiling direction has only recently been confirmed as a character of taxonomic significance (68). Porosity is not currently recognized as a diagnostic character.

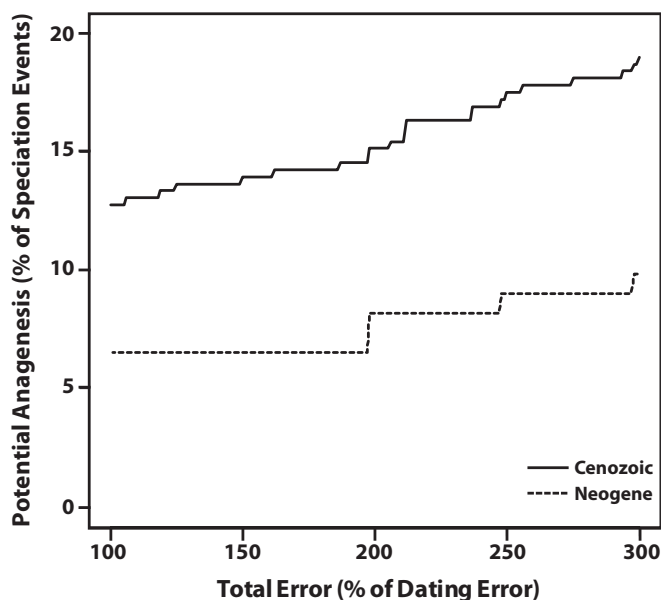




**Fig. 2.** Graphical depiction of morphological evolution at the levels of (A) morphospecies, (B) lineage, and (C) biological species, for a hypothetical lineage. (A) Darker shading corresponds to morphotypes of higher abundance in morphospace for an evolving population. Dashed vertical lines represent typological divisions of the morphospace into discrete morphospecies. Based on assumptions *i–iv*, we identify the transitions between morphospecies 1 and 2, and between morphospecies 2 and 3, as potentially anagenetic, because the FAD and LAD values overlap within our error margins. Typological errors may broaden the estimated temporal overlap between ancestor and descendant, and must therefore also be accounted for. As the LAD of morphospecies 3 and FAD of morphospecies 4 do not overlap, and the two forms occupy distinguishable peaks in morphospace, they represent distinct morphospecies, implying that a cladogenetic event occurred. (B) All four morphospecies are assigned to the same lineage, despite morphospecies 3 and 4 having distinguishable frequency distributions in morphospace, because adjacent frequency distributions exhibit some overlap. Had there been unoccupied morphospace between morphospecies 3 and 4, this would have represented a cladogenetic event at the lineage level (24). (C) Selected “cryptic” biological species within morphospecies are represented as thin lines. Molecular evidence indicates that a morphospecies generally comprises multiple cryptic biological species (30). Consequently, the cladogenetic event at the morphospecies level may occur at the same time as the cladogenetic event at the biological species level or may occur later, as depicted in the figure.

64 of 337 events), including seven putative anagenetic events, and the Neogene estimate from 18.9 to 9.8% (95% CI: 4.9–15.6%; 12 of 122 events), including one putative anagenetic event (Table S1).

These estimates are sensitive to the assumed magnitudes of typological errors (Fig. 3). For example, if morphological and phylogenetic evidence are considered, but typological errors are



**Fig. 3.** Estimated fractions of speciation events that are potentially anagenetic for Cenozoic (solid lines) and Neogene (dashed lines) planktic Foraminifera, assuming varying levels of total error. Total errors are expressed as percentages of the dating errors (100% =  $\pm 0.1$  Ma for the Neogene,  $\pm 0.4$  Ma for Paleogene). Estimates take into account morphological and phylogenetic evidence (Table S1).

assumed to be negligible, the estimated fraction of speciation events that are potentially anagenetic is reduced further to 12.8% for Cenozoic events (95% CI: 9.2–16.3%; 43 events), and to 6.6% for Neogene events (95% CI: 2.5–11.5%; 8 events).

Before discussing the potential implications of our findings, it is important to first consider the effects of typological error. In the absence of data to directly estimate typological error, we estimate the total errors based on the maximum observed range overlap for ancestor–descendant pairs shown to be the result of anagenesis. Our estimates are necessarily crude given the paucity of detailed investigations on particular speciation events (Table S1). However, even if we allow the total errors to exceed our crude estimates by 50% (0.6 Ma range overlap for the Neogene, 2.4 Ma overlap for the Paleogene), our central conclusion that cladogenesis is the predominant mode is qualitatively unaffected (Fig. 3). Thus, our study highlights that range overlap may aid in distinguishing cladogenesis from anagenesis, particularly when combined with other forms of evidence. It also highlights the need for further inquiry into the issue of typological error, which will require detailed morphometric investigations (e.g., refs. 50, 55, 56).

Our estimates for the fractions of speciation events that are potentially anagenetic is quantitatively consistent with that of Wagner and Erwin (14), who estimated that anagenesis comprised at most 12.5% of Foraminifera speciation events using a phylogeny that was much smaller (40 morphospecies of globigerinids) and less precisely dated than the one used here. It is also consistent with more recent, detailed analyses of putative anagenetic events, which have subsequently been shown to result from cladogenesis. Hull and Norris (21), for example, demonstrated that the sequence between *G. plesiotumida* and *G. tumida*—which was once upheld as an example of phyletic gradualism (19)—included a third previously unrecognized morphotype, which arose through cladogenesis and gave rise to *G. tumida* over an exceptionally short time period (~45,000 y). Our analysis also identifies the emergence of *G. tumida* as a cladogenetic event, because *G. plesiotumida* persists at sites



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