

ORIGINAL ARTICLE

Interactions between Thaumarchaea, *Nitrospira* and methanotrophs modulate autotrophic nitrification in volcanic grassland soil

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Ammonium/ammonia is the sole energy substrate of ammonia oxidizers, and is also an essential nitrogen source for other microorganisms. Ammonia oxidizers therefore must compete with other soil microorganisms such as methane-oxidizing bacteria (MOB) in terrestrial ecosystems when ammonium concentrations are limiting. Here we report on the interactions between nitrifying communities dominated by ammonia-oxidizing archaea (AOA) and *Nitrospira*-like nitrite-oxidizing bacteria (NOB), and communities of MOB in controlled microcosm experiments with two levels of ammonium and methane availability. We observed strong stimulatory effects of elevated ammonium concentration on the processes of nitrification and methane oxidation as well as on the abundances of autotrophically growing nitrifiers. However, the key players in nitrification and methane oxidation, identified by stable-isotope labeling using ¹³CO₂ and ¹³CH₄, were the same under both ammonium levels, namely type 1.1a AOA, sublineage I and II *Nitrospira*-like NOB and *Methylomicrobium*-/*Methylosarcina*-like MOB, respectively. Ammonia-oxidizing bacteria were nearly absent, and ammonia oxidation could almost exclusively be attributed to AOA. Interestingly, although AOA functional gene abundance increased 10-fold during incubation, there was very limited evidence of autotrophic growth, suggesting a partly mixotrophic lifestyle. Furthermore, autotrophic growth of AOA and NOB was inhibited by active MOB at both ammonium levels. Our results suggest the existence of a previously overlooked competition for nitrogen between nitrifiers and methane oxidizers in soil, thus linking two of the most important biogeochemical cycles in nature.

The ISME Journal (2014) 8, 2397–2410; doi:10.1038/ismej.2014.81; published online 23 May 2014

Subject Category: Microbial ecology and functional diversity of natural habitats

Keywords: ammonia-oxidizing archaea; methane-oxidizing bacteria; nitrification; methanotrophy; mixotrophy; stable-isotope probing

Introduction

Ammonia-oxidizing archaea and bacteria (AOA and AOB, respectively) play a key role in the global nitrogen cycle. They convert relatively immobile ammonium into more mobile nitrite that is then further oxidized into nitrate by nitrite-oxidizing bacteria (NOB). Thereby, these two functional microbial guilds (that is, AOA/AOB and NOB), each referred to as nitrifiers, produce a reactive form of inorganic nitrogen with numerous environmental consequences (Gruber and Galloway, 2008). In nitrogen-limited soils, AOA and AOB must compete

with ammonium-assimilating organisms such as plants and heterotrophic microorganisms, and it has been shown that AOB are poorer competitors for limiting amounts of ammonium in grassland soils (Verhagen *et al.*, 1995). Only recently, indications of interactions between AOA and other microorganisms, which are relevant to both structure and activity of archaeal ammonia oxidizer communities, have been predicted in the ocean and described for a geothermal spring (Boyett *et al.*, 2013; Hamilton *et al.*, 2014). However, the ability of AOA to compete with ammonium-assimilating organisms in ammonium-limited, terrestrial environments has not been studied.

The necessity of AOA to compete for ammonium may vary between species, as the cultured soil AOA ‘*Ca. Nitrososphaera gargensis*’ and ‘*Ca. Nitrosotalea devanaterrea*’ have a preference for low ammonium levels, whereas *Nitrososphaera viennensis* EN76 and ‘*Ca. Nitrosoarchaeum koreensis*’ seem to be

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Received 14 January 2014; revised 10 April 2014; accepted 15 April 2014; published online 23 May 2014

adapted to substantially higher ammonium concentrations (Hatzenpichler *et al.*, 2008; Jung *et al.*, 2011; Lehtovirta-Morley *et al.*, 2011; Tourna *et al.*, 2011). Considering that the archaeal ammonia-monooxygenase (AMO) is reported to have a higher affinity for its substrate than known for most other microbial enzymes (Martens-Habbena *et al.*, 2009; Jung *et al.*, 2011; Lehtovirta-Morley *et al.*, 2011; Kim *et al.*, 2012) and ammonium uptake is likely facilitated by high densities of $\text{NH}_4^+/\text{NH}_3$ transporters (Urakawa *et al.*, 2011), AOA are possibly strong competitors for ammonium. However, it is currently unclear whether their competitive strength and ammonia oxidation activity varies depending on the availability of ammonium. Field and microcosm studies show either enhanced abundance and activity of AOA or inhibition in response to nitrogen (N) fertilization (Di *et al.*, 2009; Stopnisek *et al.*, 2010; Pratscher *et al.*, 2011; Verhamme *et al.*, 2011; Levicnik-Hofferle *et al.*, 2012; Ke *et al.*, 2013).

Interactions between nitrogen- and methane-cycling microbes are very likely, as the process of methane oxidation and the growth of methane-oxidizing bacteria (MOB) are influenced by ammonium (Bodelier and Laanbroek, 2004; Bodelier, 2011). Furthermore, MOB share many physiological, structural and ecological features with AOA such as dependence on monooxygenase reactions catalyzed by enzymes of the copper-containing membrane-bound monooxygenase superfamily, intracellular membrane systems, sensitivity to the same inhibitors, possession of hydroxylamine oxidoreductase systems, reliance on copper and the occupancy of oxic environments (Stein *et al.*, 2012).

Whether MOB are inhibited or stimulated by ammonium in soil largely depends on the ratio of methane-C to ammonium-N and on the genetic potential of the MOB community to deal with ammonia directly as well as indirectly via the ammonia oxidation products, hydroxylamine and nitrite (Stein *et al.* 2012). Given a steady methane availability enabling strong microbial growth, such as in wetlands and other high methane environments, ammonium is more likely to serve as a nutrient than as an inhibitor to methanotrophy (Bodelier and Laanbroek, 2004). In environments with such conditions, the regulation of carbon (C) and N cycling is tightly linked to nitrifiers and MOB competing for mineral N as a source of growth and/or metabolic activity. The balance in activity between these three functional guilds (that is, ammonia, nitrite and methane oxidizers) will determine whether N will be lost through gaseous N oxides and mineral leaching (Lipschultz *et al.*, 1981; Freitag and Bock, 1990; Stein and Klotz, 2011; Jung *et al.*, 2013) and whether the soil is a source or sink of methane. Moreover, competition for ammonium may be an important modulator of niche differentiation and the selection of nitrifier and MOB ecotypes in oxic, N-limited soils. Despite important ecological implications for ecosystem functioning that are

tied to the coupling of C and N cycles in soil, we have a rather limited understanding of the involved microorganisms and how they interact *in situ*.

Hypothesizing that ammonium-assimilating MOB inhibit nitrification by competition with AOA depending on the amount of generally available ammonium, we performed a microcosm study with two controlled levels of ammonium. By incubating a soil with limited ammonium availability and in which no AOB could be detected (Daebeler *et al.*, 2012) we utilized an ideal model system to study the importance of AOA for nitrification and their (competitive) interactions with other microorganisms. The nitrifier and MOB communities were targeted by a combination of stable-isotope probing (SIP) and deep sequencing. This approach allowed us to determine the autotrophically growing nitrifiers and active methanotrophs and enabled us to study (competitive) interactions and niche adaptation of nitrifiers and methanotrophs to different ammonium levels.

Materials and methods

Soil microcosms

The effects of ammonium concentration and methane availability on nitrification, methane oxidation and the respective ammonia- and methane-oxidizer community dynamics were determined in microcosms containing grassland soil taken from Grændalur valley (64° 1' 7" N, 21° 11' 20" W), Iceland, in August 2012. The soil is a Histic Andosol with two tephra layers in the top 30 cm, has a pH of 6.7 ± 0.1 (as determined in replicated 1:5 soil/demineralized water solutions on day 0 of the incubation) and an *in situ* ammonium content of 0.5–1.1 $\mu\text{g NH}_4^+\text{-N}$ per g dry soil (determined in August 2010 and May 2012 as described below). It has experienced stable, elevated temperatures of 33 °C through geothermal activity from at least 2005 and is cooling down since an earthquake nearby in 2009 with a recorded temperature of 14 °C in August 2012. Details of sampling and determination of water content and water-holding capacity can be found in the Supplementary Information.

The incubation was conducted in triplicate microcosms consisting of sterile 250 ml bottles containing 14 g of moist, sieved soil (corresponding to 5 g of dry soil). Loosely capped bottles were preincubated for 14 days to reduce the flux of CO_2 released from soil due to respiration. Conditions of the preincubation are given in the Supplementary Information.

On day 14 of preincubation, microcosms were amended with ammonium sulfate solutions to final additions of 15 (further referred to as LAT or low ammonium treatment) and 150 $\mu\text{g NH}_4^+\text{-N}$ per g dry soil (further referred to as HAT or high ammonium treatment). We determined the ammonia losses due to adsorption to soil particles with a ^{15}N -tracing

approach (see Supplementary Information) and could therefore estimate that ammonia availability was elevated by the ammonium treatments by at least 4.5 and 45 $\mu\text{g NH}_4^+\text{-N}$ per g dry soil in LAT and HAT microcosms, respectively, after the first addition.

Microcosms were sealed by rubber stoppers and aluminum caps and incubated for 28 days in the dark at 25 °C. HAT and LAT were refreshed weekly with ammonium sulfate solutions by addition of 7.5 and 75 $\mu\text{g NH}_4^+\text{-N}$ per g dry soil to the LAT and HAT microcosms, respectively, on days 7, 14 and 21 by injection through the rubber stopper. Soil moisture content was monitored by weight loss and kept constant at 70% water-holding capacity by addition of demineralized water or respective volumes of ammonium sulfate solutions.

The soil microcosms were amended with 5% (vol/vol) $^{12}\text{C-CO}_2$ or $^{13}\text{C-CO}_2$ (99% at excess, Campro Scientific BV, Veenendaal, The Netherlands) by injection through the rubber stopper on days 0, 7, 14 and 21. Before every gas addition microcosms were flushed with synthetic air (21% O_2 in N_2) for 5 min to ensure oxic conditions. Half of the bottles were additionally amended with 1% (10 000 p.p.m.v.) $^{12}\text{C-CH}_4$ or $^{13}\text{C-CH}_4$ (99% at excess, Campro Scientific BV) after 14 days of incubation. For an overview of the amendments made to the microcosms, see Supplementary Table A in the Supplementary Information. Microcosms receiving both $^{13}\text{C-CH}_4$ and $^{13}\text{C-CO}_2$ from day 14 were amended with $^{12}\text{C-CO}_2$ only until day 14 to ensure that a ^{13}C -signal in genomic DNA would only be found in microbes active in the presence of methane. Concentrations of methane and carbon dioxide in the headspace were monitored regularly as described in the Supplementary Information.

Triplicate LAT and HAT microcosms with either $^{12}\text{C-CO}_2$ or $^{13}\text{C-CO}_2$ were sampled destructively after 0, 14, 21 and 28 days of incubation. In the same manner, triplicate LAT and HAT microcosms with either $^{12}\text{C-CO}_2$ and $^{12}\text{C-CH}_4$ or $^{13}\text{C-CO}_2$ and $^{13}\text{C-CH}_4$ were sampled at day 28. At sampling, soil for extraction and analysis of DNA was mixed and immediately stored at -20°C . The remaining soil of each microcosm (~ 12 g) was taken for measurements of inorganic nitrogen as described in the Supplementary Information.

Nucleic acid extraction and SIP gradient fractionation

Total nucleic acids were extracted from 0.5 g soil according to Lueders *et al.* (2004) and the concentration was determined using a NanoDrop ND 1000 (Thermo Fisher Scientific, Wilmington, DE, USA) before storing at -20°C until further analysis of DNA.

Density gradient centrifugation and gradient fractionation was performed as described in Lu and Jia (2013) with 6 μg DNA extracted from duplicate LAT and HAT microcosms that had been incubated with $^{13}\text{CO}_2$ for 14 and 28 days without methane as well as from $^{13}\text{CO}_2$

+ $^{13}\text{CH}_4$ and $^{12}\text{CO}_2$ + $^{12}\text{CH}_4$ microcosms amended with methane on day 14 and sampled on day 28.

Quantitative PCR

Quantification of archaeal *amoA* genes was performed as described previously (Daebeler *et al.*, 2012). Primers used for all quantitative PCR (qPCR) assays as well as details of cycling conditions and standards used can be found in the Supplementary Information. For quantification of AOB and MOB, the bacterial *amoA* and the *pmoA* genes were targeted. NOB-based qPCR assays were not performed because of polyphyly of NOB, phylogenetically nonconserved copy numbers of the nitrite-reductase gene cluster and nonspecificity of primers. For both the archaeal and the bacterial *amoA* assays, the detection limit of qPCR was 2.85×10^3 per g dry soil; for *pmoA* it was 4.35×10^3 per g dry soil.

Pyrosequencing and sequence analysis

To characterize microbial communities in the gradient fractions from ^{13}C and from ^{12}C -control microcosms obtained after density gradient centrifugation, we carried out barcoded pyrosequencing of fractions 3 to 12 of each gradient targeting the 16S rRNA gene using a Roche 454 GS FLX Titanium sequencer (Roche Diagnostics Corporation, Branford, CT, USA) as reported previously (Xia *et al.*, 2011; Lu and Jia, 2013). The raw, demultiplexed SFF files were deposited into the EMBL-ENA SRA database (<https://www.ebi.ac.uk/ena/>) and can be found under study accession number PRJEB5248 at the following site: <http://www.ebi.ac.uk/ena/data/view/PRJEB5248>.

A total of 1 138 939 raw sequences were processed with mothur v1.31.1 (Schloss *et al.*, 2009) according to the standard operating procedure published by Schloss *et al.* (2011; detailed description accessible at http://www.mothur.org/wiki/454_SOP#OTUs) until the extraction of representative sequences for each operational taxonomic unit (OTU). Only the alignment was performed differently than described by Schloss *et al.* (2011). Sequence alignment was generated with SINA (<http://www.arb-silva.de/aligner/>) and the alignment was then imported back to mothur for subsequent analysis steps according to the mothur standard procedure referenced above. OTUs were clustered with mothur using the average neighbor algorithm at 99% sequence divergence. The OTUs containing < 50 sequences over the entire data set were discarded to increase computation speed. Finally, we standardized to 5426 sequences per sample, obtaining a final number of 868 160 sequences of which 254 676 were unique. Details of the phylogenetic analysis are given in the Supplementary Information.

Network analysis was performed using the R-package qgraph (Epskamp *et al.*, 2012) based on a Spearman's rank correlation matrix generated for

the 10 most abundant AOA, NOB and MOB phylotypes in samples from LAT and HAT $^{13/12}\text{C}$ microcosms incubated with and without methane. Only correlation pairs with a ρ -value above 0.6 and significance level below 0.05 were used. The network layout was generated using a modified version of the force-directed algorithm proposed by Fruchterman and Reingold (1991) as implemented in qgraph.

Statistical analysis

To test for effects of treatments and time as well as their interaction on AOA and MOB abundance, methane oxidation rate and mineral N concentrations, we performed analysis of variance (ANOVA) followed by Tukey's HSD (honest significant difference) *post hoc* tests. Abundance data of AOA and MOB genes were log-transformed to meet homoscedasticity assumptions. We calculated the balance between expected and observed nitrification by subtracting the total amount of accumulated NO_3^- -N on day 28 from all added and initially present NH_4^+ -N. Differences in community compositions of nitrifiers and methane oxidizers based on Bray-Curtis distances of weighted phylotype profiles were assessed using the *adonis* function of the R package *vegan* v2.0-8 (Oksanen *et al.*, 2010) that compares dissimilarities of variance in groups and uses permutation tests to assess significance.

Results

Nitrification and methane oxidation in soil microcosms
Concentrations of ammonium in the soil microcosms were barely measurable throughout the incubation, except for initial concentrations because of ammonium additions to the LAT and HAT microcosms (Figure 1a). Concentrations close to zero were measured for days 14, 21 and 28 in all microcosms before the weekly reapplication of ammonium took place.

Significant net nitrification assessed via production of $\text{NO}_2^- + \text{NO}_3^-$ was only observed in HAT and HAT + CH_4 microcosms over the time frame of the entire incubation (Figure 1b; Tukey's HSD, $p_{\text{day}0:28} < 0.001$, and 0.033 respectively). Net nitrification was significantly affected by the ammonium treatment (ANOVA, $p_{\text{low:high}} = 0.013$) but methane-amended microcosms did not differ from nonmethane-amended microcosms of the same ammonium treatment. Both LAT microcosms without and with methane showed a negative balance between expected nitrification and measured nitrification by -5.3 and $-0.9 \mu\text{g N per g dry soil}$, respectively. In contrast, the balance between expected and measured nitrification in HAT microcosms without and with methane was positive with 336.2 and $393.0 \mu\text{g N per g dry soil}$, respectively.

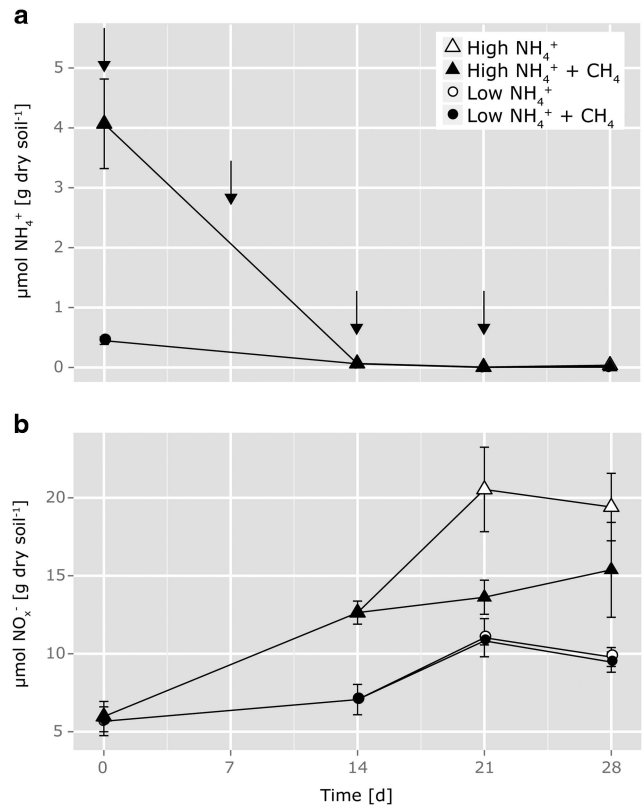


Figure 1 Changes in ammonium (a) and nitrite plus nitrate (b) concentrations in soil microcosms incubated at 25 °C for 0, 14, 21 or 28 days at low ($15 \mu\text{g NH}_4^+$ -N per g dry soil added initially; $7.5 \mu\text{g NH}_4^+$ -N per g dry soil on days 7, 14 and 21) and high ($150 \mu\text{g NH}_4^+$ -N per g dry soil added initially; $75 \mu\text{g NH}_4^+$ -N per g dry soil on days 7, 14 and 21) levels of ammonium and concentrations of 0% and 1% methane in the headspace. Arrows in (a) indicate ammonium addition. Data plotted are mean values and s.e. from 3 to 6 replicate microcosms destructively sampled at each time point, with some symbols nonvisible because they are masked by others.

Consumption of methane was observed upon the first addition of methane on day 14 with rates of $161.3 (\pm 13.3)$ and $146.37 (\pm 23.98)$ nmol oxidized methane per g dry soil per h after a lag time of 20 h and 50 h in the LAT and HAT microcosms, respectively (Figure 2). With the repeated addition of 1% methane to the headspace, methane oxidation rates increased throughout the 14 subsequent days of incubation and were significantly higher in the HAT than in the LAT microcosms (ANOVA, $p_{\text{low:high}} = 0.003$) reaching values of $3702.3 (\pm 134.6)$ and $4058.5 (\pm 192.4)$ nmol oxidized methane per g dry soil per h, respectively, at the end of the incubation.

Quantification of ammonia- and methane-oxidizing communities

Archaeal *amoA* abundance increased significantly only in HAT + CH_4 microcosms (Tukey's HSD, $p_{\text{day}0:28} = 0.006$; Figure 3). The addition of methane significantly elevated archaeal *amoA* gene abundance (ANOVA, $p_{\text{CH}_4: + \text{CH}_4} = 0.025$) and we found

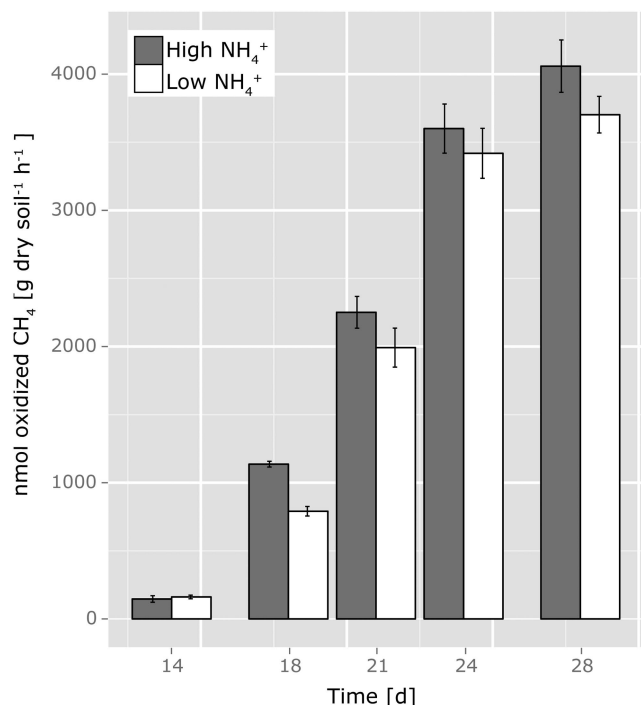


Figure 2 Methane oxidation rates on days 14, 18, 21, 24 and 28 of microcosms incubated at 25 °C at low (15 µg NH₄⁺-N per g dry soil added initially; 7.5 µg NH₄⁺-N per g dry soil on days 7, 14 and 21) and high (150 µg NH₄⁺-N per g dry soil added initially; 75 µg NH₄⁺-N per g dry soil on days 7, 14 and 21) levels of ammonium and repeated addition of 1% methane to the headspace starting on day 14. Methane addition was repeated every 3 to 4 days. Data plotted are mean values and s.e. from triplicate nonlabeled microcosms.

significantly more *amoA* genes in high than in low ammonium microcosms (ANOVA, $p_{\text{high:low}} < 0.001$). Ammonia-oxidizing bacteria were undetectable by qPCR.

The community size of methane-oxidizing bacteria was not significantly different at the start of the incubation between LAT and HAT microcosms (Figure 4). In the absence of methane, *pmoA* gene abundance dropped below detection limit in LAT microcosms, whereas the abundance in HAT microcosms on day 28 did not differ significantly from day 0. In the presence of methane, *pmoA* gene abundance drastically increased within 2 weeks in the LAT microcosms from below detection limit on day 14, the day of the first methane amendment, to 1.13×10^6 copies per g dry soil on day 28. The *pmoA* gene abundance in HAT + CH₄ microcosms reached a rather similar level constituting significant growth upon methane addition (Tukey's HSD, $p_{\text{day14:28}} = 0.039$).

DNA-SIP analysis of nitrifying and methane-oxidizing communities

In order to reveal which part of the nitrifying and methane-oxidizing community was actively growing

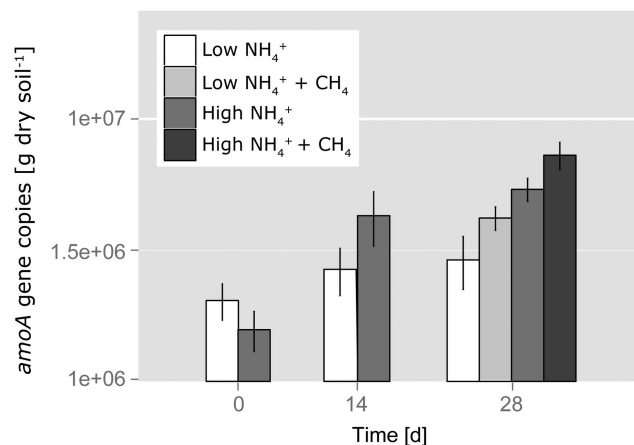


Figure 3 Changes in archaeal *amoA* gene copy numbers in soil microcosms incubated at 25 °C for 0, 14 and 28 days at low (15 µg NH₄⁺-N per g dry soil added initially; 7.5 µg NH₄⁺-N per g dry soil on days 7, 14 and 21) and high (150 µg NH₄⁺-N per g dry soil added initially; 75 µg NH₄⁺-N per g dry soil on days 7, 14 and 21) levels of ammonium as well as 0% and 1% methane in the headspace. Repeated methane addition every 3 to 4 days was started on day 14. Data plotted are mean values and s.e. from triplicate nonlabeled microcosms destructively sampled at each time point.

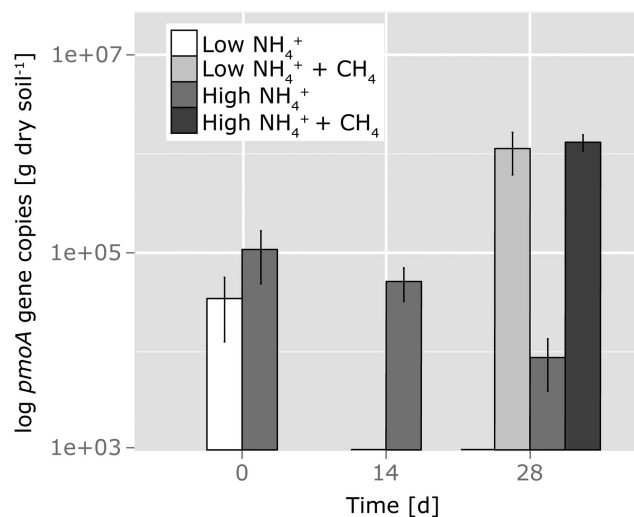


Figure 4 Changes in bacterial *pmoA* gene copy numbers in soil microcosms incubated at 25 °C for 0, 14 and 28 days at low (15 µg NH₄⁺-N per g dry soil added initially; 7.5 µg NH₄⁺-N per g dry soil on days 7, 14 and 21) and high (150 µg NH₄⁺-N per g dry soil added initially; 75 µg NH₄⁺-N per g dry soil on days 7, 14 and 21) levels of ammonium as well as 0% and 1% methane in the headspace. Repeated methane addition every 3 to 4 days was started on day 14. Data plotted are mean values ($n=2$ or 3) and s.e. from nonlabeled microcosms destructively sampled at each time point.

by assimilating CO₂/CH₄, we separated 'heavy' (¹³C-enriched) DNA from 'light' (¹²C) DNA by isopycnic, CsCl gradient ultracentrifugation of DNA from microcosms incubated with ¹³CO₂ for 14 and 28 days as well as from microcosms incubated with ¹³CO₂ and ¹³CH₄ from day 14 to 28, and their respective ¹²C controls (Figures 5 and 6).

Total active communities. Relative abundances of AOA, AOB, NOB and MOB 16S rRNA genes in purified fractions 3 to 12 of the 15 CsCl gradient fractions were determined by pyrosequencing. The 16S rRNA gene sequences of AOB were mostly not detectable in the gradient fractions and if present never constituted >0.68% of the total number of sequence reads in any fraction, whereas AOA and NOB reads were found in all fractions and comprised up to 9.13% and 12.75% of total reads in a fraction, respectively (Figure 5). Because of the rareness of detected AOB sequences, no specific pattern of distribution along the CsCl gradients could be obtained, and consequently we did not find evidence for the incorporation of ^{13}C into genomic DNA of AOB.

The relative numbers of AOA-related 16S rRNA genes tended to increase in fractions above 1.73 g CsCl per ml in all microcosms without methane amendment, irrespective of the ammonium level applied. This distribution pattern was rather weak and remained at a low level of AOA-related 16S rRNA gene enrichment in the 'heavy' end of the CsCl gradient. We did however find a clear enrichment of NOB-related 16S rRNA genes in all 'heavier' fractions of the CsCl gradients (Figures 5a, b, d and e). In addition, we could detect different

distributions of NOB-related 16S rRNA genes between the two ammonium treatment levels and sampling days with a trend to a larger, actively CO_2 -incorporating NOB community in the HAT microcosms as seen by a peak shift in relative sequence abundance toward the 'heavy' end of the gradient, especially for day 14 (Figure 5d).

In the presence of methane, we could not detect an increase in relative 16S rRNA gene sequence abundances along the gradient for either of the nitrifiers (Figures 5c and f). In contrast, we observed a large enrichment of relative MOB 16S rRNA gene abundance in the fractions above 1.73 g CsCl per ml from the ^{13}C -labeled microcosms, but not from ^{12}C microcosms (Figures 5c and f), indicative of a very active, methanotrophic community. Furthermore, there was an unmistakable effect of ammonium treatment level on MOB as apparent from an increase of ^{13}C -labeled MOB DNA in the 'heavy' fractions by nearly one-third in the microcosms with high compared with those with low NH_4^+ treatment (Figures 5c and f).

In addition, we quantified archaeal *amoA* and bacterial *pmoA* genes in purified CsCl fractions 2 to 13 by qPCR which confirmed the pyrosequencing results and is shown in Supplementary Figure A in the Supplementary Information.

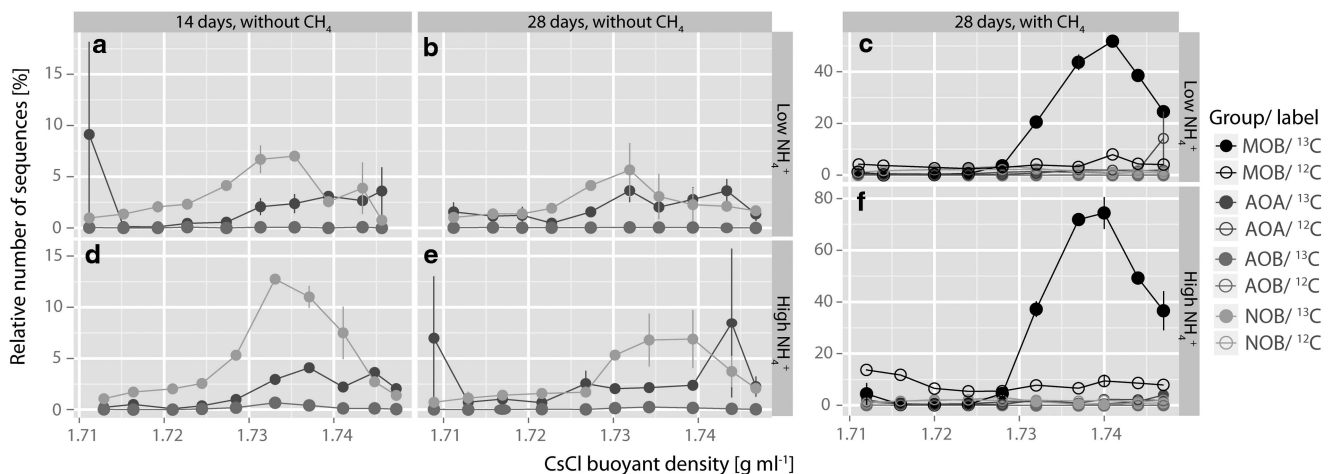
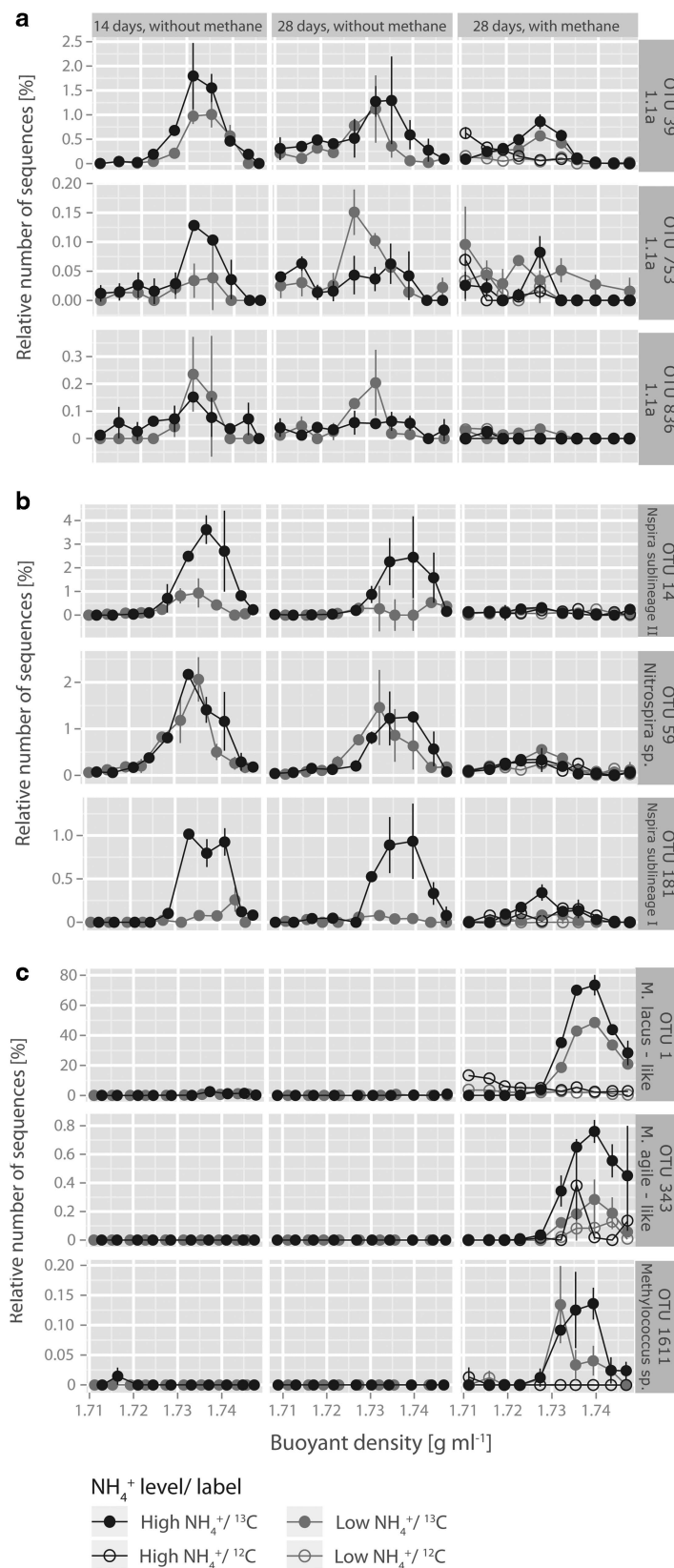


Figure 5 Average relative abundances of AOA, AOB, NOB and MOB 16S rRNA genes in fractionated DNA from microcosms regularly amended with low ((a–c) $15 \mu\text{g NH}_4^+ \text{-N}$ per g dry soil added initially; $7.5 \mu\text{g NH}_4^+ \text{-N}$ per g dry soil on days 7, 14 and 21) and high ((d–f); $150 \mu\text{g NH}_4^+ \text{-N}$ per g dry soil added initially; $75 \mu\text{g NH}_4^+ \text{-N}$ per g dry soil on days 7, 14 and 21) levels of ammonium and 5% $^{13}\text{CO}_2$ sampled on day 14 (a, d) and day 28 (b, c, e, f) of incubation. Microcosms were incubated in the absence (a, b, d, e) and presence (c, f) of 1% $^{13}\text{CH}_4$ from day 14 onwards. The relative frequencies are expressed as the percentage of 16S rRNA genes from the respective functional group to total 16S rRNA genes of each DNA fraction. Data plotted are mean values and s.e. from DNA obtained from gradient fractions of duplicate nonlabeled ^{12}C (open symbols) and labeled ^{13}C (closed symbols) microcosms.

Figure 6 Changes in the relative frequencies of 16S rRNA gene sequence reads related to the three most abundant and labeled phylotypes of the archaeal ammonia-oxidizing (a), nitrite-oxidizing (b) and methane-oxidizing (c) community in fractionated DNA from microcosms regularly amended with low ($15 \mu\text{g NH}_4^+ \text{-N}$ per g dry soil added initially; $7.5 \mu\text{g NH}_4^+ \text{-N}$ per g dry soil on days 7, 14 and 21) and high ($150 \mu\text{g NH}_4^+ \text{-N}$ per g dry soil added initially; $75 \mu\text{g NH}_4^+ \text{-N}$ per g dry soil on days 7, 14 and 21) levels of ammonium sampled on days 14 and 28. Microcosms were incubated with and without 1% methane in the headspace for the last 14 days of incubation. The relative frequencies are expressed as the percentage of 16S rRNA genes from the respective OTU to total 16S rRNA genes of each DNA fraction. Data plotted are mean values and s.e. from DNA obtained from gradient fractions of duplicate nonlabeled ^{12}C (open symbols) and labeled ^{13}C (closed symbols) microcosms.



Community composition and active nitrifier and methanotroph phylotypes. The most abundant AOA phylotype in all samples fell within the 1.1a-associated *Nitrosotalea* cluster of the Thaumarchaeal phylum (Supplementary Figures B and E). In addition, we found four phylotypes belonging to the 1.1a cluster as well as five phylotypes of the 1.1b cluster. Generally, the 1.1b phylotypes were of much lower relative abundance than the 1.1a phylotypes. The AOA community composition in 'heavy' and 'light' fractions was significantly different (adonis, $R^2=0.17$, $p=0.025$). In addition, 'heavy' fractions from day 28 of methane-amended ^{13}C -microcosms were significantly different from ^{12}C -microcosms of the same day (adonis, $R^2=0.29$, $p=0.034$). We found no significant difference between AOA communities in 'heavy' fractions of replicates between the two ammonium treatment levels, time points or between methane-amended and non-amended microcosms.

For the majority of AOA phylotypes there was no clear evidence for incorporation of CO_2 as they showed similar abundances in the 'heavy' fractions from ^{13}C and ^{12}C microcosms. Only three phylotypes of the 1.1a cluster (that is, 39, 753 and 836) were enriched in the 'heavy' end of the CsCl gradient from ^{13}C but not ^{12}C microcosms (Figure 6a), although at a rather low level of maximally 1.80% relative sequence abundance. Their distribution pattern is indicative of active CO_2 incorporation and was significantly lower, but not absent in samples from methane-amended microcosms. For all three phylotypes, there was no significant difference of relative abundance in 'heavy' fractions between the LAT and HAT microcosms, except in the incubations without methane on day 28. Here, the enrichment of phylotypes 753 and 836 in 'heavy' fractions was larger in samples from LAT microcosms.

The low abundance of AOB in our microcosms was coupled to a low diversity. We only obtained three AOB phylotypes in total. They were affiliated with *Nitrosomonas oligotropha* and a *Nitrospira*-associated cluster (Supplementary Figure F). Community analysis of AOB based on relative abundances was not possible because of their absence from most samples and the extremely low sequence counts in the samples where we did detect them.

Analysis of phylotype-based community composition of NOB showed that the whole community consisted exclusively of the genus *Nitrospira*. It was foremost a *Nitrospira* sublineage II-, a *Nitrospira* sp.- and a *Nitrospira* sublineage I-like phylotype that were enriched in the 'heavy' fractions from the ^{13}C microcosms (Figure 6b and Supplementary Figures C and G). These three NOB phylotypes strongly dominated the active autotrophic NOB community (Supplementary Figure C). There was no significant difference in NOB community composition between replicated 'heavy' fractions, but there was a significant difference between NOB communities in 'heavy' and 'light'

fractions (adonis, $R^2=0.33$, $p=0.001$), showing that NOB did not only actively incorporate CO_2 in our microcosm incubations, but that the autotrophically active NOB community was different from the nonautotrophically active one. This difference was significantly related to the ammonium treatment level (adonis, $R^2=0.31$, $p=0.003$), the methane amendment (adonis, $R^2=0.27$, $p=0.003$) as well as the time of sampling (adonis, $R^2=0.15$, $p=0.006$). These factors jointly explained 74% of the observed variance. There was no significant difference in NOB communities between 'heavy' fractions of ^{12}C control and ^{13}C -labeled, methane-amended microcosms on day 28 of the incubation, but a large difference in 'heavy' fractions between methane-amended and nonamended microcosms of the same day. Moreover, we observed a substantial and significant decrease in relative abundance of the actively CO_2 -incorporating NOB phylotypes 14, 59 and 181 in the 'heavy' end of the CsCl gradient in samples from methane-amended microcosms as compared with microcosms without methane (Figure 6b).

The methanotrophic community comprised eight phylotypes with the two most dominant ones closely related to *Methylosarcina lacus* (Supplementary Figures D and H). Adonis analysis of variances for MOB communities showed that there was no significant difference between replicates, but a significant difference between communities in 'heavy' fractions from microcosms with and without methane amendment (adonis, $R^2=0.87$, $p=0.002$) and between communities of ^{12}C and ^{13}C microcosms in the presence of methane on day 28 (adonis, $R^2=0.84$, $p=0.014$). Furthermore, MOB communities in 'heavy' fractions differed significantly from the ones in 'light' fractions in microcosms to which 1% ^{13}C -methane was added (adonis, $R^2=0.57$, $p=0.019$), whereas there was no difference in community composition related to ammonium level. It was foremost a single *M. lacus*-like MOB phylotype that was responsible for these results and that was enriched from 0 to up to 73.40% relative abundance in 'heavy' fractions of samples from methane-amended microcosms (Figure 6c).

Network analysis

In order to elucidate interactions between AOA, NOB and MOB in our microcosms, we performed a network analysis based on co-occurring phylotypes of the three functional guilds in all incubations. The results show exclusively positive, stronger and more frequent correlations between the AOA and NOB than between MOB and one of the other guilds (Figure 7). The actively CO_2 -incorporating AOA phylotypes 39, 753 and 836 of the 1.1a cluster (see Figure 6a) were all strongly and positively correlated with NOB phylotypes, among which were the autotrophically growing phylotypes 14, 59 and 181 (see Figure 6b). Interestingly, two of these NOB

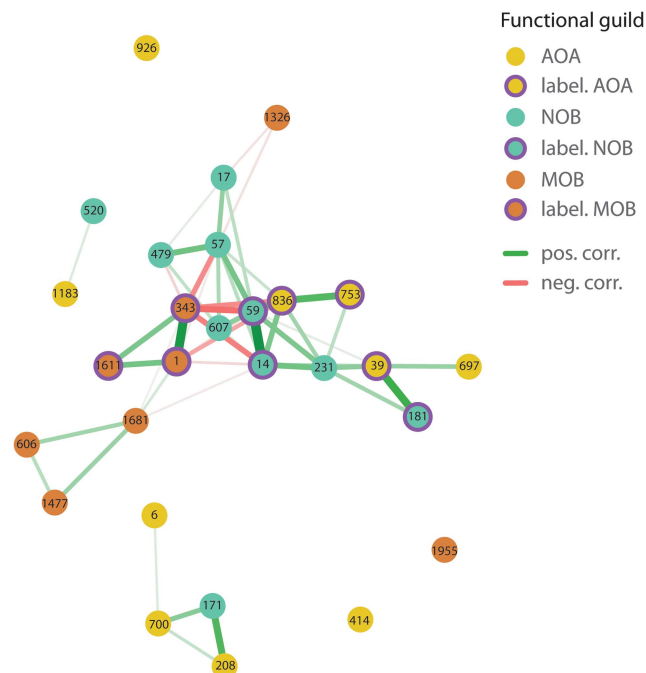


Figure 7 Network analysis of co-occurring phylotypes of AOA, NOB and MOB in microcosms incubated at low ($15 \mu\text{g NH}_4^+\text{-N}$ per g dry soil added initially; $7.5 \mu\text{g NH}_4^+\text{-N}$ per g dry soil on days 7, 14 and 21) and high ($150 \mu\text{g NH}_4^+\text{-N}$ per g dry soil added initially; $75 \mu\text{g NH}_4^+\text{-N}$ per g dry soil on days 7, 14 and 21) levels of ammonium for 14 and 28 days as well as with and without 1% methane for the last 14 days of incubation. Nodes represent the 10 most abundant OTUs of each of the three functional guilds and connecting lines stand for strong (Spearman's $\rho > 0.6$) and significant (p -value of < 0.05) positive (green) or negative (red) correlations where a thicker and shorter line represents a stronger correlation. Purple circles around nodes mark phylotypes that had actively incorporated the ^{13}C label.

phylotypes were negatively correlated to OTU 1, which is closely affiliated with *M. lacus*, the dominant and active methane-oxidizer phylotype (see Figure 6c). Another interesting correlation suggesting negative interactions is the one between the 1.1a AOA phylotype 836, whose autotrophic growth was suppressed in methane-amended microcosms, and the actively methane-oxidizing phylotypes 1 and 343. OTU 6, which fell within the *Nitrosotalea* cluster and was the most abundant phylotype of AOA in our microcosms, only showed a positive correlation with OTU 700, which is associated with the *Nitrososphaera* cluster.

Discussion

Dominance of AOA and Nitrospira among nitrifiers

The ammonia-oxidizing community was clearly dominated by archaea as *amoA*-targeted qPCR as well as pyrosequencing of the total microbial community could demonstrate. Although we cannot exclude that AOB are a substantial fraction of the rare biosphere, this confirms the finding of a

previous study conducted with soils of the same geothermal area (Daebeler *et al.*, 2012).

An adaptation to low substrate best explains the numerical dominance of AOA in this soil that has the typically high cation exchange capacity of an Andosol, resulting in low ammonia availability. AOA of the 1.1a and the *Nitrosotalea* lineages, which comprise the vast majority of the AOA community detected in our incubations, are often found under oligotrophic conditions and may be adapted to low ammonium environments (Martens-Habbena *et al.*, 2009; Martens-Habbena and Stahl, 2011; French *et al.*, 2012; Nakagawa and Stahl, 2013; Gorman-Lewis *et al.*, 2014). *Nitrosotalea*-like AOA have been demonstrated to be adapted to low pH environments, but are also abundant among ammonia oxidizers in many other terrestrial ecosystems and not exclusively found in low pH soils (Leininger *et al.*, 2006; Gubry-Rangin *et al.*, 2011; Lehtovirta-Morley *et al.*, 2011; Pester *et al.*, 2011). Their relatively high abundance in our soil may hence reflect their adaptation to low substrate availability rather than to low pH itself. On the other hand, cultivated members of the 1.1b lineage, which were a minority among the AOA observed in our incubations, have been shown to prefer higher substrate concentrations (Jung *et al.*, 2011; Tourna *et al.*, 2011).

Dominance of *Nitrospira* sp., the proposed k-strategist nitrite oxidizer with high substrate affinity and hence adaptation to low nitrite concentrations (Ehrich *et al.*, 1995; Schramm *et al.*, 1999; Koops and Pommerening-Röser, 2001; Attard *et al.*, 2010), may provide further evidence of selection for low substrate-adapted nitrifiers in this volcanic grassland soil.

Activity and growth of nitrifiers and methane oxidizers

We observed an increase in archaeal *amoA* gene abundance in HAT microcosms without and with methane supply. This demonstrates that AOA were indeed actively growing despite the fact that evidence for autotrophic growth was only found at a rather low level and only for three specific 1.1a AOA phylotypes. These results of limited autotrophic growth by AOA are similar to those of Jia and Conrad (2009) and Pratscher *et al.* (2011) who concluded AOB to be more important for ammonia oxidation. However, in our soil AOB represent such low numbers that they are unlikely responsible for the observed ammonia oxidation. Accordingly, net nitrification rates were comparable to those found previously (Zhang *et al.*, 2010; Verhamme *et al.*, 2011), where the authors reported ammonia oxidation to be mainly driven by AOA. The detected CO_2 incorporation by 1.1a AOA phylotypes, which jointly represented 23% of the total AOA community, suggests that at least some AOA did indeed grow autotrophically, likely coupled to ammonia oxidation. Such preferential activity of 1.1a AOA

confirms findings of previous studies (Tourna *et al.*, 2008; Offre *et al.*, 2009; Zhang *et al.*, 2010).

As no cultured AOA has been found to contain more than one copy of the *amoA* gene (Hallam *et al.*, 2006; Walker *et al.*, 2010; Blainey *et al.*, 2011; Spang *et al.*, 2012) we may assume that copy numbers of archaeal *amoA* can be directly related to cell abundance. This is not the case for AOB, as several AOB harbor several copies of the *amoC/amoA/amoB* gene complex (Chain *et al.*, 2003; Stein *et al.*, 2007; Norton *et al.*, 2008) and it is therefore more reliable to infer AOB cell abundance from 16S rRNA gene quantification. Even if the estimated number of labeled AOA cells (23% of the archaeal *amoA* abundance) was the same throughout the entire incubation period as on day 28 and if all AOA cells were growing coupled to ammonia oxidation with maximum specific activity rates similar to reported values for 1.1a AOA (Jung *et al.*, 2011), they could have only oxidized a maximum of $0.5 \mu\text{M NH}_3$. We however measured for example an oxidation of $67.1 \mu\text{M NH}_3$ in HAT microcosms without methane, suggesting heterotrophic or mixotrophic ammonia oxidation performed by AOA. The increase in *amoA* gene abundance suggests that the AOA community consisted of 75% newly grown cells by the end of the incubation period. Such growth, if supported autotrophically, would have been detected by DNA SIP to a much larger extent, analogous to previous studies (Zhang *et al.*, 2010; Xia *et al.*, 2011; Lu and Jia, 2013). Mixotrophy has been suggested for AOA in soil before (Jia and Conrad, 2009; Musmann *et al.*, 2011; Tourna *et al.*, 2011) and through genomic and metagenomic studies, the mechanisms and abilities of lineage 1.1a Thaumarchaeota to utilize small organic substances have been identified (Ouverney and Fuhrman, 2000; Ingalls *et al.*, 2006; Walker *et al.*, 2010; Tully *et al.*, 2012). Nevertheless, it can be imagined that some of the observed ammonia oxidation was not coupled to growth of AOA or AOB, as gene abundance was one and four orders of magnitude lower than expected from relating observed net nitrification to reported per cell activity rates for AOA and AOB, respectively (Prosser, 1989; Laanbroek and Gerards, 1993; Jung *et al.*, 2011).

Of the detected *Nitrospira* sp. only a minority was assimilating CO_2 , suggesting a high importance of specific *Nitrospira* spp. for autotrophic nitrite oxidation which is consistent with previous reports (Freitag *et al.*, 2005; Xia *et al.*, 2011). Differences in autotrophic growth of the *Nitrospira* sublineage I OTU 181 between HAT and LAT microcosms supports the report on a preference for higher nitrite concentrations of this sublineage (Maixner *et al.* (2006). In contrast we found the autotrophic activity of a *Nitrospira* sublineage II phylotype increased upon addition of more NH_4^+ . We may therefore conclude, that substrate-driven niche differentiation

exists within sublineage II *Nitrospira* and that at least some nitrite oxidizers of this sublineage are selected by elevated nitrite concentrations as well.

The methane-oxidizing community in our incubations responded rapidly and vigorously to the availability of CH_4 , suggesting methanotrophy to be an important metabolic lifestyle in the Icelandic grassland soil studied here. MOB abundance rapidly dropped in LAT but not HAT microcosms without methane. This is indicative of N limitation and nitrogen dependence, but not in relation to methanotrophy. Despite the origin of the soil material used, being top soil from a grassland site, availability of biologically derived methane from deeper layers as a substrate for MOB is highly likely as anoxic microcosm incubations produced significant amounts of methane at a rate of $1000 \text{ nmol per g dry soil per d}$ after a lag phase of 2 weeks (A. Daebeler, unpublished results). Furthermore, steams of geothermal areas in Iceland carry average methane concentrations of $0.44 \text{ mmol kg}^{-1}$ leading to estimated methane emission rates of 1300 t per year from Icelandic geothermal sites (Etiope *et al.*, 2007). In combination with the readily responsive MOB community, this suggests the mitigation of methane emission as an important ecosystem service for these grasslands.

Surprisingly, over 98% of the observed $\text{CH}_4\text{-C}$ assimilation was carried out by only a single phylotype closely related to *M. lacus*. Although we cannot exclude growth of other MOB through alternative carbon sources (Semrau *et al.*, 2011), the increase in relative abundance of this *M. lacus*-like MOB under methane-amended conditions leads to the conclusion that the observed growth of the MOB community was to a large extent because of the growth of this organism.

Putative interactions of nitrifiers and MOB

The design of our study allowed for analysis of responses of CO_2 - and CH_4 -incorporating microorganisms to low and elevated ammonium availabilities. As only HAT but not LAT microcosms showed more expected than observed nitrification, we conclude that significant amounts of ammonia were immobilized only in HAT microcosms. This further leads to the conclusion that this immobilization was likely not due to adsorption of ammonium to soil particles, but to assimilation by heterotrophic microorganisms. Therefore, we are convinced that the ammonium additions led to conditions of more and less N limitation, resulting in different forces of competition for the common nutrient.

Both nitrifying guilds studied were responsive to ammonium amendment: *amoA* gene abundances of AOA were elevated and incorporation of $\text{CO}_2\text{-C}$ by NOB was stronger in HAT than in LAT microcosms. In addition, net nitrification was stimulated by ammonium amendment, suggesting that both nitrifying guilds may be N limited under *in situ* conditions and are adapted to tolerate and utilize

higher substrate concentrations. We cannot exclude, however, that specific groups with a higher tolerance and lower affinity for ammonium than the *in situ* populations were selected by the incubation conditions.

Likely, the AOA will have transmitted the higher substrate availability in HAT microcosms to NOB by providing them with more nitrite. Conclusively, the network analysis showed that phylotypes of AOA and NOB had exclusively positive correlations. This is especially interesting in the case of autotrophically growing NOB and AOA as it may display their occupancy of the same niche space as well as the syntrophic relationship that is apparent from their energy metabolism as nitrifiers. We are therefore convinced that strong and positive links between autotrophically growing AOA of the 1.1a cluster and nitrite-oxidizing *Nitrospira* exist in the *in situ* soil environment, comparable to co-occurrence patterns of *amoA*-encoding archaea and *Nitrospina* in the ocean (Mincer *et al.*, 2007). In addition, positive co-occurrence patterns between AOA and NOB phylotypes which were not actively incorporating CO₂ leave room for the possibility of interrelations of AOA and NOB with metabolic lifestyles other than chemolithoautotrophic ammonia and nitrite oxidation.

In the presence of methane, autotrophic growth of AOA and NOB ceased at both ammonium levels, more so in case of the NOB, and the network analysis showed exclusively negative correlations between active MOB and autotrophically growing nitrifiers. Interestingly, net nitrification was not significantly affected by the addition of methane, possibly suggesting a metabolic switch of nitrifiers. However, it is possible that only nitrite oxidation was in fact inhibited by the presence of active MOB. Unfortunately, the absence of data on nitrite accumulation separately from nitrate failed to resolve this possibility. Little is known about interactions between nitrifiers and methanotrophs, but our data show inhibition of autotrophic ammonia and nitrite oxidation by an active, methane-oxidizing community coupled to a shift in community composition of AOA and NOB. This result supports our hypothesis of a competition for common substrates between active methane oxidizers and nitrifiers. Perhaps, the suppression of autotrophic nitrification by active MOB was further a direct consequence of oxygen consumption by MOB causing suboxic conditions which limited the growth of strictly autotrophic NOB, but not to the same extent the growth of AOA who are known to tolerate low oxygen concentrations (Martens-Habbena *et al.*, 2009; Park *et al.*, 2010; Jung *et al.*, 2011; Martens-Habbena and Stahl, 2011). When assuming a 1:1 CH₄/O₂ stoichiometry for methane oxidation, MOB must have consumed ~0.51 and 0.55 mmol oxygen in the LAT and HAT microcosms, respectively, within the last week of the incubation and may thus have created oxygen-limited conditions within parts of the soil matrix.

Under such low oxygen conditions, MOB equipped to use nitrite as electron acceptor alternative to oxygen as reported by Nyerges *et al.* (2010) would likely have competed with *Nitrospira* for nitrite, resulting in detrimental circumstances for the nitrite oxidizers. This effect would be more unfavorable in HAT microcosms and may have caused the observed pattern of higher methane oxidation, ceased NOB activity, but less affected AOA growth with no detectable decrease in oxidized mineral N.

Summary

In the Icelandic grassland soil studied, syntrophic communities of autotrophic 1.1a AOA and autotrophic *Nitrospira* likely drove chemolithotrophic nitrification, and mixotrophic AOA may also have contributed to nitrification. Our results challenge the perception of archaeal ammonia oxidizers and nitrite oxidizers as physiologically uniform guilds and hint at metabolic plasticity of AOA and NOB, such as active selection for different modes of growth as well as uncoupling of metabolic activity from growth. Both nitrifying guilds showed positive responses to inorganic N supply and we found indications of niche differentiation within the genus of *Nitrospira* as a result of adaptation to N availability. Likely because of strongly nitrogen-limited conditions in the soil, bacterial ammonia oxidizers constituted a negligible fraction of the microbial community and remained unresponsive throughout the incubation period. The availability of mineral nitrogen played a pivotal role not only for the regulation of nitrification and abundance activity patterns of soil nitrifiers, but also for the extent to which the soil was capable of methane oxidation. Our results show that nitrogen stimulation of active methanotrophs can lead to suppression of autotrophic nitrification and of autotrophic growth of AOA and NOB, perhaps through competition for mineral nitrogen and oxygen. This previously unrecognized competition between nitrifiers and methanotrophs may therefore cause uncoupling of the N cycle and will have to be accounted for when addressing the impacts of N deposition.

Conflict of Interest

The authors declare no conflict of interest.

Acknowledgements

We thank Bjarni D Sigurdsson and the Agricultural University of Iceland for access to the sampling site and Bas Dingemans for collecting the soil samples. We also thank Roos Keijzer and Marion Meima-Franke for excellent technical assistance. The Schure-Beijerinck-Popping fund is acknowledged for funding the stay of AD as a guest researcher in the ZJ group. This study was partially supported by the Ministry of Science and Technology of China (2010DFA22770). This is publication number 5603 of The Netherlands Institute of Ecology (NIOO-KNAW).

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