

Distribution, Activities, and Interactions of Methanogens and Sulfate-Reducing Prokaryotes in the Florida Everglades

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To gain insight into the mechanisms controlling methanogenic pathways in the Florida Everglades, the distribution and functional activities of methanogens and sulfate-reducing prokaryotes (SRPs) were investigated in soils (0 to 2 or 0 to 4 cm depth) across the well-documented nutrient gradient in the water conservation areas (WCAs) caused by runoff from the adjacent Everglades Agricultural Area. The methyl coenzyme M reductase gene (*mcrA*) sequences that were retrieved from WCA-2A, an area with relatively high concentrations of SO₄²⁻ (\geq 39 µM), indicated that methanogens inhabiting this area were broadly distributed within the orders *Methanomicrobiales*, *Methanosarcinales*, *Methanocellales*, *Methanobacteriales*, and *Methanomassiliicoccales*. In more than 3 years of monitoring, quantitative PCR (qPCR) using newly designed group-specific primers revealed that the hydrogenotrophic *Methanomicrobiales* were more numerous than the *Methanomicrobiales* in WCA-3A (with relatively low SO₄²⁻ concentrations; \leq 4 µM). qPCR of *dsrB* sequences also indicated that SRPs are present at greater numbers than methanogens in the WCAs. In an incubation study with WCA-2A soils, addition of MOO₄²⁻ (a specific inhibitor of SRP activity) resulted in increased methane production rates, lower apparent fractionation factors [α_{app} ; defined as (amount of δ^{13} CO₂ + 1,000)/(amount of δ^{13} CH₄ + 1,000)], and higher *Methanosaetaceae mcrA* transcript levels compared to those for the controls without MOO₄²⁻. These results indicate that SRPs play crucial roles in controlling methanogenic pathways and in shaping the structures of methanogen assemblages as a function of position along the nutrient gradient.

"he Florida Everglades is a large freshwater subtropical wetland at the southern end of the Florida peninsula (see Fig. S1 in the supplemental material), and it was estimated to harbor at one time the largest single body of organic soils in the world, covering over $8,000 \text{ km}^2$ (1). Wetlands, including the Everglades, are the primary source of natural global CH4 emissions, producing more than 150 Tg of CH₄ annually (roughly 20% of global annual emissions) (2, 3). The Everglades ecosystem was historically limited in nutrients, particularly phosphorus (P); however, the discharge of agricultural drainage from the adjacent Everglades Agricultural Area (EAA) led to elevated nutrient levels in the northern Everglades, particularly in Water Conservation Area 2A (WCA-2A), which is characterized by a well-documented gradient in soil and water P concentrations (4-7). The alleviation of P limitation resulted in many changes to the WCA-2A ecosystem; primary productivity was significantly increased, and the dominant plant species changed from saw grass to cattail. In addition, organic matter mineralization to CO_2 and CH_4 was greatly increased (8, 9).

Numerous studies on the impacts of nutrient additions to WCA-2A soils have been conducted, including analyses of methanogen community structures (10-12) and methanogenesis rates (8, 9). However, the detailed mechanisms controlling methanogenic pathways and the development of methanogenic guilds in response to shifting nutrient limitations are poorly understood.

In freshwater wetlands, CH₄ is primarily produced via two pathways: hydrogenotrophic methanogenesis (HM; CO₂ + $4H_2 \rightarrow CH_4 + 2H_2O$) and acetoclastic methanogenesis (AM; CH₃COOH \rightarrow CH₄ + CO₂). From the stoichiometry of glucose fermentation, the relative proportions of the two pathways are 33% HM and 67% AM (13). The natural distribution of the two pathways generally follows this proportion; however, some notable exceptions have been reported in the literature. Recent work (14) indicated that the relative contributions of the methanogenic pathways in WCA-2A may be related to nutrient status in the Everglades; relatively greater HM is observed in nutrient-impacted soils, and the predicted proportions are observed in nutrient-unimpacted soils.

Various factors have been suggested to be responsible for those cases in which the relative contributions of the HM and AM pathways deviate significantly from those predicted by stoichiometry, including soil depth (15, 16), nutrient type (17), seasonal conditions (18), pH (19, 20), and vegetation type (21). However, competition between methanogens and other functional groups of microorganisms has not been extensively studied as a possible mechanism responsible for shifts in the relative proportions of HM versus AM pathways. Methanogens and microorganisms that use more energy-yielding alternative terminal electron acceptors, such as SO_4^{2-} , Fe(III), NO_3^{-} , or O_2 , may compete for substrates such as acetate and H₂.

Competition between sulfate-reducing prokaryotes (SRPs)

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and methanogens for the common substrates acetate and H_2 has been well documented in various environments, including marine sediments (22, 23), freshwater sediments (24, 25), and bioreactors (26); however, little attention has been given to the possibility that this competition may impact the dominant methanogenic pathway. The interactions between SRPs and methanogens may be complex, depending on the availability of both electron donors and electron acceptors. The interaction may be competitive when sufficient sulfate is available to serve as a terminal electron acceptor for SRPs; however, it may be more likely to be cooperative via a syntrophic relationship when SO_4^{2-} is limiting (27, 28). SO_4^{2-} has been recognized to be an important contaminant in the northern water conservation areas (WCAs) (29). Hence, the SRPmethanogen interaction may be a crucial factor determining the methanogenic pathway in the WCAs.

The objective of this research was to determine the distribution and population dynamics of methanogens and SRPs across nutrient gradients in WCAs and to evaluate the potential interactions between SRPs and methanogens as a driving force shaping methanogenic community structures and pathways. For this study, we selected sites distinct from each other with respect to the concentrations of P and SO_4^{2-} , key geochemical parameters that have been shown to affect methanogenesis and SO_4^{2-} reduction in Everglades wetland soils (9, 10, 29). This study extends our knowledge of the mechanisms related to methane production and of the interplay between methanogens and SRPs in freshwater wetlands.

MATERIALS AND METHODS

Sampling and sample processing. Replicate soil cores (≥3 cores for each site within approximately 25 m²) were obtained from sites F1, F4, and U3 within WCA-2A (in October 2009, April 2010, August 2011, and January, August, and December 2012) and site W3 within WCA-3A (in February 2012 and March and April 2013). Soil cores were sectioned to an interval of 2 cm or 4 cm from the top after removing floc to minimize inclusion of the major O₂ interface regions. A portion of each soil sample (approximately 50 to 100 g) was immediately frozen on dry ice and transported to the laboratory in Gainesville, FL, where the soils were stored at −80°C until the isolation of nucleic acids or geochemical analysis. The remaining soils were used for incubation studies within 1 week. In addition, 5 to 10 liters of surface water was collected from each site in 10-liter polypropylene bottles, which served as the source water for incubation studies. Pore water samples were collected from each sampling location and stored as described by Holmes et al. (14).

Nucleic acid isolation, PCR, clone library construction, and sequence analyses. DNA was isolated from 0.2 g (wet weight) of soil using PowerSoil DNA isolation kits (MO BIO Laboratories, Carlsbad, CA). Total RNA was isolated from 2.0 g of soil using a MO BIO PowerSoil total RNA isolation kit. The residual DNA in the RNA extracts was removed using a MO BIO RTS DNase kit. RNA was converted to cDNA using SuperScript III first-strand synthesis SuperMix including random hexamers as reverse transcriptase PCR primers (Invitrogen, Carlsbad, CA). Nucleic acids were stored at -80°C until use.

Clone libraries were constructed for analysis of the methyl coenzyme M reductase (*mcrA*) gene and the dissimilatory (bi)sulfite reductase (*dsrB*) gene. The *mcrA* gene in soil DNAs from sites F1, F4, and U3 (sampled in October 2009; depth, 0 to 2 cm) was amplified using primers mlas/mcrA-rev as previously described by Steinberg and Regan (30). Reverse transcription-PCR (RT-PCR) was performed to amplify *dsrB* from cDNA derived from F4 soils (August 2012; depth, 0 to 4 cm) using primers DSRp2060F/DSR4R as described by Foti et al. (31). The PCR products were cloned and subsequently transformed into *Escherichia coli* TOP10 by use of a TOPO TA cloning kit for sequencing (Invitrogen, Carlsbad, CA). The transformants were randomly selected on Luria-Bertani (LB) agar

plates containing kanamycin (50 μ g · ml⁻¹) and sent to the University of Florida Sequencing Core Laboratory (http://www.biotech.ufl.edu/) for sequencing of the inserts.

All DNA sequences determined in this study were converted *in silico* into the corresponding amino acid sequences by use of the BioEdit (v.7.1.3) program (32). For phylogenetic analysis, reference sequences of *mcrA* or *dsrB* with a high similarity to the sequences recovered in the present study on a BLAST search against the sequences in the NCBI database (http://www.ncbi.nlm.nih.gov/) were collected. Reference sequences of a variety of taxa were also obtained on the FunGene website (http://fungene.cme.msu.edu/), and environmental sequences were obtained from previously published literature. Selected sequences representing operational taxonomic units (OTUs) with 5% differences in amino acid sequences were pooled with the reference sequences and aligned by use of the ClustalX (v.2.0) program (33). The alignment was used as the input file for phylogenetic analysis in MEGA (v.5.2) software (34). The phylogenetic tree was constructed using the maximum likelihood method with bootstrap analysis (1,000 resamplings).

The deduced amino acid sequences were assigned to OTUs on the basis of the percent differences between the sequences (e.g., 1%, 5%, 10%) using the furthest-neighbor method in the program mothur (v.1.31.2) (35). The mothur program was also used to estimate the diversity of OTUs and the coverage of OTUs sampled within each clone library and to create a Venn diagram showing the number of OTUs shared between clone libraries. Fast UniFrac online analysis (36) was performed for principal coordinates analysis (PCoA), phylogenetic test (37), UniFrac significance test (38), and hierarchical cluster analysis (38).

(RT-)qPCR for mcrA. The numbers of mcrA copies were estimated using quantitative PCR (qPCR) with the universal primer set targeting total methanogens (T-M), mlas/mcrA-rev (39). Three forward primers, primers MM-F (5'-CAA GTW YGG MGG ATT CGC CAA GG-3'), MST-F (5'-CAA GTW YGG MGG ATT CGC CAA GG-3'), and MB-F (5'-AAG CAC CWA ACA MCA TGG AHA CHG T-3'), were designed to enumerate the organisms in the groups Methanomicrobiales, Methanosaetaceae, and Methanobacteriales, respectively. The conserved sequence region for each group was used for primer design (see Fig. S2 in the supplemental material). These forward primers made a pair with the universal primer mcrA-rev in PCRs (see Table S1 in the supplemental material). The specificity of the primers was verified by analysis of sequences amplified by the group-specific primers. A total of 21, 28, and 26 clones were randomly selected from clone libraries constructed from PCR products using the primers MM-F, MST-F, and MB-F combined with mcrArev, respectively. All sequences from selected clones were matched to the target groups.

All qPCRs were conducted using iQ SYBR green supermix (Bio-Rad Laboratories, Hercules, CA) in a StepOnePlus real-time PCR system (Applied Biosystems, Foster City, CA). The reaction mixture contained 10 μ l of iQ SYBR green supermix, 2 μ l of primers (concentration of each primer, 10 pmol $\cdot \mu$ l⁻¹), and 2 μ l of DNA (cDNA for RT-qPCR) in a 20- μ l final volume. qPCR cycling parameters were 3.5 min at 95°C, followed by 6 cycles of touchdown PCR (30 s at 95°C, 45 s with a 1°C-per-cycle decrement from 60°C to the final annealing temperature, 30 s at 72°C) and 34 cycles of the main PCR (denaturation at 95°C for 30s, annealing at 55°C for 45 s, and extension at 72°C for 30 s, image capture at 80°C for 15 s, and a final extension at 72°C for 7 min). All qPCR runs included an image capture step (15 s at 80°C) after the final extension step of each cycle and a melt curve analysis (in which the temperature was increased from 60 to 95°C in 0.5°C increments every 10 s) when the PCR amplification was completed.

Copy numbers of *dsrB* were estimated using primers DSRp2060F/ DSR4 under the cycling conditions described by Foti et al. (31). The reaction mixture was prepared as described above for the *mcrA* qPCR, except that *dsrB*-specific primers were used.

For all (RT-)qPCRs, (c)DNA from each soil sample was applied in triplicate to a 96-well PCR plate with a 10-fold dilution series of a plasmid

carrying the gene fragment of interest (plasmid DNA standard). The plasmid DNA standard was prepared by cloning the target gene fragment amplified from soil samples using the same primer set used for qPCR. The insertion of the correct gene fragments of the plasmid DNA standard was confirmed by sequence analysis. The prepared plasmid DNA standards were stored in aliquots at -80° C, and a separate standard was used for each qPCR run. The PCR efficiency (*E*) was calculated using the formula $-1 + 10^{(-1/slope)}$. The PCR efficiency measured using plasmid DNA standards under the above-described conditions for *mcrA* and *dsrB* ranged from 93.4 to 96.7% (see Table S2 in the supplemental material).

Soil incubation experiments. Several soil incubation experiments were conducted to measure methane production rates, sulfate reduction rates, and the isotopic compositions of CH_4 and CO_2 . Each treatment was conducted in triplicate. All incubations were performed at 28°C in the dark. Bottles were shaken by inversion a few times every 1 or 2 days (for determination of the CH_4 production rate) or on a shaking incubator at 125 rpm (for determination of the SO_4^{2-} reduction rate).

(i) CH₄ production rate. Ten grams of surface soil (depth, 0 to 4 cm) was mixed with 10 ml of site water in triplicate 60-ml serum bottles closed with rubber stoppers and aluminum seals. The headspace gas of the bottles was exchanged by flushing N₂ through syringe needles for 10 min. The bottles were supplemented with 4 mM acetate or an H₂-CO₂ mixture (80%-20%, vol/vol) to 50 kPa. Bottles with no substrate addition were used as controls. The CH₄ concentration was analyzed on days 3, 7, and 14, as described below.

(ii) SO_4^{2-} reduction rate. Sulfate reduction rates were measured according to previously described methods (40, 41) with a slight modification as described by Castro et al. (42). Sulfate reduction rates were calculated as described by Fossing and Jørgensen (43).

(iii) Isotope composition of CH₄ and CO₂. Ten grams of soil was incubated with 10 ml of surface water in 60-ml serum bottles closed with a rubber stopper and an aluminum seal. MOO_4^{2-} (20 mM) was added to half of the incubation vials. Headspace gas was taken on incubation days 5, 8, and 12 to measure the CH₄ production rates. The stable isotopes of CH₄ and CO₂ were determined as described below. The headspace of vials containing samples from sites F1 and U3 was sampled on day 12. The headspace of vials containing samples from site F4 was sampled on day 27 because there was not enough CH₄ in the headspace for isotopic measurement until then.

(iv) Potential SAO activity. Potential syntrophic acetate-oxidizing (SAO) activity was measured as described by Hori et al. (44). Briefly, 5 g of soil was anaerobically incubated with 15 ml of surface water in 60-ml serum bottles with acetate labeled with ¹³C at either the C₁ or C₂ position or both the C₁ and C₂ positions (Sigma Biochemicals) to a final concentration of 0.5 mM in separate incubations. After 6 days of incubation (12 days for site U3 soils), the concentrations of ¹²CH₄ and ¹³CH₄ in the headspace gases were analyzed as described below. Briefly, this method is based on the fact that the syntrophic acetate oxidation and acetoclastic pathways yield methane from carbons at different positions in the acetate molecule.

Analytical methods. Total phosphorus (TP), total carbon (TC), and total nitrogen (TN) concentrations were determined according to methods described by White and Reddy (45).

The CH₄ concentrations in soil incubation and pore waters were measured from the headspace using a Shimadzu 8A gas chromatograph equipped with a Carboxen 1000 column (Supelco, Bellefonte, PA) and a flame ionization detector operating at 110°C as described previously (10). The CH₄ concentration in the aquatic phase was calculated using Henry's law constant for CH₄ ($1.3 \times 10^{-3} \text{ mol} \cdot \text{liter}^{-1} \cdot \text{atm}^{-1}$ at 298 K) (46; http://henrys-law.org).

The H₂ concentrations in pore waters were measured from the headspace of the bottles using a Peak Performer 1 gas analyzer (Peak Laboratories, Mountain View, CA) with a reducing compound photometer. The H₂ concentration in the aquatic phase was calculated using Henry's law constant for H₂ (7.8 × 10⁻⁴ mol · liter⁻¹ · atm⁻¹ at 298 K) (46). Acetate was derivatized with 2-nitrophenyl hydrazide, and the derivative was separated by use of a high-pressure liquid chromatography system (Waters 2695; Waters Corp., Milford, MA) equipped with a Platinum EPS C₈ column (1.6 by 250 mm; Alltech, Deerfield, IL) under a gradient profile composed of two mobile phases, as described by Albert and Martens (47). The derivative was detected at 400 nm with a UV absorbance detector (Waters 2487; Waters Corp.).

The composition of δ^{13} CH₄ and δ^{13} CO₂ in the headspace of the incubation bottles was determined using a Finnigan Mat Delta V isotope ratio mass spectrometer coupled to a gas chromatograph, as described by Merritt et al. (48).

Statistical analyses. The significances in the differences in gene copies, microbial activities, and chemical data between study sites or treatments were computed using one-way analysis of variance (ANOVA) or Student's *t* test with JMP (v.10) software (SAS Institute Inc., Cary, NC, USA), followed by the *post hoc* Tukey-Kramer honestly significant difference (HSD) test for ANOVA. *P* values of <0.05 were considered significant. For exploring the relationships between variations in gene copy numbers and geochemical parameters, a redundancy analysis (RDA) was implemented in Canoco (v.4.5 for Windows) software (49). The statistical significances of axis and individual parameters were evaluated using a Monte Carlo permutation full model with 999 unrestricted permutations.

Nucleotide sequence accession numbers. The GenBank accession numbers for the sequences determined in this study are KR075171 to KR075423 for *mcrA* and KR075424 to KR075502 for *dsrB* transcript sequences.

RESULTS

Site descriptions and biogeochemical characteristics. Sites F1, F4, and U3 are located within WCA-2A of the northern Everglades (see Fig. S1 in the supplemental material), where agricultural drainage from EAA has resulted in a north-to-south gradient in soil P concentrations (50). The soil P concentrations in the samples used for this study were $1.28 \text{ g} \cdot \text{kg}^{-1}$ in site F1, the northernmost site; $0.59 \text{ g} \cdot \text{kg}^{-1}$ in site F4; and $0.25 \text{ g} \cdot \text{kg}^{-1}$ in site U3, the southernmost site (Table 1). These values are in good agreement with previously reported results (42, 51). In contrast to the soil P concentrations gradient, a sharp gradient in SO₄²² concentrations was not observed in surface waters (176 to 200 μ M) or soil pore waters (39 to 56 μ M) (52). Acetate and H₂ concentrations were highly variable, and average values are presented in Table 1. No significant differences were observed between the sites due to the high variability among the samples.

Site W3 is located in the interior region of WCA-3A (see Fig. S1 in the supplemental material), such that it is removed from the direct influence of surface water discharges. The site harbors relatively low soil P concentrations (0.38 g \cdot kg⁻¹), similar to site U3 in WCA-2A; however, this site is distinguished from the WCA-2A sites by a much lower SO₄²⁻ concentration ($\leq 4 \mu$ M) in surface waters and pore waters (52). Site W3 is therefore valuable for comparison with the WCA-2A sites. The sites selected for this study are distinct from each other with respect to the concentrations of soil P and surface water SO₄²⁻: site F1 has high P and high SO₄²⁻ concentration, site F4 has an intermediate P concentration and a high SO₄²⁻ concentration, and site W3 has low P and SO₄²⁻ concentration, and site W3 has low P and SO₄²⁻ concentration.

Diversity and distribution of McrA sequences. A total of 255 *mcrA* sequences were obtained from three clone libraries derived from site F1, F4, and U3 soils in WCA-2A and were translated *in silico* to amino acid sequences (henceforth referred to as McrA

		Soil chemistry	q^{L}		Pore water chem	listry ^c		CH4 produc soil ⁻¹ · day	ction rate (μr -1) ^e	nol · g	$SO_4^{2^-}$ reduct $day^{-1})^e$	ion rate (μmol ·	g soil ⁻¹ .
Site	Coordinates	TP concn (g/kg)	TN concn (g/kg)	TC concn (g/kg)	Acetate concn (μM)	$H_2 \ concn (\mu M)$	SO_4^{2-} concn $(\mu M)^d$	Intact soils	Acetate	H_2 -CO ₂	Intact soils	Acetate	H_2 -CO ₂
FI	26°21'39"N, 80°22'10"W	1.28 ± 0.24	31 ± 3	442 ± 31	19.5 ± 9.9	0.33 ± 0.10	56 ± 86	3.1 ± 3.2	5.3 ± 1.4	11.6 ± 3.5	0.5 ± 0.4	1.4 ± 0.3	1.2 ± 1.0
F4	26°18'59"N, 80°23'01"W	0.59 ± 0.26	34 ± 3	424 ± 29	25.9 ± 33.8	0.28 ± 0.20	74 ± 110	0.9 ± 0.9	2.7 ± 0.5	5.6 ± 2.9	<0.01	0.1 ± 0.1	<0.01
U3	26°17'15"N, 80°24'41"W	0.25 ± 0.04	27 ± 5	385 ± 69	17.1 ± 16.4	0.28 ± 0.14	39 ± 35	<0.1	0.4 ± 1.2	2.1 ± 1.0	0.03 ± 0.04	0.07 ± 0.07	0.14 ± 0.17
W3	26°02′35″N, 80°49′39″W	0.38 ± 0.02	36 ± 1	424 ± 4	ND ^f	ND	≤4	ND	ND	ND	ND	ND	ND
^{<i>a</i>} Dat ^{<i>b</i>} Soil ^{<i>c</i>} Acet	The transform of the transform of the transformation of transformatio	Ds. (0 to 4 cm) sample ions were measure	ad in April 2010 : ad in pore waters	and August and \$(0 to 4 cm) sam	December 2012 fror apled in April 2010 a	n the WCA-2A sitt nd August 2011 (<i>i</i>	es $(n = 9$ from tripl $n = 6$).	icates of each sa	mple) and Feb	ruary 2012 and A	April 2013 from si	te W3 ($n = 6$) was	assayed.

in incubations with soils (depth, 0 to 4 cm) sampled in August 2012. Acetate (4 mM) and H_2 - CO_2 (50 kPa) were added in triplicate.

reduction were measured

The rates of methane production and SO_4^{2-}

ND, not determined

sequences) for further analysis. The McrA sequences were grouped into 96 OTUs defined by a 5% difference in amino acid sequences. The Chao1 richness estimator predicted the presence of 58 to 74 OTUs, with the highest number being observed for the sequences from site F1. The Shannon diversity indices (range, 3.4 to 3.6) did not differ significantly between sites. Coverage statistics indicated that 71 to 78% of the OTUs in each library were accounted for in this study (see Table S3 in the supplemental material).

The 96 OTUs were distributed within the orders *Methanomicrobiales*, *Methanosarcinales*, *Methanocellales*, *Methanobacteriales*, and *Methanomassiliicoccales*. The isolate of the *Methanomassiliicoccales* is related to recently isolated *Methanomassiliicoccus luminyensis* B10 (53) and to "*Candidatus* Methanomethylophilus alvus" Mx1201 retrieved from genomes derived from a human gut (54) (Fig. 1). Detailed phylogenetic affiliations for the OTUs are provided in Table S4 in the supplemental material.

One hundred forty-four sequences (accounting for 56% of all sequences) made up 50 OTUs belonging to the order Methanomicrobiales. Of the OTUs belonging to the Methanomicrobiales, 24 OTUs clustered within the family Methanomicrobiaceae (Methanomicrobiales subclades I [9 OTUs] and II [15 OTUs]), and 13 OTUs clustered within the Methanoregulaceae (Methanomicrobiales subclades V [6 OTUs] and VII [7 OTUs]). The remaining 13 OTUs were divided into novel Methanomicrobiales clades III (2 OTUs), IV (7 OTUs), and VI (4 OTUs), which did not include previously taxonomically defined methanogens. Sixty sequences (accounting for 25% of all sequences) made up 24 OTUs belonging to the order Methanosarcinales; of these, 22 OTUs were assigned to the Methanosaetaceae and 2 OTUs were assigned to the Methanosarcinaceae. Twenty sequences (7.8% of all sequences) made up eight OTUs belonging to the order Methanobacteriales, and *Methanobacteriales* were present as a minor group ($\leq 10.4\%$) at all sites. One mrtA sequence (Methanobacteriales clade I) which is usually detected in Methanobacteriales as an isozyme of McrA (55) was found in the library of sequences from site F1. Four OTUs belonging to the Methanomasiliicoccales were included as a minority (\leq 5.7%) in the libraries of clones from sites F1 and U3, but no sequence belonging to the Methanomasiliicoccales was found at site F4. Nine OTUs belonging to the Methanomasiliicoccales group were included in the libraries of clones from sites F1 (accounting for 6.9% of all site F1 sequences) and U3 (14.9% of all site U3 sequences), and this group was further divided into subgroups I and II, with subgroup II including strains B10 and Mx1201.

The differences in the McrA assemblages between sites F1, F4, and U3 were clearly illustrated by the compositions and relative abundances (in percent) of subclades for each site, as depicted in the stacked-column graph on the right in Fig. 1. Of 96 OTUs, only 4 OTUs were shared between the three sites, and less than 15 OTUs were shared between two sites (see the Venn diagram in Fig. S3 in the supplemental material). Significant differences between assemblages were statistically verified ($P \leq$ 0.001) by the phylogenetic test (37) and UniFrac significance test (38) in the UniFrac implementation. In a scatter plot of an unweighted PCoA, all assemblages were clearly separated by either principal component 1 (explaining 62.57% of the variation) or principal component 2 (explaining 37.43% of the variation), which is in good agreement with the largely separate assemblages represented in a hierarchical tree (see Fig. S3 in the supplemental material).



FIG 1 (Left) Maximum likelihood tree representing the phylogeny of amino acid sequences (seqs) deduced from the sequences of *mcrA* genes retrieved from soils taken from sites F1, F4, and U3 in October 2009. Bootstrap values of \geq 50% from 1,000 reassemblages are placed at the branch points. Gray, clades targeted by the group-specific primers designed in the present study; white, clades containing only sequences from this study; black, clades with reference sequences and not containing sequences from this study. (Right) Relative percentages of clades generated from the maximum likelihood tree and the clades targeted by the currently designed primers. MMC, *Methanomasiliicoccales*; MB, *Methanobacteriales*; MCEL, *Methanocellales*; MSR, *Methanosarcinaceae*; MM, *Methanomicrobiales*; MST, *Methanosaetaceae*.

qPCR enumeration of mcrA and dsrB copies. (i) mcrA copy numbers. The numbers of mcrA copies were estimated for total methanogens and for the groups Methanomicrobiales, Methanosaetaceae, and Methanobacteriales. Since the mcrA sequences from WCA-2A were largely unique and diverse, new primers MM-F, MST-F, and MB-F were designed to enumerate these groups. Degenerative primer MM-F targets the major groups of the Methanomicrobiales included in subclades II, III, IV, and V, which included 73% of the WCA-2A Methanomicrobiales sequences (Fig. 1). Primer MST-F targets Methanosaetaceae sequences on branches that include 70% of the WCA-2A Methanosaetaceae sequences, with the exception of one branch closely related to Methanosaeta harundinacea. Primer MB-F targets the Methanobacteriales, and its sequence was highly matched to all the WCA-2A Methanobacteriales sequences. Even though there was some variation among the sites, the relative proportions of sequences targeted by primers MM-F, MST-F, and MB-F were consistent with the proportions of *Methanomicrobiales*, *Methanosaetaceae*, and *Methanobacteriales* in the order *Methanomicrobiales* > *Methanosaetaceae* > *Methanobacteriales* for each clone library (Fig. 1). Therefore, qPCR using these primers is believed to be appropriate for estimating the relative number of target groups within and between sites.

qPCR for *mcrA* in soil samples collected in different seasons from 2009 to 2013 was conducted (see Table S5 in the supplemental material). Three subsamples were taken at each time point, and then the amounts were averaged for analysis. A one-way ANOVA blocked by time was run for each of the sites and for the different response variables. When significant differences in the numbers of genes were found, the Tukey-Kramer HSD test was used determine which genes differed in number (see Fig. S4 in the supplemental material).



FIG 2 Box-and-whisker plot (top) and temporal profile (bottom) of the copy numbers of *mcrA* from *Methanomicrobiales*, *Methanosaeta* (*Methanosaetaceae*), and *Methanomicrobacteriaceae* (*Methanobacteriales*) in WCA soils. Data in the temporal profile are presented in order of sampling: from the left, October 2009, April 2010, August 2011, and January, August, and December 2012 for WCA-2A sites F1, F4, and U3 and February 2012 and March and April 2013 for site W3. Error bars in the bar graph represent ± 1 SE (n = 3). Box-and-whisker plots were generated from the pooled data from the temporal profile. Boxes depict the medians (horizontal lines in the boxes) and the lower and upper quartiles (bottoms and tops of the boxes, respectively). The vertical bars (whiskers) show the highest and the lowest values, excluding outliers. The different letters indicate a significant difference among the groups *Methanomicrobiales*, *Methanosaetaceae*, and *Methanobacteriales* (P < 0.05 by the Tukey-Kramer HSD test).

qPCR for the groups Methanomicrobiales, Methanosaetaceae, and Methanobacteriales revealed that the Methanomicrobiales and/or Methanosaetaceae were dominant in all WCA sites; however, the Methanomicrobiales were dominant in WCA-2A and the Methanosaetaceae were dominant in WCA-3A. Within all WCA-2A sites, the Methanomicrobiales mcrA copy numbers (2.2 to 9.6 \times $10^8 \cdot \text{g soil}^{-1}$, on average) were higher than the *Methanosaetaceae mcrA* copy numbers $(1.7 \text{ to } 5.7 \times 10^8 \cdot \text{g soil}^{-1}, \text{ on average})$ (Fig. 2; see also Table S5 in the supplemental material). A one-way ANOVA blocked by time revealed that the Methanomicrobiales mcrA copy numbers were higher than the Methanosaetaceae mcrA copy numbers in site F1 (P < 0.05) but were not significantly different from the numbers in the other sites, due to the temporal variation within each site (Fig. 2, bottom). Comparison of the percentages of each group within individual samples ([number of mcrA copies for one group]/[total numbers of mcrA copies for the Methanomicrobiales, Methanosaetaceae, and Methanobacteriales combined] \times 100) more clearly showed that the percentage of Methanomicrobiales was significantly higher than the percentage of *Methanosaetaceae* (Tukey-Kramer HSD test, P < 0.0001) in all the WCA-2A sites (Fig. 2). In contrast, for site W3 it was found that the Methanosaetaceae significantly outnumbered the Methanomicrobiales according to both the absolute number of mcrA copies $(7.8 \times 10^7 \cdot \text{g soil}^{-1} \text{ versus } 1.9 \times 10^7 \cdot \text{g soil}^{-1}$, on average) and the relative percentage of mcrA copies (73.9% \pm 11.6% versus

14.2% \pm 9.0% [mean \pm standard deviation {SD}]) (*P* < 0.0001 for both comparisons).

(ii) *dsrB* copy numbers. The number of *dsrB* copies estimated in WCA soils ranged from 5.1×10^7 to $4.9 \times 10^9 \cdot \text{g soil}^{-1}$ (see Fig. S5 and Table S5 in the supplemental material). WCA-2A sites showed significant variations in *dsrB* copy numbers by season (one-way ANOVA, P < 0.001 within each site), while site W3 in WCA-3A did not (one-way ANOVA, P = 0.8855) (see Fig. S5 in the supplemental material). In all WCA-2A sites, the highest *dsrB* copy numbers were obtained in January, while the lowest copy numbers were observed in April. Comparison of the pooled temporal data from each site indicated that the *dsrB* copy numbers were significantly different among the sites (one-way ANOVA, P < 0.001) (see Fig. S5 in the supplemental material). Site F1 had the highest number $(1.88 \times 10^9 \cdot \text{g soil}^{-1}$, on average), followed by sites F4 $(1.52 \times 10^9 \cdot \text{g soil}^{-1})$, U3 $(6.72 \times 10^8 \cdot \text{g soil}^{-1})$, and W3 $(8.55 \times 10^7 \cdot \text{g soil}^{-1})$, in that order.

The *dsrB* copy numbers were significantly higher than the T-M *mcrA* copy numbers in the WCA-2A sites (Student's *t* test, P < 0.05) but were not significantly different from the copy numbers in site W3 (Student's *t* test, P = 0.86) (Fig. 3).

Relatedness of *mcrA* and *dsrB* copy numbers to geochemical parameters. Redundancy analysis (RDA) was performed to evaluate the relationships between the abundances of methanogens and SRPs and the geochemical concentrations in the WCAs. T-M



FIG 3 Box-and-whisker plot (top) and temporal profile (bottom) of the *dsrB* copy numbers compared with the T-M *mcrA* copy numbers estimated in the same sample. Data in the temporal profile are presented in order of sampling: from the left, April 2010, August 2011, and January, August, and December 2012 for WCA-2A sites F1, F4, and U3 and February 2012 and March and April 2013 for site W3. Error bars in the bar graph represent ± 1 SE (n = 3). The box-and-whisker plot was constructed from the pooled data from the temporal profile. Boxes depict the medians (horizontal lines in the boxes) and the lower and upper quartiles (bottoms and tops of the boxes, respectively), while the whiskers show the highest and the lowest values, excluding outliers. The different letters indicate a significant difference between *dsrB* and T-M *mcrA* copy numbers (P < 0.05 by Student's *t* test).

copy numbers were positively correlated with TP concentrations (Fig. 4). *dsrB* copy numbers were also positively correlated with SO_4^{2-} and TP concentrations. The positive correlations of T-M and *dsrB* copy numbers with those geochemical parameters were due to the elevated numbers in the presence of higher P and SO_4^{2-} concentrations by site in the order F1 > F4 > U3 > W3, as described above.

RDA was also applied to the relative proportions of *Methanomicrobiales*, *Methanosaetaceae*, and *Methanobacteriales* within each sample to observe the relationships between the compositions of these methanogenic groups and geochemical factors. The percentage of *Methanomicrobiales* was positively correlated with TP and SO_4^{2-} concentrations, while the percentages of *Methanosaetaceae* and *Methanobacteriales* were negatively correlated with those parameters.

Levels of *mcrA* and *dsrB* gene expression. In order to assess the extent to which the genes were transcribed, RT-qPCR for *mcrA* and *dsrB* and potential activity for methanogenesis and sulfate reduction (to indirectly measure the levels of enzymes encoded by the corresponding genes) were determined in the samples from the WCA-2A sites collected in August 2012. The average *mcrA* mRNA copy numbers were as follows: $1.7 \times 10^7 \cdot \text{g soil}^{-1}$ in site F1, 1.1×10^6 in site F4, and $9.3 \times 10^4 \cdot \text{g soil}^{-1}$ in site U3. The average *dsrB* mRNA copy numbers were 1.6×10^6 in site F1, $9.0 \times 10^4 \cdot \text{g soil}^{-1}$ in site F4, and $5.9 \times 10^4 \cdot \text{g soil}^{-1}$ in site U3. The mRNA copy numbers were 2 to 4 orders of magnitude lower than the gene copy numbers; however, they followed the order of the gene copy numbers according to the SO_4^{2-} and P gradient, i.e., F1 > F4 > U3 (Fig. 5). Likewise, the numbers of copies of cDNA from the *Methanomicrobiales*, *Methanosaetaceae*, and *Methanobacteriales* were consistent with the relative abundances of the gene copy numbers within sites and among the sites.

Potential methane production rates were 3.1 μ mol \cdot g soil⁻¹ · day⁻¹ in site F1 soils, 0.9 μ mol \cdot g soil⁻¹ · day⁻¹ in site F4 soils, and less than 0.1 μ mol \cdot g soil⁻¹ · day⁻¹ in site U3 soils (Table 1). Potential SO₄²⁻ reduction rates were 0.5 μ mol \cdot g soil⁻¹ · day⁻¹ in F1 soils, which were higher than those measured in F4 and U3 soils ($\leq 0.03 \mu$ mol \cdot g soil⁻¹ · day⁻¹). These potential activities for CH₄ production and SO₄²⁻ reduction were in good agreement with the order of the levels of the *mcrA* and *dsrB* copy numbers measured along nutrient gradients (F1 > F4 > U3); hence, the gene copy numbers measured throughout this study are likely to predict the level of gene expression in each site.

Impact of SRP activities on methanogenesis. Soil incubation studies were conducted to evaluate the potential influence of SRP activities on methanogenic pathways and, correspondingly, on the shaping of the methanogen assemblage structure. The incubation was done using the WCA-2A soils with relatively high $SO_4^{\ 2^-}$ concentrations in the presence or absence of $MOO_4^{\ 2^-}$, a specific inhibitor of SRP activity. Soil incubation studies showed higher methane production rates in soils treated with $MOO_4^{\ 2^-}$ (a specific inhibitor of SRP activity) for all sites (Fig. 6A): $1.9 \cdot \text{g}$ soil⁻¹ · day⁻¹



FIG 4 Results of RDA presenting the correlation between *mcrA* and *dsrB* copy numbers and geochemical parameters obtained for the samples from sites F1, F4, and U3 collected in April 2010 and August and December 2012 and the samples from site W3 collected in February 2012 and April 2013. Arrows pointing in the same direction indicate positive correlations, and arrows pointing in opposite directions indicate negative correlations. The arrow length corresponds to the variance explained by the environmental variable. The first two axes explain 89.9% of the total canonical eigenvalues with a significant Monte-Carlo test value (P < 0.05).

for treated soils versus 1.3 μ mol \cdot g soil⁻¹ \cdot day⁻¹ for untreated soils in site F1, 1.3 versus 0.6 μ mol \cdot g soil⁻¹ \cdot day⁻¹ in site F4, and 0.5 versus 0.05 μ mol \cdot g soil⁻¹ \cdot day⁻¹ in site U3 (Fig. 6A). These findings indicate that methanogens and SRPs are in competition for common substrates.

 ${\rm MoO_4}^{2-}$ treatment resulted in increases in the levels of δ^{13} CH₄: -67‰ to -55‰ in site F1 and -76‰ to -57‰ in site F4. The small amount of CH₄ produced in U3 soil incubations without added MoO₄² prohibited measurement of δ^{13} C, but in incubations with Mo treatment, the level of δ^{13} CH₄ was -57‰ in site U3 (Table 2; Fig. 6B). In contrast, there was no significant difference in the amount of δ^{13} CO₂ observed between the control soils and the MoO₄²⁻-treated soils. The apparent fractionation factor (α_{app}) quantifies the isotopic difference between CH₄ and CO₂ and is a generally accepted index used to estimate the contribution of a particular methane production pathway to a methane pool (56, 57). MOQ_4^{2-} treatment reduced α_{app} from 1.054 to 1.040 in site F1 and 1.060 to 1.041 in site F4, indicating a shift toward the acetoclastic pathway.

Inhibition of SRP activity using MOQ_4^{2-} resulted in an increase in the relative *mcrA* transcript level for the *Methanosaetaceae* from 27 to 43% in site F1 soils and from 48% to 55% in site F4 soils (Fig. 7), consequently reducing the relative percentage of *Methanomicrobiales mcrA* transcripts. With MOQ_4^{2-} treatment, the *Methanobacteriales* increased in relative percentage, from 3 to 9% in F1 soils and from 1 to 2% in F4 soils, even though this group still appeared to be a minor group. The changes in the *mcrA* transcript level in U3 soil incubations were not accurately determined due to the recovery of low levels of RNA from the incubated soils.



FIG 5 Numbers of copies of the *mcrA* and *dsrB* genes and their transcripts measured from surface soils (0 to 4 cm depth) sampled from WCA-2A sites F1, F4, and U3 in August 2012. Error bars represent ± 1 SE (n = 3).



FIG 6 CH₄ production rate (A) and composition of the δ^{13} CH₄ and δ^{13} CO₂ produced (B) during incubation of soils sampled in August 2012. Error bars represent ±1 SD (n = 3; note that the control incubation of site U3 soils did not produce a detectable amount of δ^{13} CH₄). CT, control without MoO₄²⁻; Mo, addition of MOO₄²⁻ (20 mM).

DISCUSSION

The nutrient gradient in WCA-2A soils provides an excellent opportunity to study the impacts of nutrient additions to naturally P-poor wetlands. A very large body of work has been published on changes that the P additions have brought to the greater WCA-2A ecosystem, ranging from the distribution of endangered vertebrates to changes in biogeochemical cycling (49, 58). The present study builds on previous studies on the distribution and function of methanogenic and sulfidogenic guilds (10, 42, 51) and investigates controls on methanogenic pathways as a function of position along the nutrient gradient.

The methanogenic assemblage structure based on *mcrA* sequences revealed distinct features reflecting the nutrient status of WCA-2A. One of the key features was the numerical dominance and diversity of the hydrogenotrophic *Methanomicrobiales*. *Methanomicrobiales* sequences accounted for >49% of the total sequences retrieved from WCA-2A soils, and the phylogeny of the *Methanomicrobiales* was broadly distributed across 7 distinct clades, *Methanomicrobiales* clade I to clade VII. *Methanomicrobiales* have been shown to dominate in at least some acidic bogs or rice fields (30, 59); however, WCA *Methanomicrobiales* members are distinct from *Methanomicrobiales* sequences referred to as to the "fen cluster" (60) or the "rice cluster" (55) (Fig. 1).

Another feature is that the acetotrophic order *Methanosarcina*les was dominated by the *Methanosaetaceae* (\geq 95% of *Methano*sarcinales sequences). In general, the *Methanosaetaceae* exhibit a low threshold for acetate (61), and WCA-2A pore waters harbor low concentrations of acetate (<0.03 mM) (Table 1), as would be expected for a system dominated by the *Methanosaetaceae*. The low concentrations of acetate might result from SRP activities competing for acetate, thereby selecting for this type of acetotrophic methanogen in WCA-2A. Along with the *Methanomicrobiales* (49% to 65%), the *Methanosaetaceae* (22% to 32%) were a dominant methanogenic group in WCA-2A, such that these two groups play a significant role in determining the pathways of methanogenesis in WCA-2A.

Recently, Holmes and colleagues (14) reported that AM is the dominant pathway (50% to 75%) over HM (25% to 50%), based on the differences in CH4 production rates in soils with and without incubation with methyl fluoride (an inhibitor of AM) and the δ^{13} CH₄ and δ^{13} CO₂ values in pore waters collected from the same sites used in this study. Those results do not appear to be consistent with our observation that the Methanosaetaceae were outnumbered by the Methanomicrobiales in all sites of WCA-2A. This paradoxical result might be explained by the higher free energy of formation in HM (4H₂ + CO₂ \rightarrow CH₄ + 2H₂O, for which $\Delta G^{\circ'}$ is equal to $-135 \text{ kJ} \cdot \text{mol} \,\text{CH}_4^{-1}$) than in AM ($\text{CH}_3 \text{COOH} \rightarrow \text{CH}_4 +$ CO₂, for which ΔG° is equal to $-33 \text{ kJ} \cdot \text{mol CH}_4^{-1}$ (62). This allows HM to produce larger amounts of biomass even if less CH4 is produced by this pathway. We did not calculate ΔG for these reactions in situ, however, and there may be alternative explanations for these observations. Most hydrogenotrophs are able to grow with additional substrates (e.g., formate, methyl amines, methanol) other than H₂ and CO₂. For example, members of the numerically dominant group in WCA-2A, the Methanomicrobia*les*, utilize acetate as a carbon source, although they do not use it for methanogenesis (63).

Even though AM is the dominant methanogenic process overall in sites of WCA-2A, Holmes et al. (14) reported that HM became relatively more important at site F1, accounting for almost 50% of the total methane produced. These results are relatively consistent with the group-specific qPCR results reported here, where the *Methanomicrobiales* accounted for 60.3% of the organisms in site F1, 58.6% in site F4, and 55.0% in site U3, consequently decreasing the percentage of *Methanosaetaceae* in the order F1 (35.7%) < F4 (38.6%) < U3 (42.1%).

One of primary aims of this study was to evaluate the influence of SRP activities on the methanogenic pathways and the methanogen community as controlling forces in response to nutrient gradients within the WCAs. The qPCR results for the *Methanomicrobiales* and *Methanosaetaceae* across the SO_4^{2-} gradients provide evidence that SRP activity is likely involved in shaping methanogen assemblage structure and activity. In our long-term monitoring, WCA-2A sites representing SO_4^{2-} -rich environments consistently showed a dominance of the hydrogenotrophic *Methanomicrobiales*

TABLE 2 δ^{13} CH₄ and δ^{13} CO₂ concentrations and α_{app} values for incubated Everglades soils

		Concn (‰) ^b		
Site	Treatment ^a	$\delta^{13}CH_4$	$\delta^{13}CO_2$	α_{app}^{c}
F1	CT	-67.00 ± 2.16	-17.02 ± 1.20	1.054
	Мо	-55.02 ± 1.17	-17.16 ± 0.84	1.040
F4	CT	-75.92 ± 1.00	-20.12 ± 0.88	1.060
	Мо	-57.35 ± 2.81	-18.92 ± 2.28	1.041
U3	СТ	ND^d	-17.46 ± 0.58	NA ^e
	Мо	-56.88 ± 1.67	$-15.87 \pm .053$	1.043

^{*a*} CT, control without MoO_4^{2-} ; Mo, addition of MoO_4^{2-} (20 mM).

^{*b*} Data represent means \pm SDs (n = 3).

 $^{c}\,\alpha_{app}$ is defined as (amount of $\delta^{13}CO_{2}$ + 1,000)/(amount of $\delta^{13}CH_{4}$ + 1,000).

^d ND, not determined.

^e NA, not applicable.



FIG 7 mcrA transcript copy numbers, estimated using RT-qPCR, from the incubation of soils from sites F1 and F4 sampled in August 2012. For this RT-qPCR analysis, the RNA was isolated from the soils sampled on the same date that the gas samples were collected for the δ^{13} CH₄ and δ^{13} O₂ analysis. CT, control without MoO₄²⁻; Mo, addition of MoO₄²⁻ (20 mM). Error bars represent ±1 SE (*n* = 3).

(58% on average) over the acetotrophic *Methanosaetaceae* (39%), while site W3, representing a $SO_4^{2^-}$ -poor environment, revealed the opposite relationship (the *Methanomicrobiales* comprised 14% of the organisms, on average, whereas the *Methanosaetaceae* comprised 74%) (Fig. 2; see also Table S5 in the supplemental material). An RDA plot shows the positive correlation of the percentage of *Methanomicrobiales* organisms with the $SO_4^{2^-}$ concentration, while the percentage of *Methanosaetaceae* concentrations might cause an enrichment of sulfidogenic SRPs, which are typically thought to outcompete methanogens for acetate (26, 64), consequently decreasing the percentage of *Methanomicrobiales* of *Methanomicrobiales* of *Methanomicrobiales*.

The soil incubation study using MoO_4^{2-} as an SRP inhibitor provides evidence that SRPs control, at least to some extent, the methanogenic pathways and drive an enrichment of hydrogenotrophs, specifically, the Methanomicrobiales group, in WCA-2A. The addition of MoO_4^{2-} resulted in increasing values of $\delta^{13}\text{CH}_4$ and lower α_{app} values (Fig. 6B and Table 2), consistent with competition between SRPs and acetotrophic methanogens for acetate. Since the production of CH₄ by AM is generally associated with lower $\alpha_{\rm app}$ values and, often, with less negative $\delta^{13} CH_4$ values than HM (56, 65, 66), those increments imply that AM was enriched by blocking SRP activity; in other words, SRP activity most likely inhibited the activities of the Methanosaetaceae in these soils. The increased proportion of Methanosaetaceae mcrA mRNA observed in the MoO_4^{2-} treatments (Fig. 7) supports this contention, which is in good agreement with the aforementioned increase in AM caused by SRP inhibition.

The specific interactions between SRP and methanogens can

be quite complex. It is possible that some syntrophic fermentation of primary fermentation products, such as short-chain fatty acids or alcohols, occurred in our incubations with $MOO_4^{2^-}$, which may have contributed to the acetate used by the *Methanosaetaceae*. Wu et al. (67) reported that $MOO_4^{2^-}$ inhibited syntrophic fermentation by SRPs to varying degrees (97% for propionate, 24% for ethanol) in a wastewater bioreactor.

An additional sink for acetate and a corresponding source of H_2 might have been via syntrophic acetate oxidation to H_2 ; however, we determined in separate experiments without MOO_4^{2-} that SAO was not significant in our samples (data not shown).

It should be noted that other factors, such as differences in organic carbon quality, may also impact the relative proportions of AM and HM (68). We also expected that the P concentration may be an important factor governing the methanogen composition regarding AM and HM, likely through increased primary productivity and carbon input into the soil. Increases in P concentrations correlated with increases in the population size of methanogens, such that a positive correlation between *mcrA* copy numbers and P concentrations was observed (Fig. 4). However, the P concentration gradient was not related to the relative compositions of hydrogenotrophs and acetotrophs to the degree that it was observed for the SO_4^{2-} gradient (e.g., between WCA-2A and WCA-3A). Thus, sulfate concentrations and the activities of SRPs appear to be the most dominant controllers of the methanogenic pathway in the WCAs of Everglades.

SRPs are important coinhabitants with methanogens (12, 42, 52). Our qPCR results indicated that SRPs outnumbered methanogens in WCA-2A and revealed that their numbers were similar even in WCA-3A (Fig. 2). Even though WCA-2A has higher concentrations of $SO_4^{2^-}$ than many other freshwater marshes, such as

WCA-3A (Table 1), the SO₄²⁻ concentration may not be high enough to support such high numbers of SRPs relative to the numbers of methanogens in WCAs. In a recent study, we found that *dsrB* transcripts from syntrophic SRPs belonging to the *Syntrophobacterales* comprised \geq 75% of total *dsrB* transcripts found in the soils of sites F1, U3, and W3 (52). F4 soils, which were not tested in the previous study, also showed similar proportions of syntrophs in the present work (76%) (see Fig. S6 in the supplemental material). The high proportion of syntrophic SRPs likely explains the relatively high number of SRPs that were observed in our studies.

In conclusion, the numbers of copies and structures of *dsrB* and *mcrA* and their respective activities vary with nutrient status in the water conservation areas of the Florida Everglades. Depending on the available SO_4^{2-} concentration, SRPs are involved in controlling the methanogenic pathways, shaping methanogen assemblage structure, and controlling the CH₄ emission rate and pathway.

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