

In Situ Hydrogen Dynamics in a Hot Spring Microbial Mat during a Diel Cycle

Niels Peter Revsbech,^a Erik Trampe,^b Mads Lichtenberg,^b David M. Ward,^c Michael Kühl^{b,d}

Department of Bioscience, Aarhus University, Aarhus, Denmark^a; Marine Biological Section, Department of Biology, University of Copenhagen, Helsingør, Denmark^b; Department of Land Resources and Environmental Sciences, Montana State University, Bozeman, Montana, USA^c; Plant Functional Biology and Climate Change Cluster, University of Technology Sydney, Sydney, NSW, Australia^d

ABSTRACT

Microbes can produce molecular hydrogen (H₂) via fermentation, dinitrogen fixation, or direct photolysis, yet the H₂ dynamics in cyanobacterial communities has only been explored in a few natural systems and mostly in the laboratory. In this study, we investigated the diel *in situ* H₂ dynamics in a hot spring microbial mat, where various ecotypes of unicellular cyanobacteria (*Synechococcus* sp.) are the only oxygenic phototrophs. In the evening, H₂ accumulated rapidly after the onset of darkness, reaching peak values of up to 30 μmol H₂ liter⁻¹ at about 1-mm depth below the mat surface, slowly decreasing to about 11 μmol H₂ liter⁻¹ just before sunrise. Another pulse of H₂ production, reaching a peak concentration of 46 μmol H₂ liter⁻¹, was found in the early morning under dim light conditions too low to induce accumulation of O₂ in the mat. The light stimulation of H₂ accumulation indicated that nitrogenase activity was an important source of H₂ during the morning. This is in accordance with earlier findings of a distinct early morning peak in N₂ fixation and expression of *Synechococcus* nitrogenase genes in mat samples from the same location. Fermentation might have contributed to the formation of H₂ during the night, where accumulation of other fermentation products lowered the pH in the mat to less than pH 6 compared to a spring source pH of 8.3.

IMPORTANCE

Hydrogen is a key intermediate in anaerobic metabolism, and with the development of a sulfide-insensitive microsensor for H₂, it is now possible to study the microdistribution of H₂ in stratified microbial communities such as the photosynthetic microbial mat investigated here. The ability to measure H₂ profiles within the mat compared to previous measurements of H₂ emission gives much more detailed information about the sources and sinks of H₂ in such communities, and it was demonstrated that the high rates of H₂ formation in the early morning when the mat was exposed to low light intensities might be explained by nitrogen fixation, where H₂ is formed as a by-product.

Cyanobacterial mats are found in various limnic, marine, and hypersaline environments. The thickest and most homogeneous mats are found in extreme environments such as hypersaline habitats and geothermal springs with a scarcity or total absence of grazers. Such microbial mats are important natural model systems for studying microbial interactions and the fundamental links between structure, diversity, and function in microbial communities (1, 2). Another reason for the extensive interest in these mats is their strong apparent similarity to ancient microbial mats inhabiting the early Earth before grazers evolved and now preserved as stromatolites (3). Knowledge about the functioning of modern microbial mats may thus give insights into the functioning of early microbial ecosystems on planet Earth. Hydrogen (H₂) presumably played a major role in the metabolism of primitive microbial communities (4), and outgassing of H₂ originating or escaping from these microbial communities might have played a major role in the gradual oxidation of our planet (5).

Coastal and hypersaline cyanobacterial mats are known to be rather vigorous producers of H₂ during darkness (5–10), and some H₂ production may also be observed in the beginning of a light period due to photolysis (10). *Microcoleus* sp., which was a dominating cyanobacterium in the mat H₂ evolution studies listed above, is known to be a vigorous H₂ producer during fermentation (11). Burow et al. (7) showed that the H₂ production in the dark in a hypersaline microbial mat might be ascribed to cyanobacterial fermentation, whereas significant H₂ production from

N₂ fixation as a by-product of nitrogenase activity (N₂ + 8H⁺ + 8e⁻ → 2NH₃ + H₂) was not verified. A detailed analysis of an intertidal cyanobacterial mat dominated by *Microcoleus* sp. using sulfide-insensitive H₂ microsensors showed that the maximum H₂ concentrations reached 40 μmol H₂ liter⁻¹ (depending on the previous light exposure history of the mat) within the first hour after darkening (10), whereafter the H₂ concentration decreased to near zero within about 7 h. A main reason for this decrease was consumption by sulfate-reducing bacteria (SRB), as addition of the sulfate analogue and SRB inhibitor molybdate led to even higher H₂ concentrations that persisted for a much longer time period. In another coastal microbial mat, SRB were also identified as the major hydrogenotrophic prokaryotes (12).

The hot spring microbial mats in the siliceous alkaline Mushroom Spring and nearby Octopus Spring have been extensively studied for more than 40 years as natural model systems for inves-

Received 3 March 2016 Accepted 29 April 2016

Accepted manuscript posted online 6 May 2016

Citation Revsbech NP, Trampe E, Lichtenberg M, Ward DM, Kühl M. 2016. *In situ* hydrogen dynamics in a hot spring microbial mat during a diel cycle. *Appl Environ Microbiol* 82:4209–4217. doi:10.1128/AEM.00710-16.

Editor: A. M. Spormann, Stanford University

Address correspondence to Niels Peter Revsbech, revsbech@biology.au.dk.

Copyright © 2016, American Society for Microbiology. All Rights Reserved.

tingating *in situ* interactions of metabolic processes and microbial diversity (1, 2). In these hot springs, microbial mats growing at water temperatures of $>54^{\circ}\text{C}$ show a restricted diversity, with various ecotypes of the cyanobacterium *Synechococcus* (13) being the only oxygenic phototrophs. Metagenomic and metatranscriptomic analyses have demonstrated a relatively low diversity of taxonomic groups, including several photoheterotrophic or photomixotrophic *Chloroflexi* (*Roseiflexus*-, *Chloroflexus*-, and *Anaerolineae*-like taxa), photoheterotrophic acidobacteria (*Chloracidobacterium*), *Chlorobi* (“*Candidatus* Thermochlorobacter”), and two novel aerobic heterotrophic taxa (14–17). Such relatively limited diversity has made these hot spring mats attractive for the study of metabolic interactions among taxa under *in situ* conditions and under a natural diel cycle. In the slightly alkaline 50 to 70°C Octopus Spring mat, for example, it was shown that the two most abundant taxa, *Synechococcus* and *Roseiflexus*, produce glycogen (18) (and in *Roseiflexus*, polyhydroxyalkanoic acid as well [17, 19]), which they ferment at nighttime to produce large amounts of acetate and propionate that are photoassimilated by filamentous, *Chloroflexus*-like organisms (probably related to *Roseiflexus* [20]) the next morning (19, 21, 22). Both Anderson et al. (22) and van der Meer (23) reported that H_2 was formed during darkness in the mats, but the accumulation was restricted by diffusion into the overlying water (19) and by H_2 consumption by SRB (24, 25) and methanogenic bacteria (21, 26) in the 50 to 60°C mat. The rates of H_2 accumulation were much higher in the 65°C mat, which did not exhibit any methanogenesis.

Genomic and metagenomic analyses of microbial mats in Mushroom Spring and Octopus Spring have demonstrated that *Roseiflexus* (15, 27), but not *Synechococcus* (28), possesses hydrogenases and might thus produce H_2 by fermentation. The lack of hydrogenases in *Synechococcus* means that any H_2 produced by *Synechococcus* nitrogen fixation is lost to the environment where other mat organisms may use it as an electron donor in anoxygenic photosynthesis or respiration. Metatranscriptomic studies conducted on a 60°C mat in Mushroom Spring demonstrated that the expression of *Synechococcus* nitrogenase genes occurred in darkness and that N_2 fixation peaked in the early morning under dim light conditions below the compensation irradiance, keeping the mat largely anoxic (16, 29, 30). Based on these previous observations, we hypothesized that H_2 formation in Mushroom Spring and Octopus Spring cyanobacterial mats would exhibit a peak in the early morning due to N_2 fixation and a peak after sunset that might be caused by fermentation and N_2 fixation supplied with ATP from fermentation.

In this study, we tested this hypothesis by quantifying the diel H_2 dynamics in a Mushroom Spring microbial mat at a hitherto unreached spatiotemporal resolution using a newly developed sulfide-insensitive H_2 microsensor (31). Earlier studies showed the potential for H_2 production in dark-incubated mats (22), including a few, albeit uncalibrated, H_2 microsensor measurements (23). The sensitivity of the previous H_2 microsensors to hydrogen sulfide strongly limited their applicability for studying the zonation and dynamics of H_2 metabolism in microbial mats. By use of the improved microsensor for H_2 in combination with microsensors for O_2 , H_2S , and pH, it was possible to describe the chemical microenvironment in the mat in great detail, enabling new insights into the regulation of H_2 dynamics.

MATERIALS AND METHODS

Mushroom Spring is a siliceous, slightly alkaline hot spring (2) located in the Lower Geyser Basin of Yellowstone National Park (Wyoming). The spring has a large source pool that drains 68 to 70°C water through a single effluent channel that gradually widens, while the spring water is cooling down. Our study site had a water temperature throughout the diel cycle of $58 \pm 3^{\circ}\text{C}$ and harbored a microbial mat with an ~ 1 -mm-thick green cyanobacterial layer on top of an orange undermat. Field measurements were conducted on 12 to 13 August 2014, and the associated laboratory experiments were done a few days earlier.

Most measurements were done *in situ*, but a few cores (8-mm diameter, about 1-cm deep) were sampled and brought back to the laboratory for preliminary analysis. Samples were transported to the laboratory in a 1-liter stainless steel thermos filled with *in situ* water and were then immediately placed in a 58°C water bath filled with stirred *in situ* water. A fiberoptic halogen lamp equipped with a collimating lens (KL-2500; Schott, Germany) illuminated the mat samples with a downwelling photon irradiance of $1000 \mu\text{mol photons m}^{-2} \text{s}^{-1}$, as determined with a photosynthetically active radiation (PAR) quantum irradiance sensor (MQS-B; Heinz Walz GmbH, Effeltrich, Germany).

Concentrations of dissolved H_2 and O_2 and pH were measured in the mat and overlying spring water using microsensors. The H_2 microsensors (Unisense A/S) had tip diameters of $70 \mu\text{m}$ and were made H_2S insensitive by placing a ZnCl_2 -propylene carbonate-based H_2S trap in front of the H_2 microsensor tip. The addition of the sulfide trap, furthermore, lowered the baseline current by lowering the H_2O influx to the anode so that the sensors could be used at temperatures up to about 60°C (31). The sensors optimally have a detection limit of about $20 \text{ nmol H}_2 \text{ liter}^{-1}$ (31), but at the high temperature of the spring and lack of continuous signal recording, the detection limit was $0.2 \mu\text{mol H}_2 \text{ liter}^{-1}$, corresponding to a 1-pA change in signal. The sensors were four-point calibrated (0 to $69 \mu\text{mol H}_2 \text{ liter}^{-1}$) in the laboratory at 58°C , verifying the linearity of the signal, and two-point calibrated (0 to $34 \mu\text{mol H}_2 \text{ liter}^{-1}$) in the field. The standards were prepared using dilutions of water saturated with H_2 at 20°C ($805 \mu\text{mol H}_2 \text{ liter}^{-1}$ at sea level pressure, but $693 \mu\text{mol H}_2 \text{ liter}^{-1}$ at the barometric pressure of Bozeman, MT, where the bottle was prepared) (32). This procedure may induce some inaccuracy in the calibration, as the temperature of our supposedly $693 \mu\text{mol H}_2 \text{ liter}^{-1}$ solution containing a H_2 headspace varied under field conditions, causing changes in the equilibrium between gas-phase and dissolved H_2 , and the temperature of the $34 \mu\text{mol H}_2 \text{ liter}^{-1}$ standard was also affected by the dilution. We estimate that the potential error in the calibration might be as high as 15%. The sensors gave a linear response of about 4 pA per $\mu\text{M H}_2$ at the *in situ* temperature of 58°C . The sensors needed about 30 s of equilibration at each depth and concentration profile; we therefore only measured with 0.5-mm depth resolution.

The O_2 microsensors (33) had a tip diameter of $10 \mu\text{m}$ and were made in our laboratory at Aarhus University (AU). They were linearly calibrated in the field based on sensor readings in anoxic parts of the mat (zero oxygen) and in spring water (58°C) equilibrated with air by vigorous shaking of a 50-ml centrifuge tube that was about 70% filled with spring water and with intermittent placement in the spring to keep the temperature near 58°C . At the atmospheric pressure for the 2,227-m elevation of Mushroom Spring, it is necessary to multiply the saturations from standard tables (<http://www.unisense.com/files/PDF/Diverse/Seawater%20&%20Gases%20table.pdf>) by a factor of 0.77, resulting in an O_2 concentration of $118 \mu\text{mol liter}^{-1}$. Due to nonperfect temperature control during air equilibration, we estimate a potential error in the calibration of up to 15%.

The pH microelectrode was made in our laboratory at AU and had a 250- μm -long pH-sensitive glass tip and a tip diameter of $40 \mu\text{m}$ (34). It was calibrated in pH 6.96 and pH 9.76 buffers at 58°C using a commercial reference electrode (REF201; Radiometer, Denmark) and yielding a sensitivity of 58 mV per pH unit.

Laboratory profiles of H_2S were measured with an amperometric H_2S

microsensor (35) having a tip diameter of 30 μm . The sensor was made in our laboratory at AU and was painted black to avoid light interference on the signal (35). The sensor was linearly calibrated from signal readings in H_2S -free water and in a standard prepared by dissolving a washed $\text{Na}_2\text{S} \cdot 9\text{H}_2\text{O}$ crystal in 10 $\mu\text{mol liter}^{-1}$ anoxic HCl (10 mmol liter^{-1}) to a final concentration of 200 $\mu\text{mol H}_2\text{S liter}^{-1}$. The calibration was done within a few minutes of preparation of the standard.

The signals from the O_2 , H_2 , and H_2S microsensors were read by a custom-built, battery-operated pA meter with a resolution of 10^{-13} A. The signal from the pH sensor was read with a custom-built, battery-operated high-impedance voltmeter with an internal resistance of 10^{14} ohms. Both meters were enclosed in plastic bags with silica gel during field measurements. In the laboratory and during daytime field measurements, the instruments were connected to a battery-operated strip-chart recorder (SE-110; Gossen-Metrawatt, Germany), while signals were read directly from the meter display during nighttime as steam from the hot spring water caused excessive condensation of water on all surfaces.

The microsensors were mounted and advanced vertically into the microbial mat with manually operated micromanipulators that were read with better than 10- μm accuracy. The micromanipulators were attached to heavy metal stands.

The field measurements were initiated in the late afternoon, but in the afternoon and evening only single profiles for all three parameters were measured. In the morning, this was changed to triplicate measurements for H_2 , but still only single profiles for O_2 and pH. The three locations for measurement of H_2 were chosen at random within the 2.5- by 2.5-cm square covered by the x - y positioning of the micromanipulator.

Simultaneous with the diel microsensor measurements, 1-liter bottles were filled to capacity with 58°C spring water, and an 8-ml headspace of atmospheric air was introduced into the bottles through a butyl rubber septum. After equilibration of the water with the gas phase via 30 s of vigorous shaking, the H_2 gas concentration in the headspace was immediately analyzed on-site using a portable gas chromatograph (GC) (490 micro GC; Agilent Technologies, Santa Clara, CA, USA) using a thermal conductivity detector connected to a 10-m unheated MolSieve 5-Å column, running with a flow of 99.9999% pure argon as the carrier gas. A standard curve for the analysis was obtained by adding various volumes of H_2 -saturated water to spring water sampled during daytime when there was no H_2 in the spring water. The detection limit of the GC was 4 $\text{nmol H}_2 \text{ liter}^{-1}$.

Temperature loggers (iButton DS1922L; Maxim Integrated Products, Inc., San Jose, CA, USA) were placed in the spring on the biofilm surface, logging the spring water temperature at 5-min intervals. Downward photon irradiance of PAR (400 to 700 nm) was monitored continuously in air throughout the field measurements with a PAR quantum irradiance sensor (MQS-B; Heinz Walz GmbH, Effeltrich, Germany) placed next to the spring and connected to a light meter (ULM-500; Heinz Walz GmbH).

Based on the *in situ* concentration profiles measured over a diel cycle, isopleths of pH, O_2 , and H_2 concentrations were generated using the software package Ocean Data View (<http://odv.awi.de>) based on weighted-average gridding.

RESULTS

The temperature of the spring water flowing above the mat was relatively constant over time, but lower air temperatures at night reduced the water temperature to $\sim 55^\circ\text{C}$ and solar exposure during daytime caused maximum temperatures of up to 60°C at noon (Fig. 1). The incident photon irradiance was fluctuating in the afternoon prior to 19:00 due to shifting cloud cover but exhibited a very smooth curve in the interesting regions around sunset and sunrise. A photon irradiance of $\sim 1\%$ of maximum noontime irradiance ($1,500 \mu\text{mol photons m}^{-2} \text{ s}^{-1}$) was reached at 20:20 in the evening and at 6:28 in the morning.

The chemistry of the Mushroom Spring mat changed dramati-

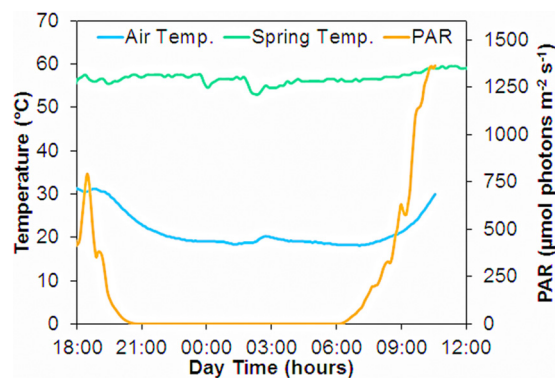


FIG 1 Photon irradiance (400 to 700 nm), air temperature 10 cm above the water, and water temperature at the mat surface measured throughout the diel experimental period. The dip in irradiance at 18:00 was due to cloud cover.

cally throughout the diel cycle. To illustrate this, pH, O_2 , and H_2 concentration profiles measured at noon and at sunrise (7:00) down to 4- to 6-mm depth in the mat are shown in Fig. 2. At sunrise, the irradiance was too low ($60 \mu\text{mol photons m}^{-2} \text{ s}^{-1}$) (Fig. 1) to cause a net production of O_2 , and O_2 only penetrated to about 0.3-mm depth. At noon, oxygenic photosynthesis caused an oxygen peak in the surface layer of $>800 \mu\text{mol O}_2 \text{ liter}^{-1}$, and O_2 was present in high concentrations down to the deepest analyzed layer at 5.25 mm. The pH values showed similar large changes with the water pH increasing from 7.4 at 7:00 to 8.5 at noon and pH at 1-mm depth increasing from 6.1 to 9.4 in the same period. Hydrogen was present down to 3-mm depth at 7:00, and a maximum concentration of $46 \mu\text{mol H}_2 \text{ liter}^{-1}$ was found at 1.5-mm depth. At noontime, no H_2 was detected in the mat.

Isopleths of pH, O_2 , and H_2 concentrations in the mat through the diel cycle (18:00 to 12:00) are shown in Fig. 3. The data points used for construction of the H_2 isopleths are the average values for three measurements in the period from 4:00 to 10:00.

The diel variation in the H_2 concentrations within the mat exhibited two peaks (Fig. 4): a broader peak with maximum concentrations of about $30 \mu\text{mol H}_2 \text{ liter}^{-1}$ after sunset (20:30 to 22:15) that faded away to reach about $11 \mu\text{mol H}_2 \text{ liter}^{-1}$ just before sunrise; and a sharp peak in the early morning at 7:00 under a low diffuse photon irradiance of $\sim 60 \mu\text{mol photons m}^{-2} \text{ s}^{-1}$. The average maximum H_2 concentration from the three profiles measured at 7:00 was $46 \mu\text{mol H}_2 \text{ liter}^{-1}$, but the profiles were very different with maximum concentrations at 1- to 1.5-mm depth of 49, 27, and $64 \mu\text{mol H}_2 \text{ liter}^{-1}$, respectively. The depth penetration of H_2 was also variable, with two profiles extending down to 2 mm, whereas one profile had a H_2 penetration of 4 mm. The deep H_2 penetration seen in Fig. 3 at 7:00 is thus due to one atypical profile and should not be overinterpreted. Similar variations among sites were also found in the other morning H_2 concentration profiles as evident from the large standard deviations in Fig. 4.

The first trace of H_2 ($0.7 \mu\text{mol H}_2 \text{ liter}^{-1}$) in the evening was found at 19:25 when there was still an irradiance of $170 \mu\text{mol photons m}^{-2} \text{ s}^{-1}$. An O_2 profile was recorded 10 min later (19:35; at an irradiance of $100 \mu\text{mol photons m}^{-2} \text{ s}^{-1}$), and at that time the O_2 peak had disappeared and the mat was anoxic below 0.5-mm depth. A similar high photon irradiance was needed to cause a significant buildup of O_2 in the mat in the morning, where

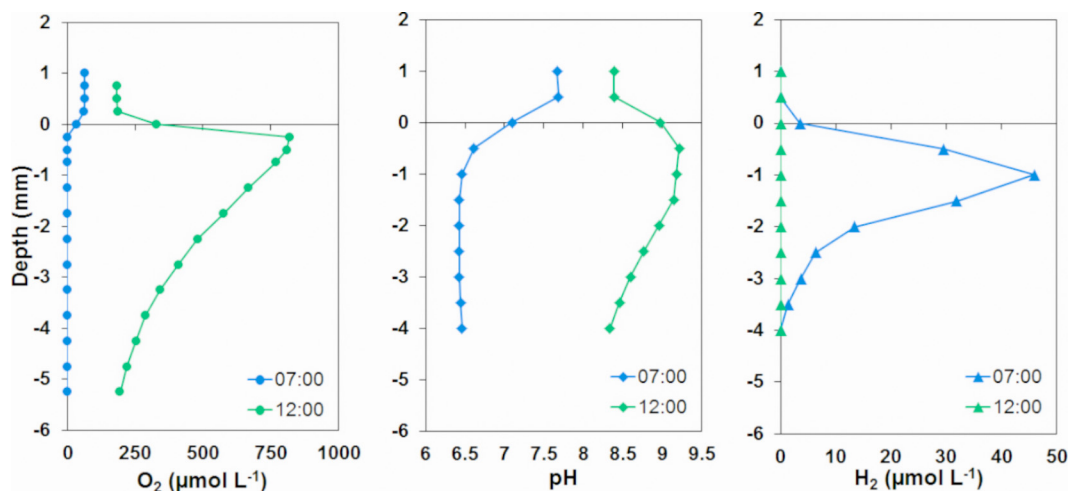


FIG 2 *In situ* vertical profiles of O_2 , pH, and H_2 levels in the Mushroom Spring mat measured during the period with maximum H_2 concentrations (7:00) and at noontime (12:00). The conversion factor from micromolar H_2 to parts per million by volume (ppmv) is $1 \mu\text{mol } H_2 \text{ liter}^{-1} = 1,613 \text{ ppmv}$ (at the elevation of Mushroom Spring).

a photon irradiance of $250 \mu\text{mol photons m}^{-2} \text{ s}^{-1}$ at 7:58 did not cause significant change. However, the following O_2 concentration profile measured at 8:22 ($350 \mu\text{mol photons m}^{-2} \text{ s}^{-1}$) exhibited an O_2 peak and a 0.75-mm increase in O_2 penetration depth compared to that for the dark profiles. The H_2 concentration decreased to one-third from 7:00 to 7:36, although we did not detect a change in the O_2 concentration profile until 7.58, where no H_2 was detected in the mat. The depth resolution of our measurements was, however, rather crude with O_2 being measured only every 0.25 mm and pH and H_2 every 0.5 mm. While a slight increase in O_2 penetration between 7:00 and 7:58 that went undetected cannot be ruled out, the O_2 concentration in the overlying water remained constant at 61 to $63 \mu\text{mol } O_2 \text{ liter}^{-1}$ during the measurements at 5:53, 6:26, and 6:53, before showing a slight increase to $91 \mu\text{mol } O_2 \text{ liter}^{-1}$ at 7:58. Consequently, there cannot have been any change in the O_2 profile until after 6:53 (and possibly considerably later), while the morning peak in H_2 concentrations in the mat was measured from 6:32 to 7:40.

The standard deviations on the measured H_2 concentrations are substantial, and based on the data from the H_2 microprofiles alone, we cannot conclude that the two peaks in H_2 accumulation are representative. However, the efflux of H_2 from the mat was investigated by GC analysis, and the data from these measurements also showed a two-peak pattern (Fig. 4). However, the H_2 content in the spring water was probably strongly affected by the H_2 emission from mats further upstream and thus presents some kind of integration of the H_2 efflux from mats growing at 58 to 70°C .

Concentration profiles of O_2 and H_2S were measured in mat cores that were brought back to the laboratory. We also attempted H_2S measurements in the field, but insufficient optical insulation of the H_2S microsensors (35) destroyed these before H_2S might have accumulated during the night. The laboratory measurements showed a maximal concentration of $70 \mu\text{mol } H_2S \text{ liter}^{-1}$ (nondissociated H_2S ; total sulfide is higher) at ~ 1.5 mm below the mat surface after 2 h of dark incubation, decreasing to $40 \mu\text{mol } H_2S \text{ liter}^{-1}$ at 6-mm depth (data not shown). An earlier study of the Mushroom Spring mat found even lower free H_2S concentrations

(23), while Dillon et al. (25) reported maximum concentrations of $250 \mu\text{mol } H_2S \text{ liter}^{-1}$ and substantial sulfate reduction rates in the upper mat layers. The H_2 sensor applied in the field was only tested for H_2S insensitivity up to $\sim 200 \mu\text{mol } H_2S \text{ liter}^{-1}$, but H_2S interference on the H_2 signal would show up as a high signal increasing with depth (9, 10), while all H_2 profiles showed zero concentration in deeper layers.

DISCUSSION

Earlier studies have also reported H_2 formation in hot spring microbial mats, where H_2 production may be due to fermentation (18, 19, 22, 36) or to formation as a by-product of nitrogenase activity during N_2 fixation (30). The relative importance of these processes for H_2 dynamics in hot spring cyanobacterial communities has remained unclear, while studies of hypersaline and coastal cyanobacterial mats have identified fermentation as the major H_2 source (7–9). In this study, we show that the H_2 production/accumulations in the Mushroom Spring mat correlate well with the patterns of nitrogenase expression and activity described in previous studies.

Most nonheterocystous cyanobacteria are prone to O_2 inhibition of the nitrogenase enzyme catalyzing N_2 fixation (37), although some can perform the process under fully oxic conditions (38–40). In cyanobacterial mats found at ~ 55 to 65°C in Mushroom Spring and nearby Octopus Spring, N_2 fixation is limited to conditions of hypoxia or anoxia in the mat during low irradiance and nighttime (29, 30). Dinitrogen fixation is a very energy-demanding process, and the finding of the highest levels of nitrogenase activity at sunrise was explained by the presence of light for ATP production, while the mat was still largely anoxic due to the consumption of O_2 within the mat being higher than the production (30).

We observed a peak in H_2 production exactly in the period where the cyanobacterial layer of the mat was still largely anoxic under a low incident photon irradiance of about $60 \mu\text{mol photons m}^{-2} \text{ s}^{-1}$. However, it cannot be ruled out that the H_2 production continued to be high later in the morning also, but that autotrophic or mixotrophic anoxygenic phototrophy by *Chloro-*

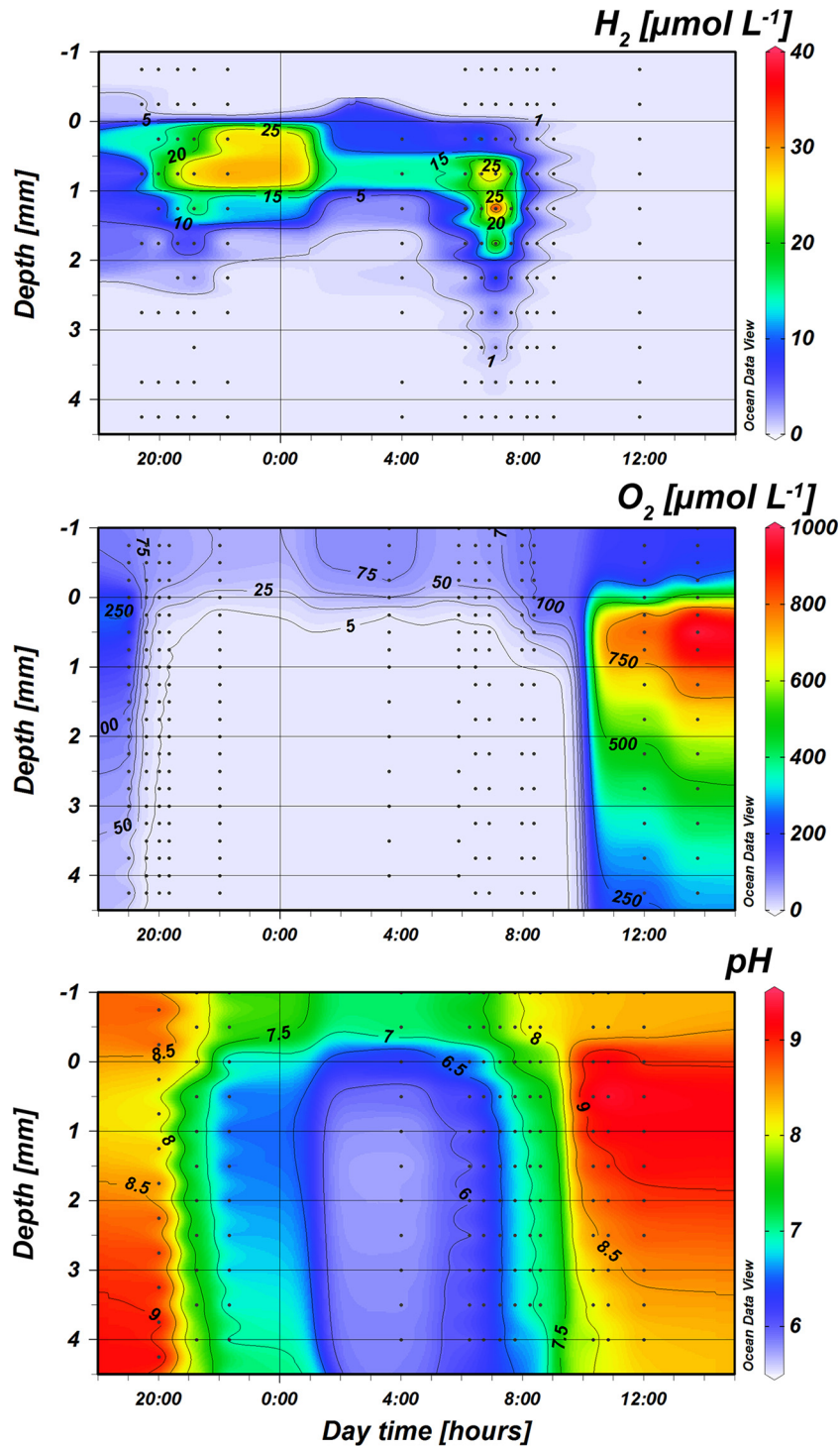


FIG 3 Isopleths of O_2 , pH, and H_2 depth distribution over a diel cycle in the 58°C Mushroom Spring mat. The measured profiles on which the interpolations were based are shown as dots. The last O_2 profile in the morning was recorded at 8:22 when the first peak of O_2 was recorded, the upper 1 mm of the mat became oxic, and H_2 became nondetectable. Notice that the interpolation for H_2 between about 23:00 and 3:00 is uncertain due to a lack of measuring points during this time interval.

flexus-like organisms (17, 23) and oxic metabolism based on cryptic oxygenic photosynthesis consumed this H_2 . Prokaryotes having high-affinity terminal oxidases are not restricted in their O_2 consumption until they experience concentrations of a few nano-

moles of O_2 per liter (41), and an environment with the simultaneous presence of H_2 and cryptic O_2 production by oxygenic photosynthesis would select for knallgas bacteria (i.e., bacteria oxidizing H_2 with O_2) having such high-affinity terminal oxidases.

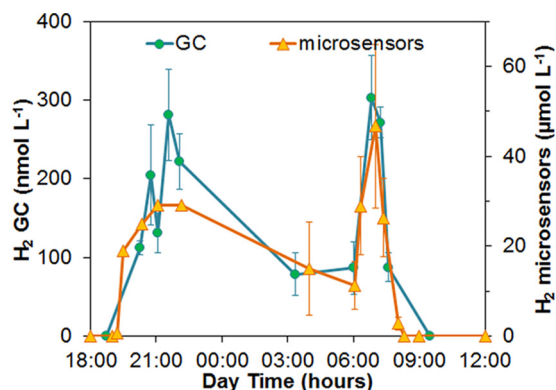


FIG 4 Maximum *in situ* H₂ concentrations in the 58°C mat as measured with microsensors, and H₂ concentrations in the overlying spring water as measured by GC during the diel cycle. Only one profile was measured for each time point in the evening, whereas three profiles at random positions were measured in the morning. Standard deviations are thus shown for the morning values.

High levels of nitrogenase gene expression (16, 29, 30) and N₂ fixation activity (30) have been observed at sunset in the Octopus Spring and Mushroom Spring mats. While we cannot rule out the possibility that all H₂ production in the afternoon-evening transmission was due to nitrogenase activity, the massive accumulation of fermentation products shown in dark-incubated mats from Mushroom Spring (19, 22) and the very low nighttime pH values measured in this study provide evidence that fermentation might be another source of H₂ accumulation in the mat. Previously measured accumulations of acetate plus propionate in dark-incubated vials containing a 1-cm-deep core of the mat (22) thus indicate that volatile fatty acids may build up to tens of millimolar concentrations in the upper millimeters of the mat during the night. Furthermore, metagenomic analysis of the mat has demonstrated the presence of many taxa that possess hydrogenase genes and should be capable of producing H₂ during fermentation (42).

The balance between nitrogenase- and fermentation-mediated H₂ production in a hypersaline mat dominated by *Microcoleus* sp. was investigated by Burow et al. (7), and they only found evidence of fermentation as a source of H₂, which was further supported in a recent study of H₂ dynamics in intertidal cyanobacterial mats (9). Diel changes in N₂ fixation in Mushroom Spring were investigated by Steunou et al. (30), who found early morning peak rates of nitrogenase activity (i.e., acetylene reduction) in the illuminated but largely anoxic mat that were >10 times higher than the nitrogenase activity measured just before sunset, indicating a pronounced potential stimulation of H₂ formation by low light. There are no known mechanisms whereby light stimulates fermentation, but low irradiance might provide energy for cyanobacterial N₂ fixation (and associated H₂ production) in three different ways: (i) by cryptic oxygenic photosynthesis producing both ATP and NADH, but where O₂ does not build up as the consumption in the mat is larger than the production, (ii) by respiratory use of the O₂ produced by cryptic oxygenic photosynthesis, and (iii) by photosystem I (PSI)-driven ATP formation without O₂ evolution by use of the Mehler reaction (43). We conclude that the high H₂ concentration measured in the Mushroom Spring mat during the morning was most likely due to light-driven alleviation of energy limitation fueling intense N₂ fixation. Diazotrophs other than

Synechococcus might in principle produce H₂, but the light stimulation of H₂ formation indicates that the dominating N₂-fixing organisms should be phototrophs, and the abundant anoxygenic phototrophs in the mat (*Roseiflexus* and *Chloroflexus*) have not been shown to express genes for N₂ fixation (30). Instead of producing H₂, these organisms may use H₂ as an electron donor (30). The quantitative contribution of fermentation to H₂ production in the Mushroom Spring mat remains unclear. Steunou et al. (30) found that N₂ fixation (nitrogenase activity) in the evening just after sunset was only 20% of the peak morning values at low light intensities, whereas we found only 1 1/2 times higher (H₂ profiles, Fig. 4) or similar (overlying water H₂, Fig. 4) H₂ concentrations in the morning peak compared to those in the evening. Just before sunrise, the nitrogenase activity was <5% of the peak rates found just 2 h later. The discrepancy between the nitrogenase activity pattern and H₂ accumulation indicates that fermentation might be a major factor responsible for the H₂ accumulation in the evening and night.

High morning rates of N₂ fixation and associated H₂ production might also be significant in other types of cyanobacterial mats. Diel variations in N₂ fixation in marine cyanobacterial mats have been subjected to investigation (44), but the expression of nitrogenase genes and rates of N₂ fixation were not found to correlate, and no morning peak in N₂ fixation was detected. However, if high morning rates were present, it is unlikely that they would result in massive H₂ accumulation as H₂ produced by nitrogenase activity in the early morning may be consumed rapidly by SRB that are much more abundant in marine and hypersaline mats and that may be very active after several hours of anoxia. It was thus shown that high concentrations of H₂ were found in a marine cyanobacterial mat throughout an 8-h dark period when molybdate was added as an inhibitor of sulfate reduction, while a mat without inhibitor addition was depleted of H₂ within a few hours (10). We note that SRB might potentially also contribute significantly to N₂ fixation in such mats, but a recent study only showed a minor contribution of SRB to N₂ fixation in a hypersaline mat (45).

We only found measurable H₂ levels down to 4.5 mm below the mat surface, and H₂ never accumulated in the deeper parts of the Mushroom Spring mat. High rates of sulfate reduction have been measured in the Mushroom Spring mat (25), and sulfate reduction might thus cause depletion of H₂ in deeper layers. In nearby Octopus Spring, methanogenesis was shown to be the major sink for H₂ in dark-incubated 55°C mats (22), where addition of the methanogenesis inhibitor bromoethane sulfonic acid (BES) led to massive efflux of H₂ from mat samples. An efflux of 0.2 µmol H₂ cm⁻² h⁻¹ can be calculated from the profile in Fig. 2 using Fick's first law of diffusion and the diffusion coefficient for H₂ at 58°C in water (9.7 × 10⁻⁵ cm² s⁻¹) (tables at www.unisense.com), and that is essentially identical to the H₂ formation rate after BES addition found earlier. Similar rates of methanogenesis (0.04 µmol CH₄ cm⁻² h⁻¹ corresponding to a consumption of 0.16 µmol H₂ cm⁻² h⁻¹) were measured in the 60°C Mushroom Spring mat (46), i.e., close to the site analyzed in this study (58°C). Both sulfate reducers and methanogens might thus have contributed to keeping the H₂ concentrations below the detection limit in deeper mat layers (>4.5-mm depth). High-sensitivity gas chromatography on anoxic mat material from deeper layers would probably have shown H₂ concentrations of 1 to 10 nmol liter⁻¹ as typically found in methanogenic and sulfate-reducing environ-

ments (47), but such concentrations are more than 1 order of magnitude lower than the detection limit of our sensors, which at the high temperature of the mat was $0.2 \mu\text{mol liter}^{-1}$. The mat was oxic far below 4.5-mm depth during the daytime (Fig. 3), but analysis of mat samples (from Octopus Spring) that were oxic during the daytime showed abundant presence of methanogenic archaea (10^7 to 10^8 cells g^{-1} wet weight), and pure cultures isolated from the samples survived hours of oxic conditions (24; N. P. Revsbech and D. M. Ward, unpublished results).

We found early morning pH levels ($<\text{pH } 6$) in the mats that were lower than the pH minimum of 7 measured earlier in Octopus Spring (48) and Mushroom Spring (49) mats, and we speculate whether a change in sensor calibration during the measuring period might be the explanation. A change in the reference potential might cause such effects, while it is not possible to obtain an increase in pH microsensor sensitivity significantly above the 58 mV per pH unit determined in the calibration. The measured pH at 2- to 3-mm depth was 1.4 pH unit below the water reading at 4:00 and as other investigations (48–50) found early morning pH values in the overlying water from 7.7 to 7.9 in similar settings, it can be concluded that the pH in the mat must have been below 6.5. Our site was far outside the main flow channel, and the overlying water pH might have been further lowered as seen in our data (pH 7.1 at 4:00) by a long contact time with the mat, resulting in a pH minimum in the mat considerably below 6.5. The extensive effects of mat-water exchange on the overlying water chemistry are also evident from the O_2 measurements showing minimum O_2 concentrations of $61 \mu\text{mol O}_2 \text{ liter}^{-1}$ from 6:00 to 7:00 and up to $185 \mu\text{mol O}_2 \text{ liter}^{-1}$ at 12:00.

Although we did not observe any change in O_2 profile in the mat until after 8:00 in the morning, with the first significant change observed at 8:22, a substantial change in the pH profile was observed. At 4:00, the overlying water was at pH 7.1, and the minimum pH at 1.5-mm depth was ~ 5.7 . At 7:12, coinciding with the maximum H_2 accumulation, the pH in the water had increased to pH 7.7, and the pH at 1.5-mm depth was pH 6.4. At 7:43, the pH in the water was 8.0, and the pH at 1.5-mm depth was 6.8. These rather large deviations in overlying water pH compared to that for the source water (pH 8.3) (23) indicate that there was high photosynthetic activity in the mat prior to 8:22, although this did not result in an accumulation of O_2 as the incident photon irradiance was at or below the compensation irradiance, where O_2 production and consumption processes balanced each other. It is not possible to estimate how much of this photosynthetic activity was due to the oxygenic cyanobacteria or the anoxygenic phototrophs in the mat, but photomixotrophic anoxygenic metabolism has been documented for the filamentous anoxygenic phototrophic bacteria, *Chloroflexus* spp. and *Roseiflexus* spp., that are abundant in the Mushroom Spring mat (19), and these organisms may thus both fix CO_2 and photoassimilate volatile fatty acids (VFAs) (22).

Both the N_2 fixation by *Synechococcus* and the photoassimilation of fermentation products are facilitated by the diel changes in chemistry and the light regime of the mat. While *Chloroflexus aurantiacus* and *Roseiflexus castenholzii* can be cultured in laboratory media supplied with, for example, acetate (27, 51), high concentrations of acetate are usually not found in static environments. Microbial specialization to utilize the metabolic opportunities created by cyclic changes in environmental conditions (e.g., 19) may be more common than generally realized. Other

well-known examples are the accumulation of polyphosphate by, e.g., “*Candidatus Accumulibacter phosphatis*” during the oxic part of oxic/anoxic cycles (52) and the accumulation of elemental sulfur in the beginning of a light period by purple sulfur bacteria (53).

In conclusion, our *in situ* measurements of the diel H_2 , O_2 , and pH dynamics in the Mushroom Spring microbial mat showed strong evidence for nitrogenase activity being a major source of H_2 accumulating in the mat during the evening and especially in the early morning, when dim irradiance alleviated energy limitation in the upper mat layers. Photosynthetic activity at such times with low irradiance may have profound effects on microbial communities by supplying energy to anoxygenic types of metabolism and possibly also oxic metabolism occurring at nondetectable O_2 concentrations.

ACKNOWLEDGMENTS

We thank Preben Sørensen and Lars Borregaard Pedersen for the construction of the microsensors, Emilio Garcia-Robledo for help with illustrations, and the US National Park Service and personnel from Yellowstone National Park for their permission to conduct this work and their helpful assistance.

This study was supported by the European Research Council (“Oxygen,” grant 267233; N.P.R.), the Grundfos Foundation (N.P.R.), an Elite Researcher Travel Stipend from the Danish Council for Independent Research (E.T.), an instrument grant from the Carlsberg Foundation (M.K.), and a Sapere-Aude Advanced Grant from the Danish Council for Independent Research/Natural Sciences (M.K.). D.M.W. appreciates the support from the Montana Agricultural Experiment Station (project 352).

FUNDING INFORMATION

This work, including the efforts of Niels Peter Revsbech, was funded by Grundfos Foundation. This work, including the efforts of Erik Trampe, was funded by Danish Council for Independent Research. This work, including the efforts of Michael Kühl, was funded by Danish Council for Independent Research. This work, including the efforts of Niels Peter Revsbech, was funded by EC | European Research Council (ERC) (267233). This work, including the efforts of David M. Ward, was funded by Montana Agricultural Experiment Station (MAES) (352). This work, including the efforts of Michael Kühl, was funded by Carlsbergfondet (Carlsberg Foundation).

The funders had no role in study design, data collection and interpretation, or the decision to submit the work for publication.

REFERENCES

1. Ward DM, Bateson MM, Ferris MJ, Nold SC. 1998. A natural view of microbial biodiversity within hot spring cyanobacterial mat communities. *Microbiol Mol Biol Rev* 62:1353–1370.
2. Ward DM, Bateson MM, Ferris MJ, Kühl M, Wieland A, Koeppel A, Cohan FM. 2006. Cyanobacterial ecotypes in the microbial mat community of Mushroom Spring (Yellowstone National Park, Wyoming) as species-like units linking microbial community composition, structure and function. *Philos Trans R Soc Lond B Biol Sci* 361:1997–2008. <http://dx.doi.org/10.1098/rstb.2006.1919>.
3. Awramik SM. 1984. Ancient stromatolites and microbial mats, p 1–22. *In* Cohen Y, Castenholz RW, Halvorson HO (ed), *Microbial mats: stromatolites*. Liss, New York, NY.
4. Rees DC, Howard JB. 2003. The interface between the biological and inorganic worlds: iron-sulfur metaloclusters. *Science* 300:929–931. <http://dx.doi.org/10.1126/science.1083075>.
5. Hoehler T, Bebout BM, Des Marais DJ. 2001. The role of microbial mats in the production of reduced gases on the early Earth. *Nature* 412:324–327. <http://dx.doi.org/10.1038/35085554>.
6. Skyring GW, Lynch RM, Smith GD. 1989. Quantitative relationships between carbon, hydrogen, and sulfur metabolism in cyanobacterial mats,

- p 170–179. In Cohen Y, Rosenberg E (ed), Microbial mats, physiological ecology of benthic microbial communities. American Society for Microbiology, Washington, DC.
7. Burow LC, Woeckel D, Bebout BM, McMurdie PJ, Singer SW, Pett-Ridge J, Prufert-Bebout L, Spormann AM, Weber PK, Hoehler TM. 2012. Hydrogen production in photosynthetic microbial mats in the Elkhorn Slough estuary, Monterey Bay. *ISME J* 6:863–874. <http://dx.doi.org/10.1038/ismej.2011.142>.
 8. Lee JZ, Burow LC, Woeckel D, Everroad RC, Kubo MD, Spormann AM, Weber PK, Pett-Ridge J, Bebout BM, Hoehler TM. 2014. Fermentation couples *Chloroflexi* and sulfate-reducing bacteria to cyanobacteria in hypersaline microbial mats. *Front Microbiol* 5:61. <http://dx.doi.org/10.3389/fmicb.2014.00061>.
 9. Hoffmann D, Maldonado J, Wojciechowski MF, Garcia-Pichel F. 2015. Hydrogen export from intertidal cyanobacterial mats: sources, fluxes, and the influence of community composition. *Environ Microbiol* 17:3738–3753. <http://dx.doi.org/10.1111/1462-2920.12769>.
 10. Nielsen M, Revsbech NP, Kühl M. 2015. Microsensor measurements of hydrogen gas dynamics in cyanobacterial microbial mats. *Front Microbiol* 6:726. <http://dx.doi.org/10.3389/fmicb.2015.00726>.
 11. Kothari A, Potrofká R, Garcia-Pichel F. 2012. Diversity in hydrogen evolution from bidirectional hydrogenases in cyanobacteria from terrestrial, freshwater and marine intertidal environments. *J Biotechnol* 162: 105–114. <http://dx.doi.org/10.1016/j.jbiotec.2012.04.017>.
 12. Burow LC, Woeckel D, Marshall IPG, Singer SW, Pett-Ridge J, Prufert-Bebout L, Spormann AM, Bebout BM, Weber PK, Hoehler TM. 2014. Identification of *Desulfobacterales* as primary hydrogenotrophs in a complex microbial mat community. *Geobiology* 12:221–230. <http://dx.doi.org/10.1111/gbi.12080>.
 13. Ward DM, Castenholz RW, Miller SR. 2012. Cyanobacteria in geothermal habitats, p 39–63. In Whitton BA (ed), *Ecology of Cyanobacteria*, 2nd ed. Springer Science+Business Media, Heidelberg, Germany.
 14. Bryant DA, Costas AMG, Maresca JA, Chew AGM, Klatt CG, Bateson MM, Tallon LJ, Hostetler J, Nelson WC, Heidelberg JF, Ward DM. 2007. *Candidatus Chloracidobacterium thermophilum*: an aerobic phototrophic acidobacterium. *Science* 317:523–526. <http://dx.doi.org/10.1126/science.1143236>.
 15. Klatt CG, Wood JM, Rusch DB, Bateson MM, Hamamura N, Heidelberg JF, Grossman AR, Bhaya D, Cohan FM, Kühl M, Bryant DA, Ward DM. 2011. Community ecology of hot spring cyanobacterial mats: predominant populations and their functional potential. *ISME J* 5:1262–1278. <http://dx.doi.org/10.1038/ismej.2011.73>.
 16. Liu Z, Klatt CG, Ludwig M, Rusch DB, Jensen SI, Kühl M, Ward DM, Bryant DA. 2012. “*Candidatus Thermochlorobacter aerophilum*”: an aerobic chlorophototrophic member of the phylum *Chlorobi* defined by metagenomics and metatranscriptomics. *ISME J* 6:1869–1882. <http://dx.doi.org/10.1038/ismej.2012.24>.
 17. Klatt CG, Liu Z, Ludwig M, Kühl M, Jensen SI, Bryant DA, Ward DM. 2013. Temporal metatranscriptomic patterning in phototrophic *Chloroflexi* inhabiting a mat in a geothermal spring. *ISME J* 7:1775–1789. <http://dx.doi.org/10.1038/ismej.2013.52>.
 18. van der Meer MTJ, Schouten S, Sinninghe Damsté JS, Ward DM. 2007. Impact of carbon metabolisms on ¹³C signatures of cyanobacteria and green nonsulfur-like bacteria inhabiting a microbial mat from an alkaline siliceous hot spring in Yellowstone National Park (USA). *Environ Microbiol* 9:482–491. <http://dx.doi.org/10.1111/j.1462-2920.2006.01165.x>.
 19. Kim Y-M, Nowack S, Olsen MT, Becraft ED, Wood JM, Thiel V, Klapper I, Kühl M, Fredrickson JK, Bryant DA, Ward DM, Metz TO. 2015. Diel metabolomics analysis of a hot spring chlorophototrophic microbial mat leads to new hypotheses of community member metabolisms. *Front Microbiol* 6:209. <http://dx.doi.org/10.3389/fmicb.2015.00209>.
 20. Nübel U, Bateson MM, Vandieken V, Kühl M, Ward DM. 2002. Microscopic examination of distribution and phenotypic properties of phylogenetically diverse *Chloroflexaceae*-related bacteria in hot spring microbial mats. *Appl Environ Microbiol* 68:4593–4603. <http://dx.doi.org/10.1128/AEM.68.9.4593-4603.2002>.
 21. Sandbeck KA, Ward DM. 1981. Fate of immediate methane precursors in low sulfate hot spring algal-bacterial mats. *Appl Environ Microbiol* 41: 775–782.
 22. Anderson KL, Tayne TA, Ward DM. 1987. Formation and fate of fermentation products in hot spring cyanobacterial mats. *Appl Environ Microbiol* 53:2343–2352.
 23. van der Meer MT, Schouten S, Bateson MM, Nübel U, Wieland A, Kühl M, de Leeuw JW, Sinninghe Damsté JS, Ward DM. 2005. Diel variations in carbon metabolism by green nonsulfur-like bacteria in alkaline siliceous hot spring microbial mats from Yellowstone National Park. *Appl Environ Microbiol* 71:3978–3986. <http://dx.doi.org/10.1128/AEM.71.7.3978-3986.2005>.
 24. Ward D, Beck E, Revsbech NP, Sandbeck KA, Winfrey MR. 1984. Decomposition of hot spring microbial mats, p 191–214. In Cohen Y, Castenholz RW, Halvorson HO (ed), *Microbial mats: stromatolites*. Liss, New York, NY.
 25. Dillon JG, Fishbain S, Miller SR, Bebout BM, Habicht KS, Webb SM, Stahl DA. 2007. High rates of sulfate reduction in a low-sulfate hot spring microbial mat are driven by a low level of diversity of sulfate-respiring microorganisms. *Appl Environ Microbiol* 73:5218–5226. <http://dx.doi.org/10.1128/AEM.00357-07>.
 26. Ward DM. 1978. Thermophilic methanogenesis in a hot-spring algal-bacterial mat (71 to 30 degrees C). *Appl Environ Microbiol* 35:1019–1026.
 27. van der Meer MTJ, Klatt CG, Wood J, Bryant DA, Bateson MM, Lammerts L, Schouten S, Sinninghe Damsté JS, Madigan MT, Ward DM. 2010. Cultivation and genomic, nutritional and lipid biomarker characterization of *Roseiflexus* strains closely related to predominant *in situ* populations inhabiting Yellowstone hot spring microbial mats. *J Bacteriol* 192:3033–3042. <http://dx.doi.org/10.1128/JB.01610-09>.
 28. Bhaya D, Grossman AR, Steunou AS, Khuri N, Cohan FM, Hamamura N, Melendrez MC, Bateson MM, Ward DM, Heidelberg JF. 2007. Population level functional diversity in a microbial community revealed by comparative genomic and metagenomic analyses. *ISME J* 1:703–713. <http://dx.doi.org/10.1038/ismej.2007.46>.
 29. Steunou AS, Bhaya D, Bateson M, Melendrez M, Ward DM, Brecht E, Peters JW, Kühl M, Grossman A. 2006. *In situ* analysis of nitrogen fixation and metabolic switching in unicellular thermophilic cyanobacteria inhabiting hot spring microbial mats. *Proc Natl Acad Sci U S A* 103: 2398–2403. <http://dx.doi.org/10.1073/pnas.0507513103>.
 30. Steunou AS, Jensen SI, Brecht E, Becraft ED, Bateson MM, Kilian O, Bhaya D, Ward DM, Peters JW, Grossman AR, Kühl M. 2008. Regulation of *nif* gene expression and the energetics of N₂ fixation over the diel cycle in a hot spring microbial mat. *ISME J* 2:364–378. <http://dx.doi.org/10.1038/ismej.2007.117>.
 31. Nielsen M, Larsen LH, Ottosen LDM, Revsbech NP. 2015. Hydrogen microsensors with hydrogen sulfide traps. *Sens Actuators B Chem* 215: 1–8. <http://dx.doi.org/10.1016/j.snb.2015.03.035>.
 32. Wiesenburg DA, Guinasso NL. 1979. Equilibrium solubilities of methane, carbon monoxide and hydrogen in water and sea water. *J Chem Eng Data* 24:359–360.
 33. Revsbech NP. 1989. An oxygen microelectrode with a guard cathode. *Limnol Oceanogr* 34:472–476.
 34. Revsbech NP, Jørgensen BB. 1986. Microelectrodes: Their use in microbial ecology. *Adv Microb Ecol* 9:293–352. http://dx.doi.org/10.1007/978-1-4757-0611-6_7.
 35. Kühl M, Steuckart C, Eickert G, Jeroschewski P. 1998. A H₂S microsensor for profiling sediments and biofilms: application in acidic sediment. *Aquat Microb Ecol* 15:201–209. <http://dx.doi.org/10.3354/ame015201>.
 36. Otaki H, Everroad RC, Matsuura K, Haruta S. 2012. Production and consumption of hydrogen in hot spring microbial mats dominated by filamentous anoxygenic photosynthetic bacterium. *Microbes Environ* 27: 293–299. <http://dx.doi.org/10.1264/jsme2.ME11348>.
 37. Fay P. 1992. Oxygen relations of nitrogen fixation in cyanobacteria. *Microbiol Rev* 56:340–373.
 38. Carpenter EJ, Price CC. 1976. Marine *Oscillatoria* (*Trichodesmium*)—explanation for aerobic nitrogen-fixation without heterocysts. *Science* 191:1278–1280. <http://dx.doi.org/10.1126/science.1257749>.
 39. Reddy KJ, Haskell JB, Sherman DM, Sherman LA. 1993. Unicellular, aerobic nitrogen-fixing cyanobacteria of the genus *Cyanothece*. *J Bacteriol* 175:1284–1292.
 40. Wilson ST, Foster RA, Zehr JP, Karl DM. 2010. Hydrogen production by *Trichodesmium erythraeum*, *Cyanothece* sp and *Crocospaera watsonii*. *Aquat Microb Ecol* 59:197–206. <http://dx.doi.org/10.3354/ame01407>.
 41. Gong X, Garcia-Robledo E, Schramm A, Revsbech NP. 2016. Respiratory kinetics of marine bacteria exposed to decreasing oxygen concentrations. *Appl Environ Microbiol* 82:1412–1422. <http://dx.doi.org/10.1128/AEM.03669-15>.
 42. Klatt CG, Inskeep WP, Herrgard M, Jay ZJ, Rusch DB, Tringe SG, Parenteau MN, Ward DM, Boomer SM, Bryant DA, Miller SR. 2013. Community structure and function of high-temperature chloropho-

- totrophic microbial mats inhabiting diverse geothermal environments. *Front Microbiol* 4:106. <http://dx.doi.org/10.3389/fmicb.2013.00106>.
43. Milligan AJ, Berman-Frank I, Gerchman Y, Dismukes GC, Falkowski PG. 2007. Light-dependent oxygen consumption in nitrogen-fixing cyanobacteria plays a key role in nitrogenase protection. *J Phycol* 43:845–852. <http://dx.doi.org/10.1111/j.1529-8817.2007.00395.x>.
 44. Severin I, Stal LJ. 2010. Temporal and spatial variability of *nifH* expression in three filamentous cyanobacteria in coastal microbial mats. *Aquat Microb Ecol* 60:59–70. <http://dx.doi.org/10.3354/ame01405>.
 45. Wobken D, Burow LC, Behnam F, Mayali X, Schintlmeister A, Fleming ED, Prufert-Bebout L, Singer SW, Cortes AL, Hoehler TM, Pett-Ridge J, Spormann AM, Wagner M, Weber PK, Bebout BM. 2015. Revisiting N₂ fixation in Guerrero Negro intertidal microbial mats with a functional single-cell approach. *ISME J* 9:485–496. <http://dx.doi.org/10.1038/ismej.2014.144>.
 46. Sandbeck KA, Ward DM. 1982. Temperature adaptations in the terminal processes of anaerobic decomposition of Yellowstone National Park and Icelandic hot spring microbial mats. *Appl Environ Microbiol* 44:844–851.
 47. Goodwin S, Lovley DR. 1988. Hydrogen concentrations as an indicator of the predominant terminal electron-accepting reactions in aquatic sediments. *Geochim Cosmochim Acta* 52:2993–3003. [http://dx.doi.org/10.1016/0016-7037\(88\)90163-9](http://dx.doi.org/10.1016/0016-7037(88)90163-9).
 48. Revsbech NP, Ward DM. 1984. Microelectrode studies of interstitial water chemistry and photosynthetic activity in a hot spring microbial mat. *Appl Environ Microbiol* 48:270–275.
 49. Jensen SI, Stenou AS, Bhaya D, Kühl M, Grossman AR. 2011. *In situ* dynamics of O₂, pH and cyanobacterial transcripts associated with CCM, photosynthesis and detoxification of ROS. *ISME J* 5:317–328. <http://dx.doi.org/10.1038/ismej.2010.131>.
 50. Ramsing NB, Ferris MJ, Ward DM. 2000. Highly ordered vertical structure of *Synechococcus* populations within the one-millimeter-thick photic zone of a hot spring cyanobacterial mat. *Appl Environ Microbiol* 66:1038–1049. <http://dx.doi.org/10.1128/AEM.66.3.1038-1049.2000>.
 51. Madigan MT, Petersen SR, Brock TD. 1974. Nutritional studies of *Chloroflexus*, a filamentous photosynthetic, gliding bacterium. *Arch Microbiol* 100:97–103. <http://dx.doi.org/10.1007/BF00446309>.
 52. Nguyen HTT, Nielsen JL, Nielsen PH. 2012. “*Candidatus* Halomonas phosphatis,” a novel polyphosphate-accumulating organism in full-scale enhanced biological phosphorus removal plants. *Environ Microbiol* 14:2826–2837. <http://dx.doi.org/10.1111/j.1462-2920.2012.02826.x>.
 53. van Gemerden H. 1974. Coexistence of organisms competing for the same substrate: an example among the purple sulfur bacteria. *Microb Ecol* 1:104–119. <http://dx.doi.org/10.1007/BF02512382>.