

Aggregate sound velocities and acoustic Grüneisen parameter of iron up to 300 GPa and 1,200 K

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Successful interpretation of available geophysical data requires experimental and theoretical information on the elasticity of solids under physical conditions of Earth's interior. Because iron is considered as major component in Earth's core, elastic properties of iron at high pressures and temperatures are very important for modeling its composition and dynamics. We use *in situ* x-ray diffraction data on ϵ -iron at static pressures up to 300 GPa and temperatures to 1,200 K to determine the Debye–Waller temperature factors and calculate aggregate sound velocities and Grüneisen parameter of ϵ -iron by using an approach that is based on Rietveld refinement at high pressures and temperatures.

The data on static compression of iron at ambient and elevated temperatures (1–4) provides important information on compressibility at the conditions present in Earth's deep interior. However, information on shear modulus, which is crucial for calculating sound velocities, is very limited. Mao *et al.* (5) published results of a study of elasticity of iron to a pressure of 220 GPa based on a novel method of determination of elastic moduli with diamond anvil cells (DAC) (6). This method utilizes the fact that in opposite anvil devices (like DAC) compression is not hydrostatic. As a result, positions of the reflections on the powder diffraction patterns depend not only on bulk modulus of the compressed material, but also on all of the elastic moduli and a value of deviatoric stress. The method cannot be used at high temperatures because deviatoric stress in DAC practically vanishes above ≈ 800 K (4). The mean sound velocities of *hcp* iron were extracted from phonon density measurements at pressures up to 153 GPa and ambient temperature (7, 8). Raman spectroscopy (9) of ϵ -Fe at pressures up to 150 GPa gave additional input to elastic properties of iron at high pressures. However, so far there are no experimental information on sound velocities of iron at pressure range of 300 GPa and high temperatures—conditions that are relevant to Earth's inner core.

The vibrations of individual atoms, as applied to intensities of x-ray diffraction peaks, control the Debye–Waller temperature factor B . This includes all wavelengths of phonons. As a result, the information on temperature factors has been used over years to determine Debye temperatures and Grüneisen parameters of different materials (10–15). The Debye temperature depends on aggregate elastic properties of a crystal—the adiabatic bulk modulus K_s and shear modulus G (16). Moreover, it was demonstrated by Anderson (16) that the Debye temperature is primarily dependent on the shear modulus, and this opens the way, as we are going to show in the present study, to extract aggregate shear and compressional sound velocities from data on temperature factors.

Following Grimvall's (12) consideration and using dispersion relation (10, 11)

$$\omega(q) = V_{\text{sound}} \left(\frac{2q_D}{\pi} \right) \sin \left(\frac{\pi q}{2q_D} \right), \quad [1]$$

where ω is the circular frequency, V_{sound} is the sound velocities, q is the wave number of a lattice mode, and q_D is the Debye cut-off wave number. We found for the Debye–Waller temperature factor (15) that

$$B = 4.105 \frac{\hbar}{M} \left(\frac{V}{N} \right)^{1/3} \left[\frac{2}{V_s} \int_0^{\pi/2} \frac{\text{cth} \left(1.471 \frac{\hbar V_s (N/V)^{1/3}}{k_B T} \sin t \right) t^2}{\sin t} dt + \frac{1}{\left(\frac{K_s}{\rho} + \frac{4}{3} V_s^2 \right)^{1/2}} \int_0^{\pi/2} \frac{\text{cth} \left(1.471 \frac{\hbar \left(\frac{K_s}{\rho} + \frac{4}{3} V_s^2 \right) (N/V)^{1/3}}{k_B T} \sin t \right) t^2}{\sin t} dt \right], \quad [2]$$

where \hbar is the normalized Planck constant, k_B the Boltzmann constant, T the temperature, N the Avogadro's number, V the molar volume, ρ the density, K_s the adiabatic bulk modulus, and M the molar mass.

At high temperatures Eq. 2 could be significantly simplified:

$$B = 7.205 \frac{k_B T}{M} \left(\frac{V}{N} \right)^{2/3} \left(\frac{1}{\left(\frac{K_s}{\rho} + \frac{4}{3} V_s^2 \right)} + \frac{2}{V_s^2} \right). \quad [3]$$

Eq. 3 makes clear that there is a correlation between shear sound velocities and the Debye–Waller temperature factor.

In Fig. 1, we compare the aggregate V_s for 30 metals with *bcc*, *fcc*, and *hcp* structures obtained from direct observations and calculated with Eq. 2 from temperature factors (10–12, 16–27). The standard deviation of calculated V_s from the experimentally observed shear sound velocities is less than 5%. For *hcp* metals (Fig. 1), the unit cell parameters (c/a) ratio covers the range from 1.55 to 1.65, which is the whole range expected for ϵ (*hcp*) iron (1, 4). For metals Zn and Cd ($c/a = 1.88$), V_s calculated with Eq. 2 differ from experimental ones by 30%. For α -Fe, aggregate shear moduli at pressures to 5 GPa were found (15), in good agreement with ultrasonic measurements. Therefore, if pressure and temperature dependence of the temperature factor, adiabatic bulk modulus, and molar volume are known, it is possible to estimate the aggregate sound velocity (both shear and compressional), as a function of P and T. Recently (4), we determined the equation of state for ϵ -Fe based on experimentally measured P–V–T data (pressures to 300 GPa and temperatures to 1,400 K). By combining those data with the data on the temperature factor of ϵ -Fe, as presented here, we can calculate the aggregate sound velocities of *hcp* iron at pressures corresponding to the physical conditions of Earth's core.

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Abbreviation: DAC, diamond anvil cells.

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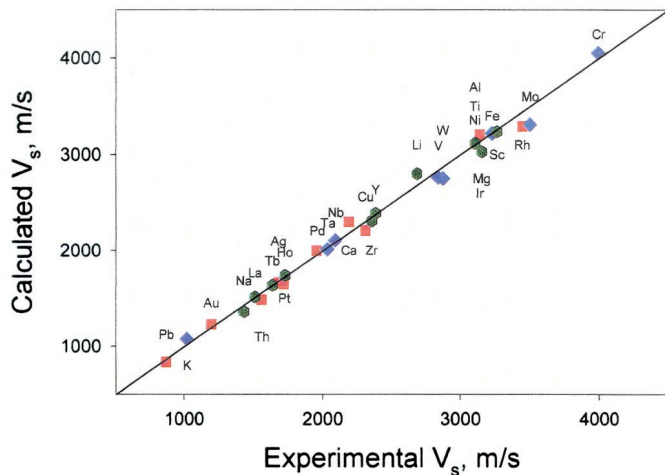


Fig. 1. Comparison of aggregate V_s obtained from direct observations and calculated with Eq. 2, using temperature factors for 30 metals with *bcc* (diamonds), *fcc* (squares), and *hcp* (hexagons) structures (10–12, 16–27). The standard deviation of calculated V_s from those experimentally observed is less than 5%. For *hcp* metals shown in the figure the c/a ratio covers the range from 1.55–1.65, or all reasonable values expected for ϵ (*hcp*) iron (1, 4).

Experimental Procedures

In situ x-ray experiments were performed on the beamline ID 30 at the European Synchrotron Radiation Facility (ESRF, Grenoble, France) and in Uppsala by using electrical (external and internal) heating. We also cross-checked volume determinations at high pressures and temperatures at Brookhaven National Laboratory (BNL, the beamline X-17C).

In Uppsala we obtained powder x-ray diffraction data with a Siemens x-ray system consisting of a direct-drive rotating anode generator (18 kW), a charge-coupled device (CCD) area detector (512×512 pixels), a zoom video system for sample visualization and alignment, and a capillary focusing system. MoK_α radiation is focused on a sample position of 30–40 μm full width half maximum (FWHM). Sample to detector distance varied from 170–310 mm.

At the ID30 beamline at ESRF two focusing mirrors provide a bright monochromatic (λ from 0.3738–0.4245 Å) x-ray beam in a $8 \times 10 \mu\text{m}^2$ spot at the sample location from two phased undulators and a channel-cut Si (111) monochromator. Spectra are collected with an image-plate system (FastScan or MAR345) placed 250–350 mm from the sample.

At BNL we used polychromatic x-ray radiation collimated to $10 \times 12 \mu\text{m}^2$ in FWHM. Powder diffraction data were collected with a Ge detector placed at the 2θ angle 8° .

In all experiments pressure was determined from the ϵ -Fe thermal equation of state (4) and, when possible, at pressures below 45 GPa, cross-checked from the equation of state of a pressure medium (Ar, NaCl, CsCl, etc.) (49, 50).

We heated the samples externally in a Mao–Bell-type DAC. Details of experimental set-up and a procedure are described elsewhere (4, 28). Examination of the samples following complete decompression after experiments did not reveal any signs of oxidation and/or contaminations.

Results and Discussion

The integrated intensities of a Bragg reflection is related to the Debye–Waller factor: $I_{\text{Bragg}} \approx \exp(-2B \cdot \sin^2 \theta_B / \lambda)$ (10). An experimental determination of $\ln(I_{\text{cor}})$ versus $(\sin^2 \theta_B / \lambda)$ (θ_B is the Bragg diffraction angle, λ the wavelength of radiation) should yield a straight line whose slope is $-2B$ (“Wilson plot”) (10). We define I_{cor} as an experimental intensity corrected for a

number of factors [polarization, absorption, extinction, thermal diffusive scattering (TDS), preferred orientation, etc.]. All of the corrections in experimental intensities can be introduced simultaneously for a whole diffraction pattern in the procedure of the Rietveld refinement, which allows accurate determination of the temperature factor (Method I) at elevated temperatures and pressures (29–31). However, the Rietveld refinement for diamond anvil cell experiments, especially at very high pressures, is associated with considerable difficulties (32). One problem is poor crystallite statistics (“spotty lines”). We used submicron iron powder, which allowed us to get smooth continuous lines even at pressures over 300 GPa and temperatures above 1,000 K (Fig. 2). Another problem is that a strong preferred orientation usually developed in *hcp* metals, including ϵ -Fe (1, 4). It leads to a decrease of intensity or even disappearance of some reflections [for example (002), reflection of ϵ -Fe; refs. 1, 4, and 32]. Although the Rietveld refinement program can handle the effect of preferred orientation, we found a strong correlation between temperature factor parameters and preferred orientations for ϵ -Fe. Therefore, determination of temperature factor from powder diffraction data for the samples with strong effects of texture is unreliable. Fortunately, if the initial sample consists of submicron particles and the phase α is transformed to ϵ phase at temperature around 800 K, preferred orientation in ϵ -Fe is almost absent (Fig. 2).

The most difficult problem in the Rietveld refinement is the existence of a deviatoric stress in DAC, because it can influence both the position and the shape of the reflections (5, 6, 32). At comparably low pressure (below 45 GPa), we conducted experiments in a soft pressure medium (Ar, CsI, CsCl, etc.), but at multimegabar pressures even helium could not provide hydrostatic conditions, and we did not use any pressure medium above 100 GPa. However, homogeneous heating above 1,000 K drastically reduces stresses (4, 32, 33) (Fig. 2b). To illustrate this statement we can use the equations of theory of diffraction at nonhydrostatic conditions (48). According to ref. 48, for a hexagonal sample under deviatoric stress condition, studied in the angle-dispersive mode with a loading direction parallel to the incident beam, one can write

$$d_{hk0} = H_{hk0} a \left[1 + \frac{t}{6 \langle G_{hk0} \rangle} (3 \sin^2 \theta_{hk0} - 1) \right], \quad [4a]$$

$$d_{00l} = H_{00l} c \left[1 + \frac{t}{6 \langle G_{00l} \rangle} (3 \sin^2 \theta_{00l} - 1) \right], \quad [4b]$$

where d_{hkl} is the d-spacing, a and c are the lattice parameters, θ_{hkl} is the Bragg diffraction angle, t is the component of deviatoric stress, $\langle G_{hk0} \rangle$ and $\langle G_{00l} \rangle$ are the average shear modulus in ($hk0$) and (00l) planes, respectively, H_{hkl} is the coefficient: $H_{100} = \sqrt{3}/2$, $H_{110} = 1/2$, $H_{002} = 1/2$, $H_{004} = 1/4$.

The expressions $\Delta_{hk0} = d_{100}/d_{110} - H_{100}/H_{110}$ and $\Delta_{00l} = d_{002}/d_{004} - H_{002}/H_{004}$ could be used to evaluate the stress state in the sample—if Δ_{hk0} and Δ_{00l} deviate from 0, material is under stress. For the sample compressed at ambient temperature to 273(2) GPa we obtained $\Delta_{hk0} = 0.0082(3)$ and $\Delta_{00l} = 0.0094(3)$; upon heating at 274(2) GPa and 640(10) K $\Delta_{hk0} = 0.0084(3)$ and $\Delta_{00l} = 0.0091(3)$; at 273(2) GPa and 840(10) K $\Delta_{hk0} = 0.0011(3)$ and $\Delta_{00l} = 0.0014(3)$, and at 273(2) GPa and 1,115(10) K $\Delta_{hk0} = 0.0003(3)$ and $\Delta_{00l} = 0.0002(3)$. In other words, with temperature increase at 273(2) GPa and 1,115(10) K, stresses in the sample vanish and an accurate structural refinement is possible (Fig. 2b).

Thermal vibration in *hcp* metals could be completely characterized by two anisotropic temperature factors, B_a and B_c (10). Our attempts to refine B_a and B_c separately for ϵ -Fe led to the same values within the accuracy of calculations. This result could mean that atomic thermal vibrations in ϵ -Fe are really close to isotropic, or that our powder diffraction data are not sensitive

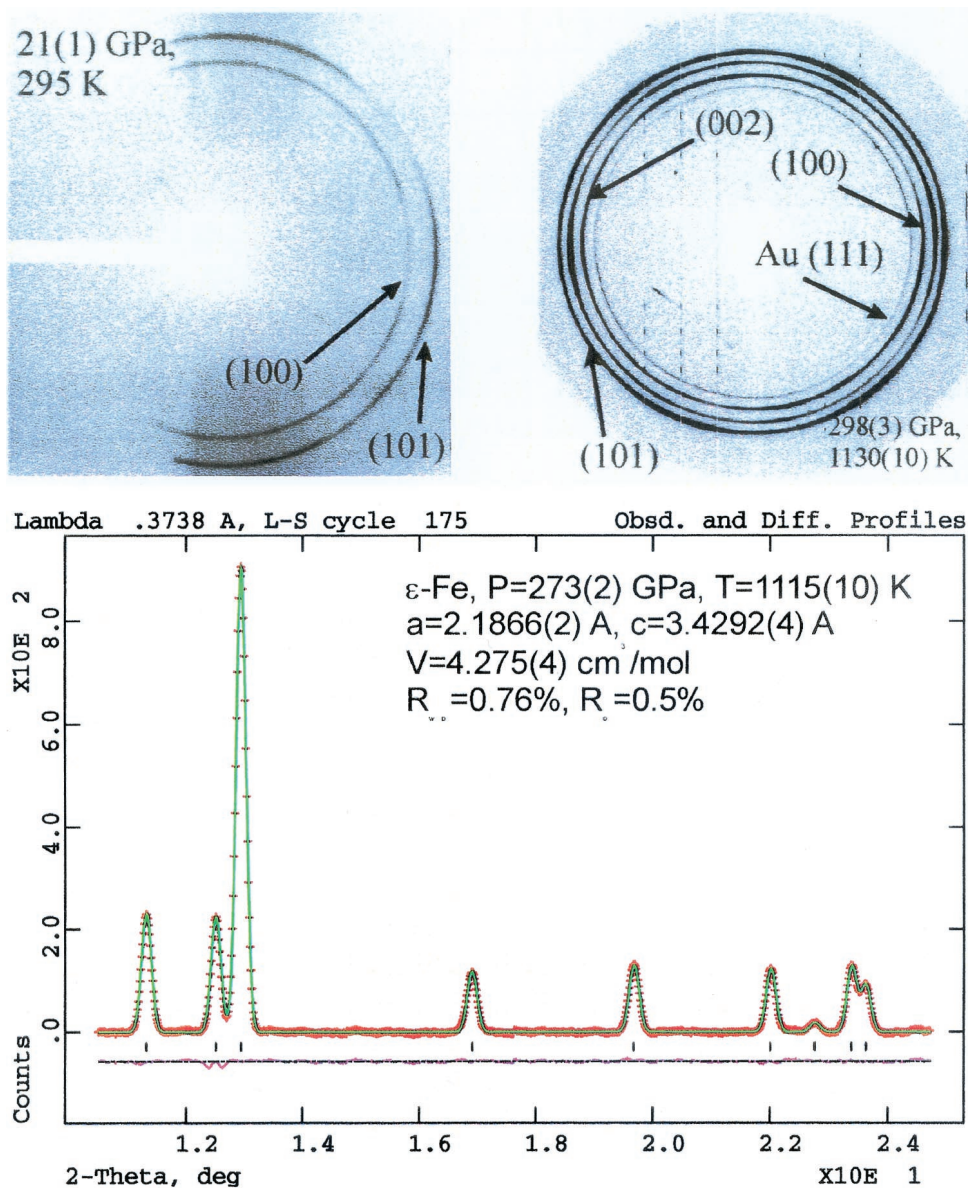


Fig. 2. (a) Examples of images collected on the ID30 beam line at the European Synchrotron Radiation Facility with monochromatic 0.3738 Å radiation with an image plate, demonstrating the dependence of the diffraction pattern on methods of samples preparation. (Left) ϵ -Fe was synthesized under compression at room temperature in fluid pressure medium from 2- μ m-thick iron foil. At 21(1) GPa the (002) diffraction line is absent. (Right) ϵ -Fe at 298(3) GPa and 1,130(10) K was synthesized from submicron particles and transformed from α to ϵ phase at temperature 800–850 K. The preferred orientation in ϵ -Fe is almost absent and the reflections (100), (002), and (101) are presented by continuous lines. Spots are due to diamonds. (b) Typical example of analyzed integrated patterns of the spectrum collected at 1,115(10) K and 273(2) GPa. The GSAS program package (30) was used. The background is subtracted.

enough for the Rietveld refinement with anisotropic temperature factors.

Because the Rietveld refinement or the Wilson plot is sensitive to the deviatoric stress, one may determine the temperature factor from the temperature dependence of the integrated intensity of a reflection I_{hkl}^T at a given pressure (10, 34, 35) (Method II):

$$\ln\left(\frac{I_{hkl}^T}{I_{hkl}^{T_0}}\right) = -\frac{2}{\lambda^2} (B(T)\sin^2\theta_B^T - B(T_0)\sin^2\theta_B^{T_0}) \quad [5]$$

(intensities are corrected for thermal diffuse scattering and T_0 is the reference temperature). If the temperature factor is obtained by the Rietveld refinement method at a certain temperature T_0 (for example, at high pressure and sufficiently high

temperature, when deviatoric stresses are low), the temperature factor at another temperature could be found by using Eq. 2. We applied this method for (100), (002), and (101) reflections of ϵ -Fe (for those lines we were able to record full Debye rings) and found that the temperature factors, determined from any of those lines, are the same within the precision of measurements (Fig. 3a). Therefore, anisotropy of ϵ -Fe does not change significantly with increasing temperature (at least in the temperature range of our study up to $\approx 1,300$ K).

The temperature factor of ϵ -Fe was determined by the two methods for 46 different P–T points (Fig. 3b). The phonon density measurements of ϵ -Fe at pressures up to 42 GPa and ambient temperature conducted by Lübbbers *et al.* (7) allowed for the extraction of information on mean-squared thermal displacements $\langle x^2 \rangle$, related to the Debye–Waller temperature factor $B =$

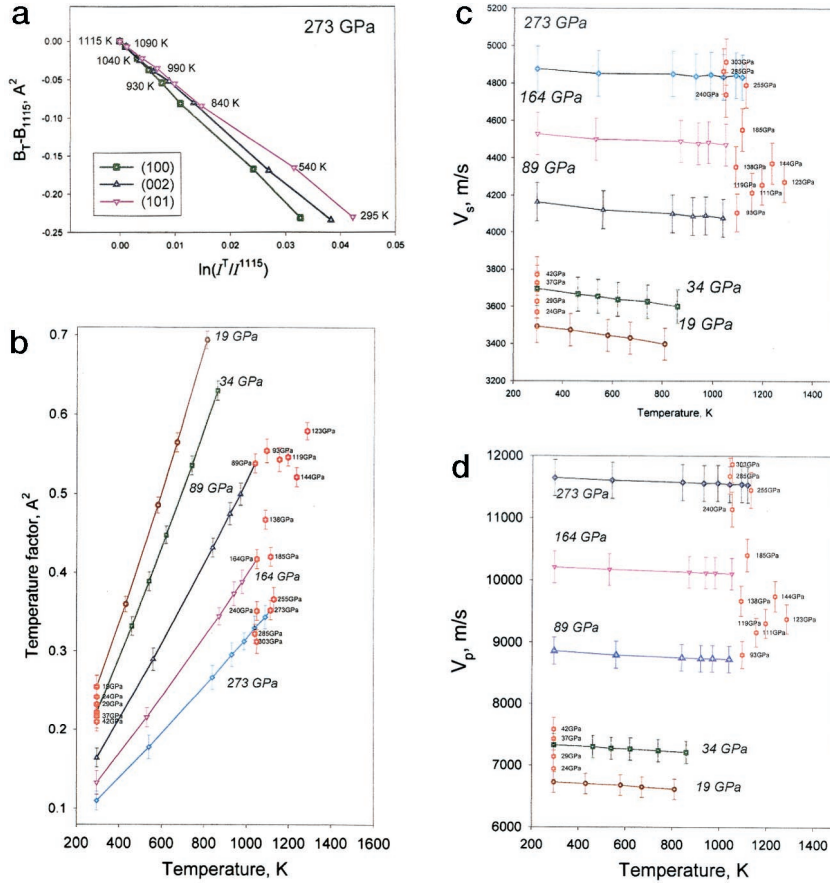


Fig. 3. (a) Dependence of the difference between temperature factor at temperature T (B_T) and 1,115 K (B_{1115}) on $\ln(I_T^2/I_{1115}^2)$ for the (100), (002), and (101) reflections of ϵ -Fe (for these lines we registered full Debye rings). (b) Temperature factor of ϵ -Fe determined by different methods (I or II) for 46 different P-T points. Red hexagons present B obtained from the Rietveld refinement (Method I). Continuous lines correspond to B obtained for constant pressures by Method II. (c) Shear waves V_s of ϵ -Fe found from temperature factors (symbols are the same as in b). (d) Compressional waves V_p of ϵ -Fe found from temperature factors (symbols are the same as in b).

$8\pi^2\langle x^2 \rangle$. It is remarkable that mean-squared thermal displacements, which are usually difficult to reproduce by different methods (18, 20), came out reasonably close; at 42 GPa, for example, Lübbbers *et al.* (7) reported 0.00203(7) \AA^2 , whereas our value at the same pressure is 0.0019(1) \AA^2 . [Note that the value given by Lübbbers *et al.* (7) for α -Fe at ambient conditions, 0.00413(7) \AA^2 , significantly deviates from the value recommended by the Neutron Diffraction Commission of the International Union of Crystallography, 0.00443 \AA^2 (20).]

Measured at different pressures and temperatures, temperature factors of ϵ -Fe were used for calculating the V_s from Eq. 2 (Fig. 3c). Together with the data on K_T , α , and dependence of V_s on molar volume (4, 16), we may also determine V_p (Fig. 3d). We should note that the thermal equation of state (4) provides the isothermal bulk modulus K_T , which is related to the adiabatic bulk modulus as $K_s = K_T(1 + \alpha\gamma T)$ (α is the thermal expansion, γ the Grüneisen parameter). Numerical estimates show that differences between K_T and K_s for ϵ -Fe (36) are less than 2% at temperatures lower than 1,400 K, introducing a negligible error in sound velocities calculated from Eq. 2. If the dependence of the Grüneisen parameter γ on volume has the form of $\gamma = \gamma_0(\rho_0/\rho)^q$ (ρ is the density, index 0 refers to the reference state 1 bar and 295 K, q is the constant) (16, 21), one can get the following equation for the sound velocity V_i (“i” could be “p” or “s” for compressional or shear waves, correspondingly)

$$V_i = V_{i0} \left(\frac{\rho_0}{\rho} \right)^{1/3} \exp \left[- \frac{\gamma_{i0}}{q_i} \left(\left(\frac{\rho_0}{\rho} \right)^{q_i} - 1 \right) \right]. \quad [6]$$

In Eq. 5 we used the acoustic γ_{ac} approximation to the Grüneisen parameter γ : $\gamma_{ac} = 1/3(\gamma_p + 2\gamma_s)$ (16, 37). The results of determination of parameters of Eq. 5 by fitting all of the available data (Fig. 3) on sound velocities as a function of volume are presented in Table 1. We found that Eq. 5, with parameters from Table 1, reproduces data both on V_s and V_p with accuracy better than 1% and we used it for further extrapolation of the sound velocities to the conditions of Earth’s inner core. At the temperatures of Earth’s core (5,000–7,000 K) the electronic contribution together with lattice vibrations could affect the total Grüneisen parameter. However, according to the modern *ab initio* theoretical calculations (45) the lattice contribution dominates the total Grüneisen parameter at high compressions. For example, it was found (45) that the total and the lattice Grüneisen parameters were practically the same at 3,000 K and 280 GPa (extreme point probed by calculations in ref. 45). Based on the

Table 1. Parameters of Eq. 5 as obtained by fitting all available data on sound velocities as a function of volume

	Shear sound velocity	Compressional sound velocity
γ_{i0}	1.549 (8)	2.144 (38)
Q_i	0.852 (18)	0.656 (62)
V_{i0} m/s	3142 (28)	5787 (26)
ρ_0^* , g/cm ³	8.298	8.298

*Fixed on value for $P = 1$ bar and $T = 295$ K (4).

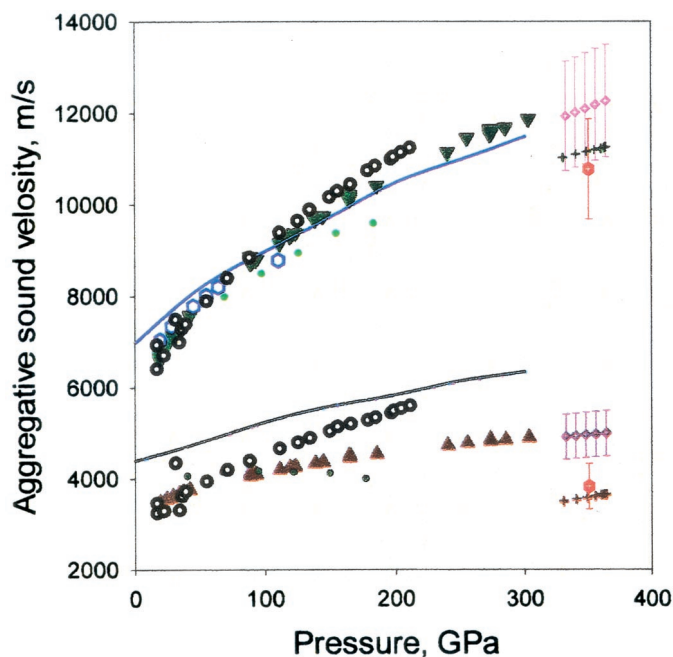


Fig. 4. Comparison of the present results (triangles) with theoretical (continuous lines) (8–40), static compression (open circles) (5, 9), inelastic x-ray scattering (open hexagons) (47), shock-waves studies (dotted lines) (41, 42), and with seismic observations (PREM) (43) in the inner core (crosses). Extrapolations of sound velocities to the inner core pressure at ambient temperature (based on Eq. 5, with parameters from Table 1) are shown by diamonds with errors bars. Estimation of sound velocities for the inner core conditions (350 GPa and 6,000 K), using $1/V_p \cdot \Delta V_p / \Delta T$ and $1/V_s \cdot \Delta V_s / \Delta T$, are shown by red hexagons.

theoretical data in refs. 45 and 46 and our data for γ_{ac} , we estimated maximum possible corrections in sound velocities at 330 GPa and 5,000 K as 2.5%. This value is well below the uncertainties associated with the theoretical calculations in refs. 45 and 46 or with the determinations of sound velocities from Eq. 2.

The longitudinal wave velocities in iron to 110 GPa measured by inelastic x-ray scattering (47) are in quantitative agreement with our data (Fig. 4). At ambient temperature, aggregate sound velocities, obtained in the present study, are close to those found in refs. 5 and 8 to moderate pressures (≈ 100 GPa), but at higher

pressures our value of V_s is lower and for V_p higher than the values in refs. 5 and 8. Such differences could be, for example, due to the effect of nonhydrostatic conditions (5, 8), which should be not so significant in our experiments, because samples were thermally relaxed. Still, although the mean values of sound velocities obtained in the present study and ref. 5 are different, the intervals of uncertainties are overlapping. For example, at 298 K and 211 GPa, the lowest estimated value of V_s in ref. 5 is 5.05 km/s, and the upper limit according to our data is 5.17 km/s. Theoretical calculations at 0 K (38–40) predict higher V_s , in comparison with our room- or high-temperature data, but our results (Fig. 3a) support comparably low anisotropy of ϵ -Fe as found by *ab initio* studies (40). The high temperature shock-wave Hugoniot sound velocities for ϵ -Fe to 200 GPa (41, 42) are close to our data (Fig. 4), but the difference becomes greater at high pressure because of high temperature along Hugoniot. Combining the Hugoniot data on V_p and V_s at 4,407 K and 200 GPa with low-temperature data from the present study, we estimated $1/V_p \cdot \Delta V_p / \Delta T = -2.3 \cdot 10^{-5} \text{ K}^{-1}$ and $1/V_s \cdot \Delta V_s / \Delta T = -3.9 \cdot 10^{-5} \text{ K}^{-1}$ in good agreement with phenomenological calculations by Stacey (36). Moreover, the acoustic Grüneisen parameter γ_{ac} calculated from our data (Table 1) for 5,000 K and 330 GPa is equal to 1.207, which is close to the value of the Grüneisen parameter obtained directly from considerations of the physics of Earth's core and the Preliminary Reference Earth Model (PREM) (36, 43, 44). The large ($4\text{--}6\times$) extrapolation of our data in temperature to the inner core conditions requires caution, because when close to the melting temperature the properties of a solid could change significantly. However, by using $1/V_p \cdot \Delta V_p / \Delta T$ and $1/V_s \cdot \Delta V_s / \Delta T$ estimated above, we can find $V_p = 10.8(1.1) \text{ km/s}$ and $V_s = 3.9(4) \text{ km/s}$ at 350 GPa and 6,000 K; these are close to the PREM values of $V_p = 11.18 \text{ km/s}$ and $V_s = 3.60 \text{ km/s}$. Considering the uncertainties in temperature and PREM values of sound velocities, and a possible change in the state of iron [in the inner core it could be β - or θ -phase (4, 10, 44)], agreement between PREM sound velocities and measured values is remarkable.

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