

REPLY TO GIRESSE ET AL.:

No evidence for climate variability during the late Holocene rainforest crisis in Western Central Africa

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Giresse et al. (1) criticize both our paleoclimatic reconstruction and our inferred anthropogenic origin of the late Holocene rainforest crisis (LHRC) (2). However, their argumentation, which is combined with alleged evidence for a climatic change during the LHRC, lacks strong support.

Citing studies describing both brief (weeklong) and limited periods of leaf wax production in deciduous trees, Giresse et al. (1) conclude that leaf waxes cannot record the environmental variability of a full season. However, this argument is flawed, as evergreen and subtropical deciduous trees produce leaf waxes over much longer timescales—their δD (δD_{wax}) values have been shown to capture environmental variability on even seasonal timescales (3, 4). In tropical Africa, the apparent decadal-scale covariation of sedimentary δD_{wax} records and instrumental records of mean annual precipitation (2, 5) proves that δD_{wax} accurately records hydrological changes. Any change in hydrological conditions (precipitation amount and/or seasonality) that would trigger the prominent vegetation change observed in Lake Barombi (Cameroon) during the LHRC is expected to be reflected in sedimentary δD_{wax} . Although our δD_{wax} record from Lake Barombi reveals a prominent drying trend during the Holocene, it shows no evidence for abrupt hydrological changes during the LHRC.

The arguments put forward by Giresse et al. (1) against human settlements in the Lake Barombi catchment during the LHRC contradict their own prior work. In their Letter (1), they posit that charcoal is absent in topsoils, but they previously reported on the abundance of charcoal fragments in the upper meters of soil sections in that area (6). Furthermore, the inferred absence of pottery fragments in topsoil samples (1) is merely a speculative argument against human impact as no rigorous archaeological survey has been carried out to date in the Lake Barombi catchment.

Giresse et al. (1) additionally point out the presence of savannah patches located 30 km southwest of Lake Barombi and use these to argue for a climatic driver for the LHRC. These presumably drier vegetation areas and the adjacent rainforest in the Lake Barombi area coexist under a rainfall regime with similar mean annual rainfall (of \sim 2,500 mm·y⁻¹), but with a slightly shorter rainy season (9 mo >50 mm·mo⁻¹) for the former compared with the latter (10 mo >50 mm·mo⁻¹) (7). Based on the existence of these savannah patches, Giresse et al. (1) infer that the replacement of rainforest by a forest-savannah mosaic during the LHRC was associated with a shorter rainy season. However, the authors (1) ignore the fact that these savannah patches have been cultivated and impacted by anthropogenic activity for a long time (8), which has involved repeated burning of the vegetation cover. This makes these savannah patches modern analogs of human impact instead of climate change.

Finally, Giresse et al. ignore the regional hydrological and sea-surface temperature reconstructions derived from continental (9) and marine (10, 11)

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records. Similarly to our δD_{wax} record from Lake Barombi, these records do not display abrupt changes at the onset of the LHRC,

which therefore puts the hypothesis of a regional climate crisis into question.

- **1** Giresse P, et al. (2018) Paleoclimatic changes are the most probable causes of the rainforest crises 2,600 y ago in Central Africa. *Proc Natl Acad Sci USA* 115:E6672–E6673.
- 2 Garcin Y, et al. (2018) Early anthropogenic impact on Western Central African rainforests 2,600 y ago. Proc Natl Acad Sci USA 115:3261-3266.
- 3 Sachse D, Dawson TE, Kahmen A (2015) Seasonal variation of leaf wax *n*-alkane production and δ^2 H values from the evergreen oak tree, *Quercus agrifolia*. *Isotopes Environ Health Stud* 51:124–142.
- **4** Huang X, Zhao B, Wang K, Hu Y, Meyers PA (2018) Seasonal variations of leaf wax *n*-alkane molecular composition and δD values in two subtropical deciduous tree species: Results from a three-year monitoring program in central China. *Org Geochem* 118:15–26.
- 5 Niedermeyer EM, et al. (2016) The stable hydrogen isotopic composition of sedimentary plant waxes as quantitative proxy for rainfall in the West African Sahel. Geochim Cosmochim Acta 184:55–70.
- 6 Giresse P, Maley J, Zogning A (2014) Lake level changes of Barombi Mbo (Cameroon) during the Late Quaternary: Compared catchment and crater lake records. New Studies on Former and Recent Landscape Changes in Africa, Palaeoecology of Africa 32, ed Runge J (CRC, London), pp 91–122.
- 7 Suchel JB (1972) La Répartition des Pluies et les Régimes Pluviométriques au Cameroun: Contribution à l'Etude des Climats de l'Afrique Tropicale (Centre d'Etudes de Géographie Tropicale, Talence, France; Centre de Recherches Africanistes, Yaoundé, Cameroon).
- 8 Letouzey R (1985) Notice de la Carte Phytogéographique du Cameroun au 1:500 000 (Institut de la Carte Internationale de la Végétation, Toulouse, France).
- 9 Shanahan TM, et al. (2009) Atlantic forcing of persistent drought in West Africa. Science 324:377–380.
- **10** Weldeab S, Lea DW, Schneider RR, Andersen N (2007) Centennial scale climate instabilities in a wet early Holocene West African monsoon. *Geophys Res Lett* 34:L24702.
- 11 Collins JA, et al. (2017) Rapid termination of the African Humid Period triggered by northern high-latitude cooling. Nat Commun 8:1372.