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Identifying and Quantifying the Intermediate Processes during Nitrate-Dependent Iron(II) Oxidation

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Abstract

Microbially driven nitrate-dependent iron (Fe) oxidation (NDFO) in subsurface environments has been intensively studied. However, the extent to which Fe(II) oxidation is biologically catalyzed remains unclear because no neutrophilic iron-oxidizing and nitrate reducing autotroph has been isolated to confirm the existence of an enzymatic pathway. While mixotrophic NDFO bacteria have been isolated, understanding the process is complicated by simultaneous abiotic oxidation due to nitrite produced during denitrification. In this study, the relative contributions of biotic and abiotic processes during NDFO were quantified through the compilation and model-based interpretation of previously published experimental data. The kinetics of chemical denitrification by Fe(II) (chemodenitrification) were assessed, and compelling evidence was found for the importance of organic ligands, specifically exopolymeric substances secreted by bacteria, in enhancing abiotic oxidation of Fe(II). However, nitrite alone could not explain the observed magnitude of Fe(II) oxidation, with $60-75\%$ of overall Fe(II) oxidation attributed to an enzymatic

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ASSOCIATED CONTENT

Supporting Information

The Supporting Information is available free of charge on the ACS Publications website at DOI: 10.1021/acs.est.8b01122. Tables including the stoichiometry of reactions used in the model, parameter uncertainty, potential enzymatic NDFO reactions, and calculated electron balances of select studies as well as additional figures including enzymatic Fe(II) oxidation rate for all strains investigated, solid green rust-Fe(II) oxidation rate calibration, simulation of Fe(II) concentrations during NDFO in Acidovorax sp. incubations, and model simulation results for nitrous oxide for Acidovorax sp. strain TPSY (PDF)

pathway for investigated strains: Acidovorax (A.) strain BoFeN1, 2AN, A. ebreus strain TPSY, Paracoccus denitrificans Pd 1222, and Pseudogulbenkiania sp. strain 2002. By rigorously quantifying the intermediate processes, this study eliminated the potential for abiotic Fe(II) oxidation to be exclusively responsible for NDFO and verified the key contribution from an additional, biological Fe(II) oxidation process catalyzed by NDFO bacteria.

Graphical Abstract

INTRODUCTION

Iron (Fe) minerals play an important role in the attenuation of contaminants within aquifer environments. The significant influence microorganisms have on Fe cycling has led to a burgeoning of research into bioremediation technologies that stimulate the formation and/or transformation of Fe minerals to remove metal(loid)s and radionuclides from groundwater. $1-8$ One of the most promising approaches for bioremediation relies on enhancing the activity of nitrate-dependent Fe oxidizing (NDFO) bacteria that couple Fe(ll) oxidation to nitrate reduction (Reaction 1): $9,10$

$$
10Fe(II) + 2NO_3^- + 12H^+ \rightarrow 10Fe(III) + N_{2(g)} + 6H_2O \quad 1
$$

These NDFO bacteria play an important role in the generation of Fe(III) and mixed valence minerals in subsurface environments, particularly in systems low in other oxidants such as manganese and nitrite.¹¹ Understanding how NDFO bacteria oxidize Fe(II) and how widespread this metabolic capability is among bacteria has been a focus of research since it was first observed in enrichment cultures.^{9,12} All neutrophilic NDFO bacterial strains isolated are mixotrophic, requiring an organic cosubstrate for growth. $13-15$ This complicates our understanding of NDFO as conflicting results in the literature have failed to resolve whether NDFO bacteria utilize an unknown enzymatic process to directly oxidize dissolved Fe(II) or if reactive intermediate compounds, such as nitrite or nitric oxide, abiotically oxidize Fe(II) during mixotrophic growth.^{16–18} While lithoautotrophic growth by neutrophilic NDFO pure-culture isolates have been described, their capacity for autotrophy has subsequently been questioned.^{19,20} Deciphering the NDFO mechanism using model enrichment cultures such as KS^9 , which is unambiguously capable of enzymatic NDFO, may also not necessarily resolve these outstanding questions as the interspecies interactions

and metabolic interdependencies²¹ are equally challenging to untangle. Without a model NDFO autotroph and the scarcity of genetic information explaining electron transfers pathways,18 quantitatively determining the contributions from interconnected abiotic and potential biotic reactions during mixotrophic NDFO is a substantial challenge.

Although the abiotic oxidation of dissolved $Fe(II)$ by nitrate is slow in the absence of a catalyst, 2^2 there are pathways involving nitrite that can facilitate abiotic Fe(II) oxidation. Bacteria capable of NDFO have been found to produce appreciable quantities of nitrite only when dissolved Fe(II) is supplemented in the basal growth medium.^{13,14,17,23,24} This nitrite can chemically oxidize dissolved Fe(II) at a relatively fast rate, which produces ferric (oxyhydr)oxides along with nitric (NO) and nitrous oxide (N_2O) gases, a process known as chemodenitrification (Reaction 2): $25-28$

 $4Fe(II) + 2NO_2^- + 5H_2O \rightarrow 4FeO(OH) + N_2O_{(g)} + 6H^+$ 2

Chemodenitrification experiments performed in anoxic sterile homogeneous media, using initial nitrite concentrations that fall within the range (1–4 mM nitrite) of what many NDFO incubation experiments accumulate, 13,14,23,29 showed extensive Fe(II) oxidation over 24 h.²⁵ The potential of these abiotic reactions to oxidize Fe(II) has raised questions of whether enzymatic NDFO exists and to what extent biotic Fe(II) oxidation contributes to their overall metabolism. Measuring the rate of oxidation can be a means of determining the effect of $(fast)$ abiotic Fe (II) oxidation. However, such measurements are often complicated by the presence of the mixed valence solid substrates, which make it difficult to determine the exact extent of oxidation and their catalytic activity.16,26,30,31 Nitrous oxide is also seldom analyzed in the headspace of NDFO cultures, which could otherwise be used to infer the extent of chemodenitrification or combined with stable isotope systematic measurements to discriminate between abiotic and biotic processes.25,32 Furthermore, it is difficult to critically review and interpret previous studies due to the fact that experimental conditions, such as the basal growth media, often differ between studies, which can influence Fe (trans)formations.17,33,34

While it is challenging to quantify NDFO-related processes experimentally due to the sensitivity of these geomicrobiological systems, developing biogeochemical models of NDFO experiments provides an effective means of quantifying the contribution of individual reaction pathways and evaluating the potential for strictly biological Fe(II) oxidation. So far, while a substantial number of experimental NDFO studies have generated comprehensive data sets, no attempts have been made to interrogate those data sets through process-based biogeochemical modeling, specifically quantifying the extent to which reactive intermediates contribute to overall dissolved Fe(II) oxidation.

The main objective of this study was therefore to derive a rigorous, data-constrained quantification of the relative contributions of biotic and abiotic processes during NDFO by integrating measured Fe(II) oxidation kinetics with suitable numerical modeling approaches based on previously generated data and conceptual models. While there is no clear evidence for the existence of an enzymatic pathway, assumptions were made on its theoretical

mechanism to specifically test whether it is required to explain the overall Fe oxidation observed or if chemodenitrification alone could be exclusively responsible. Extensive data sets generated from incubations with different NDFO bacteria were sourced from the literature and used to constrain the development of the numerical implementations of different conceptual models.

METHODS

Conceptual Models of Biogeochemical Processes and Reaction Network Implementation.

To quantify the rates of chemical and biological Fe(II) oxidation, a biogeochemical model was developed using the knowledge of the physiology and respiratory processes of NDFO bacteria and translating different conceptual models into process-based numerical models. The reactive processes that were considered in the modeling framework were limited to heterotrophic nitrate reduction to nitrite (Reaction 3) and nitrite reduction to nitrogen gas (Reaction 4),

$$
CH_3COO^- + 4NO_3^- \rightarrow 4NO_2^- + H_2O + HCO_3^- + CO_2(g)
$$
 3

$$
3CH_3COO^- + 8NO_2^- + 8H^+ \rightarrow 4N_2(g) + 7H_2O + 3HCO_3^- + 3CO_2(g) \quad 4
$$

In the case of enzymatic NDFO, microbial respiration couples dissolved Fe(II) oxidation and nitrate reduction, where nitrate is assumed to be exclusively reduced to nitrite:

$$
2Fe(II) + NO_3^- + 2H^+ \rightarrow 2Fe(III) + NO_2^- + H_2O \quad 5
$$

With a single exception,³⁵ NDFO bacteria have not yet demonstrated the capability of utilizing solid-phase Fe(II), and consequently, this source of Fe(II) was ignored. The nitrite produced from Reactions 3 and 5 can react with dissolved Fe(II) as per Reaction 2.

Two plausible conceptual models were used to explore NDFO: the first (S1) assumes that dissolved Fe(II) is oxidized through both an abiotic and a biotic pathway. S1 simulations were performed as three separate simulations (S1a, S1b, and S1c) using the three separate chemodenitrification rates determined from abiotic experiments reported under differing geochemical conditions. S1a represented conditions under which Fe(II) was complexed to the greatest extent by both inorganic and organic chelators; S1b represented conditions where $Fe(II)$ was complexed only by inorganic chelators, and S1c was where no $Fe(II)$ complexation occurred. The second conceptual model (S2) assumed that dissolved Fe(ll) oxidation is exclusively controlled by biogenic nitrite (i.e., chemodenitrification only). Conceptual models S1a,b,c and S2 were then translated into numerical models consisting of a mix of equilibrium and kinetically controlled reactions (Table 1).

Literature Data for Developing the Biogeochemical Model.

To develop the numerical model, important biogeochemical reactions affecting dissolved Fe(II) oxidation within NDFO growth cultures were identified and relevant data was collated from three comprehensive studies by Klueglein et al.,¹⁷ Carlson et al.,²⁴ and Chakraborty et al.14 Separately, the kinetic quantification of abiotic dissolved Fe(II) oxidation by biogenic nitrite was explored from chemodenitrification experiments performed by Kopf et al.,³⁶ Klueglein and Kappler, 16 and Jones et al. 25

Model results were compared against the measured data reported in the NDFO growth culture studies. The study by Klueglein et al.¹⁷ investigated four bacterial strains including the Acidovorax strain BoFeNl, Paracoccus denitrificans, and Pseudogulbenkiania sp. strain 2002, all grown simultaneously under balanced electron donor—acceptor conditions. The data sets from studies by Carlson et al. and Chakraborty et al. explored NDFO using the strains A delafieldii strain $2AN^{14}$ and A. ebreus strain TPSY²⁴ from the genus Acidovorax and were included here for a more robust model validation.

General Rate Law for Microbial Respiration.—Microbial respiration was modeled using the chemiosmotic model³⁷ and irreversible thermodynamics.³⁸ In this model, the overall microbial respiration rate, v (mol L⁻¹ s⁻¹), is described by eq 6:

$$
v = k[X]F_{\rm D}F_{\rm A}F_{\rm T} \quad 6
$$

where k is a rate constant (mol s⁻¹ mol biomass⁻¹), F_D and F_A are unitless kinetic factors, F_T is a unitless thermodynamic factor, and X is the biomass concentration (mol L⁻¹), in accordance with the formulation proposed by Jin and Bethke. ³⁸ Factors F_D and F_A (eqs 7) and 8) control the kinetics of the electron donating $([D]/[D^+])$ and accepting half reactions $([A]/[A⁻])$ as reactant and product concentrations change over time with constants K_D and K_A reflecting the standard free energy changes for each of the half reactions. Exponents β_D and β_A were assumed to equal 1 for all investigated microbial respiratory pathways.

$$
F_{\rm D} = \frac{{\rm [D]}^{\beta_{\rm D}}}{\rm [D]}^{\beta_{\rm D}} + K_{\rm D} {\rm [D^+]}^{\beta_{\rm D}} \quad 7
$$

$$
F_{A} = \frac{{[A]}^{\beta_{A}}}{[A]^{\beta_{A}} + K_{A}[A^{-}]^{\beta_{A}}}
$$
 8

Thermodynamic calculations (eq 9) used for evaluating the term F_T in the rate expression were based on commonly used thermodynamic data³⁹ and studies on aqueous $Fe^{2+}{}_{aq} - Fe^{3+}{}_{aq}$ oxide redox couples.⁴⁰

$$
F_{\rm T} = 1 - \exp\left(\frac{\Delta G + m\Delta G_{\rm p}}{\chi RT}\right) \quad 9
$$

where \tilde{G} is the Gibbs free energy of the reaction, R is the gas constant, T is absolute temperature, χ is the average stoichiometric number, m is the number of ATP molecules synthesized per electron transferred (assuming 3 protons are consumed per ATP synthesized), and G_p is the phosphorylation potential and assumed to be 50 kJ mol⁻¹, as in Jin and Bethke.⁴¹ The two key parameters that need to be defined for F_T are χ and m. However, as the mechanism for enzymatic NDFO is currently unknown, plausible values of χ and m were derived on the basis of the functioning of the anaerobic respiratory cycle.⁴² Values for both χ and m are proportional to the number of electrons transferred and in this instance were equal to 1 and 1/3, respectively, for the enzymatic NDFO reaction (Table S1).

Model Assumptions for Denitrification Processes.—The primary objective of the biogeochemical model development was to elucidate the contribution of biotic and abiotic processes on dissolved Fe(II) oxidation, not to provide a versatile, general tool to predict denitrification. Therefore, several reasonable assumptions allowed for simplification of the quantitative description of chemical and biological denitrification. (a) Microbial denitrification was modeled as a two-step process where nitrate is reduced to nitrite (Reaction 3) and nitrite, to dinitrogen gas (Reaction 4). Alternatively, nitrite could undergo chemical reduction via facile chemodenitrification.^{28,36,43} (b) Due to the overwhelmingly favorable thermodynamics of nitrate reduction, with acetate as the electron donor, the F_T term could safely be assumed to be unity in this study. Consequently, the generalized rate law (eq 6) was simplified to Reactions 10 and 11:

$$
v_1 = k_1[X] \frac{[\text{ Acetate }^-]}{[\text{ Acetate }^-] + K_D[\text{HCO}_3^-]} \frac{[\text{NO}_3^-]}{[\text{NO}_3^-] + K_A[\text{NO}_2^-]} \quad 10
$$

$$
v_2 = k_2[X] \frac{\text{[Acetate^-]}}{\text{[Acetate^-]} + K_{\text{D}}[\text{HCO}_3^-]} \frac{\text{[NO}_2^-]}{\text{[NO}_2^-]} + K_{\text{A}} \qquad 11
$$

where v_1 and v_2 are the microbial respiration rates for each step of denitrification, k_1 and k_2 represent the rate constants (mol s⁻¹ mol biomass⁻¹) for either nitrite or dinitrogen, K_D and K_A are unitless kinetic factors in v_1 but mol L⁻¹ for KA in v_2 (i.e. half saturation constant) for electron donating and accepting reactions, X is the biomass concentration (mol L^{-1}), and [NO₂⁻] represents the nitrite concentration (mol L^{-1}).

An electron balance calculation for experimental studies with BoFeN1 and Paracoccus denitrificans demonstrated that there is not enough reducing equivalents supplied from the growth medium to achieve both the observed extent of nitrate reduction and expected

increases in cell density using dissolved Fe(II) oxidation alone (Table S2).^{13,17} It was therefore concluded that an endogenous carbon source must have been utilized and was therefore included in the conceptual/numerical modeling framework. It was assumed that, (i) once exogenous carbon (i.e., acetate) is exhausted, bacteria are capable of switching to their stored endogenous carbon source and (ii) that this stored carbon source consists of polyhydroxybutyrate (PHB), a common product for bacteria synthesizing polyhydroxyalkanoates.44 The amount of stored PHB was calculated for each strain on the basis of the amount of nitrate reduced in a nongrowth medium, where no exogenous carbon was supplied (Figure S1; PHB quantities for BoFeN1 were based off *Acidovorax* sp. strain 2AN, as no literature data was available). The rate of PHB driven denitrification was similarly determined from eqs 10 and 11, using calibrated parameters (Table 2). For growth experiments, where it was apparent that the bacteria did not utilize any potentially stored endogenous carbon (e.g., where nitrate is not completely consumed but all acetate is exhausted), this process was deactivated. For results presented in Figure 2, only the simulations for BoFeN1 and Pd 1222 included this process.

The stoichiometric reactions for all denitrification reactions and associated biomass growth were derived using thermodynamic and bioenergetic literature data^{14,23,45,46} while applying the principles described in Rittmann and McCarty (Table S2). Where the specific growth yield coefficient was unknown, it was determined from the measured net growth in biomass per mole of acetate supplied. Biomass was represented using the generic formula, $C_5H_7O_2N$, ⁴⁷ and only assumed to increase while acetate, which was the electron donor and organic carbon source in all cases, was oxidized to carbon dioxide and bicarbonate during nitrate reduction. None of the bacteria investigated are capable of lithoautotrophic growth. Therefore, changes to biomass due to enzymatic NDFO was excluded. Biomass growth was described by eq 12:

$$
\frac{\text{d}[X]}{\text{d}t} = Yv - D_{\text{E}} \quad 12
$$

where t is time (s), Y is a biomass yield coefficient (mol cells per mole of acetate), v is the microbial respiration rate (mol L⁻¹ s⁻¹), and D_E is an encrustation term (mol L⁻¹ s⁻¹). A factor of 10^{-13} g cell⁻¹ was used to convert from cells mL⁻¹ to mol L⁻¹ using the formula for biomass given above. Due to the short duration of the investigated incubation experiments,^{14,16,24} biomass decay was considered to be insignificant. However, cells in batch incubations capable of NDFO quickly became heavily encrusted by Fe(III) or mixed valence mineral precipitates, $13,17$ especially within the periplasm, inhibiting cellular activity. ⁴⁸ Consequently, an encrustation inhibition term was employed to account for this process (eq 13):

$$
D_{\rm E} = v_2 \frac{[\rm{GR}-\rm{CO}_3] + [\rm{FeOOH}}{([\rm{GR}-\rm{CO}_3] + [\rm{FeOOH}]) + K_d} \quad 13
$$

Chemodenitrification.—Chemodenitrification has been detailed in many literature reports, where Fe^{2+} _{aq} reduces nitrite to nitrous oxide with a 2:1 stoichiometry (Reaction 2), and ferric oxyhydroxides^{25,28} or mixed valence Fe minerals are produced.^{28,30,32} The kinetics of chemodenitrification was described using a second order rate law (eq 14): 25,31

$$
-\frac{\mathrm{d[Fe(II)]}}{\mathrm{d}t} = k_3 \left[\mathrm{NO}_2^- \right] \left[\mathrm{Fe(II)}_{\mathrm{aq}} \right] \quad 14
$$

where k_3 is the pH-dependent rate constant (L mol⁻¹ s⁻¹), [Fe(II)] is the dissolved Fe(II) concentration (mol L⁻¹), t is time (s), and [NO₂⁻] is the nitrite concentration (mol L⁻¹). In this modeling study, chemodenitrification was defined as an extracellular abiotic oxidation process. The rate constants applied for S1a, S1b, and S1c were based on recent studies investigating chemodenitrification as a single process at pH 7.0. S1a represented conditions where dissolved Fe(II) was complexed by organic chelators and/or bicarbonate species by using rate constants reported in chemodenitrification experiments amended with 0.5 mM citrate in a bicarbonate basal medium (BBM) ;³⁶ S1b represented conditions without organic chelators where chemodenitrification experiments were performed in sterile BBM solution, ¹⁶ and S1c was used to explore the chemodenitrification rates for cases where no complexation of Fe(II) occurred and where experiments were not performed in BBM but instead buffered with PIPES.²⁵ Rate constants for S1b and S1c were derived from data sets using nonlinear regression based on the decay equations suggested by Kopf et al., 36 where S1a used the rate reported in the same study for 0.5 mM citrate. Model fits using the reported and derived rate constants are given in Figure 1 and range from 2.43×10^{-4} to 1.86 \times 10⁻³ mol L⁻¹ s⁻¹. For conceptual model S2, the chemodenitrification rate used in the simulations was the same as for S1a. These chemodenitrification experiments were all performed in anoxic, sterile environments at pH 7.0, similar to the NDFO growth experiments.16,17,25,36

Enzymatic NDFO.—Alongside chemodenitrification, enzymatic NDFO provided an alternative biological pathway for dissolved Fe(II) oxidation. Because the existence of a dedicated Fe(II) oxidoreductase for NDFO has not been demonstrated, the enzymatic NDFO pathway in this modeling study was broadly defined as any Fe(II) oxidation pathway dependent on cells. The mechanisms outlined by Carlson et al. 24 were primarily used to inform the numerical implementation of this separate respiratory process, where dissolved Fe(II) oxidation is potentially catalyzed by respiratory complexes (e.g., nitrate reductase, Nar, nitrite reductase, Nir).

Enzymatic NDFO was assumed to be suppressed until most of the acetate was consumed ϵ (<0.5 mM). This represented the point at which electrons donated by acetate to the quinone pool were thought to no longer be able to keep respiratory complexes reduced and continue

dissolved Fe(II) efflux pumping from the periplasm.²⁴ Consequently, dissolved Fe(II) accumulated and was oxidized via one or more of the potential mechanisms mentioned previously. Nitrate was assumed to be exclusively reduced to nitrite, supported by NDFO causing encrustation of the nitrite reductase enzyme within the periplasm, diminishing its activity and leading to the accumulation of nitrite.^{14,24,50} Direct oxidation of Fe^{2+} _{aq} to Fe^{3+} _{aq} by nitrate was not considered as it is not thermodynamically favorable (i.e., G^{0} _f > 0). Instead, for thermodynamic calculations within the F_T term, green rust carbonate (GR-CO3) was included as the Fe end product for enzymatic NDFO (see further explanation below, Table S3).

As the biochemical mechanism for NDFO is unknown, the parameters suggested by Jin and Bethke³⁸ for anaerobic respiration were used. The average stoichiometric number, χ , was taken to equal the number of times the rate-determining step occurs in the overall reaction, which is typically the proton translocation step during each instance of the quinone cycle. 38,42 Each instance of the quinone cycle consumes a pair of electrons and translocates two protons.⁴² Therefore, in the case of NDFO, where dissolved Fe(II) donates a single electron, χ and m are equal to 1 and 1/3, respectively (eq 15). Conceptually, this mechanism is the equivalent of assuming extracellular electron transport through a series of enzymes capable of direct dissolved Fe(II) oxidation to drive electron flow along the cell respiratory chain to reduce the quinone pool and ultimately generate proton motive force.^{18,21} The microbial respiration rate of NDFO was described by

$$
v_3 = k_4[X] \frac{\text{[Fe(II)]}}{\text{[Fe(II)]} + K_{\text{d}} \text{[NO}_3^-]} \left[1 - \exp\left(-\frac{\left(-\Delta G + \frac{1}{3} \times 50 \text{kJ} \cdot mol^{-1}\right)}{1 \times \text{RT}}\right) \right]
$$
 15

where v_3 is the microbial respiration rate for enzymatic NDFO and all parameters and units are consistent with eqs 10 and 11. The rate constants for enzymatic $Fe(II)$ oxidation were calibrated using data from chemolithotrophic growth experiments of Acidovorax spp. strains BoFeN1, 2AN, and TPSY, where dissolved Fe(II) and stored PHB were the only available electron donors (Figure S3). Quantities of stored PHB were estimated on the basis of the amount of nitrate consumed (Table S2).

Iron Mineral Formations and Transformations.—Due to its extremely low solubility at circumneutral pH, any Fe^{3+} _{aq} produced from Fe(II) oxidation would immediately precipitate in the solution medium. Green rust is known to be the initial precipitate in NDFO cultures.⁵¹ Green rust carbonate (GR-CO₃) was therefore included in the model and assumed to be an equilibrium phase. Furthermore, previous studies have found $GR-CO₃$ to be an intermediate product during microbial Fe(II) oxidation, $30,51$ which actively transforms to more stable Fe mineral products (e.g., goethite). In our study, structural Fe(II) in GR-CO₃ was modeled to abiotically reduce nitrite to ammonium while transforming to goethite.^{30,51} The reduction of nitrite by $GR-CO₃$ was modeled using an overall first order kinetic rate law:

$$
\frac{\text{d[NH}_4^+]}{\text{d}t} = k_5[\text{Fe(II)}_{\text{GR}}] \quad 16
$$

where k_5 is the rate constant (s⁻¹) and [Fe(II)_{GR}] is the moles of Fe(II) within green rust. Nitrate reduction by $[Fe(II)]_{GR}$ was not included as it only competes with microbial respiration at higher pH.⁵² The rate constant k_5 was calibrated using literature data from studies that investigated the rate of nitrite reduction by green rust sulfate,⁵³ but it should be noted that rates between different green rusts are not necessarily comparable.⁵⁴

Modeling Tools and Calibration.—Aqueous and solid phase biogeochemistry for all model variants were simulated with PHREEQC-3.55 Conceptual models S1a, S1b, S1c, and S2 contained 12 adjustable parameters and were initially calibrated using the heuristic particle swarm optimization^{56,57} method due to the severe nonlinearity common to similar geochemical models.58,59 The resulting parameter estimates were subsequently used as initial values for the Gauss-Levenberg-Marquardt method contained in $PEST++^{60}$ for final calibration refinement and sensitivity analysis; this two-step procedure is consistent with Rathi et al.58 The sum of squared residuals between l65 measurements and their associated model-simulated results was used as the objective function and minimized during both steps of the calibration procedure. The observation data used to constrain the models consisted of measurements of Fe(II), acetate, nitrate, nitrite, and biomass concentrations taken from studies by Klueglein et al.,¹⁷ Klueglein and Kappler,¹⁶ and Carlson et al.²⁴ Table 2 provides calibrated parameters used for S1a simulations for all bacteria with further pertinent statistical information provided in Table S1.

RESULTS AND DISCUSSION

Quantification of Chemodenitrification Rate.

To determine the contributions of both abiotic and biotic Fe oxidation that occur simultaneously, it is critical to properly estimate the chemodenitrification rate. The rate constant reported by Kopf et al.³⁶ for a chemodenitrification experiment amended with 0.5 mM citrate (36 mg C L⁻¹, broadly equivalent to previously reported concentrations of exopolymeric substances (EPS) of ~100 mg C L⁻¹, based on representative carbohydrate production61) provided a good model fit for the data set obtained from abiotic experiments that were sourced from an active growth experiment¹⁷ (Figure 1). Consequently, chemodenitrification was constrained using the rate constant derived for this experiment $(9.84 \times 10^{-4} \text{ mol L}^{-1} \text{ s}^{-1})$ in model S1a and provided the best overall fit for all NDFO experiments (Figure 2). Simulations of S1b, where the slower sterile BBM rate was employed ($k_3 = 2.43 \times 10^{-4}$ mol L⁻¹ s⁻¹), caused nitrite concentrations to be consistently elevated above the observed concentrations in order to compensate for the lower rate constant. Alternatively, simulated nitrite concentrations for S1c that used the faster rate constant derived from a PIPES buffered solution ($k_3 = 1.86 \times 10^{-4}$ mol L⁻¹ s⁻¹) were consistently below the observed nitrite concentrations.

The rate constant reported by Kopf et al.³⁶ and used in the S1a provided the best overall fit for a number of reasons. Chemodenitrification rates from experiments reported in Jones et al. and used here in S1c were performed in a Good's buffered medium (GBM), which differs from most NDFO growth cultures that commonly used a BBM. The different geochemical conditions between experiments, such as differences in potential mineral catalysts, are significant and likely altered the overall kinetics. The experiments performed by Klueglein and Kappler¹⁶ used a typical BBM medium, from which a chemodenitrification rate constant of 2.43 × 10⁻⁴ mol L s⁻¹ was derived. The problem with constraining chemodenitrification based on experiments utilizing sterilized BBM is that they fail to account for biologically produced Fe(II)-chelating organic ligands, $17,62,63$ that are known to exist within active NDFO cultures and to enhance dissolved $Fe(II)$ oxidation by nitrite.³⁶ Consequently, applying the chemodenitrification rate constant from Kopf et al.36 to the experiment performed by Klueglein et al., 17 utilizing spent BBM, provides valuable insight into the potential significance of dissolved organic ligands on overall Fe(II) oxidation rates in NDFO cultures (Figure 1). EPS are secreted by many bacteria, including NDFO species, and have previously been found to strongly complex with Fe(II) and enhance its oxidation rate (Figure 1).⁶²

Chemodenitrification produces nitrous oxide (Reaction 2); however, measured data were only available for Acidovorax spp. strain TPSY cultures.24 The nitrous oxide concentrations measured in Carlson et al., 24 while not included as observations during model calibration. were used as additional verification of the biogeochemical model. Simulation results of nitrous oxide concentrations in model variant S1a for strain TPSY broadly matched the observations, whereas S2 model results were significantly above the measured data (Figure S4). Further experimental work is required to investigate the yields of nitrous oxide in NDFO cultures for different bacteria to better constrain nitrous oxide production rates from chemodenitrification. Characterizing stable isotope systematics in NDFO cultures to determine the fraction of nitrous oxide produced via biological or chemical pathways^{25,32} would further elucidate the impact chemodenitrification has on nitrous oxide production in Fe rich environments.

Quantification of Iron(II) Oxidation Processes.

Model variant S2 underpredicted the rate of total Fe(II) oxidation and overpredicted nitrite concentrations by as much as 300% (Figure 2), indicating biological oxidation was contributing to Fe(II) oxidation (all model variants of S1). The maximum rate and onset of Fe(II) oxidation was better matched using variants of S1 compared to S2, lending further support to the presence of a biological Fe(II) oxidation pathway (observed max oxidation rates of 3.60, 1.22, 3.20, 4.50, and 3.60 mM d⁻¹ versus S1a simulated max oxidation rates of 4.74, 1.94, 2.37, 4.48, and 2.96 mM d−1 for BoFeN1, 2AN, TPSY, Pd1222, and 2002, respectively). The contribution of chemodenitrification was consistent across all bacteria (Figure 2). In growth media where acetate was supplied, chemodenitrification was responsible for 35%, 39%, 25%, 37%, and 40% of the overall dissolved Fe(II) oxidized for Acidovorax spp. strains BoFeN1, TPSY, 2AN, Pseudogulbenkiania strain 2002, and Paracoccus denitrificans, respectively, with enzymatic NDFO responsible for the remaining proportion. The extent of chemodenitrification was similar across all bacteria strains,

potentially indicative of biological process(es) common to most denitrifers, as has been postulated previously.²⁴ The extent of abiotic and biotic Fe(II) oxidation processes for the lithoautotrophic Pseudogulbenkiania strain 2002 yielded similar results to those found by Kopf et al.³⁶ In their experiments, abiotic oxidation accounted for 30–35% of the total dissolved Fe(II) oxidized using the *Pseudogulbenkiania* sp. strain MAI-1, which is closely related to Pseudogulbenkiania strain 2002. Acidovorax spp. Strain BoFeN1 cultures oxidized \sim 35% of Fe(II) via chemo-denitrification, consistent with previous studies that used Fe(II)-EDTA to discern biotic and abiotic dissolved Fe(II) oxidation contributions and found chemodenitrification played a major role.⁶⁴

In all of the simulations for model variant S2, the thermodynamic drive of enzymatic Fe(II) oxidation never reached zero as long as the dissolved $Fe(II)$ and nitrate concentrations remained above zero (Figure 2). Similarly, the concentrations of nitrate, the electron acceptor, had little impact on the overall rate, with the kinetic factor F_A being consistently at or close to 1 for much of the experiment, while falling rapidly when concentrations approached zero. Consequently, enzymatic Fe(II) oxidation was controlled predominantly by mineral encrustation, particularly for strains TPSY and 2AN, as well as the kinetic factor F_D . Given NDFO was never thermodynamically constrained, in the natural environment where substrate concentrations are significantly lower, enzymatic NDFO could still be an active metabolic pathway for nitrate reducing species.

Green Rust Carbonate (GR-CO3) as a Reactive Intermediate.

 $GR-CO₃$ was assumed to precipitate as the initial Fe mineral product and was an equilibrium phase in the model simulations for all bacteria.^{30,51} Structural Fe(II) within green rust (Fe(II)_{GR}) was oxidized by nitrite and transformed GR-CO₃ into goethite^{30,51,65} at a rate calibrated using experimental data collected by Hansen et al.⁵³ The calibrated rate constant for Fe(II)_{GR} oxidation by nitrite (4.12 × 10⁻⁵ s⁻¹) was faster than the rate determined by Weber et al.³⁵ for Fe(II) in magnetite and biologically reduced goethite (3.38 \times 10⁻⁵ and 3.18×10^{-6} s⁻¹, respectively) and almost 1 order of magnitude faster than Fe(II)_{GR} oxidation by nitrate $(0.4–6.6 \times 10^{-6} \text{ s}^{-1})$.⁵⁴ GR was responsible for reducing between ~15% and 39% of all the nitrite formed in simulation S1a, demonstrating that it can strongly control nitrite accumulation in the growth culture.

In all model simulations, the reduction of nitrite to ammonium by $GR-CO₃$ provided the best fit to the observed nitrite data. However, a number of different nitrogen species can be produced from the oxidation of $Fe(II)_{GR}$ by nitrite, including nitrous oxide, dinitrogen gas, or ammonium and is a function of the solution pH, redox potential, and concentration of phosphorus.30,54,66 While previous studies investigating the reduction of nitrite by green rust have found significant quantities of nitrous oxide, $26,28$ inclusion of this reaction within the S2 simulation causes excessive nitrous oxide production, significantly above the limited number of observed nitrous oxide concentrations. Etique et al.³⁰ found total Fe(II) oxidation by biogenic nitrite produced stoichiometric amounts of ammonium in a heterotrophic nitrate reducing growth culture, geochemically similar to the cultures investigated in this study. Moreover, ammonium production during the partial reduction of nitrate by $GR\text{-}CO₃$ has previously been found to be greater than 70% at circumneutral pH, where Fe^{2+} _{aq} was

supplied in excess.⁵⁴ Understanding the products of secondary Fe mineral reactions with biogenic nitrogen species would be improved if a full nitrogen mass balance could be calculated, but few NDFO studies have been published that include such comprehensive data.

lmplications of the Model-Derived Findings.

Although several studies had already investigated the role of biotic and abiotic Fe(II) oxidation, $16,24,36,64$ experimentally determining their relative contributions is challenging as both are intrinsic-cally coupled during NDFO.^{16,33} By translating different conceptual models into process-based numerical models, this study was able to isolate and quantify the rates of abiotic and biotic dissolved Fe(II) oxidation. Furthermore, through the model development and application, several processes were identified that require further investigation through additional targeted experiments to reduce model conceptual and parameter uncertainty. Figure 3 summarizes the conceptual model of NDFO processes that were considered in the biogeochemical model in this study. It also summarizes alternative pathways that are currently omitted from the model which could, however, potentially warrant inclusion if future experiments would demonstrate their existence and relevance.

The most pressing issue identified through our biogeochemical modeling study that demands attention is the characterization of EPS secreted by NDFO bacteria in pure cultures.17 This is a crucial knowledge gap given that EPS is commonly a combination of macromolecules that includes polysaccharides⁶² which are known to complex with $Fe(II)$ and enhance chemodenitrification.36 Undertaking experiments similar to those performed by Norman et a_{n} ,⁶² where the role of EPS on the oxidation rate and solubility of Fe were assessed, will be beneficial to understand the extent to which components within EPS complex with Fe(II) and affect its biogeochemical cycling within NDFO cultures. These experiments could also examine whether changing the initial Fe concentrations or varying the electron-donor balance affects the composition of EPS and hence the chemodenitrification rate. Furthermore, quantifying the species specific oxidation rate of Fe carbonate species by nitrite or nitric oxide is yet to be investigated in the same vein as its oxidation rate by molecular oxygen.⁶⁷ The development of a more comprehensive chemodenitrification model would improve the utility of any future biogeochemical modeling of NDFO and would allow the identification of key reactive species to assess their prevalence in natural sediments likely to inhabit NDFO bacteria. These contributions would improve the understanding of NDFO and its significance on the cycling of Fe in natural environments.

Determining the rate of enzymatic NDFO could be improved by performing NDFO growth cultures using prestarved cells. This will mitigate the potential interference of internally stored carbon within NDFO cells driving significant nitrate reduction to nitrite and causing the interfering chemodenitrification reaction even in the absence of an exogenous carbon source. The same set of prestarved batch experiments could also clarify whether nitrite is formed as a result of biological Fe(II) oxidation. Reaction kinetics may potentially be slower in prestarved culture experiments which would be beneficial for identifying the presence of reactive intermediate Fe mineral phases and their influence on Fe(II) oxidation rates.

Finally, few studies report gaseous nitrogen products nitric oxide, 24 nitrous oxide, 15,24 or dinitrogen gas^{9,29} within NDFO cultures prohibiting the calculation of the nitrogen mass balance, which would also assist in determining the potential major biotic and abiotic Fe(II) oxidation processes. Combined with stable isotope experiments, the fate of the nitrogen could be well traced to refine the conceptual model.⁶⁸ Extending this model framework to a controlled well-characterized field trial, such as the one reported by Smith et al.,⁶⁸ would be valuable to further understand the interplay between chemical and biological NDFO processes and the production on various nitrogen products.

Supplementary Material

Refer to Web version on PubMed Central for supplementary material.

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Figure 1.

Chemodenitrification rates for experiments by Klueglein et al.17 (left); Klueglein and Kappler¹⁶ (middle); Jones et al.²⁵ (right) (used in simulations S1a, S1b, and S1c, respectively). Experiment 4 mM $NO₂$ —8 mM Fe(II) is replotted on the left panel with 4 mM $NO₂$ —8 mM Fe(II)-EPS for reference. Chemodenitrification kinetics were described using eq 14. Observed Fe(II) concentrations are presented as a percentage of initial concentration (C/C_0) .

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Figure 2.

Simulation and experimental results for three Acidovorax strains BoFeNl, 2AN, and TPSY as well as Paracoccus denitrificans strain Pd 1222 and Pseudogulbenkiania strain 2002. Symbols represent observed concentrations for total Fe(II) (red circle), nitrate (blue down triangle), acetate (yellow up triangle), nitrite (yellow square), and biomass (filled green diamonds) from Klueglein and Kappler,¹⁶ Chakraborty et al.,¹⁴ and Carlson et al.²⁴ Simulations Sla (solid black line), Slb (dashed black line), and Slc (dash dot black line) were compared against S2 (solid red/gold line). All model results are similar for parameters presented in rows 2 and 4 to 6 and are therefore excluded to improve figure clarity. The fifth row presents the kinetic factors F_A and F_D (red and blue, respectively) and the thermodynamic potential factor F_T (green). The sixth row presents the contribution of chemodenitrification (red hatches) and enzymatic NDFO (blue hatches) to overall Fe(II) oxidation. Biomass subject to encrustation is provided (solid black line) as well as uninhibited biomass for reference (solid green line).

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Figure 3.

Conceptual models of NDFO with varying levels of complexity. Model (3a) represents the processes included in the biogeochemical model developed for this study. Dashed gray arrows represent processes that are either unknown or potential processes active in NDFO cultures that require ongoing investigation. (1) A dedicated Fe(II) oxidoreductase exclusively responsible for NDFO. (2) Respiratory complexes catalyze NDFO (enzymatic NDFO) with additional Fe(II) oxidation occurring extracellularly via chemodenitrification. (3a) Identical to (2) but with the inclusion of green rust oxidation by nitrite, producing

goethite and ammonium. (3b) Identical to (3a) but including species-specific rates for Fe(II) carbonate/organic complexes; enhanced abiotic Fe(II) oxidation within the periplasm due to low pH.

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Table 1.

Overview of the Conceptual Models Employed in the Biogeochemical Model

Table 2.

Calibrated Model Parameters for All Model Simulations of Model Variant S1a Calibrated Model Parameters for All Model Simulations of Model Variant S1a

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 c calibrated separately using data collected by Hansen et al. (Figure S2). Calibrated separately using data collected by Hansen et al. (Figure S2).