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# Bioenergetic constraints on the origin of autotrophic metabolism

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Autotrophs form the base of all complex food webs and seemingly have done so since early in Earth history. Phylogenetic evidence suggests that early autotrophs were anaerobic, used CO<sub>2</sub> as both an oxidant and carbon source, were dependent on H<sub>2</sub> as an electron donor, and used iron-sulfur proteins (termed ferredoxins) as a primary electron carrier. However, the reduction potential of H<sub>2</sub> is not typically low enough to efficiently reduce ferredoxin. Instead, in modern strictly anaerobic and H<sub>2</sub>-dependent autotrophs, ferredoxin reduction is accomplished using one of several recently evolved enzymatic mechanisms, including electron bifurcating and coupled ion translocating mechanisms. These observations raise the intriguing question of why anaerobic autotrophs adopted ferredoxins as central electron carriers only to have to evolve complex machinery to reduce them. Here, we report calculated reduction potentials for H<sub>2</sub> as a function of observed environmental H<sub>2</sub> concentration, pH and temperature. Results suggest that a combination of alkaline pH and high H<sub>2</sub> concentration yield H<sub>2</sub> reduction potentials low enough to efficiently reduce ferredoxins. Hyperalkaline, H<sub>2</sub> rich environments have existed in discrete locations throughout Earth history where ultramafic minerals are undergoing hydration through the process of serpentinization. These results suggest that serpentinizing systems, which would have been common on early Earth, naturally produced conditions conducive to H<sub>2</sub>-dependent autotrophic the emergence of life. The primitive process of hydrogenotrophic

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methanogenesis is used to examine potential changes in methanogenesis and Fd reduction pathways as these organisms diversified away from serpentinizing environments.

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#### 1. Introduction

Ferredoxins (Fds) are ubiquitous in biology and link oxidation and reduction reactions in a diversity of metabolisms, including those of anaerobes, aerobes, phototrophs, chemotrophs, autotrophs and heterotrophs [1]. These small soluble proteins have active sites composed of iron–sulfur (Fe–S) clusters (i.e. 2Fe-2S or 4Fe-4S clusters) [2] that can accept or donate electrons with a concomitant change in the oxidation state of the iron atom(s) between + 2 and + 3. These proteins have been widely argued to be among the most primitive of electron carriers [3]. The electrochemical potential of Fe–S clusters, including those of Fds, can range from -0.60 to +0.50 V [4], thereby encompassing the array in potentials associated with oxidation–reduction reactions that fuel life in reduced environments reminiscent of those on early Earth as well as oxidized environments that characterize surface locales on modern Earth [5].

Genes encoding Fds are enriched in the genomes of anaerobes relative to aerobes and facultative anaerobes [6], a finding that is consistent with the central role of Fds in anaerobic energy metabolism [7]. This is particularly true for the genomes of strictly anaerobic, hydrogenotrophic methanogens [6], a group of organisms often advocated as having the most primitive of extant metabolisms [8,9]. Rationales for the primacy of hydrogenotrophic methanogens are manifold. Firstly, isotopic analysis of methane in fluid inclusions in ancient rocks dated to approximately 3.46 Ga suggest the presence of ancestors of methanogens at this time [10] and phylogenetic data suggest that these ancestral cells were likely dependent on  $H_2$ as electron donor [9,11]. Secondly, hydrogenotrophic methanogens can use  $CO_2$  as their sole source of carbon as well as their electron acceptor, resulting in the production of methane (CH<sub>4</sub>). This potentially alleviated a major barrier for early autotrophic life since CO<sub>2</sub> was likely widely available on a volcanically active early Earth while organic carbon and other potential oxidants (e.g. oxygen, nitrate, sulfate) capable of supporting life are likely to have been far less available [12]. Finally, methanogens rely on the reductive acetyl CoA pathway, or the Wood Ljungdahl (WL) pathway [8], for carbon fixation. The WL pathway forms acetyl CoA from  $CO_2$  in an overall thermodynamically favourable reaction [13] and has been suggested to be the most ancestral of extant carbon fixation pathways [8,9]. Thus, methanogens and other early evolving organisms that use the WL pathway (e.g. acetogens) do not expend the limited energy supplied by their primitive anaerobic metabolisms to generate reduced carbon like other autotrophs operating less efficient pathways.

The typical Fds involved in the energy metabolism of hydrogenotrophic methanogens and acetogens are highly reduced, having standard state (25°C, pH 7.0, 1 M reactant and product concentrations) reduction potentials ( $E^{\circ}$ ) relative to the standard hydrogen electrode of approximately -0.41 to -0.50 V [14,15]. This is possibly a consequence of selection to harbour an Fe–S cluster with a low enough potential to allow for the initial reduction of CO<sub>2</sub>, which is estimated to be approximately -0.50 V [16]. However, the  $E^{\circ'}$  of H<sub>2</sub>, the source of electrons supporting the energy metabolism of these cells and organic carbon synthesis via the WL pathway, is approximately -0.41 V [17,18], which is not of low enough potential to efficiently reduce the Fe–S cluster(s) in low potential Fd (Reaction 1).

$$0.5H_2 + Fd_{ox} \leftrightarrow H^+ + Fd_{red}^{1-}$$
 (Reaction 1)

This is particularly true considering that Fd in cells is generally 90% reduced [19], thus favouring preservation of reactants.

Nonetheless, numerous anaerobic, hydrogenotrophic methanogens and acetogens have been characterized in laboratory and natural settings (e.g. [17,20]). How do contemporary anaerobic cells reduce low potential Fd with H<sub>2</sub> as an electron carrier? At least two distinct enzymatic

mechanisms have evolved to overcome this thermodynamic barrier in H<sub>2</sub> oxidizing anaerobic autotrophs: electron bifurcating Fd reduction and ion (proton or sodium) translocating Fd reduction [7,14]. Electron bifurcation involves the endergonic reduction of low potential Fd by an electron donor with a higher potential. This is only possible when the same donor is also coupled to concomitant exergonic oxidation by an even higher potential acceptor [17,18]. Twelve multimeric bifurcating enzyme complexes have been described to date and all are involved in the reduction of Fd; however, only three of these multimeric complexes are known to use H<sub>2</sub> as the hydride donor [7,15,21]. These include the trimeric/tetrameric [FeFe]-hydrogenases, the hexameric [FeFe]-hydrogenase-CO2 reductase and the hexameric [NiFe]-hydrogenaseheterodisulfide reductase (Mvh-Hdr). In ion translocating Fd reduction, the unfavourable thermodynamics associated with Fd reduction with  $H_2$  are overcome by coupling the chemical reaction to depletion of an electrochemical ion (i.e. sodium ion or proton) gradient across a membrane [7]. Two classes of membrane-bound enzyme complexes are known to catalyse this function: the energy converting [NiFe]-hydrogenases (Ech or Eha/Ehb) and the Fd:NAD<sup>+</sup> oxidoreductase (Rnf). Both complexes couple the reversible reduction of Fd and oxidation of H<sub>2</sub> (Ech/ Eha/Ehb) or NADH (Rnf) with the translocation of sodium ions or protons into the cell [22–24].

Based on the fundamental role of electron bifurcating complexes and ion translocating Fd oxidoreductase enzymes in the energy metabolism of hydrogenotrophic methanogenic or acetogenic anaerobes [7], these enzymes or processes are often advocated as having been integrated into the energy metabolism of primitive lifeforms on early Earth [25-27]. In support of this hypothesis, the genomes of hydrogenotrophic acetogens and methanogens that could be classified based on previous physiological characterization encode one or more H2-dependent electron bifurcating or ion translocating Fd oxidoreductase enzyme homologs (electronic supplementary material, table S1). However, a recent analysis of the taxonomic distribution and evolutionary history of bifurcating enzymes revealed that none of the twelve known bifurcating complexes are likely to have been a property of the Last Universal Common Ancestor (LUCA) of Archaea and Bacteria [28]. Rather, evidence suggests that the twelve known bifurcating complexes emerged independently multiple times after the divergence of Archaea and Bacteria from the LUCA. Likewise, phylogenetic analyses of Ech/Eha/Ehb suggest that these enzymes are likely to have emerged in methanogenic Archaea after their divergence from LUCA, with acquisition among Bacteria taking place more recently [11]. Finally, Rnf enzymes are found in several lineages of anaerobic Bacteria but are only found in several late diverging methanogenic Archaea (e.g. Methanosarcinales; [23]) and are absent in more ancestral methanogens (e.g. Methanococcales, Methanobacteriales; electronic supplementary material, table S1), suggesting that they too may be recently evolved. Collectively, these observations raise key new questions regarding how Fd was reduced in ancestral, autotrophic and anaerobic cells such as methanogens prior to the advent of complex mechanisms that allow for its reduction in contemporary cells.

Here, we provide thermodynamic calculations of the potential of Fd and  $H_2$  across a range of environmental conditions to identify those that may be suitable for direct reduction of Fd with  $H_2$  in early evolving anaerobic autotrophs in the absence of bifurcating or ion-translocating hydrogenases. Results are discussed at the level of the possible ancestral pathways that enabled autotrophy in early evolving hydrogenotrophic anaerobes and the role of ion-translocating and bifurcating hydrogenases in the diversification of ancestors of anaerobic,  $H_2$ -dependent autotrophs away from environmental realms that may have favoured the direct reduction of Fd with  $H_2$ .

#### 2. Material and methods

# (a) Relationship between the mode of carbon metabolism and distribution of bifurcating enzyme homologs

A database of organisms with complete genomes (n = 4,588) created in our previous analysis of the taxonomic distribution on bifurcating enzyme homologs [28] was subjected to classification

of their ability to integrate oxygen into their energy metabolism based on information acquired from the Department of Energy's Integrated Microbial Genomes and Microbiomes (DOE-IMG). All identified anaerobic taxa were further characterized as hydrogen oxidizing autotrophic acetogens or methanogens based on the available literature (electronic supplementary material, table S1). This curated list of organisms was then cross-referenced with our published database describing the distribution of bifurcating enzyme homologs in their genomes [28]. In addition, the genomes of these organisms were screened for homologs of ferredoxin:NAD<sup>+</sup> oxidoreductase (Rnf) and the group 4 membrane-bound [NiFe]-hydrogenase (Ech/Eha/Ehb). Briefly, Rnf protein homologs from *Azotobacter vinelandii* (ACO81179-ACO81185) and a homolog of the large subunit of group 4 [NiFe]-hydrogenase from *Pyrococcus abyssi* (WP\_010867842) were subjected to analysis via phmmer available in the HMMER (v.3) software package (hmmer.org) to extract protein homologs. Extracted homologs of [NiFe]-hydrogenases were aligned using Clustal Omega [29] and filtered to contain the conserved residues that demarcate these enzymes as identified in a previous study [30].

#### (b) Redox potential calculations

The ability of a compound to donate or accept electrons is dictated by its redox potential (E), which can be calculated using equation (2.1):

$$E = E^{\circ} + 2.303 \left(\frac{RT}{nF}\right) \log \frac{\alpha_{\rm Ox}}{\alpha_{\rm Red}},\tag{2.1}$$

where  $E^{\circ}$  represents the standard redox potential of reaction in volts (V), *R* designates the ideal gas constant (approx. 8.314 J mol<sup>-1</sup> K<sup>-1</sup>), *T* is the temperature in Kelvin (K), *n* is the number of electrons transferred in the reaction, *F* is the Faraday constant (approx. 96 485 J V<sup>-1</sup> mol<sup>-1</sup>),  $\alpha_{OX}$  is the activity of the oxidized form of the compound and  $\alpha_{Red}$  is the activity of the reduced form. Redox potential calculations for H<sub>2</sub> were performed across a range of H<sub>2</sub> activities (0.0001 to 10 mM), temperatures (0 to 100°C) and pH values (4.0; 7.0; 10) using equation (2.1). In addition, the redox potential of H<sub>2</sub> in natural environments was calculated using reported (electronic supplementary material, table S2) dissolved H<sub>2</sub> concentrations, temperatures and pH values. Activity coefficients were assumed equal to one for all dissolved compounds (activity = concentration) since the ionic strength was not always reported.

#### 3. Results and discussion

The non-standard state electrochemical potential (*E*) of a half-reaction is dependent on how the  $E^{\circ}$  is modified by temperature, pressure, and the concentrations of reactants and products. In the case of the half-reaction describing the oxidation of H<sub>2</sub> (Reaction 2), the potential is dependent on pH since protons (H<sup>+</sup>) are a product of the specified reaction. Thus, the potential of H<sub>2</sub> (thermodynamic drive toward products) will be lower in environments with lower H<sup>+</sup> activity.

$$H_2 \rightarrow 2e^- + 2H^+$$
. (Reaction 2)

Using thermodynamic data, we calculated the *E* of the half-reaction describing the oxidation of H<sub>2</sub> (Reaction 2, hereafter referred to as the potential of H<sub>2</sub>) as a function of temperature, proton concentration (i.e. pH), and H<sub>2</sub> concentration to identify conditions that may be conducive to the direct reduction of low potential Fds associated with anaerobic autotrophs ( $E^\circ = -0.41$  to -0.50 V at pH 7.0, 25°C), both today and in the geological past. To facilitate comparison of Fd reduction potentials relative to Reaction 1, we also calculated the *E* of Fds across a range of temperatures and pH (figure 1). The temperature dependence of the *E* of Fds ( $E^\circ = -0.41$  to -0.50 V) was calculated by extrapolating measurements of Fd purified from *Pyrococcus furiosus* as a function of temperature (-0.00123 V/°C) [31]. The pH dependence of the *E* of Fds was calculated by extrapolating the maximum measured effect of pH on the measured *E* of Fd (0.030 V/pH unit) [32] purified from a variety of taxa (range = 0.002 to 0.030 V/pH unit) [32–34]. These calculations



**Figure 1.** Reduction potential of  $H_2$  (Reaction 1) as a function of temperature and  $H_2$  concentration in water with a pH of 4.0 (*a*), 7.0 (*b*) and 10.0 (*c*). The black line in each panel depicts the calculated midpoint potential of Fds with an  $E^{\circ\prime}$  of -0.41 V while the dashed black line in each panel depicts the calculated midpoint potential of Fds with an  $E^{\circ\prime}$  of -0.50 V. Calculations of the midpoint potential of Fds assumed a concentration of 1 M. The temperature and pH dependence of the midpoint potentials calculated for Fds were assumed to be 0.00123 V/°C [31] and 0.030 V/ pH unit [32], respectively. (Online version in colour.)

assumed a single electron transfer per Fd (Reaction 1) since ancestral Fds are thought to have comprised a single Fe–S cluster [3] and assumed a constant concentration of Fd (1 M). Importantly, ionic strength has also been shown to influence the *E* of purified Fd [34]. However, the inverse relationship between ionic strengths ranging from 0 (slightly lower than freshwater) to 1.0 (greater than seawater) and the measured *E* of Fd is minimal (0.061 V) and is therefore not considered further in the calculations presented here.

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The calculated E of H<sub>2</sub> dissolved in water with a pH of 4.0, when considered over a range of dissolved H<sub>2</sub> concentrations from 100 nM to 10 mM and temperatures from 0 to 100°C, ranged from -0.03 V to -0.22 V (figure 1*a*). More negative potentials were associated with higher concentrations of dissolved  $H_2$  and higher temperature; the effects of these parameters were additive. In water with a pH of 7.0 and over the same range of dissolved  $H_2$  concentrations and temperatures, the E of H<sub>2</sub> varied from -0.19 V to -0.44 V with more negative potentials in solutions with higher concentrations of dissolved  $H_2$  and higher temperature (figure 1b). Finally, dissolved H<sub>2</sub> in water with a pH of 10.0 and over the same range of dissolved H<sub>2</sub> concentrations and temperatures had calculated E that ranged from -0.35 V to -0.67 V (figure 1c). In water with a pH of 10.0, the calculated E of  $H_2$  at a concentration of greater than 1 mM would be low enough to efficiently reduce low potential Fd ( $E^{\circ} = -0.41 \text{ V}$ ) at temperatures greater than 25°C. The reduction of lower potential Fd ( $E^{\circ} = -0.50$  V) at pH 10.0 requires higher H<sub>2</sub> concentrations (greater than 10 mM) and temperatures (greater than 100°C). As previously mentioned, the relationship between the *E* of Fd and pH was calculated using the maximum measured inverse slope of 0.030 V/pH unit. Thus, it is possible that efficient reduction of low potential Fds could take place in a setting with a lower pH, temperature, and H<sub>2</sub> concentration than those specified above. Alternatively, it is possible that the potential of Fds vary with pH in a more Nernstian-like manner, whereby a decrease in the potential of 0.059 V/pH unit would be expected. If so, a higher concentration of H<sub>2</sub>, a higher pH, and/or temperature would be required to reduce low potential Fds than those specified above.

To identify environmental types with conditions that may support efficient reduction of low potential Fds, the dissolved H<sub>2</sub> concentrations (range = 2.3 nM–2.9 mM), temperatures (range = 0.4– $90.0^{\circ}$ C), and pH values (range = 2.9–12.4) of waters from various early Earth analogue environments that had temperatures less than 122°C, the current upper temperature limit for life [35], were compiled (electronic supplementary material, table 2). These environments were chosen because of (i) available temperature, pH and H<sub>2</sub> concentration data, (ii) their potential to host non-photosynthetic, chemosynthetic cells thought to reflect the earliest forms of life [9,11] and (iii) previous suggestions of their relevance as habitats supporting descendants of early forms of life [36–40]. The following environment types met these criteria: marine and terrestrial hydrothermal springs, subglacial environments overlying basalt bedrock, and marine and terrestrial subsurface environments where water is in contact with ultramafic rocks (i.e. serpentinites).

The calculated E of  $H_2$ , when plotted as function of temperature (figure 2a), pH (figure 2b) or H<sub>2</sub> concentration (figure 2c), from a variety of environments identified natural conditions that could conceivably favour the efficient direct reduction of low potential Fd with  $H_2$ . These include subsurface fluids in contact with ultramafic rocks in the Coast Range Ophiolite Microbial Observatory (CROMO) in the USA (green triangles), the Samail Ophiolite in the Sultanate of Oman (orange dots), and the Lost City hydrothermal field in the Atlantic Ocean (blue diamonds). Common to each of these three environmental types is a combination of alkaline pH (figure 2b) and high  $H_2$  concentration (figure 2c) that span a range of temperatures (figure 2a). The alkaline pH and elevated  $H_2$  concentrations in each of these environment types is due to hydration of ultramafic rock rich in olivine and pyroxene, which stimulates a series of reactions collectively known as serpentinization that generate hydroxide ions and molecular  $H_2$  [41]. The process of serpentinization is thought to have been more prevalent on early Earth where highly reduced mineralogy was likely more widespread in an undifferentiated crust [36,42]. In turn, and as shown here, hyperalkaline, H<sub>2</sub>-rich conditions are potentially conducive to the direct and efficient reduction of Fd in anaerobic autotrophs if suitable catalysts are available, leading to the suggestion that serpentinizing systems hosted the earliest evolving, Fd- and  $H_2$ -dependent anaerobic autotrophs.

The primary catalysts or enzymes involved in the reversible oxidation of  $H_2$  in modern cells are [FeFe]- and [NiFe]-hydrogenases. While [FeFe]-hydrogenases are unlikely to be a property of the LUCA [43], [NiFe]-hydrogenase have been suggested to have been associated with the LUCA where they likely functioned in  $H_2$  oxidation [9,11]. It is worth noting that both Fe and



■ hot springs ◆ Lost City ● Oman ▲ CROMO - subglacial

**Figure 2.** Reduction potential of  $H_2$  (Reaction 1) in early Earth analogue natural environmental samples as a function of temperature (*a*), pH (*b*) and H<sub>2</sub> concentration (*c*). H<sub>2</sub> concentrations, pH, temperatures and calculated *E* for H<sub>2</sub> are presented in electronic supplementary material, table S2. (Online version in colour.)

Ni, the metals that form the active site of [NiFe]-hydrogenase, are common components of serpentinite rocks [36]. H<sub>2</sub> oxidation via [NiFe]-hydrogenases in contemporary hydrogenotrophic methanogens occurs in the cytoplasm (figure 3) [17], and this may have also been true for ancestors of these cells. If primordial hydrogenases were also localized in the cytoplasm, the calculations presented above must be modified to accommodate shifts in the reduction potential of H<sub>2</sub> when it is oxidized within the cell. This is because contemporary cells tend to maintain or buffer the pH of the cytoplasm within a narrower range than the pH outside the cell [44] but cannot regulate temperature or H<sub>2</sub> concentration. The same is likely to be true for primitive autotrophic cells.

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**Figure 3.** A simplified depiction (*a*) of the energy metabolism of contemporary hydrogenotrophic methanogens as modified from Thauer *et al.* [17] and the linear pathway (*b*) as proposed herein for the ancestor of hydrogenotrophic methanogens. The bifurcating Mvh-Hdr and the ion translocating Eha/Ehb[NiFe]-hydrogenases are depicted, as is the proposed non-bifurcating ancestor of Mvh (Mvh<sub>Anc</sub>). Enzymes and their enzyme category numbers (EC) shown as numbers are as follows: 1, formylmethanofuran (MFR) dehydrogenase (EC: 1.2.7.12); 2, formylmethanofuran:tetrahydromethanopterin (H<sub>4</sub>MPT) formyltransferase (EC: 2.3.1.101); 3, Methenyltetrahydromethanopterin cyclohydrolase (EC: 3.5.4.27); 4, F<sub>420</sub>H<sub>2</sub>-dependent methylenetetrahydromethanopterin dehydrogenase (EC: 1.5.98.1); 5, coenzyme F<sub>420</sub>-dependent N5,N10 methylenetetrahydromethanopterin reductase (EC: 1.5.98.2); 6, methyl-H<sub>4</sub>MPT–coenzyme M methyltransferase (EC 2.1.1.86); 7, methyl coenzyme M reductase (EC 2.8.4.1); 8, coenzyme F<sub>420</sub> reducing [NiFe]-hydrogenase (EC 1.12.98.1). Enzyme abbreviations are as follows: Eha/Ehb<sub>Anc</sub>, the putative ancestor of energy converting [NiFe]-hydrogenase (Eha/Ehb); Mvh<sub>Anc</sub>, the ancestor of soluble heterodisulfide reductase (Hdr). Fds are assumed to accept a single electron in the schematics. (Online version in colour.)

Hyperalkaliphiles maintain a cytoplasmic pH that is up to 2 units lower than that of their external environment [44], which would increase the E of  $H_2$  as it diffuses into the cell. As an example, for a cell inhabiting a pH 12.0 environment (assumed cytoplasmic pH of 10.0) at 25°C to efficiently reduce low potential Fds with  $E^{\circ}$  of -0.41 V (E = -0.50 V) and -0.50 V (E = -0.59 V), H<sub>2</sub> would need to be at a concentration of 1 mM and 1 M, respectively. Again, these are likely maximum  $H_2$  concentrations required based on the assumed relationship between E and pH (0.030 V/pH unit). Environments enriched in ultramafic minerals and undergoing active lowtemperature serpentinization generate water with strongly alkaline pH, concentrations of  $H_2$  that approach 10 mM, and are often at temperatures that exceed 25°C (i.e. Oman, Lost City; figure 2). However, based on available data, it is unclear whether a cytoplasmic pH of 10.0 is physiologically feasible in an autotroph since available CO<sub>2</sub> would speciate to carbonate (CO<sub>3</sub><sup>2–</sup>) which has a low solubility and is not the carbon source for methanogenesis [17]. That said, methanogens have been grown with calcium or magnesium carbonate as their sole carbon source [45] in growth medium with a pH as high as 10.0 [46]. These organisms grow presumably by consuming  $CO_2$ that had equilibrated in solution [45]. Moreover, putatively hydrogenotrophic methanogens have been detected in several environments undergoing active serpentinization [47], including well waters sampled from the Samail Ophiolite in Oman with pH of up to 11.3 [48,49]. This indicates that the low solubility of inorganic carbon in hyperalkaline environments and constraints that it imposes on the availability of CO<sub>2</sub> can be overcome, at least in modern methanogen cells.

Primitive autotrophic,  $H_2$ -dependent cells also faced the need to innovate mechanisms to shuttle electrons that minimized energy loss and that had potentials close to the  $H_2/H^+$  redox

couple. Low potential Fds would efficiently meet this need in environments where the E of  $H_2$ was more reduced, such as in alkaline hydrothermal vents or other environments with ultramafic rocks that are undergoing active serpentinization. As a modern example, the calculated E of H<sub>2</sub> in fluids from hydrothermal vents at Lost City (pH 12.1,  $[H_2] = 7.8 \text{ mM}$ , 94°C; electronic supplementary material, table S2) is -0.73 V (figure 2) or -0.66 V at a cytoplasmic pH of 10.0. These values are close to the calculated E of a typical anaerobic Fd, which is estimated to range from -0.58 to -0.67 V at 95°C at a cytoplasmic pH of 10.0. Importantly, it is also likely that ancestral Fds comprised a far simpler protein environment than contemporary Fds [3], perhaps consisting of as few as 7 amino acids [50,51]. Fe-S clusters in such primitive Fds would presumably be more solvent exposed, which in turn would be expected to poise their E more positive than contemporary anaerobic Fds [4]. This, in turn, may have allowed early autotrophic anaerobes to exploit H<sub>2</sub> as an electron donor in a broader range of environments where reduction of these primitive Fds was favourable, including those with lower H<sub>2</sub> concentrations or with slightly less alkaline pH where inorganic carbon would have been more bioavailable. However, as depicted in figures 1 and 2, environments with temperature, pH, and  $H_2$  conditions that may have been favourable for the reduction of low potential Fds represent only a fraction of those that were likely available on early Earth. We suggest that the emergence of complex enzymes enabling electron bifurcating Fd reduction and ion (proton or sodium) translocating Fd reduction [7,14] was triggered by the diversification of anaerobic life away from environments that favoured efficient and direct Fd reduction and into environments where the thermodynamics and kinetics of Fd reduction with H<sub>2</sub> were less favourable. These would have included environments with lower  $H_2$  concentrations, lower temperatures, less alkaline pH or a combination of these factors.

The emergence of electron bifurcating and ion-translocating [NiFe]-hydrogenase enzymes was accomplished by recruitment of flavoprotein/heterodisulfide reductase [28] and ion transporter modules [11], respectively, thereby providing a mechanism to reduce Fd and meet the bioenergetic needs of early anaerobic autotrophs in environments less favourable for direct and efficient enzymatic reduction of Fd with H<sub>2</sub>. A likely consequence of these diversifications is the presence of electron bifurcating and/or ion-translocating [NiFe]-hydrogenase enzyme homologs involved in reversible H<sub>2</sub> oxidation in the genomes of all anaerobic H<sub>2</sub> oxidizing methanogens and acetogens (electronic supplementary material, table S1), suggesting that such mechanisms are obligate in these derived cells. It remains to be determined if anaerobic H<sub>2</sub>dependent methanogens or acetogens that lack genes coding for bifurcating and ion-translocating enzymes can be found in contemporary environments undergoing active serpentinization (i.e. Lost City or the Samail Ophiolite, Oman), or if approximately 3.8 Ga of evolution has erased all vestiges of these ancestral organisms. Further cultivation efforts focused on such organisms from environments with high pH and high H<sub>2</sub> concentrations are necessary to address this question directly. Alternatively, it may be possible to evaluate this question by reconstructing the metabolisms of difficult to cultivate methanogens or acetogens through assembly and binning of metagenomic sequence data followed by robust informatics and evolutionary analyses.

What might such an ancestral anaerobic  $H_2$  oxidizing methanogen or acetogen cell comprise? If the hypothesis that bifurcating and ion-translocating [NiFe]-hydrogenase enzymes were not components of the progenitors of anaerobic autotropic cells [28] turns out to be correct and that these early autotrophic cells harboured metabolisms with characteristics that are reminiscent of contemporary  $H_2$ -dependent methanogens [9], a very different picture of the metabolism of primitive hydrogenotrophic methanogens emerges. Whereas the pathway of contemporary hydrogenotrophic methanogenesis is cyclic with the first and last steps linked by the oxidation and reduction of low potential Fd, respectively (figure 3*a*) [17], it is possible that an earlier version of this pathway was not cyclic but rather linear where the first and last steps were unlinked. For example, in the last step of hydrogenotrophic methanogenesis, soluble bifurcating Mvh-Hdr functions to couple the oxidation of  $H_2$  with the simultaneous endergonic reduction of Fd and heterodisulfide in coenzyme M-coenzyme B (CoM-CoB); the reduced Fd that is generated is then used to drive the exergonic reduction of CO<sub>2</sub>

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and methanofuran to form N-formylmethanofuran [17]. Membrane-associated ion-translocating [NiFe]-hydrogenase enzymes (Ech/Eha/Ehb) function to supply additional reduced Fd for the initial reduction of CO<sub>2</sub> to form N-formylmethanofuran [17]. In an early Earth environment that favoured the direct reduction of Fd with H<sub>2</sub>, a soluble [NiFe]-hydrogenase without bifurcating functionality (termed Mvh<sub>Anc</sub>) could conceivably generate reduced Fd without coupling it to the reduction of CoM-CoB; CoM-CoB reduction could have been catalysed by a soluble heterodisulfide reductase (Hdr<sub>Anc</sub>) (figure 3*b*). A membrane bound [NiFe]-hydrogenase without ion-translocating capabilities (termed Eha/Ehb<sub>Anc</sub>) could have conceivably also supplied additional reduced Fd via direct reduction with H<sub>2</sub> (figure 3*b*), although it is unclear if such an enzyme would provide a selective advantage to a cell where Mvh<sub>Anc</sub> was also present. As such, this putative ancestral methanogenesis pathway would be linear rather than cyclic. Moreover, the eventual emergence of bifurcating enzymes would help to explain the evolution of the cyclic pathway of methanogenesis that we see perpetuated in extant hydrogenotrophic methanogenesis from a simpler linear pathway in a manner similar to that proposed for the Kreb's cycle [52].

An origin for autotrophic life in an anoxic, alkaline environment undergoing active serpentinization, as suggested by the data presented here, is consistent with other data suggesting these environment types may have been conducive to the origin of autotrophic cells. For such cells to originate, they would have needed a constant source of chemically transducible energy, a requirement that is met in the form of abundant  $H_2$  in environments undergoing active serpentinization [42] and/or in the form of naturally occurring proton gradients when serpentinized fluids (pH $\sim$ 10–11) mix with ocean waters in equilibrium with atmospheric CO<sub>2</sub>  $(pH \sim 6)$  [37]. Likewise, electrons of sufficiently reduced potential would have needed to be efficiently captured for use in the reduction of  $CO_2$  [15,17], a problem that is potentially overcome by H<sub>2</sub> oxidizing, Fd-reducing [NiFe]-hydrogenases and the low potentials associated with H<sub>2</sub> in hyperalkaline serpentinizing systems (figure 2). However, the hyperalkaline pH also presents a potential problem for primitive autotrophic cells since CO<sub>2</sub> speciates towards carbonate which is less bioavailable. Contemporary hydrogenotrophic methanogens, however, are capable of growth at pH of up to 10.0 [45,46] and have been detected in serpentinizing environments with even more alkaline pH (11.3) [48,49]; the same may have been true for their ancestors. It is possible that in such modern environments, autotrophic cells overcome limitations of inorganic carbon by using more soluble and hydrogenated forms of carbon such as carbon monoxide or formate, which can also be produced abiotically from  $H_2$  and  $CO_2/HCO_3^-$  under serpentinizing conditions [53]. Indeed, formate has been detected in natural serpentinizing environments such as Lost City [54] and in Oman [48]. Under formatotrophic methanogenic growth conditions, oxidation of low potential  $H_2$  may simply augment the metabolism of these autotrophic cells. Collectively, the characteristics of environments undergoing active serpentinization make them attractive targets for understanding the possible origin of autotrophic metabolism, and perhaps metabolism in general. The ideas presented here are testable by probing extant life inhabiting alkaline serpentinite hosted, H<sub>2</sub> rich environments via biochemical, genetic and physiological approaches and through metabolic reconstructions enabled by phylogenomic approaches.

Data accessibility. All data presented herein are available in electronic supplementary material or through communication with the authors.

Authors' contributions. E.S.B., M.J.A. and A.S.T. designed the study. S.P. conducted informatics analyses. All authors contributed to data analysis and interpretation, manuscript preparation and approve of the version to be published.

Competing interests. We have no competing interests.

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## References

- 1. Meyer J. 1988 The evolution of ferredoxins. *Trends Ecol. Evol.* 3, 222–226. (doi:10.1016/0169-5347(88)90162-0)
- 2. Bruschi M, Guerlesquin F. 1988 Structure, function and evolution of bacterial ferredoxins. *FEMS Microbiol. Lett.* **54**, 155–175. (doi:10.1016/0378-1097(88)90175-9)
- 3. Eck RV, Dayhoff MO. 1966 Evolution of the structure of ferredoxin based on living relics of primitive amino acid sequences. *Science* **152**, 363–366. (doi:10.1126/science.152. 3720.363)
- 4. Dey A, Jenney Jr FE, Adams MW, Babini E, Takahashi Y, Fukuyama K, Hodgson KO, Hedman B, Solomon EI. 2007 Solvent tuning of electrochemical potentials in the active sites of HiPIP versus ferredoxin. *Science* **318**, 1464–1468. (doi:10.1126/science.1147753)
- 5. Lyons TW, Reinhard CT, Planavsky NJ. 2014 The rise of oxygen in Earth's early ocean and atmosphere. *Nature* **506**, 307–315. (doi:10.1038/nature13068)
- Major TA, Burd H, Whitman WB. 2004 Abundance of 4Fe-4S motifs in the genomes of methanogens and other prokaryotes. *FEMS Microbiol. Lett.* 239, 117–123. (doi:10.1016/j. femsle.2004.08.027)
- 7. Buckel W, Thauer RK. 2018 Flavin-based electron bifurcation, ferredoxin, flavodoxin, and anaerobic respiration with protons (Ech) or NAD<sup>+</sup> (Rnf) as electron acceptors: a historical review. *Front. Microbiol.* **9**, 401. (doi:10.3389/fmicb.2018.00401)
- Russell MJ, Martin W. 2004 The rocky roots of the acetyl-CoA pathway. *Trends Biochem. Sci.* 29, 358–363. (doi:10.1016/j.tibs.2004.05.007)
- Weiss MC, Sousa FL, Mrnjavac N, Neukirchen S, Roettger M, Nelson-Sathi S, Martin WF. 2016 The physiology and habitat of the last universal common ancestor. *Nat. Microbiol.* 1, 16116. (doi:10.1038/nmicrobiol.2016.116)
- Ueno Y, Yamada K, Yoshida N, Maruyama S, Isozaki Y. 2006 Evidence from fluid inclusions for microbial methanogenesis in the early Archaean era. *Nature* 440, 516–519. (doi:10.1038/nature04584)
- 11. Boyd ES, Schut G, Adams MWW, Peters JW. 2014 Hydrogen metabolism and the evolution of respiration. *Microbe* **9**, 361–367. (doi:10.1128/microbe.9.361.1)
- Moore EK, Jelen BI, Giovannelli D, Raanan H, Falkowski PG. 2017 Metal availability and the expanding network of microbial metabolisms in the Archaean eon. *Nat. Geosci.* 10, 629–636. (doi:10.1038/ngeo3006)
- 13. Fuchs G. 1994 Variations of the acetyl-CoA pathway in diversely related microorganisms that are not acetogens. In *Acetogenesis* (ed. HL Drake), pp. 506–538, Chapman and Hall. (doi:10.1007/978-1-4615-1777-1\_19)
- Buckel W, Thauer RK. 2013 Energy conservation via electron bifurcating ferredoxin reduction and proton/Na(+) translocating ferredoxin oxidation. *Biochim. Biophys. Acta* 1827, 94–113. (doi:10.1016/j.bbabio.2012.07.002)
- Muller V, Chowdhury NP, Basen M. 2018 Electron bifurcation: a long-hidden energycoupling mechanism. *Annu. Rev. Microbiol.* 72, 331–353. (doi:10.1146/annurev-micro-090816-093440)
- Thauer RK. 1998 Biochemistry of methanogenesis: a tribute to Marjory Stephenson. Microbiology (Reading, England) 144, 2377–2406. (doi:10.1099/00221287-144-9-2377)
- Thauer RK, Kaster A-K, Seedorf H, Buckel W, Hedderich R. 2008 Methanogenic archaea: ecologically relevant differences in energy conservation. *Nat. Rev. Microbiol.* 6, 579–591. (doi:10.1038/nrmicro1931)
- Herrmann G, Jayamani E, Mai G, Buckel W. 2008 Energy conservation via electrontransferring flavoprotein in anaerobic bacteria. J. Bacteriol. 190, 784–791. (doi:10.1128/jb. 01422-07)
- Bennett BD, Kimball EH, Gao M, Osterhout R, Van Dien SJ, Rabinowitz JD. 2009 Absolute metabolite concentrations and implied enzyme active site occupancy in *Escherichia coli*. *Nat. Chem. Biol.* 5, 593–599. (doi:10.1038/nchembio.186)
- Müller V. 2003 Energy conservation in acetogenic bacteria. *Appl. Environ. Microbiol.* 69, 6345–6353. (doi:10.1128/aem.69.11.6345-6353.2003)
- 21. Peters JW, Miller AF, Jones AK, King PW, Adams MW. 2016 Electron bifurcation. *Curr. Opin. Chem. Biol.* **31**, 146–152. (doi:10.1016/j.cbpa.2016.03.007)

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- 22. Hess V, Schuchmann K, Muller V. 2013 The ferredoxin:NAD<sup>+</sup> oxidoreductase (Rnf) from the acetogen *Acetobacterium woodii* requires Na<sup>+</sup> and is reversibly coupled to the membrane potential. *J. Biol. Chem.* **288**, 31 496–31 502. (doi:10.1074/jbc.M113.510255)
- Biegel E, Schmidt S, Gonzalez JM, Muller V. 2011 Biochemistry, evolution and physiological function of the Rnf complex, a novel ion-motive electron transport complex in prokaryotes. *Cell. Mol. Life Sci.* 68, 613–634. (doi:10.1007/s00018-010-0555-8)
- 24. Hedderich R. 2004 Energy-converting [NiFe] hydrogenases from Archaea and extremophiles: ancestors of complex I. *J. Bioenerg. Biomembr.* **36**, 65–75. (doi:10.1023/b:jobb. 0000019599.43969.33)
- 25. Martin WF. 2012 Hydrogen, metals, bifurcating electrons, and proton gradients: the early evolution of biological energy conservation. *FEBS Lett.* **586**, 485–493. (doi:10.1016/j. febslet.2011.09.031)
- 26. Nitschke W, Russell MJ. 2012 Redox bifurcations: mechanisms and importance to life now, and at its origin. *Bioessays* **34**, 106–109. (doi:10.1002/bies.201100134)
- Sousa FL, Preiner M, Martin WF. 2018 Native metals, electron bifurcation, and CO<sub>2</sub> reduction in early biochemical evolution. *Curr. Opin. Microbiol.* 43, 77–83. (doi:10.1016/j.mib.2017.12.010)
- Poudel S, Dunham EC, Lindsay MR, Amenabar MJ, Fones EM, Colman DR, Boyd ES. 2018 Origin and evolution of flavin-based electron bifurcating enzymes. *Front. Microbiol.* 9, 1762. (doi:10.3389/fmicb.2018.01762)
- 29. Sievers F *et al.* 2011 Fast, scalable generation of high-quality protein multiple sequence alignments using Clustal Omega. *Mol. Syst. Biol.* **7**, 539. (doi:10.1038/msb.2011.75)
- Greening C, Biswas A, Carere CR, Jackson CJ, Taylor MC, Stott MB, Cook GM, Morales SE. 2016 Genomic and metagenomic surveys of hydrogenase distribution indicate H<sub>2</sub> is a widely utilised energy source for microbial growth and survival. *ISME J.* **10**, 761–777. (doi:10.1038/ismej.2015.153)
- Hagedoorn PL, Driessen MCPF, van den Bosch M, Landa I, Hagen WR. 1998 Hyperthermophilic redox chemistry: a re-evaluation. *FEBS Lett.* 440, 311–314. (doi:10.1016/ S0014-5793(98)01466-5)
- Magliozzo RS, McIntosh BA, Sweeney WV. 1982 Origin of the pH dependence of the midpoint reduction potential in *Clostridiurn pasteurianum* ferredoxin: oxidation statedependent hydrogen ion association. *J. Biol. Chem.* 257, 3506–3509.
- Stombaugh NA, Sundquist JE, Burris RH, Orme-Johnson WH. 1976 Oxidation-reduction properties of several low potential iron-sulfur proteins and of methylviologen. *Biochemistry* 15, 2633–2641. (doi:10.1021/bi00657a024)
- Lode ET, Murray CL, Rabinowitz JC. 1976 Apparent oxidation-reduction potential of *Clostridium acidi-urici* ferredoxin. Effect of pH, ionic strength, and amino acid replacements. J. *Biol. Chem.* 251, 1683–1687.
- Takai K *et al.* 2008 Cell proliferation at 122°C and isotopically heavy CH<sub>4</sub> production by a hyperthermophilic methanogen under high-pressure cultivation. *Proc. Natl Acad. Sci. USA* 105, 10949–10954. (doi:10.1073/pnas.0712334105)
- 36. Sleep NH, Bird DK, Pope EC. 2011 Serpentinite and the dawn of life. *Phil. Trans. R. Soc. B* 366, 2857–2869. (doi:10.1098/rstb.2011.0129)
- Russell MJ, Hall AJ, Martin W. 2010 Serpentinization as a source of energy at the origin of life. *Geobiology* 8, 355–371. (doi:10.1111/j.1472-4669.2010.00249.x)
- Sojo V, Herschy B, Whicher A, Camprubi E, Lane N. 2016 The origin of life in alkaline hydrothermal vents. *Astrobiology* 16, 181–197. (doi:10.1089/ast.2015.1406)
- 39. Djokic T, Van Kranendonk MJ, Campbell KA, Walter MR, Ward CR. 2017 Earliest signs of life on land preserved in ca. 3.5Ga hot spring deposits. *Nat. Commun.* **8**, 15263. (doi:10.1038/ncomms15263)
- Miyakawa S, James Cleaves H, Miller SL. 2002 The cold origin of life: A. implications based on the hydrolytic stabilities of hydrogen cyanide and formamide. *Orig. Life Evol. B.* 32, 195–208. (doi:10.1023/A:1016514305984)
- 41. Bach W, Paulick H, Garrido CJ, Ildefonse B, Meurer WP, Humphris SE. 2006 Unraveling the sequence of serpentinization reactions: petrography, mineral chemistry, and petrophysics of serpentinites from MAR 15°N (ODP Leg 209, Site 1274). *Geophys. Res. Lett.* **33**, L13306. (doi:10.1029/2006GL025681)
- 42. Schulte M, Blake D, Hoehler T, McCollom T. 2006 Serpentinization and its implications for life on the early Earth and Mars. *Astrobiology* **6**, 364–376. (doi:10.1089/ast.2006.6.364)

- Mulder DW, Boyd ES, Sarma R, Lange RK, Endrizzi JA, Broderick JB, Peters JW. 2010 Stepwise [FeFe]-hydrogenase H-cluster assembly revealed in the structure of HydA(DeltaEFG). *Nature* 465, 248–251. (doi:10.1038/nature08993)
- Slonczewski JL, Fujisawa M, Dopson M, Krulwich TA. 2009 Cytoplasmic pH measurement and homeostasis in Bacteria and Archaea. *Adv. Microb. Physiol.* 55, 1–79, 317. (doi:10.1016/s0065-2911(09)05501-5)
- Kral TA, Birch W, Lavender LE, Virden BT. 2014 Potential use of highly insoluble carbonates as carbon sources by methanogens in the subsurface of Mars. *Planet. Space Sci.* 101, 181–185. (doi:10.1016/j.pss.2014.07.008)
- Miller HM, Chaudhry N, Conrad ME, Bill M, Kopf SH, Templeton AS. 2018 Large carbon isotope variability during methanogenesis under alkaline conditions. *Geochim. Cosmochim. Acta* 237, 18–31. (doi:10.1016/j.gca.2018.06.007)
- Crespo-Medina M, Twing KI, Sánchez-Murillo R, Brazelton WJ, McCollom TM, Schrenk MO. 2017 Methane dynamics in a tropical serpentinizing environment: the Santa Elena Ophiolite, Costa Rica. *Front. Microbiol.* 8, 916. (doi:10.3389/fmicb.2017.00916)
- Rempfert KR, Miller HM, Bompard N, Nothaft D, Matter JM, Kelemen P, Fierer N, Templeton AS. 2017 Geological and geochemical controls on subsurface microbial life in the Samail Ophiolite, Oman. *Front. Microbiol.* 8, 56. (doi:10.3389/fmicb.2017.00056)
- Fones EM, Colman DR, Kraus EA, Nothaft DB, Poudel S, Rempfert KR, Spear JR, Templeton AS, Boyd ES. 2019 Physiological adaptations to serpentinization in the Samail Ophiolite, Oman. *ISME J.* 13, 1750–1762. (doi:10.1038/s41396-019-0391-2)
- 50. Gibney BR, Mulholland SE, Rabanal F, Dutton PL. 1996 Ferredoxin and ferredoxin-heme maquettes. *Proc. Natl Acad. Sci. USA* 93, 15041–15046. (doi:10.1073/pnas.93.26.15041)
- Mulholland SE, Gibney BR, Rabanal F, Dutton PL. 1999 Determination of nonligand amino acids critical to [4Fe-4S]<sup>2+/+</sup> assembly in ferredoxin maquettes. *Biochemistry* 38, 10442–10448. (doi:10.1021/bi9908742)
- Melendez-Hevia E, Waddell TG, Cascante M. 1996 The puzzle of the Krebs citric acid cycle: assembling the pieces of chemically feasible reactions, and opportunism in the design of metabolic pathways during evolution. J. Mol. Evol. 43, 293–303. (doi:10.1007/BF02338838)
- McCollom TM, Seewald JS. 2003 Experimental constraints on the hydrothermal reactivity of organic acids and acid anions: I. Formic acid and formate. *Geochim. Cosmochim. Acta* 67, 3625– 3644. (doi:10.1016/S0016-7037(03)00136-4)
- Lang SQ, Butterfield DA, Schulte M, Kelley DS, Lilley MD. 2010 Elevated concentrations of formate, acetate and dissolved organic carbon found at the Lost City hydrothermal field. *Geochim. Cosmochim. Acta* 74, 941–952. (doi:10.1016/j.gca.2009.10.045)