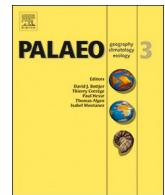




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Invited Research Article

Cyclostratigraphy and astrochronology: Case studies from China

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ABSTRACT

A high-precision geologic time scale is the essential key for understanding the Earth's evolutionary history and geologic processes. Astronomical tuning of orbitally forced stratigraphic records to construct high-resolution Astronomical Time Scales (ATS) has led to a progressive refinement of the geologic time scale over the past two decades. In turn, these studies provide new insights regarding the durations and rates of major Earth events, evolutionary processes, and climate changes, all of which provide a scientific basis for contextualizing and predicting future global change trends. South China hosts some of the best-exposed and well-dated Neoproterozoic through Mesozoic stratigraphic sections in the world; many of which are suitable for cyclostratigraphy and calibrating the geologic time scale. In North China, several Cenozoic oil-bearing basins have deep boreholes with continuous sampling and/or well logging that enable derivation of astronomically tuned time scales for an improved understanding of basin evolution and hydrocarbon generation. This Special Issue focuses on case studies of astrochronology and applied cyclostratigraphy research using reference sections within China. In this introductory overview, we: (1) summarize all existing astrochronology studies of the Neoproterozoic through Cenozoic sections within China that have been used to enhance the international geologic time scale, (2) examine briefly the astronomically forced paleoclimate information recorded in various depositional systems and the modern techniques employed to analyze the periodicity of these signals encoded within the sedimentary record, and (3) summarize the 20 contributions to this Special Issue of *Palaeogeography, Palaeoclimatology, Palaeoecology* on 'Cyclostratigraphy and Astrochronology: Case studies from China'.

1. Introduction

Astronomically induced variations in Earth's precession and obliquity modulate seasonal insolation and drive climate changes in the Earth system (Hinnov, 2018). The Earth's seasons are caused by the obliquity (tilt) of its rotation axis relative to the Sun. The Earth's orbit around the Sun is slightly elliptical (= eccentricity); therefore, the relative average summer temperature in a hemisphere depends upon whether its summer occurs during the Earth's furthest distance from the Sun (aphelion; which is the current situation for the Northern Hemisphere) or during its closest approach (perihelion). The precession of the Earth's rotation axis relative to its orbit causes a ca. 20-kyr cycle in the relative warmth of a hemisphere's summer (and coldness of its winter) as it oscillates between aphelion to perihelion conditions. The eccentricity of the Earth's orbit is influenced by the gravitational attraction of other planets, and changes from nearly circular (the current situation with a relatively low importance of precession-induced

climatic change) to a more elliptical orbit that causes greater seasonal contrasts. Eccentricity varies with a ca. 100-kyr oscillation that is further modulated by a 405-kyr cycle. Superimposed on the precession-eccentricity signal are ca. 40-kyr periodic variations of the obliquity of the Earth's axis; and these are especially important in influencing seasonal contrasts at higher latitudes.

Together, these variations in the relative differences in summer heat versus winter cooling are called "Milankovitch cycles" after Milutin Milanković, the Serbian mathematician who did the initial detailed numerical analysis of the effects on solar insolation at different latitudes (Milankovitch, 1930, 1941). These long-term orbital-induced climate cycles are also manifested as changes in many different geographic and sedimentary processes, including average seasonal rainfall and fluvial runoff, ice cap size, global sea level, types and amount of vegetation, rates and styles of continental weathering, the position and intensity of wind belts, monsoon patterns, coastal currents, and the supply and recycling of nutrients within both marine and lacustrine waters. In turn,

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sediment influxes into all depositional systems, from continental lacustrine sediments to deep-sea deposits, undergo shifts in the relative ratios of coarse-to-fine siliciclastics and of terrigenous to biogenic components, and in the relative importance of storm and flood events. Consequently, the periodic climatic changes attributable to orbital forcing are readily encoded in stratigraphic successions, and cyclostratigraphers can measure chemical and physical proxies (e.g., elemental concentrations, grain size, magnetic susceptibility, etc.) that track these cycles in order to identify astronomical signals. One challenge to this endeavor is that the recording of Milankovitch cycles in the sedimentary record is superimposed upon a wide spectrum of non-periodic variability, and sedimentary systems are prone to shifts in accumulation rates and sediment composition induced by non-Milankovitch factors such as regional tectonics, temporally irregular shifting of fluvial-delta depocenters, and other trends.

The first published application of this cyclostratigraphic "clock in the rock" concept was by G.K. Gilbert (1895), who used it to estimate elapsed time in Upper Cretaceous deposits in North America. The discipline became more widely established after the seminal work by Hays et al. (1976), which demonstrated pervasive astronomical control on climate changes across the past ~0.5 million years. Since then, the astronomical theory has been widely studied and applied. During the past two decades in particular, astronomical calibration of stratigraphic successions has been successfully used to constrain the Cenozoic and Mesozoic geologic time scales, aided by the development of numerical solutions of Earth's orbit across the past ~50 million years by Jacques Laskar et al. (1993, 2004, 2011).

China hosts superb Neoproterozoic and Mesozoic stratigraphic sections in marine facies, as well as lacustrine deposits encompassing large portions of the Permian through Neogene, and thick accumulations of wind-blown loess spanning the Quaternary. Cyclostratigraphic analyses of proxy data from outcrops and boreholes within China have been important for calibrating the geologic time scale, understanding basin evolution and oil-forming mechanisms, and for unraveling the impacts of past glacial-interglacial climate oscillations. This special issue of *Palaeogeography, Palaeoclimatology, Palaeoecology* brings together 20 papers that highlight the importance of Chinese sections for cyclostratigraphy and time scales. The studies cover a wide range of stratigraphic records in marine and terrestrial deposits within China, and incorporate aspects of biostratigraphy, lithostratigraphy, chemostratigraphy, magnetostratigraphy, and radioisotopic dating.

In this introductory review paper, we present two summaries of cyclostratigraphy studies in China. We first summarize all the cyclostratigraphy studies of sections within China that present astronomical time scales that enable high-resolution enhancement of the geologic time scale (Ediacaran through Quaternary). Six of these are also part of this Special Issue (Fang et al., 2020; Lu et al., 2019; Ma et al., 2019; Ma et al., 2020a; Sui et al., 2019; Zhong et al., 2019b). The second group are those that apply cyclostratigraphy and astronomically tuned depositional rates to better understand basin development and Earth system evolution during portions of the Neoproterozoic through Cenozoic (Chu et al., 2020; Du et al., 2020; Gong et al., 2019; Li et al., 2020; Liu et al., 2020; Peng et al., 2020; Xu et al., 2019; Xu et al., 2020; Yao and Hinnov, 2019; Zhang et al., 2019a, 2019b; Zhang et al., 2020; Zhang 2019; Zhao et al., 2019).

2. Geologic settings of the sedimentary records of China

The sedimentary history of regions within China were particularly conducive to recording climate cycles in relatively stable depositional settings. The South China Craton was a part of the Gondwana continental margin during the Cryogenian through Silurian; then, as a separate island mini-plate from the Devonian through middle Triassic, it accumulated a thick carbonate platform adjacent to eroding "oldlands". Consequently, South China hosts some of the best-exposed and well-dated Neoproterozoic, Paleozoic and Mesozoic stratigraphic sections in

the world. The exceptional nature of these fossiliferous and well-preserved marine deposits is why South China hosts nearly a dozen GSSPs (Global Stratigraphic Section and Point) of stratigraphic boundaries for the Paleozoic and early Mesozoic geologic stages. Equally, these reference sections have provided cyclostratigraphic records for calibrating the durations of biozones and other events of the international geologic time scale.

Basins within other blocks that now comprise China contained large lakes at tropical to temperate latitudes. In addition to accumulating organic-rich deposits that became important oil and gas source rocks, these lacustrine deposits were particularly conducive to recording climate-induced fluctuations in the relative influx of clay and organic carbon. Cyclostratigraphic studies of these lacustrine deposits have often utilized records from deep boreholes and range from the Permian through Neogene. In particular, North China has several Cenozoic oil-bearing basins with deep boreholes that have been continuously sampled and/or well-logged for cyclostratigraphy studies.

During the Quaternary and continuing to the present, windblown dust from the arid interior regions of central Asia accumulated as thick loess deposits in the middle watershed of the Yellow River and elsewhere. During more humid intervals, such as major interglacial episodes, reddish soil horizons developed on these loess plains; and the resulting alternation of arid-climate loess and humid-climate soil has enabled astronomical calibration of Eurasia's glacial episodes.

3. Astronomical Time Scale studies using reference sections in China

As of July 2020, there have been over 30 cyclostratigraphy studies of marine, lacustrine or loess deposits that have direct application to enhancing the age model for the geologic time scale (Table 1, Fig. 1). Studies of these are part of this Special Issue of *Palaeogeography, Palaeoclimatology, Palaeoecology*, including contributions on the Ediacaran (Sui et al., 2019), Cambrian (Fang et al., 2020) and Ordovician-Silurian (Lu et al., 2019; Zhong et al., 2019b; Ma et al., 2020b; Ma et al., 2019). Only the Cenozoic and the Jurassic of China, for which there is a lack of extensive uplifted exposures of fossiliferous marine strata, have not yielded cyclostratigraphic studies that are directly applicable to improving the geologic time scale. However, the entire Cenozoic time scale and much of the Jurassic has been cycle-tuned based on ocean drilling boreholes, uplifted oceanic sediments in the Mediterranean region, and borehole and outcrop studies in Britain and other regions.

As expected from its geologic history, nearly all of the Ediacaran, Paleozoic and Triassic cyclostratigraphy studies have been undertaken on the South China Craton (Fig. 1). The exceptions are Ordovician studies in the Tarim Basin, which included the auxiliary GSSP for the Middle/Late Ordovician boundary (Sandbian GSSP), and in Hebei Province of North China. North China has cyclostratigraphy studies of Cretaceous terrestrial basins that helped to calibrate the evolution of feathered dinosaurs (Wu et al., 2013b) and of magnetic polarity chron. The thick loess deposits in Shaanxi Province allowed calibration of the major Quaternary climate oscillations and the development of desert regions of eastern Asia (e.g., Ding et al., 2002).

Nearly all studies on pre-Cenozoic Chinese reference sections recognized the dominant 405-kyr long-eccentricity modulation of the precession cycle within marine or lacustrine deposits. In contrast to the theoretical lengthening of the periods for Earth's precession and obliquity through geologic time due to the influence of its Moon, the 405-kyr long-eccentricity cycle has had a stable period for at least the past half-billion years (e.g., Hinnov, 2018). This 405-kyr signal in the sedimentary record enables an "astronomical tuning" of a meter-scale record yielding a "floating" time scale. The resulting durations of geologic stages and of corresponding biozones, geochemical excursions, and magnetic polarity zones derived by these studies generally have a precision of ~0.1 Myr (i.e., one quarter of a long-eccentricity cycle). In the case of Songliao Basin borehole in northeastern China, radio-

Table 1

Summary of astrochronology studies on Chinese reference sections that are used to enhance the geologic time scale. Abbreviations: (1) Proxies – MS = magnetic susceptibility, GR = natural gamma-ray radiation, ARM = anhysteretic remanent magnetization, XRF = X-ray fluorescence; (2) Method – MTM = multi-taper method, Evol. spectra = evolutive spectra; (3) Cycles – E = long-eccentricity, e = short-eccentricity, O = obliquity, P = precession; (4) Other – U-Pb = uranium-lead dating (Ding et al., 1994; Lu et al., 1999; Heslop et al., 2000; Deng et al., 2019a; Sun et al., 2006; Wu et al., 2014; Wang et al., 2016; Deng et al., 2013; Wang et al., 2016; Wu et al., 2019; Xi et al., 2009; Wu et al., 2019; He et al., 2012; Wu et al., 2013b; Zhu et al., 2017a,b; Huang, 2019; Li et al., 2019; Zhang et al., 2019; Tong et al., 2017; Tong et al., 2017a,b; Litt et al., 2007a,b; Zhao et al., 2012; Mundil et al., 2004; Yuan et al., 2013c; Shen et al., 2011; Mundil et al., 2005; Wu et al., 2005; Wu et al., 2007; Wu et al., 2012; Ueno et al., 2015; Fang et al., 2015; Fang et al., 2012; Ueno et al., 2012; Ueno et al., 2019a,b; Zhang et al., 2013; Wang et al., 2019a,b; Zhang et al., 2019a,b; Ueno et al., 2019d, 2019e, 2019f; Ueno et al., 2019; Xue et al., 2015; Fang et al., 2015; Fang et al., 2012; Ueno et al., 2012; Ueno et al., 2019a,b; Zhang et al., 2019a,b; Chen et al., 2018; Chen et al., 2011; Ma et al., 2019a; Ross and Ross, 1988; Li et al., 1997; Rygel et al., 2008; Ueno et al., 2013; Bai, 1995b; Qie et al., 2019; Gong et al., 2001; Gong et al., 2005; Huang et al., 2016; Rong et al., 2019; Chen et al., 2018; Chen et al., 2011; Ma et al., 2019a; Wang et al., 2018; Ma et al., 2016; Ma et al., 2019c; Peng et al., 2009; Chen et al., 2006; Zhu et al., 2005; Condon et al., 2008; McFadden et al., 2008; Condon et al., 2005; Zhou et al., 2019a,b; Bai et al., 1982; Bai et al., 1994; Li et al., 2009; Zhong et al., 2018; Sui et al., 2017; Hu and Qi, 2017; Sui et al., 2018; Wang et al., 2009; Zhong et al., 2018).

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Table 1 (continued)

Table 1 -- Summary of astrochronology studies on Chinese reference sections used to enhance the geologic time scale

#	Author (* =Corresp. if different; # =Corresp. if different, if different, if different, if different)	Geologic System (Stage)	What was calibrated	Result	Location	Main studied facies; thickness; depos. environ.; Age control	(1) Proxies; (2) Method	Main cycle for tuning (implied span of record); Other recorded cycles	Raw data provided?
1	Ding, Z.L., et al., 2002, Palaeocean., 17: 1033.	Quaternary	Loess and paleosol cycles (arid/humid)	Loess-paleosol alternations mainly conlusive to marine delta-Oxy 16 cycles for past 1.8 Myr.	Shaanxi-Gansu border; Southern and middle Loess plateau (Baotou 34.4°N, 107.4°E; Jingchuan 35.1°N, 107.6°E; and Pingjiang 35.6°N, 106.7°E)	Silt; paleosol; ca. 160 m average; wind-blown dust; Magnetobatigraphy	(1) Grain size; (2) Assumed in-phase relation to global ice volume for tuning	Obliquity and precession (33 main paleosols (Baotou section); 2.6 Myr)	Upon request to Ding, Z.L.
Neogene -- no cyclostratigraphic studies in China applicable for enhancing geologic time scale									
2	Wu, H.C., et al., 2014, EPSL, 407: 82–95.	Cretaceous (Upper Santonian to lower Danian)	Magnetostratigraphy; spore-pollen,孢粉-孢子带; charophyte and ostracod zones	Durations C30n-C31n in 3.3 Myr). C31n (2.2 Myr) C32n (2.0 Myr). Base of Campanian Upper Cretaceous is composite of two boreholes ca. 77 km apart in the south-central part of the basin (Wang et al., 2016), then refined as 82.875 Ma (Wu et al., PNAS, in press).	Jilin Prov., Songliao Basin, SK1 north ca. 45.3°N, 125.0°E); SK1 C33n-C33r (9.0 Myr). Base of Campanian (use of chiton C33n) is assumed to be composite of two boreholes ca. 77 km apart in the south-central part of the basin.	Lacustrine dolomite-siltstone; 1542 m; deep lake to shallow lake; U-Pb dates	(1) Thorium component of GR; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (~30 m) for tuning (37 cycles; ca. 15 Myr plus ca. 3 Myr hiatus); also e, O, P, Modulations of 1.36 and 2.05 Myr	
3	Wu, H.C., et al., 2013a, Palaeo-3, 38: 55–70.	Cretaceous (lower Turonian to lower Campanian)	Magnetostratigraphy; ostracod zones	Base of Campanian base of polarity Chron C33r is 83.63 Ma (but see 2014 and 2020 revisions above). Brief Review of C33n in lower Santonian at ca. 84.95, 85, and 85.5 Ma.	Jilin Prov., Songliao Basin, SK1 south ca. 44.5°N, 125.0°E); SK1 C33n-C33r (9.0 Myr). Base of Campanian Upper Cretaceous is composite of two boreholes ca. 77 km apart in the south-central part of the basin (Wang et al., PNAS, in press).	Lacustrine dolomite-siltstone; 950 m; deep lake; U-Pb dates	(1) MS, GR; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (~35 m) for tuning (21 cycles; 8.9 Myr); also e, O, P, Modulations of 1.2 and 2.34 Myr	No
4	Wu, H.C., et al., 2013b, Palaeo-3, 38: 221–228.	Cretaceous (lower)	Jehol vertebrate beds; Magnetostrat	Feathered dinosaur interval spans only ca. 1.5 Myr. Entire normal-polarity unit is ca. 0.86 Myr; which could be M3n, ca. 120.79 Ma.	W Liaoning Prov. (Bajiaogou; Shentan outcrop, 41.59°N, 120.79°E)	Lacustrine dolomite with volcanic ash; 11.2 m; U-Pb dates (ca. 126–124 Ma)	(1) MS, ARM; (2) MTM, Evol. spectra, band-pass filters, tuning	100-kyr e (~2.1 m) for tuning (6.7 cycles; 0.7 Myr); also O, P	No
Jurassic -- no cyclostratigraphic studies in China applicable for enhancing geologic time scale									
5	Li, M.S., et al., 2017, Huang, C.J., EPSL, 474: 207–223.	Triassic (upper Norian-Rhaetian)	Magnetostrat	Verifies cyclostratigraphic data of latest 5.5 Myr of Thasic polarity scale from Newark Group. Age of the earliest dinosaur footprints in China is 2.5 My before end of Triassic (middle Rhaetian; ca. 204 Ma).	Sichuan Prov. (readouts of Xujiahe 32.42°N, 105.65°E; Zilianza 29.83°N, 106.38°E)	Siltstone and sandstone; ca. 600 m; lacustrine-fluvial; Magnetostrat (same paper)	(1) MS, GR; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (~55 m) for tuning (16.5 cycles; 6.7 Myr); also e	Suppl. Excel table ca. 3600 GR 1200 MS ps)
6	Zhang, Y., et al., 2015, Ogg, J.G., Palaeo-3, 146: 135–166.	Triassic (lower Carnian)	Magnetostrat of Early Carnian (Julian substage)	Guizhou Prov. (Mayevo outcrop, Lashike outcrop, 25.57°N, 105.49°E)	Limestone, with clay-rich uppermost Julian; ca. 150 m interbedded dolomitic-carbonate shelf-slope; Conodonts	(1) GR; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (~34 m) for tuning (6 cycles; 2.4 Myr); also e, P	Suppl. Excel table ca. 550 GR ps)	
7	Li, M.S., Huang, 2016a, EPSL, 441: 1025–	Triassic (upper Anisian)	Magnetostrat	Anisian Stage 5.3 Myr; durations of substages, of 6 condon zones (or pairs of zones), and of 9 pairs of magnetic-polarity zones.	Guizhou Prov. (Luodian, Guandao 25.61°N, 106.62°E)	Limestone; 330 m (265 m after removing breccias); slope; Conodonts	(1) GR, MS; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (~3-m) for tuning (15 cycles; 6.2 Myr); also e, O, P	Suppl. Excel table 1070 GR and 260 MS ps)
8	Li, M.S., et al., 2016a (Huang, 2016a, EPSL, 441: 1025–	Triassic (Anisian)	International geological stage (Frasnian); Conodont and magnetic-polarity zones	Guizhou Prov. (Luodian, Guandao outcrop, 25.61°N, 106.62°E)	Anhui Prov. (Chaozhou quarries with candidate GSSSP of Olenekian, and near Olenekian) limestone with much higher accumulation rate, merged composite is ca. 300 m Conodonts. Magnetostrat to Germanic Basin	(1) GR, (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (~3-m) for tuning (15 cycles; 6.2 Myr); also e, P	Suppl. Excel table total of 5000 GR ps for 3 sections)	
9	Wu, H.C., et al., 2012, Gondwana Res., 22: 748–759.	International geologic epoch (Lopingian) substages (Indian, Olenekian) and substage; Conodont zones and magnetic-polarity zones	International geologic Indian Stage 2.0 Myr (Griesbachian substage 1.4–0.1 Myr; Dierenian 0.6–0.1 Myr); Olenekian Stage 3.3 Myr (Smithian 1.7–1.0 Myr; Shantan 1.4–1.0 Myr); durations of China regional conodont zones and estimated for 10 pairs of magnetic-polarity zones.	International geologic Indian Stage 1.2 Myr (Griesbachian substage 0.5 Myr; Dierenian 0.7 Myr); durations of 6 China regional conodont zones.	Wuchapingian through middle Changhsingian stages are mainly siliciclastic limestone with cherts; Uppermost Changhsingian to lowest Indian is mainly calcareous shales with thin-platform/slope then intraplatform; Conodonts, U-Pb dates	(1) MS, ARM; (2) MTM, band-pass filters, tuning	405-kyr E (~8 m) for tuning (5 cycles; 2 Myr); also e, P	No	(1) Zhao, L.S., et al., 2005, Abderthian, 33: 113–114; Li, H., et al..., 2008, Earth Sci. - J. China Univ. Geosci., 34: 73–74.
10	Wu, H.C., et al., 2013c, Nat. Comm., 4: 2452	Permian (Lopingian)	International geologic epoch (Lopingian) substages; Conodont zones; and Permian extinction	International geologic epoch (Lopingian) substages (Indian, Olenekian) and substage; Conodont zones and Permian extinction	Zhejiang Prov. (GSSSP of Trityliferous limestone at Meishan, 31.083°N, 119.706°E); Sichuan Prov. (Shansi section, 32.333°N, 105.467°E)	Wuchapingian through middle Changhsingian stages are mainly siliciclastic limestone with cherts; Uppermost Changhsingian to lowest Indian is mainly calcareous shales with thin-platform/slope then intraplatform; Conodonts, U-Pb dates	(1) MS, ARM; (2) MTM, Evol. spectra, band-pass filters, tuning relative to Perm-Shansi; also e, O, P, Modulations of 3.45, 1.93, 1.26, 1.02 Myr (ecc) and 3.11, 0.57 Myr (obl).	No	(1) Shen, S.Z., et al., 2011, Science, 334: 1367–1372; Mundil et al., 2004, Science, 305: 1761–1763. (3) Yuan, D.X., et al., 2019, Sci. China Earth Sci., 62: 190–209.

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Table 1 (continued)

#	Author (* =Corresp. if different, † =published)	Geologic System (Stage)	What was calibrated	Result	Location	Main studied facies; thickness, depos. envirn.; Age control	(1) Proxies; (2) Method	Main cycle for tuning (implied span of record); Other recorded cycles	Raw data provided?
11	Xue, W.Q., et al., 2015, Chinese J. Geophys., 56: 611–627.	Permian (Capitanian)	International geologic stage (Capitanian); Conodont zones	Capitanian Stage spans 3.85±0.28 Myr. Capitanian conodont zones <i>Jinlongiodella postserata</i> (0.15 Myr), <i>J. shanensis</i> (0.06 Myr), <i>J. altuarensis</i> (2.30), <i>J. prexuanhanensis</i> (0.08), and <i>Clarina poststrenuensis</i> (0.03 Myr).	Guangxi Prov. (Laibin, Teqiao, Wuchapingan GSIP); 23°69'53" N, 109°32'11" E	Chert and limestone; slope-basin; Conodonts (that study)	(1) MS; (2) MTM, Fourier analysis, tuning	405-kyr E (~12 m) for tuning (15 cycles; 4 Myr); also e, O, P.	?
12	Fang, Q., Wu, H.C., et al., 2015, Palaeo-3, 474: 130–139.	Permian (Wordian, Capitanian)	International geologic stage (Capitanian); Conodont zones	Capitanian Stage spans 2.65 Myr, but with ca. 0.5 Myr no-data below. Capitanian conodont zones <i>Jinlongiodella postserata</i> (0.5 Myr, but with ca. 0.8 Myr no-data below), <i>J. altuarensis</i> (0.82), and <i>J. prexuanhanensis</i> (0.75). Wordian zone <i>J. asseraata</i> (>2.4 Myr).	Sichuan Prov. (Dukou section; 31°37'0" N, 108°30' E)	Limestone; 49 m; carbonate slope; Conodonts	(1) MS, ARM; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (~3-m) for tuning (15 cycles; 6 Myr); also e, O, P. Modulations of 1 Myr	No
13	Fang, Q., Wu, H.C., et al., 2015, Palaeo-3, 440: 848–859.	Permian (Roadian, Wordian, Capitanian)	International geologic stages (Roadian, Wordian); Conodont zones	Roadian (3.7±0.4 Myr) and Wordian (2.9±0.4 Myr) durations.	N. Sichuan Prov. (Shangsi section; 32°33' N, 105°47' E)	Limestone; 201 m; slope; Conodonts	(1) ARM; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (~7 to 15 m) for tuning (24 cycles; 4 Myr); also e, O, P. Modulations of 2.0 Myr (ecc.) and 1.0 Myr (oblique)	No
14	Ueno, K., et al., 2013, Geol. Soc. Lond. Spec. Publ., 376: 235–267.	Carboniferous (Moscowian through Gzhelian) and Early Permian	International Stages; Fusilid zones	Some major lowstand exposure surfaces correlate with major cyclothems in Midcontinent USA and Moscow basin, but not all of those cycloths are resolved.	Guizhou Prov. (Zongdi section, ca. 30 km SE of Zuny on county road connecting to Luodian; 25°38'50" N, 106°35'2" E)	Limestone; 360 m; shallow-marine; Fusilids	(1) Lowstand exposure surfaces; (2) Seq. Strat. analysis	405-kyr E (indirectly) (21 deposition sequences in Upper Pennsylvania, but the 5 recognition in Asselian may have skipped beads)	Detailed stratigraphy log diagrams
15	Wu, H.C., et al., 2018, Geology, 47: 83–86.	Carboniferous (Serpukhovian through Gzelian)	International Stages (5); regional Conodont zones (25)	Serpukhovian 7.6 Myr; Bashkirian 8.1 Myr; Moscowian 8.5 Myr; Kashinovian 2.87 Myr; Gzhelian 4.83 Myr durations.	S. Guizhou Prov. (Naging roadcut; GSSPs for Serpuhovian; Kashinovian and Gzhelian).	Limestone; 250 m; carbonate slope; Conodonts	(1) MS; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (~3-m) for tuning (837 cycles; 33.9 Myr); also e, O, P. Modulations of 1.12, 1.6, 2.4 Myr	Suppl. Excel table (5000 pps)
16	Fang, Q., Wu, H.C., et al., 2018, J. Asian Earth Sci., 156: 302–315.	Carboniferous (Serpukhovian through Moscowian)	International Stages (Serpukhovian); regional Conodont zones	Serpukhovian 7.68±0.15 Myr duration.	S. Guizhou Prov. (Luokun section; 25°31' N, 106°57' E)	Limestone; 170 m; carbonate slope; Conodonts	(1) XRF-Ca; (2) MTM, Evol. spectra, band-pass filters adjusted to changing sed-accum-rate, tuning	405-kyr E (~3.5 m) for tuning (26 cycles; 10.5 Myr); also e, O, P. Modulations of 1.12 Myr (ecc.), 1.2 and 1.6 Myr (ecc.)	No
17	Gong, Y.M., et al., 2020, Gong, Dev. Plan. Change, 193: 103267.	Devonian (Famennian)	International geologic stage (Famennian); Conodont zones	Famennian 14.4 ± 0.28 Myr duration; scaling of 20 component conodont zones and 12 bio-environmental events and/or sequences.	Guangxi Prov. (Lali section, 50 km W of Yizhou City; 24°45' N, 108°29' E)	Limestone; 180 m studied; inter-platform trough to slope; Conodonts	(1) XRF-Ca; (2) MTM, Evol. spectra, band-pass filters adjusted to changing sed-accum-rate, tuning	405-kyr E (8 to 3.5 m) for tuning (35.5 cycles; 14.4 Myr); also e, O, P	No
18	Bai, S.L., 1995b, Internat. Geol. Rev., 37: 1109–1114	Devonian (Middle through Upper)	International geological stages (Famennian); Conodont zones	Emsian 4.4 Myr; Eifelian 3.2 Myr; Famennian 9.4 Myr; "monograph publications not available at this time"	Guangxi and Hunan Prov. (several sections; La-Long not given); "monograph publications not available at this time"	Limestone; Conodonts	(1) Geochemistry (Al, NiI, etc.)	100-kyr "e" ?	?
19	Gong, Y.M., et al., 2001, Palaeo-3, 168:237–248.	Devonian (Frasnian-Famennian)	Conodont zones	Upper rhenaia Zone 0.6 Myr; <i>Inguiformis</i> Zone 0.8 Myr; <i>triangularis</i> Zone 0.9 Myr (with 3 equal subzones).	Guizhou (Yangtze and Longjiang and Lijiang outcrops) and Guangxi (Duan and Clayey limestone; slope to basin; Conodonts)	(1) Bedding hierarchy; (2) Counting Bundles, superbundles, bundle sets, correlating among sections	Superbundle as E (ca. 3 m); also e, O, P	Detailed stratigraphic diagrams	?
20	Gong, Y.M., et al., 2005, Sci. China D Earth Sci., 48: 32–41.	Devonian (Frasnian-Famennian)	International geologic stage (Frasnian); Conodont zones	Frasnian Stage 4–3 Myr. Durations of 12 condont zones; <i>falsivalvis</i> 0.4 Myr, <i>translans</i> 0.4 Myr, <i>punctata</i> 0.4 Myr, Lower and upper <i>reissi</i> ; 0.3 and 0.4 Myr, <i>meriana</i> 0.2 Myr; Lower and upper <i>imgurimoris</i> 0.8 Myr. Lower-middle-upper <i>triangularis</i> 0.3 Myr each.	Guangxi (source publications not available at this time)	carbonate-basin and slope facies (source publications not available at this time); Conodonts	(1) Bedding hierarchy; (2) Counting Bundles, superbundles, bundle sets, correlating among sections	Superbundle as E; also e, O, P	?
21	Huang, C.-J., et al., 2016, GSA Ann. Mtg., 127-2.	Ordovician-Silurian (Kaitian-Aeronian)	International geologic stage (Hirnantian); Graptolite zones	Hirnantian Stage 1.46 Myr; Rhuddanian graptolite zones <i>A. ascensus</i> (1.1 Myr) and <i>P. acuminatus</i> (1.3 Myr),	Sichuan and W-Hubei prov. (Shuang, Qiliang, Wangjiawan sections; and 4 borehole (gs))	Calcareous dolostone; deep water; Graptolites	(1) XRF; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E for tuning (ca. 35 cycles; 14 Myr)	No
(1) Associated paleontology or age control; (2) Other studies; (3) Reviews; (4) Other information									
(2) Co-authors of this study; (3) Authors of publications cited in this paper; (4) Authors of publications cited in this paper									
(1) Frasnian Stage of 4.3 Myr is less than half of its ca. 9.5 Myr duration in GTS 2012									
(4) For this study, there was only an advance abstract with the indicated information. Not yet published as full paper.									

(continued on next page)

Table 1 (continued)

#	Author (* =Coresp. if different, full reference)	Geologic System (Stage)	What was calibrated	Result	Location	Main studied facies; thickness, depots, envirn.; Age control	(1) Proxies; (2) Method	Main cycle for tuning (implied span of record); Other recorded cycles	Raw data provided?
22	Liu, Y.B., Huang, Ordovician Paleo-3, 526; 96–109.	International geologic stage (Himanian); Graptolite zones	Himanian Stage 1.74±0.4 Myr; Late Kaitian graptolite zones <i>D. complanatus</i> (0.6 Myr), <i>D. complexus</i> (0.3 Myr), <i>P. pacificus</i> (1.8 Myr).	W Hubei Prov. (Yichang, borehole EH01; 30.10°N, 110.14°E)	Black siliceous claystone with carbonaceous layers and ash beds; <1 m; low-oxygen depression; Graptolites (projected).	(1) XRF (Fe signal) at 1-cm intervals; (2) MTM, Evol. spectra, band-pass filters.	405-kyr E (0.75 m) for tuning (12.5 cycles; 5.0 Myr); also e, O, P, obliquity-modulations of 1.2 Myr.	No	(1) Associated paleontology or age control; (2) Other studies; (3) Reviews;
23	Zhong, Y.Y., Wu, H.C.* et al., 2019, <i>Paleo-3</i> , 540; 109520.	Ordovician (upper Kaitian–Himanian)	International geologic stage (Himanian); Graptolite zones	Himanian Stage 1.26 Myr; Late Kaitian graptolite zones <i>D. complexus</i> (1.3 Myr), <i>P. pacificus</i> (1.8 Myr).	NE Yunnan Prov. (Wanmei roadcut; 27.76°N, 103.48°E)	Calcareous claystone; 48.5 m; deep-water Graptolites (projected)	(1) MS; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (3 m) for tuning (18 cycles; 7.3 Myr); also e, O, P, Modulations of 2.4 and 1.4 Myr (rec.)	No
24	Zhong, Y.Y., Wu, H.C.* et al., 2019, <i>Fan, J.-X.</i> , 93; 177–180.	Ordovician (upper Sandbian–lower Kaitian)	Regional conodont and equivalent graptolite zones	Graptolite zones O, calcareous through D, gravels span 2.4 Myr.	W Hubei Prov. (Yichang, Huanghuachang borehole YH-1 near Hirnantian GSSP; 30.8°N, 111.40°E)	Limestone; Pagoda Platform; Conodonts	(1) MS, GR, Oxy-48; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (2 m) for tuning (14.5 cycles; 5.98 Myr); also e, O, P	No
25	Ma, X.Y., Deng, S.H.* et al., 2019, <i>Paleoec-3</i> , In review.	Ordovician (Sandbian–lower Kaitian)	Delta 13C excursion (GICE)	Sandbian >6.4 Myr. Latest Sandbian–early Kaitian Positive 13-C excursion (GICE) spans ca. 4 Myr: Pagoda Lns Fm spans 5.5 to 6 Myr.	NE Sichuan Prov. (Qiaoling, Bazhou; 32.46°N, 108.89°E); SE Sichuan (Xikou, Huiyang; 30.16°N, 106.68°E); Chongqing (Sanqinjian Nanchuan; 29.14°N, 107.19°E); N Guizhou Prov. (Langcun, Xishui 28.46°N, 106.35°E)	Limestone; Pagoda Platform; Conodonts	(1) MS; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (ca. 2 m) for tuning (33 cycles lower Sanbian through Ion of Kaitian Pagoda Lns; 13.6 Myr); also e, O, P	?
26	Fang, Q., Wu, H.C.* et al., 2019, <i>Glob. Plan. Change</i> , 173; 96–108.	Ordovician (Darijawan, lower Sandbian)	Regional graptolite and equivalent conodont zones	Darijawan Graptolite zones <i>Pterograptus elegans</i> (1.04 Myr), <i>Dicyograptus murichensis</i> (1.22 Myr) and <i>Stengograptus vagus</i> (0.54–0.43 N, 79.43°E) and Yangtzeograptus (1.2 Myr).	Xinjiang Prov. (N. Tarim Basin, Kalbin, Dawanganu stream cut – Sanbian Aux. GSSP; 40.72°N, 79.54°E) and Yangtzeograptus (0.54–0.43 N, 79.17°E)	Claystone (Dawango), Siliceous limestone (Yangtze); 12 and 70 m; basin cut and slopes; Graptolites, Conodonts	(1) MS, GR; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (ca. 1 m) Dawanganu, ca. 3 m Yangtze); for tuning (12 and 11 cycles; 4.9 and 6.8 Myr); also e, O, P, Modulation of obliquity (1.2 Myr) and eccentricity (2.4 Myr).	No
27	Zhong, Y.Y., Wu, H.C.* et al., 2018, <i>Paleo-3</i> , 505; 86–99.	Ordovician (Darijawan, Dapingjian)	International geologic stages (Darijawan, Dapingjian); Graptolite zones	Darijawan 8.38±0.4 Myr; Dapingjian 1.97 ±0.7 Myr.	Zhejiang Prov. (Huangqitang GSSP; adjacent Changlin borehole CJ-3; 28.67°N, 118.19°E)	Black claystone, limestone; 48 m; deep-water slope; Graptolites	(1) MS; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (ca. 2 m) for tuning (29 cycles; 12.4 Myr); also e, O, P, Modulations of 1.2 and 2.4 Myr.	No
28	Ma, X.Y., et al., 2019, <i>Gong, Y.M.</i> , <i>Paleo-3</i> , 528; 272–287.	Ordovician (Fioian)	International geologic stage; Conodont zones	Folian Stage 7.08±0.4 Myr; Conodont zones <i>S. biobulus</i> 1.0 Myr, <i>S. exensis</i> 5.1 Myr, <i>O. communis</i> 2.0 Myr, <i>O. eviae</i> 0.8 Myr, and <i>B. triangularis</i> 0.2 Myr.	W Hubei Prov. (Yichang, Huanghuachang GSSP; For Dapingjian GSSP; 40.72°N, 110.27°E) and Hebei Prov. (Langjiaoshan section; 40.10°N, 119.59°E)	Limestones; 41 m and 112 m; both shallow shelf; Conodonts	(1) XRF-Ca and Fe/Ca; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (ca. 5 m) Langjiaoshan for tuning (75 and 185 cycles; 3.0 and 7.0 Myr); also e, O, P, Modulations of 1.0 and 1.9 Myr.	Suppl. Excel table MS 3913
29	Fang, J.C., Wu, H.C.* et al., 2019, <i>Paleo-3</i> , 540; 109530.	Cambrian (Driunian–Guzhangian)	International geologic stage; Conodont zones	Folian Stage 7.08±0.4 Myr; Conodont zones <i>S. biobulus</i> 1.0 Myr, <i>S. exensis</i> 5.1 Myr, <i>O. communis</i> 2.0 Myr, <i>O. eviae</i> 0.8 Myr, and <i>B. triangularis</i> 0.2 Myr.	Hunan Prov. (Guzhang, Liuyixi (former) Wangjian) residue = 109.97°E)	Cherty dolomite and limestone; 41 m and 112 m; both shallow shelf; Conodonts	(1) MS and 13-C; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (ca. 25 m) for tuning (35 cycles; 1.4 Myr); also e, O, P	Suppl. Excel table MS 340
30	Sui, Y., Huang, C.J.* et al., 2019, <i>Paleo-3</i> , 109273.	Ediacaran	Trilobite zones	Late Driunian trilobites zones <i>G. nathorstii</i> 0.52 Myr, <i>L. annula</i> 0.15 Myr; <i>B. guizhangense</i> 1.0 Myr–2.0 Myr.	W Hubei Prov. (6 km east of Three Gorges Dam, Juilongwan roadcut; 30.80°N, 111.06°E)	Cherty dolomite and limestone; then black shale; 85 m; mid-depth shelf; 3C excursions; U-Pb dates near bases and top of this section	(1) XRF-Fe; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (ca. 1.5 m) for tuning (33 cycles; 29.7 Myr); also e, O, P, Modulations of 1.2 (eccc.)	No
31	Gong, Z., et al., 2017, <i>Precamb. Res.</i> , 289; 62–74.	Cambrian (Driunian–Guzhangian)	Trilobite zones	Shuram (EN3) excursion spans ca. 20 Myr. Onset relative to U-Pb date (551.1 ±7 Ma) at top of Doushantou Fm is ca. 571.1±0.8 Ma. Span of entire Doushantou Fm projected as 64 Myr.	E Yunnan Prov. (Dongdahai upper Dushantou Fm)	Clavate limestone, then black shale; 45 m; shallow-water basin; 13C excursions	(1) MS, ARM; (2) MTM, Evol. spectra, band-pass filters	405-kyr E (ca. 4.3 m; 11 cycles; ca. 4.4 Myr); also e, O, P, Modulations of 1.2 (eccc.)	No
32	Sui, Y., Huang, C.J.* et al., 2018, <i>Sci. Bull.</i> , 63; 1485–1494.	Ediacaran	Delta 13C excursions	Cap carbonate span projected as 1.6 Myr; My based on projected sediment accumulation rate. Onset relative to U-Pb date (551.1 ±7 Ma) at top of Doushantou Fm is ca. 560 Ma.	W Hubei Prov. (lower Dushantou Fm)	Cap-carbonate, then organic-rich claystone and dolomite; 22 m; mid-depth shelf; Doushantou formation; 22 m; 13C excursions; U-Pb dates	(1) XRF-Fe/Ti and Ca; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (ca. 0.9 m) for tuning (27 cycles; 11.2 Myr); also e, O, P, Modulations of 1.2 (obt.)	No
33	Bao, X.J., Zhang, S.H.* et al., 2018, <i>EPSL</i> , 483; 52–63.	Cryogenian	Delta 13C excursions	Cap carbonate span projected as 1.6 Myr; First large scleracthys and multi-cellular animal embryos at 3.8 My after end of cap carbonate.	NE Guizhou Prov. (Songgao, borehole ZK1909; ca. 27.5°N 108°E?)	Clavate siltstone; 292 m; deep-water basin; U-Pb and Re-Os dates on termination of Sturtian on various continents	(1) MS; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (ca. 14 m) for tuning (21 cycles plus 10 e cycles; 9.8 Myr); also e, O, P	Suppl. PDF table (MS 28.765 ds)

(1) Associated paleontology or age control; (2) Other studies; (3) Reviews;

(1) Projected from Tang, P., et al., 2017, *J. Strat.*, 41; 119–133.(1) Chen, X., et al., 2011, *Earth Env. Sci. Trans. R. Soc. Earth Env. Micro. Spec.* 35; 13–20; and other publ.(1) Ma, X.Y., et al., 2016, *Earth Sci. Front.* 29; 285–286; Ma, X.Y., et al., 2018, *Acta Microscopica Sinica* 52; 1–14; (4) This is the upper part of the section; the lower part is unpubl. (1) Liu, J.P., et al., 2016, *Earth Sci. Front.* 26; 287–291; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) This is the upper part of the section; the lower part is unpubl. (1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

(1) Chen, X., et al., 2019; (2) Chen, X., et al., 2019; (3) Chen, X., et al., 2019; (4) GICE excursion is more complete in the southern sections on the Yangtze Platform.

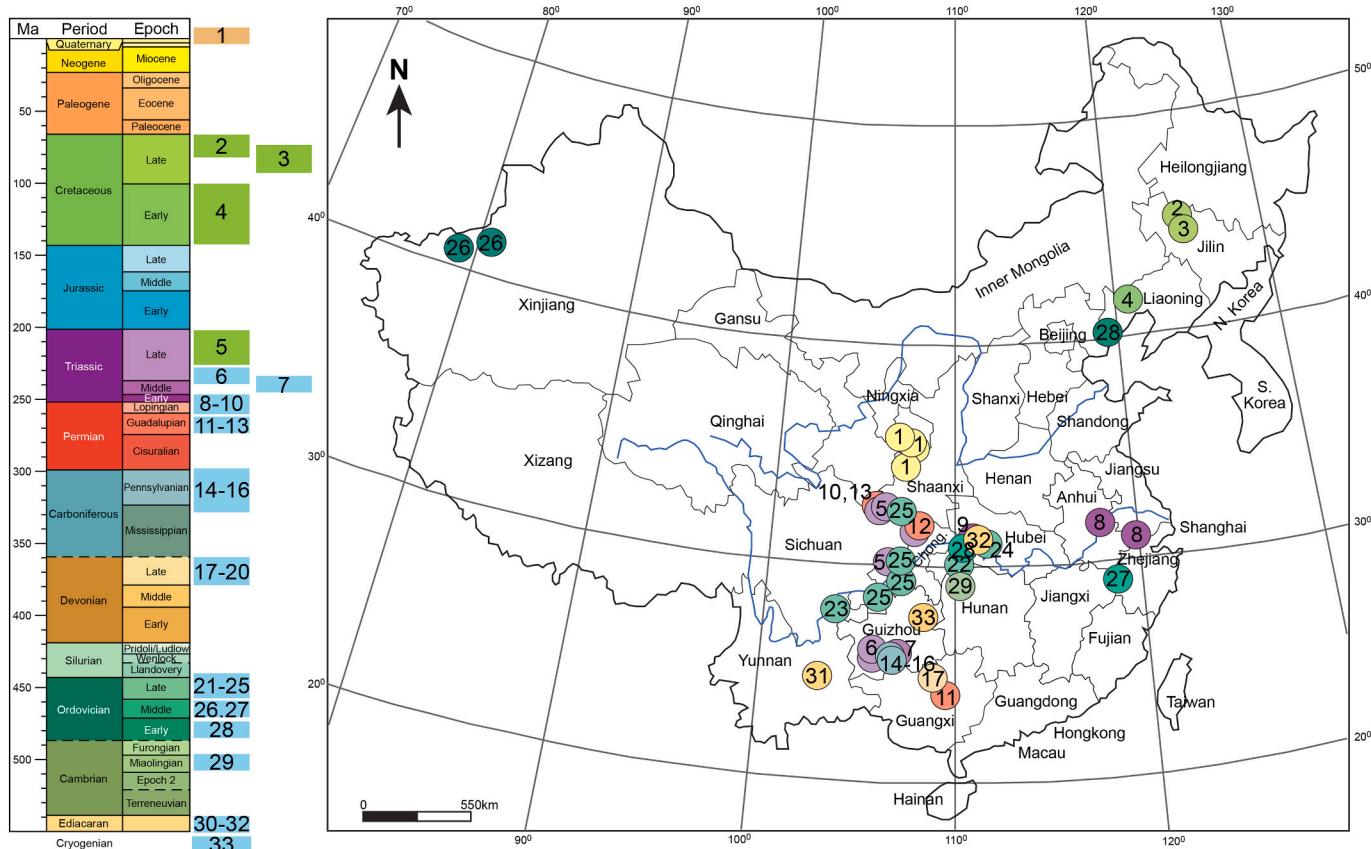


Fig. 1. Locations of cyclostratigraphy studies in China that are relevant to enhanced calibration of the geologic time scale. The color of each location dot is the same as the color of the geologic epoch. Coloring of the bars for the span of each study (next to the geologic time scale) indicates the general depositional setting (blue = marine, green = lacustrine, tan = other terrestrial). For details on each study, see the corresponding number in Table 1. (For interpretation of the references to color in this figure legend, the reader is referred to the web version of this article.)

isotopic dating has enabled the 405-kyr cycles within the Upper Cretaceous record to be anchored in absolute time by tying these to the corresponding cycles in the numerical solution of astronomical forcing (Laskar et al., 2011). This assigns an age of 82.9 Ma to the base of magnetic polarity Chron C33r, which is the candidate marker for the base of the Campanian Stage (Wu et al., 2020). Several of these cycle-scaled durations of stages and of other geologic episodes have been incorporated into the period-level syntheses of the integrated stratigraphy and timescale of China (right-hand column in Table 1) and into Neoproterozoic and Phanerozoic time scales (e.g., Ogg et al., 2016; and chapters within Gradstein et al., 2020).

4. Applications of cyclostratigraphy within China

The 20 cyclostratigraphy studies in this Special Issue (Table 2) span the Ediacaran through Pleistocene (Table 2). Of these, 6 are directly applicable to enhancing the geologic time scale, and therefore were also included in Table 1. The locations range from the Tibetan Plateau near the Himalayas to offshore Guangdong in the South China Sea (Fig. 2).

The majority of the 12 Mesozoic and Cenozoic studies in this Special Issue applied cyclostratigraphy to deposits in terrestrial basins, especially lacustrine and fluvial-delta sequences. Scientific questions include: establishing the relative timing of seismic horizons and the duration of oil- and gas-generating organic-rich source rocks (Du et al., 2020; Zhao et al., 2019; Xu et al., 2019; Peng et al., 2020; Liu et al., 2020; Chu et al., 2020; Zhang et al., 2019b), determining how glacial-interglacial oscillations influenced the influx of coarser-grained siliciclastics (Zhang et al., 2019a; Zhang et al., 2019a), and constraining the depositional history of deltas (Xu et al., 2020; Zhang et al., 2020). The

biostratigraphic constraints in many of these deposits are limited to poorly age-calibrated terrestrial faunal and floral remains. As such, some of the floating astronomically tuned time scales are anchored using the "traditional" ages based on assigning certain seismic or other horizons to geologic stage boundaries. However, with the aid of the constraints on durations and relative timing of events from cyclostratigraphy, the widespread non-marine Mesozoic-Cenozoic facies of China have the potential to be tied to contemporaneous marine rocks elsewhere (and vice versa), emphasizing the utility of cyclostratigraphy for building timescales to correlate and understand climate evolution in terrestrial environments. In the future, it may be possible to improve those age models through recognition of longer-period astronomical cycles (so-called "grand cycles").

The eight Ediacaran and Paleozoic cyclostratigraphy studies in this Special Issue are on marine successions composed of carbonate, shale or chert. Scientific questions range from durations of graptolite, conodont and radiolarian zones (Yao and Hinnov, 2019; Lu et al., 2019; Zhong et al., 2019a; Ma et al., 2019; Fang et al., 2020) to the durations of carbon-isotope excursions (Ordovician GICE by Ma et al., 2020b; Ediacaran Shuram CIE by Gong et al., 2019, and Sui et al., 2019). The two studies on the Shuram CIE yielded quite different durations, estimated at ca. 9 Myr by Gong et al. (2019) versus ca. 20 Myr by Sui et al. (2019) (note: the shorter estimate was based on extrapolating a sedimentation rate from only a portion of the entire CIE).

Table 2

Summary of cyclostratigraphy and astrochronology studies in this Special Issue of *Palaeogeography, Palaeoclimatology, Palaeoecology*. For abbreviations, see caption to Table 1. (Zhang et al., 2019b; Xu et al., 2020; Du et al., 2020; Zhao et al., 2019; Liu et al., 2017a,b; Xue et al., 2019; Peng et al., 2020; Qu et al., 2014; Chu et al., 2020; Zhang et al., 2019a,b; Liu et al., 2018; Zhang et al., 2020; Wu et al., 2017; Kamekata et al., 2009; Lu et al., 2019; Zhong et al., 2019; Tang et al., 2017; Ma et al., 2019; Fang et al., 2009; Condón et al., 2005; McFadden et al., 2019; Deng et al., 2018)

Table 2 -- Summary of cyclostratigraphy and astrochronology studies in this Special Issue of *Palaeogeography, Palaeoclimatology, Palaeoecology*

#	Author (* =Corresp.), publication	Geologic System (Stage) calibrated	What was calibrated	Result	Location	Main studied facies; depos. environ., thickness; Age control	(1) Proxies; (2) Method	Main cycle for tuning (implied span of record); Other recorded cycles	(1) Associated paleontology or age control; (2) Other information provided?
1	Rui Zhang et al., 2019b (Chunji Huang*), Palaeo-3, 530–78-89.	Quaternary (Pleistocene)	Fluvial-fan cycles	Interglacials (max of short-eccentricity cycles) are coarser-grained; interpreted as increased runoff pulses and more chemical weathering.	Xinjiang, western Tarim Basin (KT1 borehole, Kashgar region), 39°29'4"N, 76°53'9"E	Sand to clay, middle Kashgar fluvial fan, 800 m; ESR (electron spin resonance) dates and magnetostratigraphy	(1) GR, NS, Rb/Sr; (2) MTM spectral analysis, band-pass filters	100-kyr e (ca. 70 m) / 11 cycles; 1.13 Myr; O (ca. 30 m), P (ca. 14 m)	Suppl. Excel tables (9760 GR levels, 1600 MS levels, Geochem)
2	Yang Zhang et al., 2019 (James Ogg*), Palaeo-3, 527; 118-132.	Neogene-Quaternary (Pliocene-Pleistocene)	Continental slope depositional history	Coarser siliciclastic influxes at end of warm intervals (mid-Pliocene, early Galerian, early Calabrian, and at ca. 0.6 Ma).	North margin of South China Sea (SE of Hong Kong, IODP Hole 1409A), 18.409°N, 115.860°E	Bioclastic clay with clayey silt layers, 333 m; Magnetostriatigraphy (that, study and Plio-Plist polarity scale, calcareous nanofossils	(1) GR, Paleomag; (2) MTM spectra for sed-accum. rates, evolutive spectra, band-pass filters, tuning	405-kyr E (ca. 33 m); e (ca. 8 m), P (ca. 1.5 m)	Suppl. Excel tables (2700 GR levels, full Pmag 220 samples)
3	Yinchao Xu et al., 2020 (Zhongmin Li), Palaeo-3, 541; 109532.	Neogene-Quaternary (Pliocene-Pleistocene)	Delta, borehole	0 to 3.1 Ma chronology for Yangtze River delta location.	Jiangsu Prov. (north flank Yangtze River delta, core ZKA2); 32.563°N, 121.077°E	Alluvial-fluviatile sand-silt-clay, then upper intertidal silt-day; 295 m; Magnetostat (same paper)	(1) GR, MS; (2) MTM, spectral analysis, band-pass filters	405-kyr E (ca. 40 m) / 12 cycles; 4.7 Myr, then enhanced tuning with 100-kyr e (ca. 13 m); P (ca. 2.5 m)	Suppl. Excel tables (8680 GR levels, magnetic properties, etc.)
4	Wei Du et al., 2020 (Yoliang Du*), Palaeo-3, 554, 109803.	Paleogene (middle Oligocene)	Lacustrine history	Semi-deep lake and fans 33.1–31.6 Ma; shallow lake to 30.1 Ma, filling by fluvioplains to 25.5 Ma, plus ages for sand pulses (oil-gas reservoirs).	E. Olig., ca. 600 m. Hebei Prov., Bohai Bay Basin (Jizhong Depression, wells W1 and W62, 10 km E of Suning County), 38.42°N, 115.85°E	Oil shale (upper Mbr of Shahrui Fm, E. Olig., ca. 600 m). Mudstone to sandstone (Dongying Fm, Lt. Olig., ca. 800 m); Seismic correlation to wells with dated volcanics, megastratigraphic and ostracod-diage-pollen assemblages. Tuning tied to base-Neogene reflector	(1) GR; (2) MTM spectral analysis, Evol. spectra, band-pass filters, tuning	405-kyr E (ca. 70 m) / 24.5 cycles; 10.1 Myr; e (avg. ca. 20 m), O (avg. ca. 10 m), P (avg. ca. 5 m)	Suppl. Excel table (13000 GR levels well W33; 6000 well W62)
5	Ke Zhao et al., 2019 (Xuebin Du*), Palaeo-3, 528; 35-49.	Paleogene (middle Eocene)	Lacustrine laminae	Cycle-age control imply laminae are annual with Summer carbonate. Winter algal-organic clay.	east of Shandong Prov., eastern Bohai Bay Basin (NY borehole, Dongying Depression), 37.3°N, 118.4°E	Carbonate-day couplets, Shahrui Formation (upper Es4 and lower Es3 units), 350 m; Previous cyclostratigraphic study (Liu et al., 2017) on Shahrui Fm members, base ~540 Ma, top Es3/42.2 Ma	(1) GR for cycles; Geochim for laminae; (2) [Acyclic] MTM spectra, band-pass, tuning	405-kyr E (40 m, 8 cycles, ca. 3.2 Myr); e (8 m), P (2.2 m)	(1) Liu, Z.H., et al., 2017, <i>Palaeo-3</i> , 510–78-92.
6	Ke Xu et al., 2019 (Honghan Chen* & Hui Huang*), Palaeo-3, 532; 109253.	Paleogene (Paleocene-Oligocene)	Terrestrial basin history correlated to Pacific convergence rates	46-Myr cycle-scale for entire Paleogene and latest Cretaceous. Rifting episodes at 54, 50 and 33 Ma.	Henan-Hubei border, Nanjing Basin (central Biyang Depression, ca. 30 km N of Biyang, BS1 (deepest well in central China) and B270 well), ca. 32.6°N, 113.2°E	Paleocene floodplain of siliciclastics and mudstone. Eco-Olig faculstine (closed-lake) claystone, oil shale and dolomite, 6000 m and 2300 m; Green algae (Characeae), ostracoids, Spore-pollen. Astronomical tuning is relative to base-Neogene 23.03 Tuning	(1) GR for cycles; Geochim for laminae; (2) [Acyclic] MTM spectral analysis, evolutive spectra, eCO2 for sed-accum rates, peak-ratios, bandpass filtering, tuning	405-kyr E (avg. ca. 60 m); e (avg. ca. 15 m), P (avg. ca. 3 m); plus ca. 2 Myr. Modulation cycles in TOC of ca. 1.2 Myr.	Suppl. Excel tables (GR-32000 levels for each well, TOC, XRF)
7	Cheng Peng et al., 2020 (Changheun Zou*), Palaeo-3, 555; 109858.	Cretaceous (Cenomanian-Coniacian)	Fluvi-delta and lacustrine deposits	Quantuif Fm (5.5 Myr span has long-eccentricity-driven influxes of fluvial siltstone).	Heilongjiang Prov., Songliao Basin (SK-2 well, Anda City), ca. 46.5241°N, 125.3633°E	Mudstone to dolomite (fluvial-deltaic), Quantuif Fm, 860 m (studies), mudstone and oil shale (lacustrine, outer-shelf, 76 m; Plantkton, foraminifer zones, and carbon-isotope OAE2 excursion in the same formations).	(1) Resistivity imaging (redox log of HRMMS); Mud-logging; (2) [Acyclic] MTM spectral analysis, Evol. spectra, CO2 sed-accum. rate, band-pass filters, tuning	405-kyr E (ca. 14 cycles of 13 cycles of 22 m in Qingshankou Fm) implies ca. 5.5 Myr each; e, O, P	No
8	Yongqiang Li* et al., 2020, Palaeo-3, 558; 109465.	Cretaceous (Cenomanian-Turonian)	Marine carbonates.	Obliquity in this interval drove tropical-subtropical climate, in addition to at high-latitude.	southern Tibet (Gongzha section, ca. 40 km W of Tingri), 28.8°N, 86.6°E	Limestone and sandy limestone, outer-shelf, 76 m; Plantkton, foraminifer zones, and carbon-isotope OAE2 excursion in the same section.	(1) GR; litho-facies rank; (2) MTM spectral analysis, evolutive spectra, peak-ratio, band-pass filters, tuning	405-kyr E (avg. ca. 35 m) / 27.5 cycles total; 11.4 Myr; e (avg. ca. 8 m), O (avg. ca. 3 m), P (avg. ca. 1.5 m) -- Wavelengths shift among intervals	Suppl. Excel table (Rock Mag proxies; 696 levels)
9	Wei Liu et al., 2020 (Huichuan Wu*), Palaeo-3, 541, 109555.	Cretaceous (lower)	Fluvial-alluvial basin fill	Relative timing of alluvial-lacustrine dominance. Coal-rich horizons at 100-kyr intervals in upper part.	west Heliologiang Prov., Songliao Basin (50 km SE of Daqing, SS4 borehole), 46.3°N, 125.9°E	Mudstone to conglomerate with some coal horizons in upper part, Shahrui Fm, 836 m; Ar/Ar and K/Ar dating suggest Shahrui Fm spans 140 to 130 Ma	(1) GR; litho-facies rank; (2) MTM spectral analysis, evolutive spectra, peak-ratio, band-pass filters, tuning	405-kyr E (avg. ca. 35 m) / 27.5 cycles total; 11.4 Myr; e (avg. ca. 8 m), O (avg. ca. 3 m), P (avg. ca. 1.5 m) -- Wavelengths shift among intervals	Suppl. Excel table (1) Qiu et al., 2014, <i>Earth Sci. Front.</i> , 21: 234–250. (2) Analyses done in segments due to changes in sed-accum. rates
10	Runlian Chu et al., 2020 (Huachuan Wu* & Rukai Zhu*), Palaeo-3, 541, 109542.	Triassic (lower Ladinian)	Lacustrine organic-carbon-rich source rock	Chang 7 black shale spans ca. 3 Myr; Precession modulated the anoxic conditions.	Shaanxi Prov. Ordos Basin (Yaolu section outcrops, Yishu, Taogu, Tongjiahu district), 35.17°N, 108.87°E	Black shale in Chang 7 Mbr, middle Pb of 241.56 ± 0.09 Ma Base and 24.06 ± 0.12 Ma Top of black-shale in same section; plants, palynology, ostracods	(1) MS, GR, Geochem, TOC; (2) MTM spectral analysis, band-pass filters	100-kyr e (ca. 4 m) and 20-kyr e (ca. 0.8 m) semi-precession (0.1 kyr; 0.4 m)	Suppl. Excel table (1) Zhu, R.K., et al., 2019, <i>Acta Geol. Sinica</i> , 93:1823–1834. (2) But link is broken

(continued on next page)

Table 2 (continued)

#	Author (* =Corresp., publication	Geologic System (Stage)	What was calibrated	Result	Location	Main studied facies; depos. environ., thickness; Age control	(1) Proxies; (2) Method	Main cycle for tuning (implied span of record); Other recorded cycles	Raw data provided?	(1) Associated paleontology or age control; (2) Other information
11	Rui Zhang et al., 2018a (Zhijun Jin* & Quanyou Liu*), Paleao-3, 528, 87-98.	Triassic (lower Ladinian)	Lacustrine organic-carbon-rich source rock	Chang 7 Mbr spans ca. 5 Myr.; and 20.5-m oil shale is 1.7 Myr.	Shaanxi Prov. Ordos Basin (well Y101), Zhiidan County	Dark silty shale and oil shale with ash beds; Chang 7 Mbr; middle of Yanchang Fm; 65.9 m; Flora, spore-pollen assemblages, U-Pb dates from coeval intervals elsewhere	(1) MS; (2) [cycle, eCOCO] MTM spectral analysis, Ash-free MTM, ASM sed-accum rate, Evol. spectra, band-pass filters, visual tuning	405-kyr E (ave 5.4 m; 12.5 cycles, 5 Myr); e (avg. 1.6 m), O (0.5 m), P (avg. 0.27 m)	Suppl. Excel table (1256 MS levels)	(1) U-Pb; Liu, J., et al. 2018, Vert. Pal. Asia 66: 16-24; Deng, S.H., et al. 2018, Sci. China Earth Sci. 61: 1419-1439.
12	Tan Zhang et al., 2020 (Tailing Fan*), Paleao-3, 539, 109493.	Triassic (Olenevian)	Terrestrial fan-delta; 2 borehole wells	Baikouquan Fm. spans 2 Myr.	Xinjiang Prov., Junggar Basin (Huangyanguan fan on the slope of the Manu Sag at the northwestern margin), ca. 44.3°N, 85°E	Conglomerate to siltstone, Baikouquan Fm., 150–160m; fan-delta; Pollen - spore assemblages	(1) GR; (2) [cycle, eCOCO] MTM spectral analysis, Evol. spectra, band-pass filters, visual tuning	405-kyr E (ca. 33 m; 5 cycles, 2.0±0.1 Myr); e, O	Excel ?; Under embargo until 26Nov2020	
13	Xu Yao* and Linda Hinnov, 2019, Paleao-3,	Permian (Roadian-Capitanian)	Deep-marine chert	Base-Radiian to mid-Capitanian spans 5.4 Myr (Wardian 3.7 Myr). Radiian sed-accum-rate of ca. 4 m/Myr. New couplet-cycle analysis method.	Anhui Prov. (Chaohu City, Anmenkou section), 31.6°N, 117.8°E	Radiolian chert and mudstone alternations, outer-shelf Gufeng Fm., 28 m (19 m studied); ID-TIMS U/Pb 271.04±0.1 Ma at 3 m below base Gufeng Fm; condont base of assembly; then Radiolian assemblages.	(1) Chert-clay couplet coding; (2) MTM spectral analysis, Evol. spectra, band-pass filters, Oblique-Prec tuning, Oblique-Prec tuning, scaling to 270 Ma span, scaling to 270 Ma as base of Gufeng Fm	35-kyr Obliq (ca. 15 cm); P (ca. 9 cm), 405-kyr E (ca. 1.6 m)	Suppl. Excel table (175 thicknesses (267 ps))	(1) U-Pb; Wu, Q., et al. 2017, Palaeo 3, 466: 351-372; Radiolian; Kamekita et al. 2009, Island Arc, 18: 108-125.
14	Yangbo Lu et al., 2019 (Chunju Huang* and Shu Jiang*), Paleao-3, 526; 96-109.	Ordovician (u. Kalian, Hirnantian)	Marine shale, Graptolite zones and glacial sea-level curve	Lt. Kalian; <i>Di. complanatus</i> (upper), <i>D. complexus</i> , <i>Para. pacificus</i> ; = 0.6, 0.8, 1.8 Myr; Hirnantian (1.74±0.4 Myr total); Nor. <i>straordinarius</i> Per. persculptus = 1.2, 0.5 Myr ±0.1 hr	west Huber Prov. (EH1 borehole, Yichang); 30.102°N, 110.337°E (120 km NW of Hinnantian GSSP)	Marine black siliceous shale with ash beds; Wutong Formation and the lowermost Longmian formations; 9.1 m; Correlation of lithologic units to Wangliawan section outcrop with graptolite zones	(1) XRF (Fe signal) at 1-405-kyr E (avg. 0.75 m) for tuning (12.5 cycles; 5.0 Myr); e (avg. 0.5 m), O (avg. 0.4 cm); P (avg. 4 cm); obliquity-modulations of ca. 1.2 Myr (2.2 m);	405-kyr E (avg. 3 m) for tuning (1 cycles, 7.37 Myr); e (avg. 0.9 m), O (avg. 0.15 m), P (avg. 0.15 m); Modulations of 1.3 Myr (obi.), and 2.4 and 1.4 Myr (ecc.)	No	
15	Yanggang Zhong et al., 2019b (Huai-chun Wu*), Paleao-3, 540; 109520.	Ordovician (u. Kalian-Hirnantian)	Marine shale. International geologic stage (Hirnantian); Graptolite zones	Hirnantian Stage 1.26 myr; Late Kalian graptolite zones: <i>D. complexus</i> (1.3 myr), <i>P. pacificus</i> (1.8 myr). Length of day = 22.4 ±0.1 hr	NE Yunnan Prov. (Wanhe roadcut; 27.76°N, 103.48°E)	Calcareous claystone; deep-water; Calcarenous claystone; deep-water; 48.5 m; Graptolites (projected)	(1) MS; (2) MTM spectral analysis, Evol. spectra, band-pass filters, tuning	405-kyr E (avg. 3 m) for tuning (1 cycles, 7.37 Myr); e (avg. 0.9 m), O (avg. 0.15 m), P (avg. 0.15 m); Modulations of 1.3 Myr (obi.), and 2.4 and 1.4 Myr (ecc.)	Suppl. Excel table (267 ps)	(1) Graptolite zones projected from Tang, R., et al., 2017, J. Strat., 41, 119-133.
16	Xueying Ma, et al., 2020 (Shenghui Deng), Paleao-3, in review.	Ordovician (Sandbian-lower Kalian)	Marine carbonate. Delta 13C excursion (GICE)	Sandbian > 6.4 Myr. Latest Sandbian-early Kalian positive 13-C excursion ("GICE") spans ca. 4 Myr. Pagoda Lms Fm spans 5.5 to 6 Myr.	NE Sichuan Prov. (Qiaoling, Bazhou); 32.46°N, 106.89°E; SE Sichuan (Xikou, Huaying, 30.16°N, 106.69°E); Chongqing (Sanqianzhen Nanchuan, 29.14°N, 107.19°E); N Guizhou Prov. (Langcun, Xishui 28.46°N, 106.35°E)	Limestone; Pagoda Limestone, 32.2 m; Yangtze Platform; Conodonts	(1) MS; (2) MTM, Evol. spectra, band-pass filters, tuning	405-kyr E (ca. 2 m) for tuning (33 cycles lower Sanbian through top of e. Kaitan Pagoda Lms, 13.6 Myr); e, O, P	?	(1) Ma, X.Y., et al. 2019a, J. Strat., 43: 1-10; Wang, Z.H., et al. 2018, Acta Micro. Sinica, 35: 13-29; and other publ.; (2) GICE excursion is more complete in the southern sections on the Yangtze Platform.
17	Kunyuan Ma et al., 2019 (Yiming Gong*), Paleao-3, 528; 272-287.	Ordovician (Fioian)	Marine carbonate. International geologic stage; Conodont zones	Fioian 7.08±0.4 Myr. Conodont zones: <i>S. biholauis</i> 1.0 Myr. <i>S. extensus</i> 5.1 Myr. <i>O. communis</i> 1.0.37°E) and Hebei Prov. (Langjiaoshan section; 40.1°N, 119.59°E)	W Hubei Prov. (Yichang, Huanghuachang GSSP) For Dapingshan 30.86°N, 110.37°E) and Hebei Prov. (Langjiaoshan section; 40.1°N, 119.59°E)	Limestones; 41 m and 112 m, both shallow shelf; Conodonts	(1) XRF-Ca and Fe/Ca, (2) MTM spectral analysis, Evol. spectra, band-pass filters, tuning	405-kyr E (ca. 5 m) for tuning (33 cycles lower Sanbian through top of e. Kaitan Pagoda Lms, 13.6 Myr); e, O, P	?	(1) Langjiaoshan section -- conodont study is part of the cyclostratigraphic paper, (2) GICE in the southern sections on the Yangtze Platform.
18	Jichuang Fang et al., 2020 (Huai-chun Wu*), Paleao-3, 540; 109530.	Cambrian (Dumian-Guzhangjian)	Marine carbonate. Trilobite zones	Fioian 7.08±0.4 Myr. Conodont zones: <i>S. biholauis</i> 1.0 Myr. <i>S. extensus</i> 5.1 Myr. <i>O. communis</i> 0.2 Myr. <i>O. evae</i> 0.8 Myr. and <i>B. triangularis</i> 0.2 Myr.	Hunan Prov. (Guizhang, Luoyix (formerly Wanggeun) roadcut = Guizhangjian L., 28°28'28"E, 109.97°E)	Late Dumian trilobites zones; G. Mr. base; Guizhangjian L., laevigata ~0.7 Myr. Length of day = 21.6±0.2 hr.	(1) MS, 13C/MTM spectral analysis, Evol. tuning	405-kyr E (ca. 25 m) for tuning (35 cycles, 1.4 Myr); e (avg. 6.6 m), O (avg. 1.96 m), P (avg. 1.13 m)	Suppl. Excel table (MS 4340 pfs, Delta13C)	(1) Peng, S.C., et al. 2009, Episodes 32: 41-55.
19	Zheng Gong* et al., 2019, Paleao-3, 528; 232-246.	Ediacaran	Marine carbonate.	Shuram CIE onset at ca. 560 Ma (ca. 20 Myr after Gaskiers Glaciation), and spanned ca. 9 Mr. Magneostrat had been overprinted in Late Triassic.	Guizhou Prov. (Huanglianba section, Songtao county); 29.191°N, 109.266°E	Dolomite and black shale; margin-slope; Doushantuo Fm., 75 m (calystrobat on 4-m carbonate interval at 53-57 m); U-Pb date of 551.1±0.7 Ma just below top of same formation elsewhere	(1) MS; (2) MTM spectral analysis, ASM for optimal sed-accum rate; ratio fit	95-kyr e (0.8 m); 4-m section spans ca. 0.5 Myr; O (30 cm); P (18 and 14 cm); Sub-Mankovitch of ca. 6 kyr	Suppl. Excel table (200 MS levels; Pmag; China Sci., 13C, etc.)	(1) U-Pb; Condon, et al., 2005, China Sci., 308: 95-98.
20	Yu Sui et al., 2019 (Chunjie Huang*), Paleao-3, 532; 109273.	Ediacaran	Marine Carbonate, Shuram/Wonoka negative 13C	Shuram CIE onset at ca. 571 Ma, and spanned ca. 20 Myr. Mr. Magneostrat	W Hubei Prov. (5 km east of Three Gorges Dam, JuJiugong roadcut; 30.80°N, 110.6°E)	Dolomite and black shale; margin-slope; Doushantuo Fm., 154 m (upper 65 m in this study); U-Pb dates near base and top of this section; 13C excursions; U-Pb dates near base and top of this section.	(1) Fe, Si and Ca in XRF (4226 levels); (2) MTM spectral analysis, Evol. spectra, band-pass filters, tuning	405-kyr E (avg. 1.6 m) for tuning (73 cycles, 29.7 Myr); e (avg. 0.5 m), O (avg. 0.12 m), P (avg. 0.06 m); Modulations of eccentricity (ca. 2.2 Myr)	No	(1) McFadden et al., 2008, PNAS, 105: 3197-3203; Condon et al., 2005, Science, 308: 95-98; (2) The analysis was done in 4 intervals due sediment-accumulation rate changes

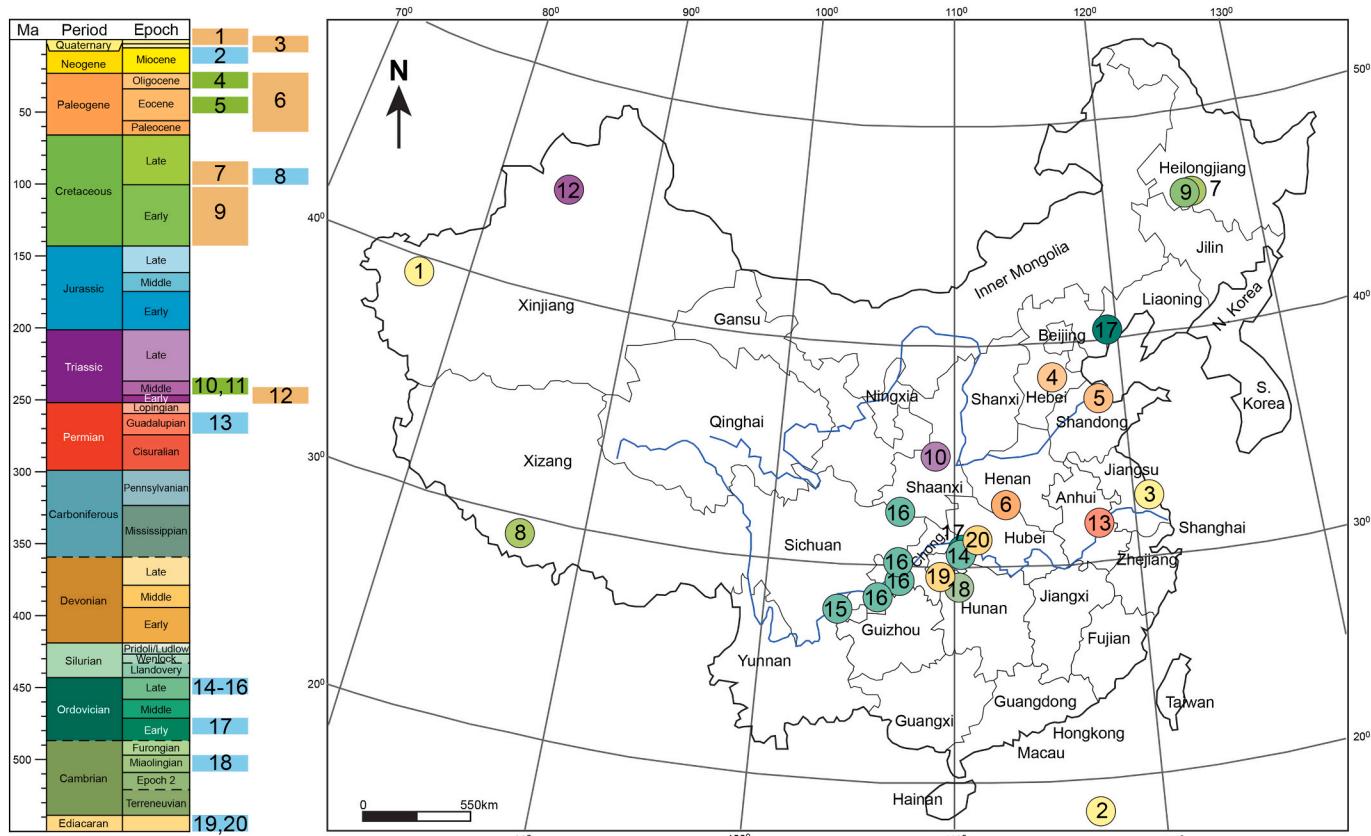


Fig. 2. Location of cyclostratigraphy studies in China in this Special Issue of *Palaeogeography, Palaeoclimatology, Palaeoecology*. The color of each location dot is the same as the color of the geologic epoch. Coloring of the bars for the span of each study (next to the geologic time scale) indicates the general depositional setting (blue = marine, green = lacustrine, tan = other terrestrial). For details on each study, see the corresponding number in Table 2. (For interpretation of the references to color in this figure legend, the reader is referred to the web version of this article.)

5. General methods

5.1. Paleoclimatic proxies

Cyclostratigraphy is dependent on recognizing orbital cycles in climatically sensitive physical, chemical or biological proxies. In order to understand how paleoclimatic proxies can record astronomical signals, cyclostratigraphers need to know how the paleoclimate changed in response to astronomically forced insolation variations, how these paleoclimatic variations influenced climate-sensitive components of the sedimentary record, and how the recording of these climate-sensitive components in the resulting deposits are linked to different proxies. Due to the Earth's rotation and its orbit around the Sun, the solar irradiance has strong daily, seasonal and annual cycles. Non-linear mechanisms rectify the seasonal modulation, thereby enhancing precessional-period and other Milankovitch cycle variability (Huybers and Wunsch, 2003). This seasonal-scale climatic transformation of the modulation of insolation forcing can amplify the astronomical frequencies and subsequently be preserved in stratigraphic records (Hinnov, 2018).

Paleoclimate proxies of widespread use in cyclostratigraphy include color, lithology, stable isotopes, rock magnetism, paleontology, gamma ray logs, elemental geochemistry, and organic carbon content (e.g., Hinnov, 2013; Li et al., 2019a). In this Special Issue, many types of paleoclimatic proxies were measured as a basis for analysis by spectral methods. Physical proxies include lithologic characteristics (i.e., chert-clay couplet coding, bed bundling, color), borehole well logs (i.e., resistivity, and natural gamma ray (GR) of Th, K and U contents), and magnetic parameters (i.e., magnetic susceptibility (MS), anhysteretic remanent magnetization (ARM), and hard Isothermal remanent magnetization/magnetic susceptibility (HIRM/χ)). Geochemical proxies

include total organic carbon (TOC), carbon isotopes ($\delta^{13}\text{C}$), and XRF scanning data of elemental concentrations and ratios (Fe, Ca, Si, Ti, Rb/Sr, Sr/Al, Zr/Al, V/Cr, Ni/Ti, Cu/Ti, Fe/Ca).

In terrestrial or lacustrine settings, variations in some of these paleoclimate proxies are responses to astronomical-forced insolation changes in weathering, precipitation and temperature. In marine depositional settings, some of these paleoclimate proxies are records of variations in the relative productivity of carbonate organisms and variable fluxes of sediments derived from continental weathering and runoff (Li et al., 2019a).

5.2. Spectral analysis

Cyclostratigraphy requires the detection and separation of astronomical signals in paleoclimate proxy data (Hinnov and Hilgen, 2012). The most common procedures involve Fourier analysis of the proxy data to establish the frequency distribution of variance. High variance within a relatively narrow band of frequencies (i.e. spectral peaks) are used to identify likely cyclic signals within the dataset that represent the astronomical forcing. These spectral analysis methods are designed primarily to isolate signals in evenly-spaced time series. In contrast, proxy data measured through a stratigraphic record are in height/depth. Therefore, depending upon the type of deposit, well-defined, rapidly deposited event beds such as debris flows or turbidites are removed from the records. Methods such as evolutive harmonic analysis (EHA) and correlation coefficient (COCO) analysis (Meyers et al., 2001; Li et al., 2018b, 2019b) can be performed to identify temporal changes in the potential dominant cyclic frequencies preserved in stratigraphic sections, to reconstruct variations in sedimentation rates, and to select a suite of subintervals to analyze separately using spectral analysis.

Associated methods of time-series analysis of paleoclimate proxies employed in cyclostratigraphy include interpolation, integral-sampling, smoothing, detrending, filtering, and demodulation, correlation and tuning (Hinnov and Ogg, 2007). Common software for such signal processing includes *KaleidaGraph* 4.0, *Analyseseries* 2.0.8, and *Acycle* 2.0. In this Special Issue, most studies utilize the multi-taper method (MTM) of spectral analysis, which offers high-resolution and statistical estimates that are independent of spectral power (Ghil et al., 2002). The MTM method can provide a well-resolved spectrum of variance within a dataset, allowing clear distinction of peaks even with small amplitude oscillations. The results are often more optimized than those based on other, more-classical spectral methods (Thomson, 1982). The statistical significance of spectral peaks, and thus whether they are likely to have been caused by astronomical cycles or perhaps relate to some random process instead, can be determined by establishing by how much the peak exceeds the expected background red-noise spectrum for an absence of cycles (e.g., Mann and Lees, 1996).

Most studies in this Special Issue also apply evolutionary spectral methods in order to visualize the time-frequency landscapes of the stratigraphic signals (e.g., Sui et al., 2019; Xu et al., 2019; Zhang et al., 2019b). Such analyses can provide important information that guide recognition of Milankovitch cycles through a section, as well as track variable sedimentation rates and the presence of hiatuses.

5.3. Astronomical tuning and construction of floating astronomical time scales

The main components of astronomical parameters are Earth's orbital cycles of precession, obliquity and eccentricity and their long-term modulation cycles. As noted above, Laskar et al. (1993, 2004, 2011) has progressively developed enhanced astronomical solutions (called La1993, La 2004, and La 2011, respectively) that now span from 250 Myr in the past to 250 Myr in the future. These numerical solutions provide an important astronomical time scale target and comparator to cyclostratigraphic results. Tuning of cyclostratigraphically identified cycles in proxy data to the astronomical solution allows for the construction of calibrated timescales anchored in absolute time. However, even the most recent solution of La2011 becomes uncertain for the high-frequency components of precession and obliquity beyond 50 Myr into the past, making any time-calibration of paleoclimatic proxy data to the Astronomical Time Scale difficult. Many stratigraphic records of paleoclimatic proxy series (e.g., Mediterranean sapropel sedimentary records, marine oxygen isotope and carbon isotope data, natural gamma ray (GR) logs, magnetic susceptibility, core-scanning XRF data, etc.) in the Cenozoic can be directly tuned to astronomical targets such as 65°N summer insolation, obliquity and eccentricity solutions (Ding et al., 2002; Lourens et al., 2004; Hilgen et al., 2012; Lisicki and Raymo, 2005; Zachos et al., 2001).

Looking "backwards" in geologic time, Earth's orbital precession and obliquity periods become shorter relative to the Present due to a reduction of the Earth-Moon distance and tidal dissipation, which caused changes in Earth's rotation rate and shape (Laskar et al., 2004). Therefore, for cyclostratigraphic analysis of deposits older than 50 Ma, the stable 405-kyr long-eccentricity cycles have emerged as the most prominent and stable orbital oscillation (Hinnov, 2018). This 405-kyr long-eccentricity cycle is caused by the combined interactions of the large mass of Jupiter and the perihelia of Venus and Jupiter on the Earth's orbit. Therefore, most cyclostratigraphy studies of Mesozoic and Paleozoic sediments tune their paleoclimatic proxy series to the 405-kyr long-eccentricity cycle (Bao et al., 2018; Boulila et al., 2014; De Vleeschouwer et al., 2015; Fang et al., 2018; Grippo et al., 2004; Huang et al., 2010a, 2010b; Husson et al., 2011; Li et al., 2016a, 2016b, 2018a; Pas et al., 2018; Ruhl et al., 2016; Wu et al., 2013c).

The predominant frequencies in rhythmic sedimentary stratigraphic records are observed from the power spectra and evolutionary spectral analysis results of paleoclimate proxies. Based on the estimates of

variable sedimentation rates from methods such as COCO and EHA analyses, it is possible to determine whether the 405-kyr long-eccentricity, ~100-kyr short-eccentricity, or obliquity cycles are the dominant orbital parameter in high-resolution paleoclimatic proxy series. These cycles become the tuning term to convert the depth-domain signals into the time-domain, thereby enabling the construction of an astronomically tuned time scale. If the stratigraphic record of interest has a calibration based on a radio-isotopically dated horizon, or if the tuned-record of magnetic polarity zones can be matched to the calibrated geomagnetic polarity time scales (or similar types of age-control), then it is possible to apply cyclostratigraphy to generate a precise age model at every level within a stratigraphic succession.

In this Special Issue, nearly every study resolves a dominant 405-kyr long-eccentricity cycle and utilizes this term to construct floating astronomically tuned time scales. However, in some cases where multiple Milankovitch signals were identified, the tuning made use of the ~100-kyr short eccentricity cycle or obliquity cycle. This was the case for Quaternary fluvial-lake deposits (Zhang et al., 2019b), Cretaceous outer-shelf marine carbonates (Li et al., 2020), Middle Triassic lacustrine deposits (Chu et al., 2020), middle Permian deep-marine chert-shale couplets (Yao and Hinnov, 2019), and an Ediacaran marine carbonate succession (Gong et al., 2019).

5.4. Some caveats and challenges

Even though astronomically induced climate signals can be recognized in nearly every depositional setting in terrestrial and marine environments, the preservation and fidelity of that signal in the geologic record are influenced by many other natural processes. Strata do not accumulate steadily, and interruptions or disruptions in sediment accumulation can result in "missing cycles", time-averaging (e.g. bioturbation), and unstable sediment accumulation rates. Tectonics can cause fault-displaced records within an outcrop or borehole, and distortion of some types of paleoclimate proxies can occur through redox processes or diagenesis. In dynamic settings such as terrestrial and shallow-marine environments, non-cyclic fluvial-delta lobe switching and facies migration can occur, obfuscating astronomical signals. Most spectral analysis methods and bandpass separation of frequencies implicitly assume a semi-stable and continuous signal in the proxy being analyzed.

As a consequence of these issues, it is important to verify the interpretations from spectral-analysis, bandpass and tuning methods. If the methods are applied to more than one type of proxy, does one obtain the same results? Are each of the assigned cycles visibly present in the proxy records, and ideally also seen as features in the outcrop? Can the same cycles be resolved in an independent coeval section from another part of the basin? These types of verification and testing procedures have enabled development of a reliable astronomically tuned time scale for the Cenozoic and portions of the Mesozoic-Paleozoic. However, it is common for a study to concentrate on only a single reference section (e.g., over a third of the studies summarized in Tables 1 and 2), and these studies therefore await future verification. Confirmation of features in multiple sections is an issue that applies not only to cyclostratigraphic analyses, but also to many studies of geochemical excursions, magnetic polarity patterns, and impact horizons, more than a few of which were initially reported only from a single reference section.

It is also important for all studies to publish/archive their raw data and exact location information, so that future techniques can be re-applied to these datasets. In compiling our summaries, it was disturbing how many important studies, both legacy and current, did not provide the supporting raw data in an accessible form or indicate the exact geographic location of the sections with precise Google-Earth-verified latitude-longitude coordinates.

5.5. Outlook

A recent international collaborative effort within the cyclostratigraphic community (The Cyclostratigraphy Intercomparison Project, CIP) has sought to lay a foundation for standardization, both in terms of approach and reporting (Sinnesael et al., 2019). The CIP workshops explored how different approaches adopted for cyclostratigraphic analysis by the community can lead to different results, and thus this community-led approach has the specific aim of developing best practice protocols. A key outcome of the CIP work was the finding that cyclostratigraphy is a trainable skill. The rapid emergence of cyclostratigraphy and timescale development in China speaks to the importance of unifying the discipline in a way that ensures optimal, reproducible results that will add data and value to the geologic time scale.

Of particular note is the recent development of new software tools; in particular the *Astrochron* package for R (Meyers, 2014) and the *Acycle* code written for Matlab (Li et al., 2019b). Both packages provide a complete set of ready-to-use tools for cyclostratigraphic analysis of proxy data. The use of these tools helps to ensure reproducibility, eases communication of results, and also guides new researchers in cyclostratigraphy. We note that all but 3 of the papers in this Special Issue utilized either the *Acycle* or the *Astrochron* package. This emphasizes the clear demand and acceptance for these tools. Moreover, 12 studies utilize the COCO correlation coefficient method for constraining sedimentation rates recently introduced by Li et al. (2018b), which is integrated into the *Acycle* code. Although all workers must be wary of adopting 'black box' approaches in the statistical treatment and analysis of data (i.e., where the tools are used but not fully understood), the provision of these tools to the community means that integration and understanding of results should ultimately improve.

6. Future applications

In the near future, it should be possible to compile a global composite of Paleozoic-Mesozoic reference sections for a complete record of the 405-kyr "metronome" relative to the Present that encompasses all major biozones and is constrained by radioisotopic dates. This will extend the current Cenozoic astronomically tuned scale back to the Ediacaran. Such a preliminary "proof-of-concept" synthesis with standardized statistical processing of numerous overlapping sections was accomplished for the entire Mesozoic by Huang (2018), and the major 405-kyr "cycloths" in the late Carboniferous through early Permian await a direct calibration to an enhanced astronomical solution.

There is also the exciting possibility of augmenting these compilations with the so-called "grand cycles" that modulate the amplitudes of the 405-kyr long-eccentricity cycles (ca. 1 to 2 Myr periodicities) and of obliquity (ca. 1.2 Myr periodicity). Such very-long-term cycles can act as additional time-scale and correlation constraints. It has already been shown that these modulations appear to influence global Phanerozoic sea level (e.g., Boulila et al., 2018) and govern the relative importance of obliquity versus eccentricity-precession cycles within basins (e.g., Li et al., 2016b). Some of these very-long-period modulations have been observed in longer sedimentary records within China (see the column with "other recorded cycles" in Tables 1 and 2). With constraints from radio-isotopic dates, it may be possible to bridge current gaps in the Mesozoic astronomically tuned time scale to begin a direct calibration of portions of Paleozoic time.

In turn, a verified master 405-kyr metronome may help to resolve the existing systematic "external uncertainty" of about 0.3 Myr or greater on most pre-Cenozoic radioisotopic dates. The analytical precision on such dates is often better than 0.1 Myr, but the added external uncertainty is a result of uncertainties in decay constants and of currently used "standards".

Looking further back in time, there is the challenge of analyzing pre-Cryogenian cyclostratigraphic successions. China hosts thick

Mesoproterozoic deposits that appear to have cyclic characteristics; but studies are inhibited by the current lack of horizons that can be dated by radioisotopic methods to constrain sediment accumulation rates and the true nature of the cycles. This will probably be the next major frontier in China-hosted cyclostratigraphic studies, and it opens the possibility of understanding climate change in the pre-Cryogenian world.

Declaration of Competing Interest

None.

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