# Bacterial Oxidation of Dibromomethane and Methyl Bromide in Natural Waters and Enrichment Cultures

K. D. GOODWIN,\* J. K. SCHAEFER, AND R. S. OREMLAND

United States Geological Survey, Menlo Park, California 94025

Received 29 May 1998/Accepted 4 September 1998

Bacterial oxidation of <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> and <sup>14</sup>CH<sub>3</sub>Br was measured in freshwater, estuarine, seawater, and hypersalinealkaline samples. In general, bacteria from the various sites oxidized similar amounts of <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> and comparatively less <sup>14</sup>CH<sub>3</sub>Br. Bacterial oxidation of <sup>14</sup>CH<sub>3</sub>Br was rapid in freshwater samples compared to bacterial oxidation of <sup>14</sup>CH<sub>3</sub>Br in more saline waters. Freshwater was also the only site in which methyl fluoride-sensitive bacteria (e.g., methanotrophs or nitrifiers) governed brominated methane oxidation. Halflife calculations indicated that bacterial oxidation of CH<sub>2</sub>Br<sub>2</sub> was potentially significant in all of the waters tested. In contrast, only in freshwater was bacterial oxidation of CH<sub>3</sub>Br as fast as chemical removal. The values calculated for more saline sites suggested that bacterial oxidation of CH<sub>3</sub>Br was relatively slow compared to chemical and physical loss mechanisms. However, enrichment cultures demonstrated that bacteria in seawater can rapidly oxidize brominated methanes. Two distinct cultures of nonmethanotrophic methylotrophs were recovered; one of these cultures was able to utilize CH<sub>2</sub>Br<sub>2</sub> as a sole carbon source, and the other was able to utilize CH<sub>3</sub>Br as a sole carbon source.

The brominated methanes, methyl bromide  $(CH_3Br)$  and dibromomethane  $(CH_2Br_2)$ , deliver bromine to the atmosphere from marine and terrestrial sources. Bromine is a catalyst for stratospheric ozone destruction, and ozone depletion could have notable consequences for biology and the climate (44). The potential of  $CH_3Br$  to deplete stratospheric ozone has led to restrictions on the use of this compound as a fumigant and to its scheduled phase-out (44). The atmospheric lifetime of  $CH_2Br_2$ (0.29 year) (45) is shorter than that of  $CH_3Br$  (~0.8 year) (38), making it less likely to deliver bromine to the stratosphere. However, shorter-lived compounds can reach altitudes higher than might be expected from their atmospheric lifetimes (20). Indeed,  $CH_2Br_2$  has been detected near the tropical tropopause (36) and in the lower stratosphere (21), indicating that it could participate in depleting stratospheric ozone.

The atmospheric burden of brominated methanes is controlled by a balance of sources and sinks. The sources of brominated methanes are both natural (10, 41) and anthropogenic (27, 44), and the sinks are both chemical (16, 19) and biotic (5, 16)9, 19). Recognition of a bacterial sink in soil led to reduced estimates of the CH<sub>3</sub>Br atmospheric lifetime and ozone depletion potential (38). The potential importance of bacterial sinks in bromine biogeochemistry has motivated recent investigations of CH<sub>2</sub>Br<sub>2</sub> and CH<sub>3</sub>Br biodegradation in saline (9, 19, 40) and hypersaline (5) environments, although investigations of estuarine and freshwater sites still appear to be lacking. Brominated methane degradation has also been reported for pure cultures of methanotrophs (9, 31), ammonia oxidizers (18), and methylotrophs (4, 37). The mechanisms, microbial ecology, and environmental impact of bacterial sinks have yet to be fully assessed.

The pattern of brominated methane degradation may provide insight into the mechanisms behind the process. Certain enzymes, such as the methane monooxygenase, can dehalogenate both  $CH_2Br_2$  and  $CH_3Br$  (1), while other enzymes dehal-

ogenate only dihalomethanes (22). In this study we directly compared bacterial oxidation of  $^{14}\text{CH}_2\text{Br}_2$  and  $^{14}\text{CH}_3\text{Br}$  in samples from a variety of aquatic environments and used an inhibitor, methyl fluoride (CH<sub>3</sub>F) (30), to determine whether bacteria, such as methane or ammonia oxidizers, play a role in brominated methane oxidation in natural waters. Measurements of oxidized products provided the first direct evidence that bacterial oxidation of brominated methanes occurs in fresh, estuarine, and marine waters. The environmental implications of bacterial oxidation of CH<sub>2</sub>Br<sub>2</sub> and CH<sub>3</sub>Br were explored by comparing rate constants to the rate constants for chemical consumption and volatilization.

#### MATERIALS AND METHODS

Field samples. Waters having different salinities and pH values were collected from a number of sites in central and northern California (Table 1). Surface water samples were collected by hand; the samples of Mono Lake water were collected at a depth of either 1 or 5 m and the samples of San Francisco Bay water were collected at a depth of 1 m with a Nisken bottle. Samples were collected in 1996, 1997, and 1998 during all seasons. Most of the San Francisco Bay water samples were collected from San Pablo Bay; the only exceptions were the samples obtained during January 1997 and April 1998, which were collected from central San Francisco Bay due to low salinity caused by heavy rain. The water was filtered on site to remove organisms that were larger than 80  $\mu$ m (larger than 120  $\mu$ m for Mono Lake water) and were stored in polyethylene bottles at 4°C for several hours or up to 1 week.

**Radiolabeling experiments.** Oxidation of  ${}^{14}\text{CH}_2\text{Br}_2$  and  ${}^{14}\text{CH}_3\text{Br}$  was measured by using a modification of the method of Connell et al. (5). Triplicate live and control samples (6 or 8 ml) were incubated without headspace in Glass-pak syringes (Becton Dickinson). The syringe tips were adapted with Teflon-silicone septa sealed with silicone to allow injection of radiolabel and removal of samples over time. The syringes were incubated statically without headspace in the dark at room temperature (23°C). Killed controls were either double filtered (pore size, 0.2 µm), autoclaved, or killed with 4% formaldehyde. Inhibited controls were used in some experiments. CH<sub>3</sub>F was added from a saturated solution to a final concentration of 9% (vol/vol) in order to inhibit methane and ammonia oxidation. Chloramphenicol (20 µg/ml) and cycloheximide (483 µg/ml) were used to inhibit prokaryotic and eukaryotic protein synthesis, respectively. The effect of adding 100 µM unlabeled CH<sub>3</sub>Br, CH<sub>3</sub>OH, CH<sub>3</sub>Cl, or trimethylamine (TMA) on the rates of  ${}^{14}\text{CH}_3\text{Br}$  oxidation was also assessed with seawater samples.

<sup>\*</sup> Corresponding author. Mailing address: United States Geological Survey, 345 Middlefield Road, MS 480, Menlo Park, CA 94025. Phone: (650) 329-4473. Fax: (650) 329-4463. E-mail: kgoodwin@usgs.gov.

Aliquots (1 ml) of water were removed from syringes at different times and centrifuged with a solution consisting of  $20 \,\mu$ l of  $1 \,$  M NaOH, 0.5 ml of  $1 \,$  M SrCl<sub>2</sub> · 6H<sub>2</sub>O, and 100  $\,\mu$ l of  $1 \,$  M Na<sub>2</sub>CO<sub>3</sub>. The Na<sub>2</sub>CO<sub>3</sub> was added to ensure adequate pellet formation. Each resulting pellet was rinsed twice with 1 ml of ethanol, suspended in a dilute NaOH solution (pH ~11.7), and collected in ScintiVerse

Site	Water type	Salinity (g/liter)	рН	Collection temp (°C)	Avg initial no. of bacteria/ml (no. of samples)	
Searsville Lake San Francisco Bay Monterey Bay, Moss Beach, Mendocino Mono Lake	Freshwater Estuarine Coastal seawater Hypersaline, alkaline	ND <sup>a</sup> 6–19 <sup>b</sup> 29–34 74–77	7.6–8.3 7.4–7.8 7.3–8.2 9.8	15–23 11–21 10–16 3–21	$3 \times 10^{6}$ (3) $2 \times 10^{6}$ (5) $1 \times 10^{6}$ (10) $1 \times 10^{7}$ (3)	

TABLE 1. Description of sampling sites

<sup>*a*</sup> ND, not detectable with a refractometer.

<sup>b</sup> The sample containing 6 g of salt per liter was obtained during a period of heavy rain in January 1997.

II counting cocktail (Fisher Scientific). The <sup>14</sup>C in the pellet, the <sup>14</sup>C in the supernatant, and the <sup>14</sup>C in the ethanol rinses were collected separately and counted with a liquid scintillation spectrophotometer (model LS 6000SC; Beckman).

The concentrations of each labeled brominated methane (nanocuries per liter) were multiplied by the specific activity of the compound to calculate the total brominated methane concentrations (nanomolar). <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> (>99% pure; Dupont NEN) was added to syringes from an aqueous stock solution that had a specific activity of 49.7 mCi/mmol. A typical final concentration was 2.4 nCi/ml (49 nM total CH<sub>2</sub>Br<sub>2</sub>). <sup>14</sup>CH<sub>3</sub>Br (99.9% pure; Dupont NEN) was added from ethanolic stock solutions that had specific activities of 54.9 and 49.0 mCi/mmol. A typical final concentration was 16 nCi/ml (~300 nM total CH<sub>3</sub>Br). The final concentrations of ethanol typically ranged from 0.14 to 8.6 mM depending on the <sup>14</sup>CH<sub>3</sub>Br stock solution and the concentration used. The oxidation rates were comparable for similar <sup>14</sup>CH<sub>3</sub>Br concentrations; therefore, ethanol concentrations in this range did not appear to affect <sup>14</sup>CH<sub>3</sub>Br oxidation rates. The levels of recovery of <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> in the supernatants ranged from 94 to 100% after processing. Chemical degradation of <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> was not significant or the supernatant of the supern

The levels of recovery of <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> in the supernatants ranged from 94 to 100% after processing. Chemical degradation of <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> was not significant on the time scale of these experiments (25). The initial levels of recovery of <sup>14</sup>CH<sub>3</sub>Br in the supernatant were >95% before processing but typically 80 to 90% after the supernatant and pellet were separated, presumably because of losses due to volatilization. The supernatant values were corrected for the amounts lost in processing. The levels of recovery of <sup>14</sup>C-labeled bicarbonate (0.0934 nCi/ml; specific activity, 54.4 mCi/mmol; Dupont NEN). The pellet was designated  $\Sigma^{14}CO_2$  to indicate that it could contain labeled cellular material, as well as carbonate products from biological oxidation which were precipitated as SrCO<sub>3</sub>(s).

Measurement of bacterial density and growth. Bacterial density was determined by using acridine orange direct counts (AODC) (15). Sterile sodium citrate (0.1 M, pH 6.6) was added dropwise during filtration of samples from Searsville Lake and Mono Lake to remove background fluorescence (12). The coefficients of variation (CV) (CV = standard deviation/average) of AODC measurements were ~5 to 30%. Bacterial densities were determined for samples taken at the beginning of each experiment. In addition, time courses for cell growth were determined in some experiments. The incubations were conducted by using unlabeled brominated methane in parallel with radiolabel experiments. The dissolved oxygen concentration was monitored during some syringe incubations by using an electrode (model 10M-4 oxygen meter; Microelectrodes, Inc.).

**Calculations.** In all of the cases tested, first-order kinetics (e.g., the rate was proportional to the concentration) were observed for the brominated methane concentrations typically used during incubations. However, a straight line generally fit time courses as well as or better than an exponential curve fit due to small changes in concentration over time and the limited number of data points. Thus, oxidation rates were calculated by linear regression over the linear portion of the time course of  $\Sigma^{14}CO_2$  production. All rates are given below in terms of bacterial oxidation. These rates were calculated by determining the difference between the regression values of live and control samples. The CV for the rates (CV = standard error/slope) were typically 3 to 9% for CH<sub>2</sub>Br<sub>2</sub> and 8 to 19% for CH<sub>3</sub>Br. Reaction rates (nanomoles/liter day<sup>-1</sup>) and half-lives were calculated by using

observed rate = 
$$k_{app}C$$
 (1)

$$t_{1/2} = \ln 2/k_{app}$$
 (2)

where  $k_{app}$  is the apparent first-order rate constant (day<sup>-1</sup>), *C* is the applied brominated methane concentration (nanomolar), and  $t_{1/2}$  is the half-life due to bacterial oxidation (days). Using these equations allowed us to directly compare our values to rate constants provided by other researchers (5, 19). However, the numbers of bacteria differed in the various types of water, and oxidation rates are expected to be a function of bacterial number. To normalize for differences in bacterial number, the rate constants for each experiment were divided by the initial bacterial density. Thus, the reaction rates (nanomoles/liter day<sup>-1</sup>) per cell per milliliter) and half-lives could be calculated by using the following equations:

observed rate = 
$$k'_{app}CN$$
 (3)

$$t_{1/2} = \ln 2/Nk'_{app}$$
 (4)

where  $k'_{app}$  is the normalized apparent rate constant (nanomoles/liter day<sup>-1</sup> per

cell per milliliter), *C* is the applied brominated methane concentration (nanomolar), *N* is the density of brominated methane-oxidizing bacteria (cells per milliliter), and  $t_{1/2}$  is the half-life due to bacterial oxidation (days). Equations 2 and 4 are equivalent (*N* cancels). However, equation 4 is more robust because normalized apparent rate constants account for different bacterial densities in natural waters. The equations given above assume that there is no threshold for bacterial uptake. The total bacterial numbers obtained from AODC measurements taken at the start of each incubation were used as a proxy for *N*. Note that the actual number of oxidizers is an unknown subset of the total population.

**Enrichment cultures.** Enrichment cultures were started by amending seawater samples (50 ml in 160-ml vials) with KH<sub>2</sub>PO<sub>4</sub> (0.02 g/liter), NH<sub>4</sub>Cl (0.5 g/liter), vitamins (including vitamin B<sub>12</sub>) (1 ml/liter) (33), and CH<sub>2</sub>Br<sub>2</sub> (10  $\mu$ M) or CH<sub>3</sub>Br (50  $\mu$ M). Cultures received several spikes of brominated methane (a total of ~500  $\mu$ M brominated methane) and were then transferred to a defined medium. Serial dilution was used for purification because growth on agar plates was not obtained. Degradation was observed at 10<sup>-8</sup> dilutions, but microscopic observation showed that there were at least two bacterial morphologies, indicating that the cultures were not pure. The CH<sub>2</sub>Br<sub>2</sub> culture was maintained on a basal medium adapted from the medium of Visscher and Taylor (42) and supplemented with NaCl (35 g/liter), vitamins (including vitamin B<sub>12</sub>) (33), trace metal solution SL-10 (43), and CH<sub>2</sub>Br<sub>2</sub> as the sole carbon source, and bicarbonate was added as the buffer after autoclaving in order to achieve a final pH of 7.0 to 7.3. The CH<sub>3</sub>Br culture was cultivated on a basal medium adapted from the medium of Doronina et al. (7) and supplemented with NaCl (10 g/liter), vitamins (including vitamin B<sub>12</sub>) (33), trace metal solution SL-10 (43), and CH<sub>3</sub>Br as the sole



FIG. 1. Time course of bacterial oxidation of  ${}^{14}\text{CH}_2\text{Br}_2$  and simultaneous formation of oxidized products in freshwater (A), estuarine water (B), coastal seawater (C), and hypersaline-alkaline water from Mono Lake (D). The error bars represent  $\pm$  standard deviation of the mean obtained with three replicate syringes.



FIG. 2. Time course of bacterial oxidation of  $^{14}CH_3Br$  and simultaneous formation of oxidized products in freshwater (A), estuarine water (B), coastal seawater (C), and hypersaline-alkaline water from Mono Lake (D). The error bars represent  $\pm 1$  standard deviation of the mean obtained with three replicate syringes.

carbon source. The enrichment cultures were maintained in serum vials capped with Teflon-butyl rubber septa. The cultures were tested for their ability to degrade CH<sub>3</sub>Br, CH<sub>2</sub>Br<sub>2</sub>, and CH<sub>4</sub>. The concentrations of these gases were measured by headspace injection (200  $\mu$ ) into a gas chromatograph equipped with flame ionization detection (model HNU gas chromatograph; 60/80 Carbopack B column [4 ft by 0.20 in.]); the detector temperature was 250°C, and the oven temperatures were 60°C for CH<sub>4</sub>, 70°C for CH<sub>3</sub>Br, and 175°C for CH<sub>2</sub>Br<sub>2</sub>. Bromide ion concentrations were determined by ion chromatograph (30). CH<sub>2</sub>Br<sub>2</sub> was added to cultures as a liquid (99% pure; Chem Service), and CH<sub>3</sub>Br was added as a gas (99.9% pure; Matheson).

### RESULTS

Brominated methane oxidation in natural samples. Biological oxidation of <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> and <sup>14</sup>CH<sub>3</sub>Br to  $\Sigma^{14}$ CO<sub>2</sub> was measured in samples from freshwater, estuarine, coastal seawater, and hypersaline-alkaline sites. No oxidation occurred in filtered controls (Fig. 1 and 2) or in samples in which bacterial activity was eliminated by formaldehyde treatment or by autoclaving (data not shown). Bacteria from the different sites oxidized similar amounts of <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> (Fig. 1) and tended to oxidize less <sup>14</sup>CH<sub>3</sub>Br (Fig. 2). Freshwater bacteria oxidized the greatest amounts of <sup>14</sup>CH<sub>3</sub>Br (Fig. 2), although the Mono Lake samples contained the highest initial cell numbers (Table 1). Oxidation of CH<sub>2</sub>Br<sub>2</sub> and CH<sub>3</sub>Br proceeded without a lag in fresh, estuarine, and hypersaline samples. A several-hour lag period (generally >10 and <24 h) was observed for production of  $\Sigma^{14}$ CO<sub>2</sub> in coastal seawater samples (Fig. 1 and 2B).

Effects of inhibitors and additions. Oxidation of <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> and <sup>14</sup>CH<sub>3</sub>Br in freshwater samples was strongly inhibited by CH<sub>3</sub>F (Table 2). <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> oxidation in Mono Lake samples and <sup>14</sup>CH<sub>3</sub>Br oxidation in San Francisco Bay samples were slightly inhibited by CH<sub>3</sub>F (Table 2). CH<sub>3</sub>F did not affect brominated methane oxidation in other experiments. Therefore, freshwater samples were the only samples in which bacteria, such as methane or ammonia oxidizers, predominately mediated brominated methane oxidation. The eukaryotic inhibitor cycloheximide was tested with Mono Lake samples and had no significant effect on <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> or <sup>14</sup>CH<sub>3</sub>Br oxidation (data not shown); therefore, eukaryotes were not significantly involved in oxidation at that site. It has been shown previously that cycloheximide does not affect CH2Br2 oxidation in seawater (9). Oxidation in unamended samples occasionally slowed during incubation (Fig. 1 and 2), suggesting that activity sometimes became limited. The dissolved oxygen concentration remained >0.4 mM when it was measured, suggesting that oxygen was not the limiting factor (34).

The supernatant in <sup>14</sup>CH<sub>3</sub>Br experiments may have contained products of <sup>14</sup>CH<sub>3</sub>Br hydrolysis (<sup>14</sup>CH<sub>3</sub>OH) or halide exchange (<sup>14</sup>CH<sub>3</sub>Cl) (19), raising the possibility that  $\Sigma^{14}$ CO<sub>2</sub> was formed indirectly from bacterial oxidation of these chemical degradation products. However, previous studies performed with Mono Lake water (5) and agricultural soils (29) indicated that  $\Sigma^{14}$ CO<sub>2</sub> resulted from direct bacterial oxidation of <sup>14</sup>CH<sub>3</sub>Br and not from oxidation of chemical degradation products. We obtained similar results with seawater samples. Addition of unlabeled CH<sub>3</sub>OH or CH<sub>3</sub>Cl (100 µM) had no significant effect on  $\Sigma^{14}$ CO<sub>2</sub> production (Table 3). In contrast, addition of unlabeled  $CH_3Br$  (100  $\mu$ M) resulted in a significant decrease in the rate of  $\Sigma^{14}$ CO<sub>2</sub> production (Table 3), as expected from isotope dilution. In a separate experiment, the rate of  $\Sigma^{14}$ CO<sub>2</sub> production was significantly decreased by adding only 2 µM unlabeled CH<sub>3</sub>Br (data not shown). These results indicate that the bacteria directly oxidized methyl bromide, which is consistent with previous findings (5, 29). Also consistent with previous results obtained with Mono Lake water (5) is the finding that adding TMA to seawater increased <sup>14</sup>CH<sub>3</sub>Br oxidation rates (Table 3).

Apparent rate constants. Apparent rate constants (Tables 4 and 5) were calculated from oxidation rates by assuming that first-order kinetics occurred with respect to brominated methane concentration. First-order kinetics were shown to occur for  $CH_2Br_2$  in seawater and for  $CH_3Br$  concentrations typically used for seawater, estuarine water, and freshwater samples. In

TABLE 2. Effect of CH<sub>3</sub>F on bacterial oxidation of <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> and <sup>14</sup>CH<sub>3</sub>Br

Water type	Amt of dibromomethane oxidized $(nmol liter^{-1} day^{-1})$			Amt of methyl bromide oxidized (nmol liter <sup>-1</sup> day <sup>-1</sup> )		
	Without CH <sub>3</sub> F	With CH <sub>3</sub> F	% Inhibition	Without CH <sub>3</sub> F	With CH <sub>3</sub> F	% Inhibition
Freshwater	$9.5 \pm 0.8$	$1.7 \pm 0.5$	82	$12 \pm 0.8$	0	100
Estuarine	$5.2 \pm 0.3$	$5.3 \pm 0.4^{a}$	0	$9.1 \pm 0.8$	$7.5 \pm 0.3$	19
Seawater	$16 \pm 1.2$	$17 \pm 1.3$	0	$0.11 \pm 0.02$	$0.14 \pm 0.05$	0
Seawater	$21 \pm 0.9$	$20 \pm 2.2$	0			
Hypersaline-alkaline	$27 \pm 1.2$	$21 \pm 1.8$	22	$0.45\pm0.05$	$0.44\pm0.07$	0

<sup>a</sup> Samples were taken during a period of heavy rain.

TABLE 3.	Effect of adding unlabeled C1 compounds on bacterial
	oxidation of <sup>14</sup> CH <sub>3</sub> Br in seawater samples <sup>a</sup>

Addition	Amt of <sup>14</sup> CH <sub>3</sub> Br added (nCi)	Amt of ${}^{14}\Sigma CO_2$ produced (nCi/day)		
None	164	$0.43 \pm 0.07^{b}$		
CH <sub>3</sub> Br (100 µM)	164	$0.25\pm0.04$		
None	85	$0.14 \pm 0.01$		
CH <sub>3</sub> OH (100 µM)	85	$0.16\pm0.01$		
None	135	$0.24 \pm 0.02$		
CH <sub>3</sub> Cl (100 µM)	135	$0.25 \pm 0.02$		
TMA (100 µM)	135	$0.34\pm0.05$		

<sup>a</sup> The data are from three separate experiments.

<sup>b</sup> Average  $\pm$  standard error.

seawater samples, the rates of <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> oxidation were proportional to concentration over the range tested (35 to 455 nM  $CH_2Br_2$ ) (data not shown), and the rates of <sup>14</sup>CH<sub>3</sub>Br oxidation were also proportional to concentration over the range tested  $(125 \text{ to } 1,035 \text{ nM CH}_3\text{Br})$  (Fig. 3). In estuarine samples, the oxidation rates were proportional to concentration at concentrations between 160 and 1,337 nM CH<sub>3</sub>Br but declined at higher concentrations (data not shown), presumably due to toxicity from the ethanol in the <sup>14</sup>CH<sub>3</sub>Br stock solution. Although true saturation was not reached, the apparent half-saturation constant (app  $K_m$ ) for estuarine water was determined to be >2,000 nM CH<sub>3</sub>Br. Oxidation of <sup>14</sup>CH<sub>3</sub>Br in Searsville Lake followed Michaelis-Menten kinetics, as demonstrated by a linear Eadie-Hofstee plot (Fig. 4). The app  $K_m$  was 234  $\pm$  32 nM  $CH_3Br$ . The maximum velocity ( $V_{max}$ ) was 48  $\pm$  2 nmol of  $CH_3Br$  liter<sup>-1</sup> day<sup>-1</sup>.

Bacterial oxidation of <sup>14</sup>CH<sub>3</sub>Br tended to be slower as the salinity of the water increased. This pattern was seen in the rate constants calculated with equation 1 (Table 4) and with equation 3 (Table 5). The rate constants for <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> calculated from equation 1 varied little (Table 4) and thus were greater in marine and estuarine waters when they were normalized to the initial cell number (Table 5). In freshwater, the rate constants for CH<sub>2</sub>Br<sub>2</sub> and CH<sub>3</sub>Br were similar (Tables 4 and 5) and resulted in half-lives of 2 days for CH<sub>2</sub>Br<sub>2</sub> and 5 days for CH<sub>3</sub>Br (Table 6). In samples from more saline sites, the rate constants for CH2Br2 oxidation were 20- to 280-fold higher than the rate constants for CH<sub>3</sub>Br oxidation (Tables 4 and 5). For example, the  $k_{app}$  values for estuarine water corresponded to a half-life of 2 days for CH<sub>2</sub>Br<sub>2</sub> and a half-life of 36 days for CH<sub>3</sub>Br (Table 6). The  $k_{app}$  values for seawater also gave a 2-day half-life for CH<sub>2</sub>Br<sub>2</sub>, but the half-life of CH<sub>3</sub>Br due to biological oxidation of CH<sub>3</sub>Br was 82 days. In Mono Lake samples, the half-life of CH<sub>2</sub>Br<sub>2</sub> was only 1 day, but the half-life of CH3Br was 298 days.

Our CH<sub>3</sub>Br  $k_{app}$  value for seawater (Table 4) was approximately sevenfold lower than the biological rate constant of King and Saltzman (19) for seawater samples incubated at 21°C. In their experiments, King and Saltzman used <sup>13</sup>CH<sub>3</sub>Br at concentrations of 50 to 800 pM, values which are up to 10<sup>3</sup>-fold lower than the <sup>14</sup>CH<sub>3</sub>Br concentrations used here. Our somewhat lower apparent rate constant suggests that the <sup>14</sup>CH<sub>3</sub>Br concentrations used in this study did not artificially increase the numbers of CH<sub>3</sub>Br-degrading bacteria. Our CH<sub>3</sub>Br  $k_{app}$  value for Mono Lake water (Table 4) was two- to fourfold lower than the values reported by Connell et al. (5) based on in situ measurements. Connell et al. used 10-fold-higher <sup>14</sup>CH<sub>3</sub>Br concentrations than the concentrations used

TABLE 4. Apparent first-order rate constants for bacterial oxidation of <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> and <sup>14</sup>CH<sub>3</sub>Br<sup>a</sup>

Watar time	$k_{\rm app} \ ({\rm day}^{-1})$ for:			
water type	Dibromomethane	Methyl bromide		
Freshwater	$0.30 \pm 0.15 \ (2/2)^b$	$0.20 \pm 0.18 (3/3)$		
Estuarine	$0.37 \pm 0.31 (2/2)$	$0.019 \pm 0.013$ (10/4)		
Coastal seawater	$0.44 \pm 0.10 (7/5)$	$0.0084 \pm 0.006(10/6)$		
Hypersaline-alkaline	$0.58 \pm 0.19$ (4/3)	$0.0023 \pm 0.0017(5/3)$		

<sup>a</sup> Values were determined by using equation 1.

<sup>b</sup> Average  $\pm$  standard deviation. The values in parentheses are number of experiments/number of separately collected water samples.

here, suggesting that the  ${}^{14}CH_3Br$  concentrations which we used were also not inhibitory.

**Bacterial growth and brominated methane oxidation.** Differences in oxidation rates appeared to reflect environmental variations at a sampling site over time. For example, brominated methane oxidation rates and bacterial numbers declined during a period of heavy rain and flooding in 1997. During the height of this flood event, <sup>14</sup>CH<sub>3</sub>Br oxidation was not detected in seawater (data not shown). In contrast, the oxidation rates and bacterial numbers were not significantly affected by storing collected water at 4°C in the dark for more than 1 week (data not shown).

Time courses of AODC measurements were used to monitor changes in bacterial number during syringe incubations. Only total cell numbers could be measured; numbers of brominated methane-degrading bacteria per se could not be measured. The total cell numbers did not increase significantly during incubation of freshwater samples and increased only slightly (1.5- to 3-fold) during incubation of estuarine water samples. The increases in cell number varied from 1.5- to 19-fold for seawater and Mono Lake samples. Figures 5A and C show the results of two separate experiments in which order-of-magnitude increases in cell number were measured during seawater incubation, as expected from bottle effects (8). Cell growth was similar in syringes that were acid washed and heat treated to remove organic material (data not shown), indicating that the syringes did not provide a significant source of substrate. Cell growth in seawater was similar whether brominated methane (Fig. 5A), ethanol (Fig. 5C), or methanol (data not shown) was added to syringes or not. No growth or oxidation occurred in filter-sterilized controls amended with <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> or <sup>14</sup>CH<sub>3</sub>Br (Fig. 5).

The total cell number was independent of the brominated methane concentration (Fig. 5); therefore, Monod type equations did not describe brominated methane oxidation. A nonlinear least-squares regression analysis of cell number and sub-

TABLE 5. Normalized apparent rate constants for oxidation of <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> and <sup>14</sup>CH<sub>3</sub>Br<sup>a</sup>

Water type	$k'_{\rm app}$ (10 <sup>8</sup> day <sup>-1</sup> per bacterium <sup>-1</sup> ) per ml)			
water type	Dibromomethane	Methyl bromide		
Freshwater Estuarine Coastal seawater Hypersaline-alkaline	$8.1 \pm 0.81 (2/2)^{b}$ $39 \pm 44 (2/2)$ $72 \pm 30 (7/5)$ $6.2 \pm 1.7 (2/2)$	$5.1 \pm 2.2 (3/3) \\ 1.3 \pm 0.7 (10/4) \\ 0.56 \pm 0.33 (10/6) \\ 0.022 \pm 0.016 (3/3)$		

<sup>*a*</sup> Values were determined by using equation 3.

<sup>b</sup> Average  $\pm$  standard deviation. The values in parentheses are number of experiments/number of separately collected water samples. Individual rate constants were normalized by initial bacterial density for each experiment. The errors in AODC measurements ranged from ~5 to 30%.



FIG. 3. Rates of  $\Sigma CO_2$  production versus CH<sub>3</sub>Br concentrations in seawater samples. The rates were calculated by linear regression after the lag period. The error bars represent ±1 standard error of the regression for the rate.

strate concentration after the lag verified that degradation was best estimated by assuming that the cell number remained constant (data not shown). Thus, the use of equation 3 was supported even in the worst-case scenario of cell growth (Fig. 5). However, a lag in cell growth (Fig. 5A and C) did accompany oxidation of  $^{14}$ CH<sub>2</sub>Br<sub>2</sub> (Fig. 5B) and  $^{14}$ CH<sub>3</sub>Br (Fig. 5D) in seawater samples, and addition of chloramphenicol to seawater samples inhibited both cell growth and oxidation of  $^{14}$ CH<sub>3</sub>Br (Fig. 5C and D). In addition, chloramphenicol added to estuarine water inhibited both the increase in cell number and the oxidation rate by a factor of about 2 (data not shown), indicating that blocking protein synthesis could affect brominated methane oxidation.

Mass balance of brominated methanes. Syringes were tested for their ability to retain radiolabeled brominated methanes. Two types of syringes, Glass-pak (Becton Dickinson) and allglass (Baxter) syringes, were tested by using either stopcocks or Teflon-silicone septa. The most reliable performance was obtained with Glass-pak syringes sealed with septa whose septum puncture was resealed with silicone after each sampling. Mass balance of radiolabel was achieved for all incubations of <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> (Fig. 1). The rates of loss of radiolabel from <sup>14</sup>CH<sub>3</sub>Br were 3 to 15% per day, but they were higher (3 to 34% per day) if the septa were not resealed with silicone. These losses and the error associated with measuring <sup>14</sup>CH<sub>3</sub>Br in the supernatant made it difficult to distinguish live samples from control samples without measuring  $\Sigma CO_2$  production (Fig. 2). The loss of <sup>14</sup>CH<sub>3</sub>Br from syringes was less pronounced when Mono Lake water or culture medium was used. To investigate this phenomenon, either the pH or the salinity (1.5, 3, or 7.5% NaCl) of filtered Searsville Lake samples was



FIG. 4. Eadie-Hofstee plot of  $CH_3Br$  oxidation in Searsville Lake samples. The error bars represent  $\pm 1$  standard error of the regression for the rate.

TABLE 6. Half-lives of dibromomethane and methyl bromide due to bacterial oxidation, chemical consumption, and volatilization

Water	Half-lives (days) of:						
	Dibromomethane			Methyl bromide			
	Bacterial <sup>a</sup>	Chemical <sup>b</sup>	Voltili- zation <sup>c</sup>	Bacterial <sup>a</sup>	Chemical <sup>d</sup>	Voltili- zation <sup>c</sup>	
Freshwater	2			5	~30		
Estuarine	2			36	8-15		
Seawater	2			82	$\sim 5$		
Hypersaline- alkaline	1			298	~3		
All waters		$4 \times 10^4$	$\sim 1-16$			~1-16	

 $^a$  Values were determined by using equation 2 and the  $k_{\rm app}$  values in Table 4.  $^b$  Data from reference 25.

<sup>c</sup> Values were determined by using the following expression:  $\ln 2/KL$ , where *K* is the transfer velocity (1 to 4.2 m day<sup>-1</sup>) (seasonal wind speeds, 0 to 20 m/s) and *L* is the mixed-layer depth (4 to 6 m for Searsville Lake, 2 to 40 m for San Francisco Bay, 11 to 23 m for Mono Lake, and 2 to 100 m for coastal seawater) (3, 14, 23, 28) (see also the following website: http://sfbay.wr.usgs.gov/access/wqdata).

wqdata). <sup>d</sup> Data from references 16 and 19. The temperature used was 23°C. Salinities are given in Table 1.

increased. The water pH was increased from 8 to 10 by adding NaOH or 50 mM NaCO<sub>3</sub>. Only the addition of NaCO<sub>3</sub> halted the loss of radiolabel from syringes (the rate of loss was 24% per day at pH 8, compared with 0.76% per day at pH 10). The increased carbonate alkalinity caused by the addition of NaCO<sub>3</sub> should have increased the hydrolysis of <sup>14</sup>CH<sub>3</sub>Br to <sup>14</sup>CH<sub>3</sub>OH due to buffer catalysis (32). Methanol is less volatile than CH<sub>3</sub>Br, and enhanced production of CH<sub>3</sub>OH was observed in Mono Lake water (5), which may have resulted in less loss of radiolabel from syringes.

Oxidation of brominated methanes by enrichment cultures. An enrichment culture designated EBr2 was obtained from seawater, and this culture could degrade 500  $\mu$ M CH<sub>2</sub>Br<sub>2</sub> when it was provided as the sole carbon source in a defined medium. Concentrations of CH<sub>2</sub>Br<sub>2</sub> greater than 3.6 mM halted cell growth and degradation activity. The culture degraded CH<sub>2</sub>Br<sub>2</sub> with concomitant cell growth (Fig. 6) and stoichiometric release of free bromide ion (data not shown). In addition, the culture oxidized <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> to  $\Sigma^{14}$ CO<sub>2</sub> (data not shown). Cell growth was observed only with samples that received CH<sub>2</sub>Br<sub>2</sub> (Fig. 6). The cell number increased exponentially during  $CH_2Br_2$  degradation, resulting in a growth yield of  $6 \times 10^6$ cells ml<sup>-1</sup>  $\mu$ mol<sup>-1</sup> and a specific growth rate of 0.18 h<sup>-1</sup> (Fig. 6). The specific growth rate was similar to that reported for a dichloromethane utilizer, strain DM11, growing on CH<sub>2</sub>Br<sub>2</sub> (37). However, unlike DM11, our enrichment culture did not degrade CH<sub>2</sub>Cl<sub>2</sub> (25 or 300 µM). Our culture also did not consume CH<sub>3</sub>Br (30 or 300 µM) or CH<sub>4</sub> (1 or 10%), nor was it inhibited by CH<sub>3</sub>F. This culture grew on methylated amines (data not shown).

CH<sub>3</sub>Br-degrading enrichment cultures were established from seawater and were maintained by 1:10 dilution in aged seawater amended with nutrients and vitamins. Cell growth was observed only in seawater transfers that received CH<sub>3</sub>Br (Fig. 7). After several transfers, an enrichment culture (designated EBr1) was grown in a defined medium supplemented with 50 ml of aged seawater per liter. Under these conditions, the culture quickly oxidized <sup>14</sup>CH<sub>3</sub>Br to  $\Sigma^{14}CO_2$  in syringe experiments performed like the experiments with natural waters (Fig. 8), and the  $k_{app}$  was 2.3 day<sup>-1</sup> for 370 nM applied CH<sub>3</sub>Br. The mass of <sup>14</sup>CH<sub>3</sub>Br lost was fully recovered as oxidized product, and the cell density remained constant over the course of the incubation (~4.5 × 10<sup>6</sup> cells/ml). This enrich-



FIG. 5. Time courses of cell growth in the presence and absence of  $CH_2Br_2$  for live and filtered seawater samples (A) with parallel  ${}^{14}CH_2Br_2$  oxidation (B) and time courses of cell growth in the presence and absence of  $CH_3BR$ , ethanol (etoh), and chloramphenicol for live and filterd seawater samples (C) with parallel  ${}^{14}CH_3Br$  oxidation (D).

ment culture did not consume  $CH_2Br_2$  (20 or 600  $\mu$ M) or  $CH_4$  (1%), nor was it inhibited by  $CH_3F$ . The culture grew on methylated amines (data not shown).

## DISCUSSION

Bacteria were able to oxidize <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> and <sup>14</sup>CH<sub>3</sub>Br in natural waters whose salinities varied from ~0 to 77 g/liter (Table 1 and Fig. 1 and 2). Freshwater was distinguished from other water types on the basis of relatively high <sup>14</sup>CH<sub>3</sub>Br oxidation rates (Fig. 2 and Tables 4 and 5). Freshwater was also the only water type in which oxidation of <sup>14</sup>CH<sub>3</sub>Br and <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> was governed primarily by CH<sub>3</sub>F-sensitive bacteria (Table 2), such as methane or ammonia oxidizers. Cooxidation in freshwater probably resulted in the relatively high rates of <sup>14</sup>CH<sub>3</sub>Br oxidation. The difference in app  $K_m$  values for CH<sub>3</sub>Br in Searsville Lake (234 nM) and San Francisco Bay (>2,000 nM) also demonstrated that the bacteria that oxidized brominated methanes in freshwater had a distinct physiology compared to the bacteria at the other sites.

In contrast to freshwater samples,  $CH_3F$  caused only slight or no inhibition in samples from other sites (Table 2); therefore, methane or ammonia oxidizers played only a limited role at these sites. Lack of  $CH_3F$  inhibition was observed previously for <sup>14</sup>CH<sub>3</sub>Br oxidation in Mono Lake (5) and for  $CH_2Br_2$ degradation in seawater (9). Methyl fluoride does not inhibit other aspects of  $C_1$  metabolism, such as methanol or formate oxidation (30); thus, other types of methylotrophs could be responsible for bacterial oxidation in such waters. For example, oxidation of <sup>14</sup>CH<sub>3</sub>Br in seawater samples was not inhibited by CH<sub>3</sub>F (Table 2), and TMA added to seawater stimulated oxi-



FIG. 6. Time course of  $CH_2Br_2$  degradation (dashed lines) and concomitant exponential growth (solid lines) for enrichment culture EBr2 that was incubated in serum vials and received  $CH_2Br_2$  as the sole carbon source in a defined medium. The error bars represent  $\pm 1$  standard deviation of the mean obtained with three replicate samples.

dation of <sup>14</sup>CH<sub>3</sub>Br (Table 3). Similar results were obtained with TMA in Mono Lake water (5) and with CH<sub>3</sub>OH in agricultural soils (29). In addition, enrichment cultures EBr1 and EBr2 were not inhibited by CH<sub>3</sub>F, and both of these cultures consumed brominated methanes that were present as sole carbon sources (Fig. 6 and 8) or when they were growing on methylated amines (data not shown). Recovery of these separate enrichment cultures, one of which was able to utilize CH<sub>2</sub>Br<sub>2</sub> and one of which was able to utilize CH<sub>3</sub>Br, suggests that the bacteria that oxidize <sup>14</sup>CH<sub>2</sub>Br<sub>2</sub> and the bacteria that oxidize <sup>14</sup>CH<sub>3</sub>Br in natural seawater samples may be distinct.

The app  $K_m$  for CH<sub>3</sub>Br in Searsville Lake samples was substantially lower than the  $K_i$  measured for CH<sub>3</sub>Br for an ammonia oxidizer (500 µM) (18) and was considerably lower than the app  $K_m$  for CH<sub>4</sub> measured in freshwater lake samples (5 to 10 µM) (11). However, the app  $K_m$  values for Searsville Lake and San Francisco Bay samples were much higher than the ambient CH<sub>3</sub>Br concentrations (13, 19), making it unlikely that ambient levels could provide enough energy for CH<sub>3</sub>Br-oxidizing bacteria to maintain themselves (6). However, higher concentrations than the concentrations indicated by steady-state measurements may be available to bacteria, particularly bacteria in the vicinity of brominated methane sources, because production and consumption occur simultaneously in aquatic



FIG. 7. Cell growth in seawater for live samples of an enrichment culture that received multiple additions of  $CH_3Br$  gas (dashed line) and samples that did not receive  $CH_3Br$ .



FIG. 8. Oxidation of <sup>14</sup>CH<sub>3</sub>Br to  $\Sigma^{14}$ CO<sub>2</sub> by enrichment culture EBr1 grown in a defined medium supplemented with 50 ml of seawater liter<sup>-1</sup>. Incubation was performed in syringes with no headspace, just like experiments performed with natural samples.

systems (6). Known aquatic sources of brominated methanes include macroalgae (10, 26) and phytoplankton (41). Terrestrial plants also have the capacity to biohalogenate (35), suggesting that submerged freshwater and estuarine plants may also be sources of brominated methanes.

It is also likely that bacteria oxidize brominated methanes in nature while supporting themselves on other C1 compounds. For example, bacteria can oxidize atmospheric concentrations of CH<sub>4</sub> with no threshold as long as they are provided with other suitable  $C_1$  compounds to sustain themselves (2, 17). In the freshwater lake, there should have been enough  $CH_4$  (28) to support methanotrophic cooxidation of brominated methanes. In the other systems, where methanotrophs were not primarily involved (Table 2), methylated amines may have been used to maintain the brominated methane-degrading population. For example, adding TMA stimulated <sup>14</sup>CH<sub>3</sub>Br oxidation in seawater (Table 3) and in Mono Lake samples (5). In addition, both EBr1 and EBr2 could consume brominated methanes while they grew on methylated amines (data not shown). Furthermore, a facultatively methylotrophic bacterium isolated from agricultural soil grows on CH<sub>3</sub>Br, as well as methylated amines (4).

In addition to removal by bacterial oxidation, brominated methanes may be removed from the water column by chemical degradation.  $CH_2Br_2$  is chemically stable in water and has a half-life due to hydrolysis on the order of hundreds of years (25). Bacterial oxidation of  $CH_2Br_2$  thus appears to be faster than chemical consumption for all of the waters tested (Table 6). In contrast,  $CH_3Br$  undergoes significant chemical degradation, particularly in more saline waters (16, 19). Bacterial oxidation of  $CH_3Br$  appears to be as fast as chemical consumption only in freshwater (Table 6).

Brominated methanes may also be removed from the water column by volatilization. Volatilization from a body of water depends on the transfer velocity and the mixed-layer depth (14). The half-lives of  $CH_2Br_2$  and  $CH_3Br$  due to volatilization from the different sites examined are on the order of 1 to 16 days depending on the conditions (Table 6). When this rough estimate was used, bacterial oxidation of  $CH_2Br_2$  had the potential to compete with volatilization for all of the natural waters sampled in this study (Table 6). This result is consistent with the conclusions of Goodwin et al. (9) for coastal seawater. In Searsville Lake, the half-life of  $CH_3Br$  due to bacterial oxidation was comparable to the half-life due to volatilization (Table 6). Therefore, bacteria may help regulate the flux of brominated methanes in this lake. Bacterial oxidation of  $CH_3Br$  appears to be slower than volatilization in San Francisco Bay, coastal seawater, and Mono Lake (Table 6).

Oxidation of CH<sub>3</sub>Br in natural seawater samples was relatively slow compared to oxidation of CH<sub>3</sub>Br in less saline samples (Tables 4 and 5). However, enrichment culture EBr1 demonstrated that marine bacteria can grow on CH<sub>3</sub>Br (Fig. 7) and can rapidly oxidize <sup>14</sup>CH<sub>3</sub>Br in syringe experiments (Fig. 8). The slow turnover observed in natural samples may have been due to a low abundance of CH<sub>3</sub>Br-degrading bacteria in the environment. For example, the  $k_{app}$  for <sup>14</sup>CH<sub>3</sub>Br oxidation in seawater was about 300 times lower than the  $k_{app}$  for EBr1  $(2.3 \text{ day}^{-1})$  (Fig. 8). The numbers of cells in the EBr1 culture were similar to the final numbers of cells in seawater samples (~6  $\times$  10<sup>6</sup> cells ml<sup>-1</sup>), but the predominant bacteria in EBr1 should have been CH<sub>3</sub>Br-degrading bacteria, unlike in natural samples. Bacteria of the type found in the EBr1 culture could thus account for the oxidation rates observed in seawater samples if they constituted only  $\sim 0.3\%$  of the final sample population. If bacteria of the type found in EBr1 had been present at a similar ratio in the initial sample population (before growth), they would have been present at a concentration of about  $10^3$  cells/ml. At that cell density, the initial oxidation rates would have been below the limit of detection of the method used in these experiments, which is consistent with the observed lag (Fig. 2B) and with the results of chloramphenicol amendment experiments (Fig. 5C and D).

The effect of chloramphenicol addition on CH<sub>3</sub>Br oxidation in estuarine and seawater (Fig. 5D) samples raises the possibility that enzyme induction and/or cell growth could cause  $k_{app}$  values to be overestimated in some cases. Overestimation of  $k_{app}$  values due to optimization of the degrading population seems to be most likely for seawater and Mono Lake water samples because the lower oxidation rates and the lag for seawater samples necessitated relatively long incubation times (>24 h) (Fig. 1 and 2). Any correction for cell growth would have lowered the apparent rate constants and would have supported the dominance of abiotic consumption and volatilization over biological oxidation of CH<sub>3</sub>Br (Table 6).

Lobert et al. (23) demonstrated that large regions of the open ocean are undersaturated with respect to  $CH_3Br$ , and recent models have implicated a significant biological sink for  $CH_3Br$  (24). Our results confirm that bacterial oxidation of  $CH_3Br$  does occur in coastal seawater and that carbonate products are produced. However, the reaction rate (Table 2) obtained in our experiments suggests that biological removal is not greater than chemical consumption. Our reaction rate was almost 10 times lower than that of King and Saltzman (at 21°C) (19), which may reflect geographical differences in microbial density and composition. This difference underscores the need to experimentally assess bacterial oxidation of  $CH_3Br$  in different ocean regions and cautions against extrapolating results based on nearshore, urban environments to the open ocean.

#### ACKNOWLEDGMENTS

This work was supported by a National Research Council postdoctoral associateship and by NASA Upper Atmosphere Research Program grant 5188-AU-0080.

We thank B. Jellison, D. Heil, and C. Culbertson for supplying Mono Lake water, T. Connell Hancock, G. Hancock, and D. Hayward for providing coastal seawater, N. Chiariello for providing access to Searsville Lake, and the Moss Beach Marine Preserve and the Long Marine Laboratory for providing access to seawater.

#### REFERENCES

 Bartnicki, E. W., and C. E. Castro. 1994. Biodehalogenation: rapid oxidative metabolism of mono- and polyhalomethanes by *Methylosinus trichosporium* OB-3b. Environ. Toxicol. Chem. 13:241-245.

- Benstead, J., G. M. King, and H. G. Williams. 1988. Methanol promotes atmospheric methane oxidation by methanotrophic cultures and soils. Appl. Environ. Microbiol. 64:1091–1098.
- Broecker, W. S., and T. H. Peng. 1974. Gas exchange rates between air and sea. Tellus 26:19–35.
- Connell, T. L., A. M. Costello, M. E. Lidstrom, and R. S. Oremland. 1998. Strain IMB-1: a novel bacterium for the removal of methyl bromide in fumigated agricultural soils. Appl. Environ. Microbiol. 64:2899–2905.
- Connell, T. L., S. B. Joye, L. G. Miller, and R. S. Oremland. 1997. Bacterial oxidation of methyl bromide in Mono Lake, California. Environ. Sci. Technol. 31:1489–1495.
- Conrad, R. 1984. Capacity of aerobic microorganisms to utilize and grow on atmospheric trace gases (H<sub>2</sub>,CO, CH<sub>4</sub>), p. 461–467. *In* M. J. Klug and C. R. Reddy (ed.), Current perspectives in microbial ecology. American Society for Microbiology, Washington, D.C.
- Doronina, N. V., V. I. Krauzova, and Y. A. Trotsenko. 1997. *Methylophaga limanica* sp. nov., a new species of moderately halophilic, aerobic, methylotrophic bacteria. Microbiology (Engl. Transl. Mikrobiologiya) 66:434–439.
- Ferguson, R. L., E. N. Buckley, and A. V. Palumbo. 1984. Response of marine bacterioplankton to differential filtration and confinement. Appl. Environ. Microbiol. 47:49–55.
- Goodwin, K. D., M. E. Lidstrom, and R. S. Oremland. 1997. Marine bacterial degradation of brominated methanes. Environ. Sci. Technol. 31:3188–3192.
- Goodwin, K. D., W. J. North, and M. E. Lidstrom. 1997. Production of bromoform and dibromomethane by giant kelp: factors affecting release and comparison to anthropogenic bromine sources. Limnol. Oceanogr. 42:1725– 1734.
- Hanson, R. S., and T. E. Hanson. 1996. Methanotrophic bacteria. Microbiol. Rev. 60:439–471.
- Harvey, R. W. 1987. A fluorochrome-staining technique for counting bacteria in saline, organically enriched, alkaline lakes. Limnol. Oceanogr. 32:993–995.
- Haung, W., and R. Gammon. Personal communication.
   Helz, G. R., and R. Y. Hsu. 1978. Volatile chloro- and bromocarbons in coastal waters. Limnol. Oceanogr. 23:858–869.
- Hobbie, J. E., R. J. Daley, and S. Jasper. 1977. Use of Nuclepore filters for counting bacteria by fluorescence microscopy. Appl. Environ. Microbiol. 33: 1225–1228.
- Jeffers, P. M., and N. L. Wolfe. 1996. On the degradation of methyl bromide in sea water. Geophys. Res. Lett. 23:1773–1776.
- Jensen, S., A. Priemé, and L. Bakken. 1998. Methanol improves methane uptake in starved methanotrophic microorganisms. Appl. Environ. Microbiol. 64:1143–1146.
- Keener, W. K., and D. J. Arp. 1993. Kinetic studies of ammonia monooxygenase inhibition in *Nitrosomonas europaea* by hydrocarbons and halogenated hydrocarbons in an optimized whole-cell assay. Appl. Environ. Microbiol. 59:2501–2510.
- King, D. B., and E. S. Saltzman. 1997. Removal of methyl bromide in coastal seawater: chemical and biological rates. J. Geophys. Res. 102:18715–18721.
- Kley, D., P. J. Crutzen, H. G. J. Smit, H. Vomel, S. J. Oltmans, H. Grassl, and V. Ramanathan. 1996. Observations of near-zero ozone concentrations over the convective Pacific: effects on air chemistry. Science 274:230–232.
- Kourtidis, K., R. Borchers, and P. Fabian. 1996. Dibromomethane (CH<sub>2</sub>Br<sub>2</sub>) measurements at the upper troposphere and lower stratosphere. Geophys. Res. Lett. 23:2581–2583.
- Leisinger, T., and R. Bader. 1993. Microbial dehalogenation of synthetic organohalogen compounds: hydrolytic dehalogenases. Chimia 47:116–121.
- Lobert, J. M., J. H. Butler, S. A. Montzka, L. S. Geller, R. C. Myers, and J. W. Elkins. 1995. A net sink for atmospheric CH<sub>3</sub>Br in the east Pacific Ocean. Science 267:1002–1005.
- Lobert, J. M., S. A. Yvon-Lewis, J. H. Butler, S. A. Montzka, and R. C. Myers. 1997. Undersaturation of CH<sub>3</sub>Br in the southern ocean. Geophys. Res. Lett. 24:171–172.
- 25. Mabey, W., and T. Mill. 1978. Critical review of hydrolysis of organic com-

pounds in water under environmental conditions. J. Phys. Chem. Ref. Data 7:383-409.

- Manley, S. L., K. Goodwin, and W. J. North. 1992. Laboratory production of bromoform, methylene bromide and methyl iodide by macroalgae and distribution in nearshore southern California waters. Limnol. Oceanogr. 37: 1652–1659.
- Mellouki, A., R. K. Talukdar, A. Schmoltner, T. Gierczak, M. J. Mills, S. Solomon, and A. R. Ravishankara. 1992. Atmospheric lifetimes and ozone depletion potentials of methyl bromide (CH<sub>3</sub>Br) and dibromomethane (CH<sub>2</sub>Br<sub>2</sub>). Geophys. Res. Lett. 19:2059–2062.
- Miller, L. G., and R. S. Oremland. 1988. Methane efflux from the pelagic regions of four lakes. Glob. Biogeochem. Cycles 2:269–277.
- Miller, L. G., T. L. Connell, J. R. Guidetti, and R. S. Oremland. 1997. Bacterial oxidation of methyl bromide in fumigated agricultural soils. Appl. Environ. Microbiol. 63:4346–4354.
- Oremland, R. S., and C. W. Culbertson. 1992. Evaluation of methyl fluoride and dimethyl ether as inhibitors of aerobic methane oxidation. Appl. Environ. Microbiol. 58:2983–2992.
- Oremland, R. S., L. G. Miller, C. W. Culbertson, T. L. Connell, and L. L. Jahnke. 1994. Degradation of methyl bromide by methanotrophic bacteria in cell suspensions and soils. Appl. Environ. Microbiol. 60:3640–3646.
- Perdue, E. M., and N. L. Wolfe. 1983. Prediction of buffer catalysis in field and laboratory studies of pollutant hydrolysis reactions. Environ. Sci. Technol. 17:635–642.
- Phenning, N. 1978. *Rhodocyclus purpureus* gen. nov. and sp. nov., a ringshaped, vitamin B<sub>12</sub>-requiring member of the family *Rhodospirillaceae*. Int. J. Syst. Bacteriol. 28:283–288.
- Rheinheimer, G. 1991. Aquatic microbiology, 4th ed., p. 144. John Wiley & Sons, Chichester, United Kingdom.
- Saini, H. S., J. M. Attieh, and A. D. Hanson. 1995. Biosynthesis of halomethanes and methanethiol by higher plants via a novel methyltransferase reaction. Plant Cell Environ. 18:1027–1033.
- Schauffler, S. M., L. E. Heidt, W. H. Pollock, T. M. Gilpin, J. F. Vedder, S. Solomon, R. A. Lueb, and E. L. Atlas. 1993. Measurements of halogenated organic compounds near the tropical tropopause. Geophys. Res. Lett. 20: 2567–2570.
- Scholtz, R., L. P. Wackett, C. Egli, A. M. Cook, and T. Leisinger. 1988. Dichloromethane dehalogenase with improved catalytic activity isolated from a fast-growing dichloromethane-utilizing bacterium. J. Bacteriol. 170: 5698–5704.
- Shorter, J. H., C. E. Kolb, P. M. Crill, R. A. Kerwin, R. W. Talbot, M. E. Hines, and R. C. Harriss. 1995. Rapid degradation of atmospheric methyl bromide in soils. Nature 377:717–719.
- Singh, H. B., L. J. Salas, and R. E. Stiles. 1983. Methyl halides in and over the eastern Pacific (40 N - 32 S). J. Geophys. Res. 88:3684–3690.
- Tanhua, T., E. Fogelqvist, and Ö. Bastürk. 1996. Reduction of volatile halocarbons in anoxic seawater, results from a study in the Black Sea. Mar. Chem. 54:159–170.
- Tokarczyk, R., and R. M. Moore. 1994. Production of volatile organohalogens by phytoplankton cultures. Geophys. Res. Lett. 21:285–288.
- Visscher, P. T., and B. F. Taylor. 1994. Demethylation of dimethylsulfoniopropionate to 3-mercaptopropionate by an aerobic marine bacterium. Appl. Environ. Microbiol. 60:4617–4619.
- 43. Widdel, F., G.-W. Kohring, and F. Mayer. 1983. Studies on the dissimilatory sulfate-reducing bacteria that decompose fatty acids. III. Characterization of the filamentous gliding *Desulfonema limicola* gen. nov., sp. nov., and *Desulfonema magnum* sp. nov. Arch. Microbiol. 134:286–294.
- 44. World Meteorological Organization. 1995. Scientific assessment of ozone depletion: 1994. World Meteorological Organization Global Ozone Research and Monitoring Project-Report no. 37. World Meteorological Organization, Geneva, Switzerland.
- Zhang, D. Q., J. X. Zhong, and L. X. Qiu. 1997. Kinetics of the reaction of hydroxyl radicals with CH<sub>2</sub>Br<sub>2</sub> and its implication in the atmosphere. J. Atmos. Chem. 27:209–215.