

Simultaneous Measurement of Denitrification and Nitrogen Fixation Using Isotope Pairing with Membrane Inlet Mass Spectrometry Analysis†

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A method for estimating denitrification and nitrogen fixation simultaneously in coastal sediments was developed. An isotope-pairing technique was applied to dissolved gas measurements with a membrane inlet mass spectrometer (MIMS). The relative fluxes of three N₂ gas species (²⁸N₂, ²⁹N₂, and ³⁰N₂) were monitored during incubation experiments after the addition of ¹⁵NO₃⁻. Formulas were developed to estimate the production (denitrification) and consumption (N₂ fixation) of N₂ gas from the fluxes of the different isotopic forms of N₂. Proportions of the three isotopic forms produced from ¹⁵NO₃⁻ and ¹⁴NO₃⁻ agreed with expectations in a sediment slurry incubation experiment designed to optimize conditions for denitrification. Nitrogen fixation rates from an algal mat measured with intact sediment cores ranged from 32 to 390 μg-atoms of N m⁻² h⁻¹. They were enhanced by light and organic matter enrichment. In this environment of high nitrogen fixation, low N₂ production rates due to denitrification could be separated from high N₂ consumption rates due to nitrogen fixation. Denitrification and nitrogen fixation rates were estimated in April 2000 on sediments from a Texas sea grass bed (Laguna Madre). Denitrification rates (average, 20 μg-atoms of N m⁻² h⁻¹) were lower than nitrogen fixation rates (average, 60 μg-atoms of N m⁻² h⁻¹). The developed method benefits from simple and accurate dissolved-gas measurement by the MIMS system. By adding the N₂ isotope capability, it was possible to do isotope-pairing experiments with the MIMS system.

Denitrification and nitrogen fixation are important counteractive processes affecting nitrogen dynamics in coastal sediments. Phytoplankton production can be limited by nitrogen availability (35), and sediments often serve as an important source of dissolved inorganic N (calculated as NH₄⁺ + NO₂⁻ + NO₃⁻). Denitrification is the only biological process that transforms combined N to gaseous forms (N₂ or N₂O) (36, 37). These gaseous end products are unavailable to most producers (e.g., phytoplankton and bacteria) unless N₂ is transformed into organic N during nitrogen fixation (12). Benthic denitrification is a significant sink for combined N in systems and may drive systems toward N limitation (37). Nitrogen fixation increases the amount of biologically available N (7, 38, 42).

Despite the important role of denitrification in coastal and open ocean systems, accurate measurement is hindered by high background levels of N₂ gas in the atmosphere and water column (1, 8, 9, 18, 20, 24, 30, 32, 39, 40). Isotope pairing and membrane inlet mass spectrometry (MIMS) techniques have improved the accuracy and precision of denitrification measurements (4, 15, 16, 25, 28, 29, 33).

The isotope-pairing technique developed by Nielsen (29) estimates denitrification by monitoring changes in nitrogen gas with different isotope compositions (²⁹N₂ = ¹⁴N + ¹⁵N, ³⁰N₂ = ¹⁵N + ¹⁵N) after enriching the overlying water with ¹⁵NO₃⁻.

The ²⁸N₂ (¹⁴N + ¹⁴N) production rate was calculated from the relative ²⁹N₂ and ³⁰N₂ production rates (4, 28, 29, 33). Avoiding the measurement of ²⁸N₂ reduced the possibility of contamination. A model simulation showed that the modified gradient created by ¹⁵NO₃⁻ addition has a minimal effect on in situ denitrification rates (26).

Quadruple mass spectrometers have been linked with a MIMS to enhance measurement of dissolved gases (15). The MIMS improved the accuracy of dissolved gas measurement and decreased chances of contamination. Observing the change in ratio between N₂ and Ar measured with the MIMS provides a sensitive and convenient determination of net denitrification rates (8, 15), but measured N₂ flux is a net result of production (denitrification) and consumption (nitrogen fixation).

Nitrogen fixation is a process mediated by microbes that convert N₂ to organic nitrogen. Benthic nitrogen fixation can increase nitrogen availability for biological production in coastal regions (7, 38, 42). Nitrogen fixation rates are low in most coastal sediments, except for areas covered by microbial mats or sea grass beds. Simultaneous measurements of denitrification and nitrogen fixation are desirable, but such data are not common. A method to measure both processes in the same sample is needed to quantify the two processes in situations where both may be important.

Available methods for nitrogen fixation measurement require certain assumptions and have shortcomings (25, 38). A common method for nitrogen fixation measurement is an acetylene (C₂H₂) reduction assay (10). Nitrogen fixers do not discriminate between C₂H₂ and N₂ as substrates during nitrogen fixation. Although simple, inexpensive, and sensitive, this tech-

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nique requires a conversion constant relating the ethylene production rate to the N₂ reduction rate (theoretical ratio = 3 mol of acetylene per 1 mol of N₂ reduction) (38). The ratio varies and is affected by environmental conditions (38).

We expanded the capability of the MIMS system (15) to measure different isotopic forms of N₂ gas (²⁹N₂ and ³⁰N₂) relative to Ar (S. An and W. S. Gardner, submitted for publication), for isotope-pairing experiments with the MIMS system. By comparing fluxes of three N₂ gas species (²⁸N₂, ²⁹N₂, and ³⁰N₂), production (denitrification) and consumption (nitrogen fixation) rates for N₂ gas could be measured at the same time. Here, we report the methodology and formulas developed for the simultaneous estimation of the two processes. Procedures for determining the different isotopic forms of N₂ gas (²⁹N₂ and ³⁰N₂) using the MIMS and testing of the method are described. In a "potential-denitrification" experiment, conditions were optimized for denitrification measurements and fluxes for the three forms of N₂ gas were observed. The sensitivities of the MIMS system for three isotopic forms of N₂ were evaluated by comparing measured fluxes with expected production rates for each gas. A second experiment measured rates in algal mat sediments where nitrogen fixation was expected. The method differentiated and quantified both denitrification and total nitrogen fixation (net N₂ change + denitrification) rates. Finally, rates of denitrification and nitrogen fixation were measured on intact sediment cores from Laguna Madre, Texas. Assumptions and problems associated with the method are discussed.

MATERIALS AND METHODS

(i) **Formulas.** The ²⁸N₂ production rate estimated from the isotope-pairing technique should match the ²⁸N₂ production rate estimated from direct methods (for example, MIMS) in cores where no nitrogen fixation occurs. The proportional changes in three N₂ species with time after ¹⁵NO₃⁻ addition can be predicted under such conditions. If nitrogen fixation occurs, however, predicted changes in the proportion of the three N₂ species with time will not match the measured changes. The difference between predicted and measured proportional changes of the three N₂ species can reflect nitrogen fixation and denitrification rates.

In the isotope-pairing technique (29), denitrification rates based on the added ¹⁵NO₃⁻ (*D*₁₅) are calculated as follows:

$$D_{15} = 2n(15,15) + n(14,15) \quad (1)$$

where *n*(15,15) is the net ³⁰N₂ production rate and *n*(14,15) is the net ²⁹N₂ production rate, and

$$D_{14} = \frac{n(14,15)}{2n(15,15)}D_{15} \quad (2)$$

where *D*₁₄ is the rate of denitrification based on ¹⁴NO₃⁻, and

$$\text{total denitrification} = D_{15} + D_{14} \quad (3)$$

The rate of denitrification based on ¹⁴NO₃⁻ (*D*₁₄) calculated by the isotope-pairing technique should equal the net ²⁸N₂ production rate [*n*(14,14)] plus the amount of ¹⁴N in ²⁹N₂ when nitrogen fixation (*f*) is zero, or

$$D_{14} = 2n(14,14) + n(14,15) \quad (4)$$

That is (from equations 2 and 4),

$$\frac{n(14,15)}{2n(15,15)}D_{15} = 2n(14,14) + n(14,15) \quad (5)$$

However, *D*₁₄ (calculated from the isotope-pairing technique [equation 2]) can be larger than 2*n*(14,14) + *n*(14,15) if nitrogen fixation occurs. Since ²⁸N₂ is the most abundant among the three N₂ gas species, the removal rate is the highest for ²⁸N₂ during nitrogen fixation. When nitrogen fixation coexists with denitrification, rates of N₂ gas species are the result of the balance of production

(denitrification) and consumption (nitrogen fixation); i.e., "gross" denitrification (*d*) is calculated as follows:

$$d = n + f \quad (6)$$

where *n* is net N₂ flux and *f* is gross nitrogen fixation (a positive number represents negative N₂ flux). The relationship can be applied to the gross denitrification of each N₂ species [*d*(14,14), *d*(14,15), and *d*(15,15)] as follows:

$$d(14,14) = n(14,14) + f(14,14) \quad (7)$$

$$d(14,15) = n(14,15) + f(14,15) \quad (8)$$

$$d(15,15) = n(15,15) + f(15,15) \quad (9)$$

If the three species of nitrogen gas (²⁸N₂, ²⁹N₂, and ³⁰N₂) are used in proportion to their concentration during nitrogen fixation (assuming no isotope fractionation), the gross nitrogen fixation based on each N₂ species [*f*(14,14), *f*(14,15), and *f*(15,15)] can be represented as follows:

$$f(14,14) = f(1 - \alpha - \beta) \quad (10)$$

$$f(14,15) = f\alpha \quad (11)$$

$$f(15,15) = f\beta \quad (12)$$

where α is the abundance of ²⁹N₂ among three N₂ gas species, β is the abundance of ³⁰N₂ among three N₂ gas species, and (1 - α - β) is the abundance of ²⁸N₂ among three N₂ gas species.

When *f* is greater than 0, equations 1 and 2 should be modified as follows:

$$D_{15}' = 2d(15,15) + d(14,15) \quad (13)$$

$$D_{14}' = \frac{d(14,15)}{2d(15,15)}D_{15}' \quad (14)$$

where *D*₁₅' is the rate of denitrification based on ¹⁵NO₃ when *f* is greater than 0 and *D*₁₄' is the rate of denitrification based on ¹⁴NO₃⁻ when *f* is greater than 0.

Here again, *D*₁₄' (calculated from the isotope-pairing technique) should equal gross ²⁸N₂ production plus ¹⁴N in gross ²⁹N₂ flux (see equation 4).

$$D_{14}' = d(14,15) + 2d(14,14) \quad (15)$$

Equations 13, 14, and 15 are combined to give an equation with regard to *d*(14,14), *d*(14,15), and *d*(15,15), as follows:

$$\frac{2[2d(14,14) + d(14,15)]d(15,15)}{d(14,15)} + d(14,15) + 2d(15,15) \quad (16)$$

The MIMS system can measure *n*(14,14), *n*(14,15), *n*(15,15), α , and β . Applying the measured numbers, we have four equations (equations 7, 8, 9, and 16) to be solved with four unknowns [*f*, *d*(14,14), *d*(14,15), and *d*(15,15)]. By solving the equations, denitrification (*D*₁₄' and *D*₁₅') and nitrogen fixation rates (*f*) can be derived. Equations 7, 8, 9, and 16 become a quadratic equation with regard to *f*.

$$Af^2 + Bf + C = 0 \quad (17)$$

where A = 4 β (1 - α - β) - α^2 , B = 4 β *n*(14,14) + 4(1 - α - β)*n*(15,15) - 2 α *n*(14,15), and C = 4*n*(14,14)*n*(15,15) - *n*(14,15)².

The solutions are

$$f = \frac{-B \pm \sqrt{B^2 - 4AC}}{2A} \quad (18)$$

The positive solution estimates nitrogen fixation. The appendix gives some sample calculations. The "no nitrogen fixation case" is illustrated in case 1. In this situation, total denitrification is a sum of three net N₂ fluxes. When nitrogen fixation exists (case 2), however, the nitrogen fixation rate is added to the net fluxes of three N₂ gases to obtain the total denitrification rate (see the appendix).

(ii) **Dissolved ²⁹N₂ and ³⁰N₂ measurements with the MIMS system.** We modified the method of Kana et al. (15) to measure dissolved ²⁹N₂ and ³⁰N₂ in addition to Ar, O₂, and ²⁸N₂. The ²⁹N₂ and ³⁰N₂ concentrations were obtained from the "excess" atomic mass unit (AMU) 29 and 30 signals, respectively, caused by the conversion of added ¹⁵NO₃⁻. The ratio between ²⁹N₂ and ²⁸N₂ is 0.00732 (0.00366 × 2) in natural samples, since the natural abundance of ¹⁵N is 0.366% (23). The relationship between ²⁹N₂ and ²⁸N₂ in samples was obtained from 30-ppt artificial seawater (standard water) held at different temperatures (21 and 30°C) (Fig. 1a). The ratio in our MIMS system was close to the theoretical value (Fig. 1a), but it was calculated in each measurement to account for

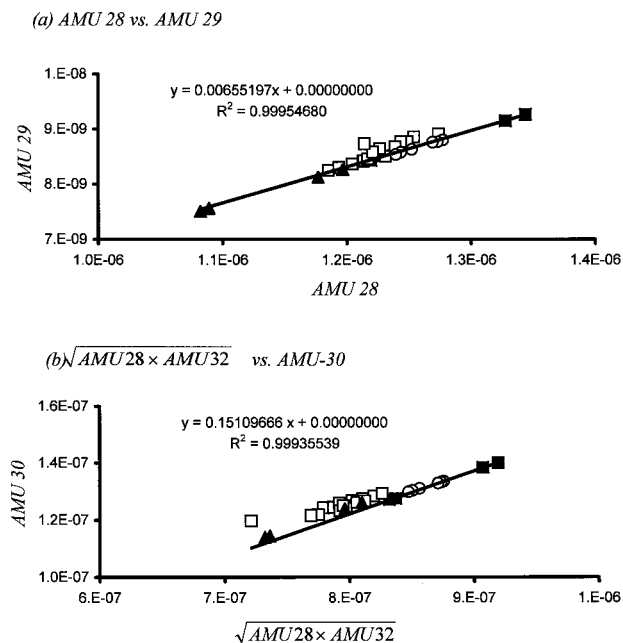


FIG. 1. Relationships between different AMU signals in liquid samples measured with the quadruple mass spectrometer. Solid symbols, standard water (■, 30 ppt, 21°C; ▲, 30 ppt, 30°C); open symbols, samples (□, outflow water; ○, inflow water). The line represents the regression among the standard water samples. The signal represents partial pressure of gases (in pascals [10^2]) inside the mass spectrometer. Data are from the work of An and Gardner (submitted).

small variability between measurements. This relationship was used to determine the excess AMU 29 signal. The feed water in a flowthrough incubation experiment does not have excess AMU 29 signal resulting from $^{29}\text{N}_2$, while outflow water has excess AMU 29 signal, indicating that $^{29}\text{N}_2$ was produced during denitrification (Fig. 1a). The excess AMU 29 signal was converted to excess $^{29}\text{N}_2$ concentration by comparing the results with those from standard water.

Interpreting the AMU 30 signal in a MIMS system is complicated by reactions between N_2 and O_2 in the ion source of the mass spectrometer. The NO^+ ions (AMU 30 signal) formed from N_2 and O^+ inside the mass spectrometer can cause a linear relationship between AMU 30 and $(\text{AMU } 28 \times \text{AMU } 32)^{0.5}$ (13). Note that $(\text{AMU } 28 \times \text{AMU } 32)^{0.5}$ exhibited a better regression with AMU 30 signal than did $(\text{AMU } 28 \times \text{AMU } 32)$ as suggested by Jensen et al. (13) (Fig. 1b). If oxygen (O) and nitrogen (N) atoms are combined randomly as N_2 , NO, and O_2 , and the proportions of O and N are p and q , respectively, then the proportions of the three products should show a binomial distribution ($\text{N}_2 = p^2$, $\text{NO} = 2pq$, $\text{O}_2 = q^2$) that explains the observed square root relationship of N_2O_2 with NO. This relationship was used to determine the excess AMU 30 signal in the water samples by comparing values with those for standard water (Fig. 1b).

The apparent N_2/Ar ratios decrease with increasing O_2 concentration because the O^- ion scavenges N_2 during the formation of NO^- inside the MIMS system. Since the O_2 concentrations of outflow water in our flowthrough system were lower than those of inflow water, uncorrected denitrification rates would be overestimated by the NO^- effect. This effect was corrected by an O_2 concentration versus N_2 concentration relationship obtained from standard water sample measurements. Standard water samples (21°C, 30 ppt) saturated with atmospheric N_2 and with various O_2 concentrations were prepared by selectively removing dissolved O_2 with known concentrations of sodium sulfite. During our flowthrough incubation experiments (see section iii, paragraphs b and c), the overestimation of denitrification rates should have been less than 3% because O_2 concentrations of outflow water were maintained at over 80% of that for inflow water during sediment incubations. The O_2 effect was not present in the closed-bottle experiments (see section iii, paragraph a), because O_2 concentrations were low and similar in both control and experimental treatments.

(iii) **Sediment incubation experiments.** The MIMS system was evaluated for measuring different N_2 species in three sediment incubation experiments. The

developed formulas were applied to data from each experiment to obtain denitrification and nitrogen fixation rates under different conditions.

(a) **Potential denitrification bottle experiment.** A potential denitrification experiment was performed in closed bottles. Sediment slurries (20 ml) from a fresh water pond (Research Park, College Station, Tex.) were placed in serum bottles (120 ml). The bottles were filled with distilled water saturated with air at 21°C before the bottles were sealed with gas tight caps. The distilled water was enriched with glucose (final concentration = 2 mM) to provide a carbon source. Five treatments with different proportions of $^{14}\text{NO}_3^-$ and $^{15}\text{NO}_3^-$ (0 and 100%, 25 and 75%, 50 and 50%, 75 and 25%, and 100 and 0%) and a control (no NO_3^- added) were prepared. Total concentrations of nitrate in each treatment bottle were the same ($32 \mu\text{g-atoms of N bottle}^{-1}$). Duplicate bottles were prepared for each treatment to give a total of 12 bottles. Each bottle was sealed so as to be gas tight with no headspace. Sediment slurries in the bottles were mixed well and then incubated at room temperature (21°C). After 24 h of incubation, dissolved gases from one bottle of each treatment (total of six) were measured with the MIMS system. Turbid sediment slurry samples were drawn into the membrane inlet tubes without filtration. Before measurements, the sediment slurry was mixed vigorously and large particles were allowed to settle out. The concentration of each nitrogen gas species was calculated as described above. After 48 h of incubation, the N_2 forms in the remaining six bottles were measured. The total production of each N_2 gas species was estimated by the concentration difference between each treatment and control.

(b) **Algal mat sediment incubation.** Undisturbed sediment cores (12-cm diameter, 30-cm length) with bottom water were collected in a shallow salt marsh area with algal mats near Port Aransas, Tex., in April 2000. The cores were transported to a temperature-controlled incubation room (21°C), and a flowthrough plunger with Teflon inlet and outlet tubes was installed over each sediment core. The flowthrough chamber setup consisted of an intake water vessel, Teflon flow tubes, a peristaltic pump, and a sample collection vessel (21). The bottom water collected from the site was passed over the core surface at a rate of 1.2 ml min^{-1} . Duplicate sediment cores were incubated under four different treatment conditions (total of eight). For the GL treatment, flowthrough water was enriched with glucose (2 mM) and incubated in the light (~ 300 microeinsteins $\text{m}^{-2} \text{ s}^{-1}$) to maximize N_2 fixation rates. GD cores were enriched with glucose and incubated without light. NL cores were incubated in the light (~ 300 microeinsteins $\text{m}^{-2} \text{ s}^{-1}$) without glucose enrichment. ND cores were incubated in the dark without glucose enrichment. Water column depth over the sediment was maintained at about 5 cm to give a water volume of ca. 570 ml in each core. Triplicate samples of feed water and outlet water were collected at intervals for dissolved-gas analysis after an initial incubation period of 1 day, to allow steady-state conditions to develop (21). Water samples were collected for analysis of dissolved inorganic nitrogen compounds (NH_4^+ , NO_3^- , and NO_2^-).

After sampling on the second day, feed water was enriched with $^{15}\text{NO}_3^-$ and concentrations of $^{28}\text{N}_2$, $^{29}\text{N}_2$, and $^{30}\text{N}_2$ were measured in inflow and outflow waters. Denitrification and nitrogen fixation rates for each sediment core were calculated using equations 17 and 18. Sediment flux of each chemical compound was calculated based on the concentration difference between feed water and outflow water, flow rate, and cross-sectional area (21). Measurements were made on days 1, 2, 3, 4, and 7.

(c) **Sediment core incubations with shallow estuarine sediments.** The flowthrough incubation experiment was repeated with sediment cores from Laguna Madre, a shallow, semienclosed estuary in the southeastern part of the Texas coast. Laguna Madre is a negative estuary where freshwater input is less than evaporation (water residence time = 1 year). The salinity is often more than 40 ppt and can vary annually up to 60 ppt (6). Laguna Madre stations had depths of 0.8 to 0.9 m and were populated with *Thalassia testudinum* (300 to 600 shoots m^{-2}) (22). During our sampling period, salinity was lower than in earlier observations (6) and did not exceed 40 ppt. Bottom water was oxygenated at most stations due to wind-driven mixing. Sediment cores (four in each station) were collected from two sites (L155 and L189). Station details are described elsewhere (An and Gardner, submitted). One half of the cores were incubated under dim light (~ 30 microeinsteins $\text{m}^{-2} \text{ s}^{-1}$), and the others were covered with aluminum foil. The dim-light condition did not cause significant light effects, so data from the two treatments were combined. The $^{15}\text{NO}_3^-$ was added after the day 1 samples were taken.

RESULTS AND DISCUSSION

(i) **Potential denitrification experiment.** Most of the added nitrate was converted to N_2 gas within 24 h. The average production rate of total N_2 ($^{28}\text{N}_2 + ^{29}\text{N}_2 + ^{30}\text{N}_2$) was 28.6

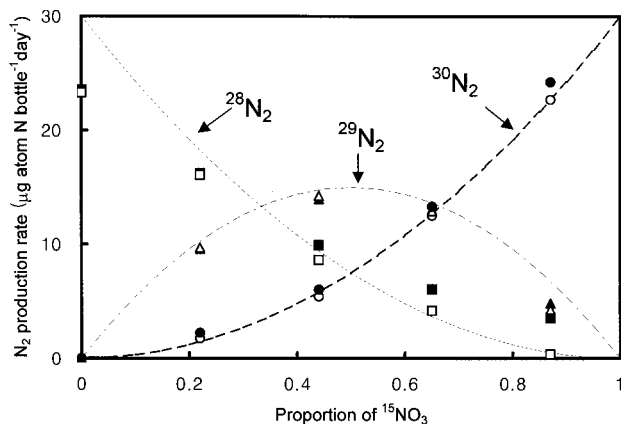


FIG. 2. Production rates of N_2 isotope species in a potential-denitrification experiment with sediments enriched with different proportions of $^{14}NO_3^-$ and $^{15}NO_3^-$. The line represents the expected production rate of the three N_2 isotope species based on the proportion of each NO_3^- species. Symbols represent measured production rates of three N_2 gases (solid symbols, day 1; open symbols, day 2). Total amounts of NO_3^- were constant in each treatment ($32 \mu\text{mol bottle}^{-1}$). About 90% of added NO_3^- was recovered as N_2 gas in this experiment.

(standard error [SE] = $1.5 \mu\text{g-atoms of N bottle}^{-1} \text{ day}^{-1}$, accounting for 90% of the added nitrate ($32 \mu\text{g-atoms of N bottle}^{-1}$). The concentrations of the three N_2 species were similar after 48 h of incubation ($27.9 \pm 1.2 \mu\text{g-atoms of N bottle}^{-1}$). For calculations, the N_2 species production was assumed to be completed at 24 h for both the 24- and 48-h-incubation samples. The production rates may have been higher than calculated because conversion of added nitrate to N_2 gas may have occurred before 24 h.

The isotopic distribution of each N_2 species end product matched the expected distribution based on the proportion of $^{14}NO_3^-$ and $^{15}NO_3^-$ species in the feed water (Fig. 2). Assuming that ^{14}N and ^{15}N bound randomly to form the N_2 gas species, the production of $^{28}N_2$, $^{29}N_2$, and $^{30}N_2$ would show binomial distribution (29). When the proportion of $^{15}NO_3^-$ is p (and that for $^{14}NO_3^-$ is $1 - p$), the production rates of the N_2 species would be calculated as follows: for $^{28}N_2$, $(1 - p)^2$; for $^{29}N_2$, $2p(1 - p)$; and for $^{30}N_2$, p^2 .

Although each treatment was designed to have 0, 25, 50, 75, and 100% $^{15}NO_3^-$ enrichment (p), the background level of $^{14}NO_3^-$ in the sediment caused the resulting $^{15}NO_3^-$ proportions (0, 22, 44, 65, and 87%) to be less than the calculated values. The $^{14}NO_3^-$ concentration in the control treatment was considered the background level of $^{14}NO_3^-$ to estimate p in each treatment. The $^{28}N_2$ flux was lower than expected at low p values and higher than expected at high p values (Fig. 2). This trend may have resulted from underestimating the background concentration of $^{14}NO_3^-$.

When the production rates of each N_2 isotope species show expected distributions, the nitrogen fixation rates can be assumed to be zero. In this case, the denitrification rates obtained from the N_2/Ar method should be identical to those from the isotope-pairing technique. The average measured nitrogen fixation rate was not exactly zero in this experiment but was a small negative number (average \pm SE = $-2.4 \pm 0.81 \mu\text{g-atoms of N bottle}^{-1} \text{ day}^{-1}$). The number of replicates was

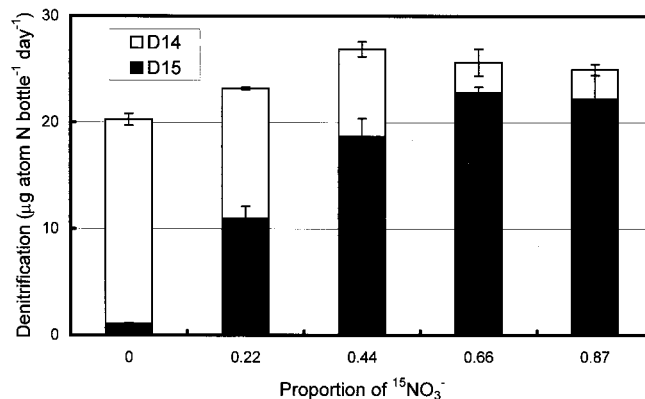


FIG. 3. Denitrification based on $^{14}NO_3^-$ (D_{14}') and $^{15}NO_3^-$ (D_{15}') during potential denitrification experiment. Error bar represents 1 SE for two incubation bottles.

not large enough to confirm statistically the zero nitrogen fixation condition observed in this experiment.

The isotope-pairing technique allowed estimation of denitrification rates produced from $^{14}NO_3^-$ (D_{14}') and $^{15}NO_3^-$ (D_{15}') (29). In our experiment, D_{14}' and D_{15}' showed the expected trends with the $^{15}NO_3^-$ -to- $^{14}NO_3^-$ proportions (Fig. 3). Denitrification rates based on the D_{15}' increased as the proportion of $^{15}NO_3^-$ increased. This result indicates that the excess AMU 29 signal and AMU 30 signal observed with the MIMS system for $^{29}N_2$ and $^{30}N_2$ concentration measurements is valid and the MIMS system has similar sensitivities for the three N_2 gas species. In another MIMS system using a quadrupole mass spectrometer, the instrumental response to $^{29}N_2$ over $^{30}N_2$ was close to 1 but larger (1.030) than the theoretical value (1.017 [13]). The sensitivity of each MIMS system for three N_2 gas species may be determined in a bioassay experiment using cultured denitrification bacteria (13).

(ii) **Sediment core incubations with algal mats.** An experiment was performed on sediments covered with an algal mat to evaluate the method in a nitrogen fixation environment (14, 43). Potential enhancers for the nitrogen fixation process, such as high light and a usable carbon source (11, 14, 25), were provided to some core samples. Before the $^{15}NO_3^-$ addition (days 1 and 2), $^{28}N_2$ fluxes into the sediment were observed with all treatments (Table 1), suggesting high nitrogen fixation rates. The lighted core without glucose (NL) had a higher negative flux than dark cores without glucose (ND). The glucose-enriched cores (GD and GL) had higher negative N_2 fluxes than treatments without glucose enrichments (ND and NL). The light effect was not obvious among glucose-enriched treatments (GD and GL). Sediment oxygen demand (SOD) was higher in dark (ND and GD) than light (NL and GL) treatments, suggesting active photosynthesis. The differences in SOD between dark and light cores were about $1,000 \mu\text{mol of O}_2 \text{ m}^{-2} \text{ h}^{-1}$ and resemble typical benthic primary production rates reported for shallow coastal environments (31). The glucose effect on SOD was not observed before the $^{15}NO_3^-$ addition.

After $^{15}NO_3^-$ additions, the sediment produced O_2 in light treatments, suggesting that photosynthesis was increased with added nitrate. Photosynthetic rates calculated from the SOD difference between light and dark treatments increased from

TABLE 1. Variations of N₂ isotope species and O₂ fluxes during a sediment incubation experiment

Gas and treatment ^a	Flux ^b on day:				
	1	2	3	4	7
²⁸ N ₂ (μg-atom of N m ⁻² h ⁻¹)					
ND	-80 (5.2)	-58 (14)	2.4 (22)	-30 (28)	-36
NL	NA	-194 (28)	-148 (20)	-140 (60)	-366 (50)
GD	NA	-176 (28)	-200 (120)	-17	-20
GL	-72 (28)	-116 (34)	-520	-420 (110)	-264 (16)
²⁹⁺³⁰ N ₂ (μg-atom of N m ⁻² h ⁻¹)					
ND			12.4 (0.4)	71 (18)	90
NL			0.6 (0.3)	2.2 (1.2)	3.7 (3)
GD			0.2 (1.5)	11	13
GL			1.2	3.4 (1.4)	3.3 (9)
O ₂ (μmol m ⁻² h ⁻¹)					
ND	-1,040 (20)	-1,300 (320)	-860 (60)	-1,200 (200)	-490
NL	NA	-360 (130)	1,200 (800)	1,400 (900)	4,500 (350)
GD	NA	-1,400 (240)	-1,050 (5)	-1,300	-1,300
GL	240 (200)	-390 (200)	460	480 (85)	160 (640)

^a See the text and Fig. 4 for explanations of the treatments.

^b Values are averages (SEs) from replicate cores (where no SE is given, replicate data were not available). Negative flux represents a flux into the sediment. Heavy nitrate (¹⁵NO₃⁻) was added after day 2, and the ²⁹⁺³⁰N₂ flux was measured from day 3. NA, not available.

1,000 μmol of O₂ m⁻² h⁻¹ before the addition to 1,500 to 5,000 μmol of O₂ m⁻² h⁻¹ after the addition.

Light and dark cores showed different ²⁸N₂ and ²⁹⁺³⁰N₂ fluxes. After the ¹⁵NO₃⁻ addition, negative ²⁸N₂ fluxes were higher in lighted cores than dark cores, suggesting light-enhanced nitrogen fixation activity. The ²⁹⁺³⁰N₂ fluxes were higher in dark cores than in light ones (Table 1). The results agree with the light enhancement effect observed for nitrogen fixation (11, 14, 25) and suggest that denitrification was inhibited by high oxygen concentrations (3, 19). The ¹⁵NO₃⁻ uptake by primary producers would be higher in lighted versus dark cores, making the ¹⁵NO₃⁻ less available for denitrification and causing the ²⁹⁺³⁰N₂ production to be lower in lighted cores. Increased oxygen penetration depth caused by benthic photosynthesis would reduce denitrification rates for the NO₃⁻ from the water column and cause reduced denitrification rates in lighted cores (33, 34). Benthic photosynthesis can enhance denitrification by increasing coupled nitrification-denitrification rates (2, 33). When O₂ production is high, however, the inhibitory effect may dominate the enhancement effect (2, 33, 34).

Figure 4 shows calculated total denitrification ($D_{14}' + D_{15}'$) and gross nitrogen fixation (f) rates after the ¹⁵NO₃⁻ additions. Denitrification rates were highest in ND cores and comparable to nitrogen fixation rates. Both denitrification and nitrogen fixation rates increased with time during the 7 days of incubation in the ND treatment. Denitrification rates in other treatments were low. Denitrification rates depended on O₂ levels in these experiments. Benthic photosynthesis would be low in dark cores, and the O₂ inhibition effect may be lower in dark than in light cores. The GD cores had higher denitrification rates than GL cores, but the denitrification rates were lower than those of ND cores (Table 1; Fig. 4).

Nitrogen fixation rates measured in this experiment (32 to 390 μg-atoms of N m⁻² h⁻¹) are comparable to rates for other cyanobacterial mats (8 to 650 μg-atoms of N m⁻² h⁻¹ [11]). The highest nitrogen fixation rate occurred in GL cores (average ± SE = 390 ± 60 μg-atoms of N m⁻² h⁻¹). The NL cores

also had high nitrogen fixation rates (220 ± 50 μg-atoms of N m⁻² h⁻¹), suggesting that nitrogen-fixing cyanobacteria were present in the cores. GD cores had higher nitrogen fixation rates (110 ± 50 μg-atoms of N m⁻² h⁻¹) than ND cores (32 ± 17 μg-atoms of N m⁻² h⁻¹). The observed glucose effect suggests the presence of heterotrophic nitrogen fixers as well as photoautotrophic cyanobacteria in the sediments (11, 27). The method presented here enables denitrification rates to be evaluated when nitrogen fixation is dominant and may provide a helpful tool for studying the complex interactions between nitrogen fixers and other microbial processes in algal mat communities (42, 43).

(iii) Sediment core incubation with shallow estuarine sediment. The cores were maintained at near natural conditions except that the illuminated ones were maintained under continuous laboratory fluorescent lighting rather than natural outdoor light conditions. Although the water depths were shallow at the sampling stations, in situ light levels at the sediment-water interface were about 10% of the surface values due to

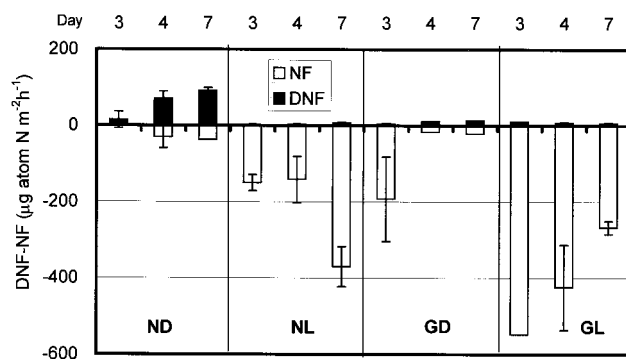


FIG. 4. Total denitrification (DNF) and nitrogen fixation (NF) rates under different light and carbon enrichment conditions. Data are averages ± SE for two sediment cores. Duplicate measurements were not available for day 4 and 7 of treatment GD and day 3 of treatment GL.

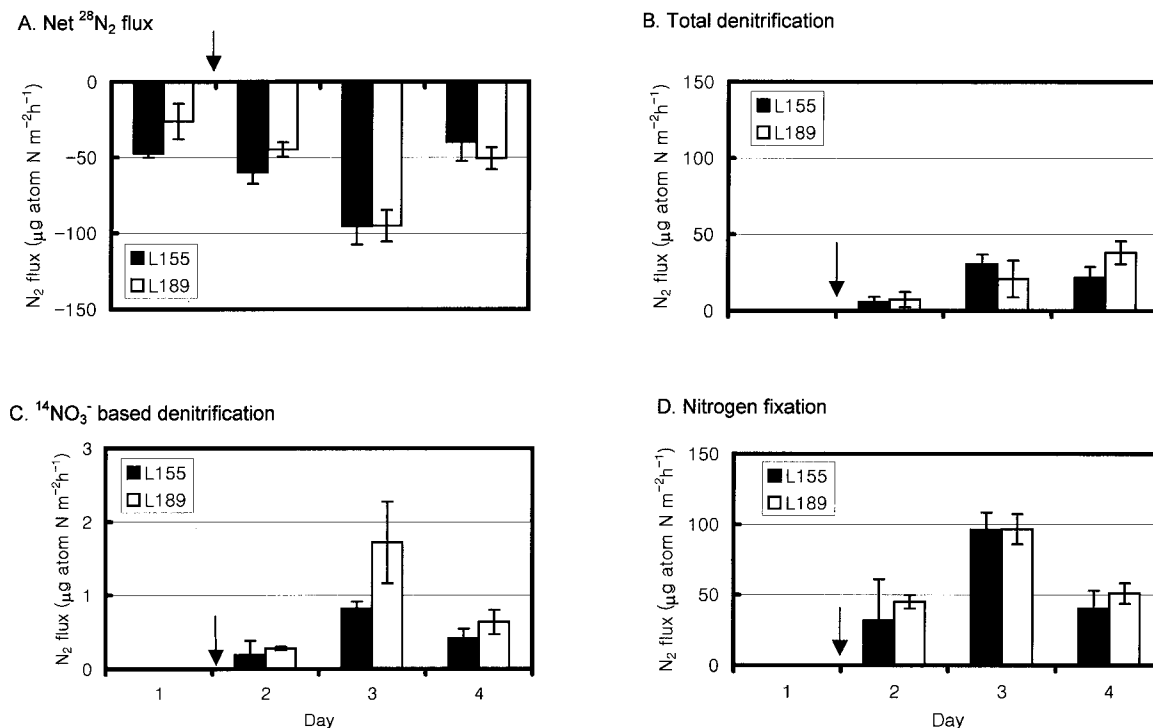


FIG. 5. Nitrogen gas flux changes during sediment incubation experiments. The arrow shows the time at which ¹⁵NO₃⁻ was added. Total denitrification is calculated as ²⁸N₂ + ²⁹N₂ + ³⁰N₂ flux. Data are averages ± 1 SE for four replicate cores. Note the scale change in panel C.

high turbidity (S. An and W. S. Gardner, Nitrogen Cycling in Laguna Madre and Baffin Bay, final report to the Texas Water Development Board). The dim light treatment during our laboratory incubation tended to lower SOD compared to dark treatments, suggesting that some benthic primary production activity occurred at the sites. However, data from the two treatments were combined because light effects were not obvious in other processes.

Negative ²⁸N₂ fluxes (37 ± 10 µg-atoms of N m⁻² h⁻¹) were observed before the ¹⁵NO₃⁻ addition, suggesting that nitrogen fixation rates were higher than denitrification rates (Fig. 5A). Denitrification rates in these sediments have been limited by organic matter supply. Temporal ²⁸N₂ flux varied with water column chlorophyll concentrations (An and Gardner, submitted; An and Gardner, report to the Texas Water Development Board). The net flux was highest in April 1999 (~130 µg-atoms of N m⁻² h⁻¹) and decreased in August 1999 (41 to 78 µg-atoms of N m⁻² h⁻¹) despite higher temperatures (An and Gardner, Nitrogen Cycling in Laguna Madre and Baffin Bay). During December 1999, a negative ²⁸N₂ flux was measured, indicating nitrogen fixation and low denitrification rates. The water column chlorophyll seemed to recover during January 2000, but denitrification rates continued to decrease in April 2000. Negative ²⁸N₂ fluxes increased from ~20 µg-atoms of N m⁻² h⁻¹ in December 1999 to 37 µg-atoms of N m⁻² h⁻¹ in April 2000 (An and Gardner, report to Texas Water Development Board).

The estimated denitrification rate was ~20 µg-atoms of N m⁻² h⁻¹, about one-third of the nitrogen fixation rate. The fixation rate may have been underestimated in these cores because of dark or dim light conditions. About 98% of the

denitrification was from the added ¹⁵NO₃⁻ rather than from ¹⁴NO₃⁻ (Fig. 5C). Considering that the sodium nitrate used in this study had 99% ¹⁵N and 1% ¹⁴N, it is possible that part of the ¹⁴NO₃⁻-based denitrification may have resulted from added ¹⁴NO₃⁻ rather than from nitrification-coupled denitrification (29). The measured nitrogen fixation rate (~60 µg-atoms of N m⁻² h⁻¹) is in the range of those reported for other sea grass beds. Sea grass beds can have higher nitrogen fixation rates (1 to 250 µg-atoms of N m⁻² h⁻¹) than uncolonized areas (0.02 to 5 µg-atoms of N m⁻² h⁻¹) (11). Potential rates of dissimilatory nitrate reduction to ammonium were also high in this region (An and Gardner, submitted). High nitrogen fixation activity together with high rates of dissimilatory nitrate reduction to ammonium may help sustain replete biota, including the Texas Brown Tide in this area (6; An and Gardner, submitted).

(iv) Sensitivity test. To evaluate measurement errors associated with ²⁹N₂ and ³⁰N₂ measurement, dissolved-gas data for seawater (30 ppt) at different temperatures (30 and 21°C) measured over 4 days were combined (total $n = 32$). The average excesses of AMU 29 signal and AMU 30 signal were -0.00007 (SE = 0.0002) and 0.001 (SE = 0.001) µM, respectively. The amount of dissolved gases interfering with the ²⁹N₂ and ³⁰N₂ measurement is different in experimental samples than in standard water, and the measurement error is larger. For example, nitric oxide (NO) produced during denitrification would interfere with the AMU 30 signal measurement and cause the ³⁰N₂ flux to be overestimated (41). The formation of ¹³CO⁺ from CO₂ also may interfere with ²⁹N₂ measurements (5). However, the excess AMU 29 signal and its variability in water samples prepared in different ways (different salinity,

TABLE 2. Sensitivity test for denitrification (DNF) and nitrogen fixation (NF) measurements^a

Flux adjustment	Sensitivity (%) in:			
	Case 1		Case 2	
	DNF	NF	DNF	NF
²⁹ N ₂				
-2%	100	100	97	34
+2%	100	100	103	168
³⁰ N ₂				
-2%	98	100	101	134
+2%	102	100	99	67
²⁸ N ₂				
-2%	100	102	100	131
+2%	100	98	100	69

^a Each N₂ isotope species flux was decreased and increased by 2% from the measured values, and resulting DNF and NF rates are presented as the percentage of the original values. In case 1, the ¹⁴NO₃⁻-based DNF rate (*D*₁₄[']) is much smaller than the ¹⁵NO₃⁻-based DNF (*D*₁₅[']) rate; in case 2, the two DNF rates are similar.

sulfite addition, or helium bubbling) were similar to those of standard water, suggesting that the interference of CO₂ in ²⁹N₂ measurement was small.

The behaviors of estimated denitrification and nitrogen fixation rates were assessed by varying the ²⁸N₂, ²⁹N₂, and ³⁰N₂ fluxes in two experiments (Table 2). The error range of the flux measurements used in this practice (2%) may overestimate variability, since the precision of ²⁸N₂ measurement of the MIMS system is less than 0.05% (15) and NO is a minor intermediate product during denitrification. When coupled nitrification-denitrification was low and denitrification depended on the added ¹⁵NO₃⁻ (*D*₁₄ ≪ *D*₁₅), the estimation of denitrification and nitrogen fixation rates was not sensitive to the ²⁹N₂ flux measurement. An overestimation of ³⁰N₂ fluxes caused an overestimation of denitrification rates but did not affect nitrogen fixation rates. The ²⁸N₂ flux affected only nitrogen fixation rates in this situation. When the amount of *D*₁₄ is similar to *D*₁₅, small measurement errors in each flux can cause large changes in estimates of denitrification and nitrogen fixation. For example, the overestimation of ²⁹N₂ by 2% would cause an overestimation of nitrogen fixation by 70%. The overestimation of ³⁰N₂ and ²⁸N₂ fluxes caused about a 30% decrease in nitrogen fixation estimates in this example.

The sensitivity of the results depends on the absolute concentrations as well as fluxes of three N₂ species and should be evaluated in each measurement. An independent measurement of interference gases may be necessary when the result is sensitive to flux data. The method relies on the accurate measurement of three N₂ gas species. When one or more flux measurements are not accurate, all estimates will be affected. A high-resolution mass spectrometer can discriminate gases with the same masses but different origins (13). Independent measurement of interference gases or parallel measurements with the high-resolution mass spectrometer could be used to calibrate the MIMS measurements.

One of the most powerful features of the MIMS system in denitrification measurements is the direct introduction of the water sample (15). The accuracy of the measurement is en-

hanced by avoiding the dissolved-gas-stripping step. The ability of the MIMS system to measure dissolved gases in turbid water samples prevents manipulation errors. By adding the capability of measuring isotopes of N₂ gas to the MIMS system, denitrification rates can be measured with an isotope-pairing technique. The simple and fast measurement procedure can be used in flowthrough and closed system sediment incubations.

APPENDIX

Some sample calculations with hypothetical ²⁸N₂, ²⁹N₂, and ³⁰N₂ production rates are presented below. See the text for explanations of each flux and constant.

Case 1. N fixation (*f*) = 0. For case 1, the following data were obtained with the MIMS system:

- n*(14,14) (net production rate of ²⁸N₂), 20 μmol of N₂ m⁻² h⁻¹
- n*(14,15) (net production rate of ²⁹N₂), 12 μmol of N₂ m⁻² h⁻¹
- n*(15,15) (net production rate of ³⁰N₂), 1.8 μmol of N₂ m⁻² h⁻¹
- α (proportion of ²⁹N₂ among three N₂), 0.00628
- β (proportion of ³⁰N₂ among three N₂), 0.00159

The equation is

$$0.006293f^2 + 7.120f = 0$$

The solutions are *f* = 0 and *f* = -1,131. Therefore, the nitrogen fixation rate is zero, and gross denitrification (*d*) is equal to net N₂ flux (*n*).

The denitrification based on ¹⁴NO₃⁻ (*D*₁₄[']) is

$$2d(14,14) + d(14,15) = 2 \times 20 + 12 \\ = 52 \text{ } \mu\text{g-atoms of N m}^{-2} \text{ h}^{-1}, \text{ or } 26 \text{ } \mu\text{mol of N}_2 \text{ m}^{-2} \text{ h}^{-1}$$

and total denitrification based on ¹⁵NO₃⁻ (*D*₁₅[']) is

$$2d(15,15) + d(14,15) = 2 \times 1.8 + 12 \\ = 15.6 \text{ } \mu\text{g-atoms of N m}^{-2} \text{ h}^{-1}, \text{ or } 7.8 \text{ } \mu\text{mol of N}_2 \text{ m}^{-2} \text{ h}^{-1}$$

Case 2. N fixation (*f*) > 0. For case 2, the following data were obtained with the MIMS system:

- n*(14,14) (net production rate of ²⁸N₂), 15 μmol of N₂ m⁻² h⁻¹
- n*(14,15) (net production rate of ²⁹N₂), 12 μmol of N₂ m⁻² h⁻¹
- n*(15,15) (net production rate of ³⁰N₂), 1.8 μmol of N₂ m⁻² h⁻¹
- α (proportion of ²⁹N₂ among three N₂), 0.00628
- β (proportion of ³⁰N₂ among three N₂), 0.00159

The equation is

$$0.006293f^2 + 7.008f - 36 = 0$$

The solutions are *f* = 5.06 and *f* = -1,131.4. Therefore, the nitrogen fixation rate is 5.06 μmol of N₂ m⁻² h⁻¹, and gross denitrification (*d*) is net N₂ flux (*n*) plus N fixation (*f*).

The denitrification based on ¹⁴NO₃⁻ (*D*₁₄[']) is

$$2d(14,14) + d(14,15) = 2[n(14,14) + f(1 - \alpha - \beta)] + [n(14,15) + f\alpha] \\ = 2[15 + 5.06(1 - 0.00628 - 0.00159)] + (12 + 5.06 \times 0.00628) \\ = 52.07 \text{ } \mu\text{g-atoms of N m}^{-2} \text{ h}^{-1}, \text{ or } 26.03 \text{ } \mu\text{mol of N}_2 \text{ m}^{-2} \text{ h}^{-1}$$

and total denitrification based on ¹⁵NO₃⁻ (*D*₁₅[']) is

$$2d(15,15) + d(14,15) = 2[n(15,15) + f\beta] + [n(14,15) + f\alpha] \\ = 15.64 \text{ } \mu\text{g-atoms of N m}^{-2} \text{ h}^{-1}, \text{ or } 7.82 \text{ } \mu\text{mol of N}_2 \text{ m}^{-2} \text{ h}^{-1}$$

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REFERENCES

1. An, S., and S. B. Joye. 1997. An improved chromatographic method to measure nitrogen, oxygen, argon and methane in gas or liquid samples. *Mar. Chem.* 59:63-70.

2. An, S., and S. B. Joye. 2001. Enhancement of coupled nitrification-denitrification by benthic photosynthesis in shallow estuarine sediments. *Limnol. Oceanogr.* **46**:42–47.
3. Anderson, T. K., M. H. Hensen, and J. Sørensen. 1984. Diurnal variation in nitrogen cycling in coastal marine sediments. I. Denitrification. *Mar. Biol.* **83**:171–176.
4. Anette, P., N. Risgaard-Petersen, and N. P. Revsbech. 1997. Denitrification and microphytobenthic NO₃ consumption in a Danish lowland stream: diurnal and seasonal variation. *Aquat. Microb. Ecol.* **12**:275–284.
5. Bender, L. M., P. P. Tans, T. Ellis, J. Orchardo, and K. Habfast. 1994. A high precision isotope ratio mass spectrometer method for measuring the O₂/N₂ ratio of air. *Geochim. Cosmochim. Acta* **58**:4751–4758.
6. Buskey, E. J., B. Wysor, and C. Hyatt. 1998. The role of hypersalinity in the persistence of the Texas 'brown tide' in the Laguna Madre. *J. Plankton Res.* **20**:1553–1565.
7. Capone, D. G. 1988. Benthic nitrogen fixation. In T. H. Blackburn and J. Sørensen (ed.), *Nitrogen cycling in coastal marine environments*. Wiley, New York, N.Y.
8. Cornwell, J. C., W. M. Kemp, and T. M. Kana. 1999. Denitrification in coastal ecosystems: methods, environmental controls and ecosystem level controls, a review. *Aquat. Ecol.* **33**:41–54.
9. Devol, A. H. 1991. Direct measurement of nitrogen gas fluxes from continental shelf sediments. *Nature* **349**:319–321.
10. Dilworth, M. J. 1966. Acetylene reduction by nitrogen-fixing preparations from *Clostridium pasteurianum*. *Biochem. Biophys. Acta* **127**:285–294.
11. Herbert, R. A. 1999. Nitrogen cycling in coastal marine ecosystems. *FEMS Microbiol. Rev.* **23**:563–590.
12. Howarth, R. W., R. Marino, and J. Lane. 1988. Nitrogen fixation in freshwater, estuarine and marine ecosystems. I. Rates and importance. *Limnol. Oceanogr.* **33**:669–687.
13. Jensen, K. M., M. H. Jensen, and R. P. Cox. 1996. Membrane inlet mass spectrometric analysis of N-isotope labeling for aquatic denitrification studies. *FEMS Microbiol. Ecol.* **20**:101–109.
14. Joye, S. B., and H. W. Paerl. 1994. Nitrogen cycling in microbial mats: rates and patterns of denitrification and nitrogen fixation. *Mar. Biol.* **119**:285–295.
15. Kana, T. M., C. Darkangelo, M. D. Hunt, J. B. Oldham, G. E. Bennett, and J. C. Cornwell. 1994. Membrane inlet mass spectrometer for rapid high-precision determination of N₂, O₂, and Ar in environmental water samples. *Anal. Chem.* **66**:4166–4170.
16. Kana, T. M., M. B. Sullivan, J. C. Cornwell, and K. Groszkowski. 1998. Denitrification in estuarine sediments determined by membrane inlet mass spectrometry. *Limnol. Oceanogr.* **43**:334–339.
17. Kasper, H. F. 1983. Denitrification, nitrate reduction to ammonium and inorganic nitrogen pools in intertidal sediments. *Mar. Biol.* **74**:133–139.
18. Koike, I., and A. Hattori. 1978. Denitrification and ammonia formation in anaerobic coastal sediments. *Appl. Environ. Microbiol.* **35**:278–282.
19. Koike, I., and J. Sørensen. 1988. Nitrate reduction and denitrification in marine sediments. In T. H. Blackburn and J. Sørensen (ed.), *Nitrogen cycling in coastal marine environments*. Wiley, New York, N.Y.
20. Lamontagne, G. M., and I. Valiela. 1995. Denitrification measurement by a direct N₂ flux method in sediment of Waquoit Bay, MA. *Biogeochemistry* **31**:63–83.
21. Lavrentyev, P., W. S. Gardner, and L. Yang. 2000. Effects of the zebra mussel on microbial composition and nitrogen dynamics at the sediment-water interface in Saginaw Bay, Lake Huron. *Aquat. Microb. Ecol.* **21**:187–194.
22. Lee, K., and K. H. Dunton. 1999. Inorganic nitrogen acquisition in the seagrass *Thalassia testudinum*: development of a whole-plant nitrogen budget. *Limnol. Oceanogr.* **44**:1204–1215.
23. Lide, D. R. 1992. *CRC handbook of chemistry and physics*. CRC Press Inc., Boca Raton, Fla.
24. Luijn, V. F., P. C. Boers, and L. Lijklema. 1996. Comparison of denitrification rates in lake sediments obtained by the N₂ flux method, the ¹⁵N isotope pairing technique and the mass balance approach. *Water Res.* **30**:893–900.
25. McGlathery, K. J., N. Risgaard-Petersen, and P. B. Christensen. 1998. Temporal and spatial variation in nitrogen fixation activity in the eelgrass *Zostera marina* rhizosphere. *Mar. Ecol. Prog. Ser.* **168**:245–258.
26. Middelburg, J. J., K. Soetaert, and P. M. J. Herman. 1996. Evaluation of the nitrogen isotope-pairing method for measuring benthic denitrification: a simulation analysis. *Limnol. Oceanogr.* **41**:1839–1844.
27. Nedwell, D., and S. Aziz. 1980. Heterotrophic nitrogen fixation in an intertidal salt marsh sediment. *Estuar. Coast. Mar. Sci.* **10**:699–702.
28. Nielson, L. P., and R. N. Glud. 1996. Denitrification in a coastal sediment measured *in situ* by the nitrogen isotope pairing technique applied to a benthic flux chamber. *Mar. Ecol. Prog. Ser.* **173**:181–186.
29. Nielson, L. P. 1992. Denitrification in sediment determined from nitrogen isotope pairing. *FEMS Microb. Ecol.* **86**:357–362.
30. Nowicki, B. L. 1994. The effect of temperature, oxygen, salinity, and nutrient enrichment on estuarine denitrification rates measured with a modified nitrogen gas flux technique. *Estuar. Coast. Shelf Sci.* **38**:137–156.
31. Pinckney, J. L., and R. G. Zingmark. 1993. Modeling the annual production of intertidal benthic microalgae in estuarine ecosystems. *J. Phycol.* **29**:396–407.
32. Risgaard-Petersen, N., L. P. Nielsen, and T. H. Blackburn. 1998. Simultaneous measurement of benthic denitrification, with the isotope pairing technique and the N₂ flux method in a continuous flowthrough system. *Water Res.* **32**:3371–3377.
33. Risgaard-Petersen, N., S. Rysgaard, L. P. Nielsen, and N. P. Revsbech. 1994. Diurnal variation of denitrification and nitrification in sediments colonized by benthic microphytes. *Limnol. Oceanogr.* **39**:573–579.
34. Rysgaard, S., P. B. Christensen, and L. P. Nielsen. 1995. Seasonal variation in nitrification and denitrification in estuarine sediment colonized by benthic microalgae and bioturbating infauna. *Mar. Ecol. Prog. Ser.* **126**:111–121.
35. Ryther, J. M., and W. M. Dunston. 1971. Nitrogen, phosphorus and eutrophication in the coastal marine environment. *Science* **171**:1008–1013.
36. Seitzinger, S. P. 1988. Denitrification in freshwater and coastal marine ecosystem: ecological and geochemical significance. *Limnol. Oceanogr.* **33**:702–724.
37. Seitzinger, S. P. 1990. Denitrification in aquatic sediments. *FEMS Symp.* **56**:301–322.
38. Seitzinger, S. P., and J. H. Garber. 1987. Nitrogen fixation and ¹⁵N₂ calibration of the acetylene reduction assay in coastal marine sediments. *Mar. Ecol. Prog. Ser.* **37**:65–73.
39. Seitzinger, S. P., S. W. Nixon, M. E. Q. Pilson, and S. Burke. 1980. Denitrification and nitrous oxide production in near shore marine sediments. *Geochim. Cosmochim. Acta* **44**:1853–1860.
40. Sørensen, J. 1978. Capacity for denitrification and reduction of nitrate to ammonia in a coastal marine sediment. *Appl. Environ. Microbiol.* **35**:301–305.
41. Sørensen, J. 1978. Occurrence of nitric and nitrous oxides in a coastal marine sediment. *Appl. Environ. Microbiol.* **36**:809–813.
42. Stal, L. J. 1995. Physiological ecology of cyanobacteria in microbial mats and other communities. *New Phytol.* **131**:1–32.
43. Stal, L. J., S. B. Behrens, M. Villbrandt, S. V. Bergekl, and F. Kruyning. 1996. The biogeochemistry of two eutrophic marine lagoons and its effect on microphytobenthic communities. *Hydrobiologia* **329**:185–198. Equations are as follows: