





# Root hydraulic phenotypes impacting water uptake in drying soils

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## Abstract

Soil drying is a limiting factor for crop production worldwide. Yet, it is not clear how soil drying impacts water uptake across different soils, species, and root phenotypes. Here we ask (1) what root phenotypes improve the water use from drying soils? and (2) what root hydraulic properties impact water flow across the soil–plant continuum? The main objective is to propose a hydraulic framework to investigate the interplay between soil and root hydraulic properties on water uptake. We collected highly resolved data on transpiration, leaf and soil water potential across 11 crops and 10 contrasting soil textures. In drying soils, the drop in water potential at the soil–root interface resulted in a rapid decrease in soil hydraulic conductance, especially at higher transpiration rates. The analysis reveals that water uptake was limited by soil within a wide range of soil water potential (–6 to –1000 kPa), depending on both soil textures and root hydraulic phenotypes. We propose that a root phenotype with low root hydraulic conductance, long roots and/or long and dense root hairs postpones soil limitation in drying soils. The consequence of these root phenotypes on crop water use is discussed.

## KEYWORDS

drought, leaf water potential, root hairs, root hydraulic conductance, root length, root water uptake, soil hydraulic conductivity, soil texture

## 1 | INTRODUCTION

Drought is an unavoidable natural hazard affects the entire ecosystem, especially agriculture (Boyer, 1982; Wilhite, 2000). Under drought conditions, plants reduce stomatal conductance to avoid an excessive decline in leaf water potential. Stomatal regulation determines the water flux across the soil–plant–atmosphere system. The mechanisms of stomatal response to water stress from the molecular and physiological aspects were intensively investigated (Buckley, 2005, 2019). Water use efficiency is positively affected by stomatal closure (Lawson & Vialet-Chabrand,

2019; Yang et al., 2021), and phenotypes with high water use efficiency have been proposed for improving drought tolerance (Tracy et al., 2020). Although stomatal conductance is a key aboveground hydraulic variable regulating water use by crops, it does not mean that belowground hydraulics have no role in it. Indeed, changes in soil–plant hydraulic conductance have been shown to drive stomatal closure (Abdalla, Carminati, et al., 2021). It has been proposed that stomatal regulation is intimately linked to a decline in soil–plant hydraulics (Carminati & Javaux, 2020; Sperry & Love, 2015). This link allows one to predict the water use based on the decline in hydraulic conductivities of key components of soil and

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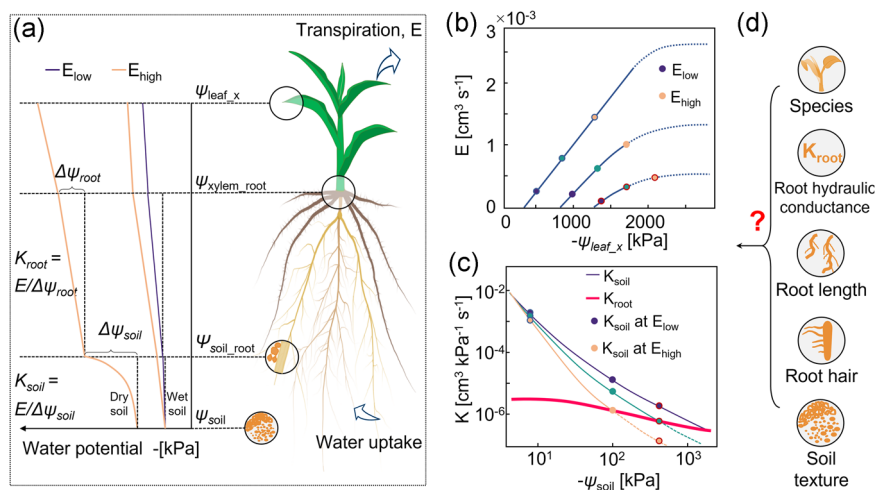
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root hydraulics (Sperry & Love, 2015). For trees, much of the attention has been placed on xylem embolism (Cochard, 2002; Savi et al., 2015; Sperry et al., 2002; Venturas et al., 2017). However, recent studies have shown that, both in trees (Albuquerque et al., 2020; Anderegg et al., 2017) and crops (Corso et al., 2020), stomatal closure before xylem embolism takes place. For crops, Vadez (2014) proposed in an inspiring paper that root hydraulics is the key to improve drought tolerance. He proposed that rather than root length, root hydraulics might determine the water use of crops. However, experimental evidence linking root hydraulics to stomatal regulation is lacking. In this context, the following questions are still open: what root hydraulic properties impact the ability of roots to extract water from drying soil? And what is the hydraulic limit across the soil–plant–atmosphere continuum? Is it the soil (as proposed by Carminati & Javaux, 2020), the root–soil interface (as measured for olive trees by Rodriguez-Dominguez & Brodribb, 2020), the root (Bourbia et al., 2021), an increase in soil–root resistance (Abdalla, Carminati, et al., 2021) and/or the outer xylem tissues (in roots [Cuneo et al., 2021, 2016] or leaves [Scoffoni & Sack, 2017])? In this review, we presented an analysis of soil–plant hydraulic measurements across species and soils and proposed a link between root hydraulic phenotypes and water use during drought.

Water flow along the soil–plant continuum is determined by a series of hydraulic conductivities and gradients in water potential (Figure 1a). In wet conditions, the water potential gradient at the soil–root interface is minimum, and root hydraulic conductance ( $K_{root}$ ,  $cm^3 kPa^{-1} s^{-1}$ ) at a given transpiration rate determines the difference in water potential between soil and leaf (Nobel & Cui, 1992; Sperry et al., 1998). The root hydraulic conductance depends on different

root traits, for example, root architecture (Doussan et al., 2006), root length (Kato & Okami, 2011), root anatomy (e.g., metaxylem vessel size, [Strock et al., 2021]), root hairs (Carminati et al., 2017) and aquaporin expression (Grondin et al., 2020; McLean et al., 2011). During soil drying, the root hydraulic conductivity ( $k_{root}$ ,  $K_{root}$  normalized by root surface area) might decrease due to lacunae formation in fine root cortical cells (Cuneo et al., 2016) and downregulation of aquaporin (AQP) activity (Caldeira et al., 2014; Rodríguez-Gamir et al., 2019). Similarly, soil hydraulic conductivity decreases and the water potential gradient around the roots increases markedly, particularly at high transpiration rates (Figure 1a) (Gardner, 1960). Additionally, root shrinkage (Carminati et al., 2009; Nobel & Cui, 1992) further aggravates the drop in water potential in drying soil. As a result, root water uptake might be constrained at the soil–root interface due to poor soil–root contact, especially in well-structured soils, as speculated by White and Kirkegaard (2010).

Direct measurements of the hydraulic conductance of the soil–plant continuum and its elements are challenging. In their literature review, Draye et al. (2010) compared root and soil conductivity across tree species and soils. Despite the importance of such information, their estimation is based on individual excised roots, including roots grown in hydroponics, and the assumption of soil water flow toward the root. In this way, it is not possible to investigate to what extent root hydraulic properties change with soil texture. Therefore, there is the need to collect information across species and soils on intact root systems growing in soils. Using the root pressure chamber (Passioura, 1980), root hydraulic conductance and the drop in water potential at the soil–root interface could be



**FIGURE 1** Schematic of water flow in the soil–plant continuum and corresponding variation of root and soil hydraulic properties during soil drying. (a) Water potential gradients from soil to leaf at low and high transpiration rates ( $E$ ) in wet and dry conditions.  $\psi_{soil}$ ,  $\psi_{soil\_root}$ ,  $\psi_{xylem\_root}$ , and  $\psi_{leaf\_x}$  are the water potentials in the soil, at the soil–root interface, root xylem and leaf xylem, respectively.  $\Delta\psi_{soil}$  and  $\Delta\psi_{root}$  are potential differences between soil and soil–root interface, and between soil–root interface and root xylem.  $K_{soil}$  and  $K_{root}$  are the soil and root hydraulic conductance. (b) Relation between  $E$  and  $\psi_{leaf\_x}$  in drying soil (for a hypothetical crop ca. 1-month-old). Dashed lines are from soil hydraulic nonlinearity. (c) Hypothetical relation between  $K_{soil}$  and  $K_{root}$  at different transpiration rates in drying soil. Declines in  $K_{soil}$  and  $K_{root}$  during soil drying cause nonlinearity in the relation between  $E$  and  $\psi_{leaf\_x}$ . Circle symbols with blue and red edges are for wet and dry conditions. (d) In this study, we explored how soil textures and root hydraulic phenotypes (e.g., root hydraulic conductance, root length, and root hairs from different species) impact root water uptake in drying soils

identified in intact plants during soil drying. The method provided accurate measurements of the relationship between transpiration rate ( $E$ ) and leaf xylem water potential ( $\psi_{\text{leaf}_x}$ ) during soil drying (Abdalla, Carminati, et al., 2021; Cai, Ahmed, Reth, et al., 2020; Cai et al., 2021; Carminati et al., 2017). Root hydraulic conductance could be obtained from transpiration rate and water potential gradient along the pathway between soil and leaf. The  $E(\psi_{\text{leaf}_x})$  relation is typically linear in wet soil and becomes nonlinear as the soil dries (Figure 1b), especially at higher transpiration rates, mainly due to the increasing gradient in water potential at the soil–root interface (Carminati et al., 2017; Carminati & Javaux, 2020). The method was successfully used to investigate the response of various species to soil drying (Abdalla et al., 2021b; Cai, Ahmed, Dippold, et al., 2020; Cai et al., 2021; Carminati et al., 2017; Hayat et al., 2020; Passioura, 1980). However, a direct comparison across species and soils has never been done.

In this study, the main objective was to propose a hydraulic framework to investigate the interplay between soil and root hydraulic properties on water uptake. To this end, we investigated how soil and root hydraulic properties determine the decline in soil water availability to crops. We compared the decline in soil and root hydraulic conductances for varying species and soils to find out at what soil water potential, one of these elements started to limit root water uptake. Finally, we discussed how given root hydraulic phenotypes impacted water use by crops in different soil textures.

## 2 | OVERVIEW OF DATA COLLECTION AND PROCESS

The selected studies used the same method (plant pressurization) and included  $E(\psi_{\text{leaf}_x})$  measurements from eight cereal and three Solanaceae varieties, covering different root traits (e.g., root hairs, root length) and soil textures (from sand to loam). The corresponding plant and soil information is summarized in Table 1. The plants were 3- to 6-week old with differences in shoot (leaf) and root growth. Soil hydraulic parameters were inversely fitted using the Brooks and Corey (1964) model. The water retention and hydraulic conductivity curves are shown in Figure 2. The soil hydraulic properties differed markedly even though some soils were in the same texture category, for example, sandy loam for maize showed a steeper slope than sandy loam for barley, millet and tomato. Taken together, this allows doing a systematic analysis of the interactive impact of soil and root characteristics on plant response to water stress.

The  $E(\psi_{\text{leaf}_x})$  relations were reproduced for individual soil moisture using a soil–plant hydraulic model (Carminati & Javaux, 2020) by adjusting root hydraulic conductance and root length active in water uptake (Abdalla, Carminati, et al., 2021; Cai, Ahmed, Dippold, et al., 2020, 2021). In the model, water flow from soil towards root, xylem and leaf is driven by the gradients in water potential and regulated by soil hydraulic conductivity ( $k_s$ ,  $\text{cm s}^{-1}$ ) and plant hydraulic conductance (including root and aboveground xylem). The detailed description of this model is in Carminati and Javaux (2020) and the Methods S1. Briefly,

a uniform water flux ( $q$ ,  $\text{cm s}^{-1}$ ) is assumed to flow into a fraction of total root length from surrounding soil:

$$q(r_0) = E / (2\pi r_0 L_{\text{act}}), \quad (1)$$

where  $r_0$  (cm) is the root radius, and  $L_{\text{act}}$  (cm) is a model parameter and is defined as the root length active in water uptake. No flow is assumed out of the casing soil,  $q(r_b) = 0$ , where  $r_b$  (cm) is the radius of the casing soil and calculated from soil volume ( $V$ ,  $\text{cm}^3$ ),  $r_b = (V/\pi L_{\text{act}})^{0.5}$ . The water flux in soil  $q$  depends on  $k_s$  and the gradient in soil matric potential ( $\psi_{\text{soil}}$ , kPa). The nonlinear equations can be solved analytically (for simplified parameterization and boundary conditions) or numerically yield the gradient in soil matric potential around roots (Gardner, 1960; Passioura, 1980). The overall soil hydraulic conductance ( $K_{\text{soil}}$ ,  $\text{cm}^3 \text{kPa}^{-1} \text{s}^{-1}$ ) is defined as:

$$E = K_{\text{soil}}(\psi_{\text{soil}} - \psi_{\text{soil\_root}}), \quad (2)$$

where  $\psi_{\text{soil\_root}}$  (kPa) is the matric potential at the soil–root interface.

The water flow in the root system is described by:

$$E = K_{\text{root}}(\psi_{\text{soil\_root}} - \psi_{\text{xylem\_root}}), \quad (3)$$

where  $K_{\text{root}}$  ( $\text{cm}^3 \text{kPa}^{-1} \text{s}^{-1}$ ) is the root hydraulic conductance,  $\psi_{\text{xylem\_root}}$  (kPa) is the matric potential at the root collar. The water flow in the xylem is defined as:

$$E = K_x(\psi_{\text{xylem\_root}} - \psi_{\text{leaf}_x}), \quad (4)$$

where  $K_x$  ( $\text{cm}^3 \text{kPa}^{-1} \text{s}^{-1}$ ) is the aboveground xylem conductance, a function of  $K_{\text{root}}$  and  $\psi_{\text{soil}}$ .  $\psi_{\text{leaf}_x}$  (kPa) is the water potential in leaf xylem. The detailed derivation of  $\psi_{\text{soil\_root}}$ ,  $\psi_{\text{xylem\_root}}$ ,  $K_x$ , and  $\psi_{\text{leaf}_x}$  is described in Methods S1. The measured and modelled parameters are described in the appendix.

## 3 | ROOT HYDRAULIC PHENOTYPES AND WATER UPTAKE

Root phenotypes differed not only between species but also within the same species grown in different soil types (Table 1). In the following sections, we discuss how the interactions between root phenotypes and soil textures impact key soil and root hydraulic properties and thus the water uptake process.

### 3.1 | Variation in root hydraulic conductance across species and soils

The maximum root hydraulic conductance,  $K_{\text{root\_max}}$ , is defined as the slope of  $E(\psi_{\text{leaf}_x})$  in wet soils.  $K_{\text{root\_max}}$  was soil texture independent and hence was positively related to root length across species and textures (Figure 3a), which was in line with Maurel et al. (2010) and Judd et al. (2016). For cereals,  $K_{\text{root\_max}}$  of maize in sand and loam, barley in sandy

**TABLE 1** Plants and soil textures used in this study

Mono-/dicot (M/D)	Family	C3/4	Root hairs	Plant age (day)	Leaf area (cm <sup>2</sup> )	Mean root length (cm)	Mean root length density (cm <sup>-3</sup> )	r <sub>0</sub> (cm)	Soil texture (bulk density, g cm <sup>-3</sup> )	References <sup>a</sup>
Wheat M	Poaceae	C3	+	21–30	30–190	1531	5.2	0.015	Loam (1.27)	Deery et al. (2013),
						9294	8	0.014	Clay loam (Un <sup>a</sup> , 1.6)	Deery (2008),
						6971	6	0.016	Clay loam (Re <sup>b</sup> , 1.3)	Passioura (1980)
						3776	3.25	0.015	Sand (Re <sup>b</sup> , 1.6)	
Barley M	Poaceae	C3	±	21–28	/	3660	3.35	0.016 (WT)	Sandy loam (0.93)	Carminati et al. (2017)
						3310	3.71	0.013 (Mut)	Red Chromosol <sup>c</sup> (0.9)	
Maize M	Poaceae	C4	+	40–50	/250–350	28 000	21.1	0.06	Sandy loam (1.4)	Hayat et al. (2020)
				26–36		2461 (WT)	2.91 (WT)	0.019 (WT, Mut)	Sand (1.47)	Cai et al. (2021)
				26–36		2865 (Mut)	3.38 (Mut)	0.012 (WT, Mut)	Loam (1.26)	
						5414 (WT)	6.39 (WT)			
		3764 (Mut)	4.44 (Mut)							
Millet M	Poaceae	C4	+	30–45	250–380	27 617	13.5	0.04	Sandy loam (1.3)	Cai, Ahmed, Dippold, et al. (2020)
Tomato D	Solanaceae	C3	+	22–30	846–1116	3700 (long root) <sup>c</sup>	1.81	0.05	Sandy loam (1.3)	Abdalla, Ahmed, et al. (2021)
				22–30	943–1145	1700 (short root)	0.83			Abdalla, Carminati, et al. (2021)
				30–40	960–2016	7540 (normal)	3.68			

Abbreviations: Mut, mutant; WT, wild type.

<sup>a</sup>The data of wheat were extracted from the literature using the software WebPlotDigitizer (Rohatgi, 2020). The other data were obtained from the authors directly.

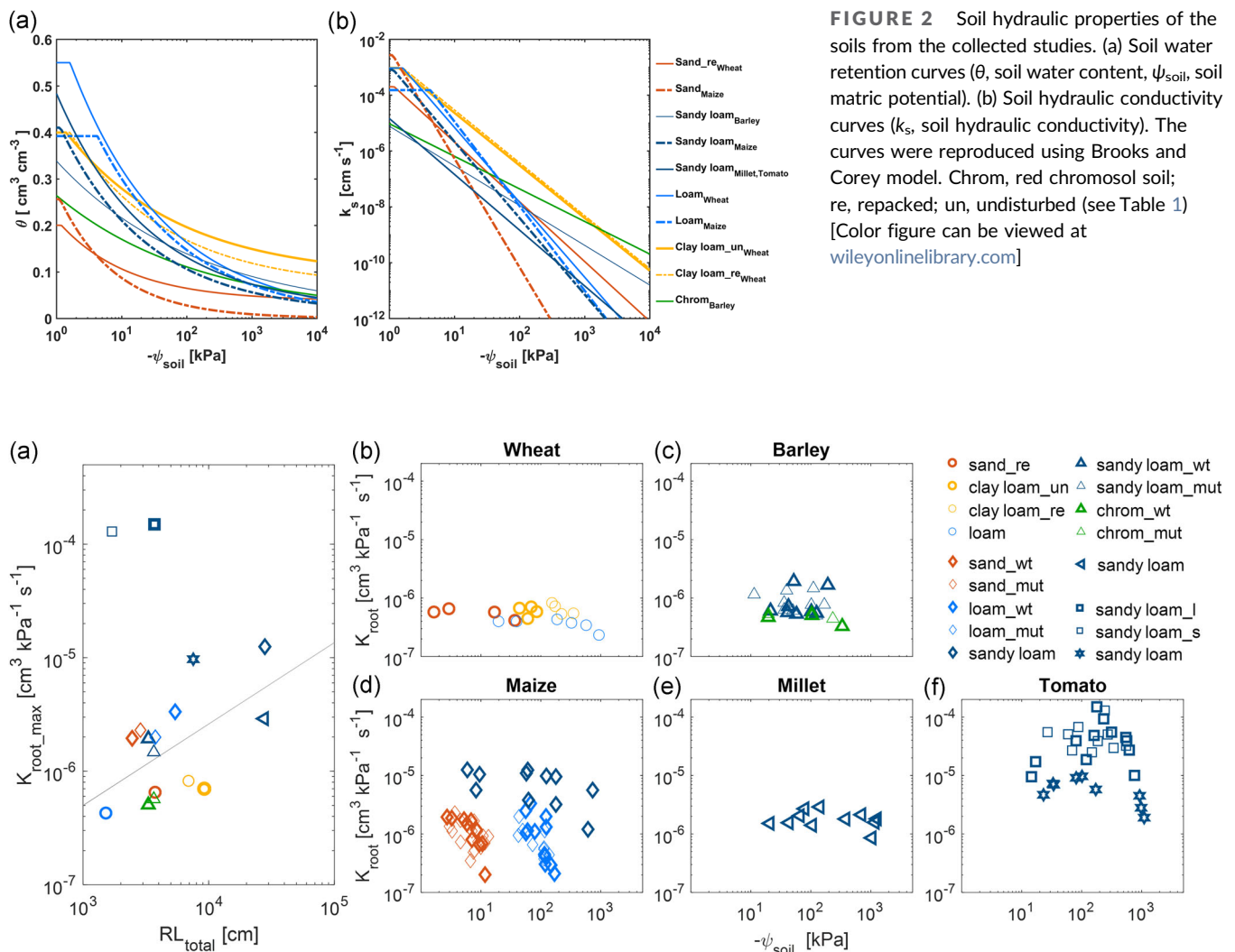
<sup>b</sup>Un, undisturbed soil; Re, repacked soil.

<sup>c</sup>Red chromosol is from southeast Australia and is characterized by a strong contrasting texture (including aggregates and macropores) with low bulk density (ca. 0.9 g cm<sup>-3</sup>).

<sup>d</sup>‘Long root’, ‘short root’ and ‘normal’ were different phenotypes. Tomato with a long and short root system was grafted with an identical shoot.

loam, and millet was similar. The value was around  $10^{-6} \text{ cm}^3 \text{ kPa}^{-1} \text{ s}^{-1}$ , which was four times higher than wheat and barley in red chromosol (green triangle in Figure 3c) but one-fifth of the older maize with longer roots (Table 1). The magnitude of  $K_{\text{root,max}}$  was in line with plants at similar ages using other methods (e.g., measuring exudate from immersed roots via pressurization), for example, wheat by Zhao et al. (2005), maize by Sunita et al. (2014) and barley by Kodama et al. (2021). Tomato with long and short root systems (grafted with an identical shoot) had surprisingly the highest  $K_{\text{root,max}}$  (around  $10^{-4} \text{ cm}^3 \text{ kPa}^{-1} \text{ s}^{-1}$ ), which was similar to the tomato at a similar age in commercial soil (Tsuda & Tyree, 2000), and ten times higher than the genotype without grafting (hexagram symbol).

$K_{\text{root}}$  declined as the soil dried and the decline differed between species and textures (Figure 3b–f). Comparing textures,  $K_{\text{root}}$  decreased at less negative soil water potentials in sand ( $\psi_{\text{soil}} > -10 \text{ kPa}$ ) compared to loam ( $\psi_{\text{soil}} > -60 \text{ kPa}$ ) (Figure 3d). This is expected considering the sharp drop in hydraulic conductivity in sand (Figure 2), which would cause a steeper gradient in soil water potential around roots. Similarly, a great root length reduces the decline in water potential gradient at the soil–root interface and softens the reduction in  $K_{\text{root}}$ , for example, maize in loam (Figure 3d, dark blue diamond), and tomato with long roots (Figure 3f, thick square). Compared with other species, the decrease of  $K_{\text{root}}$  was steeper in maize and tomato.



## 3.2 | Water potential gradient at the soil–root interface

Gradients in matric potential occur between the bulk soil and the soil–root interface once plants start to transpire. During soil drying, how quickly and how far  $\psi_{\text{soil\_root}}$  deviates from  $\psi_{\text{soil}}$  with increasing transpiration depends on soil textures and root hydraulic phenotypes.

### 3.2.1 | Soil texture

$\psi_{\text{soil\_root}}$  deviates at less negative  $\psi_{\text{soil}}$  in sand than in loam because of the sharp decrease in hydraulic conductivity of sand. This is visible in Figure 4a,i,j, which shows that, in sand,  $\psi_{\text{soil\_root}}$  deviated remarkably from the 1:1 line. The comparison of  $\psi_{\text{soil\_root}}$  between sand and loam for different crops is shown in Figure 5a,b. These results are consistent with the findings of Dodd et al. (2010). Additionally, roots might shrink and form gaps between roots and the surrounding soils earlier in sand (Koebernick et al., 2018), resulting in a steeper decline in conductivity at the soil–root interface (North & Nobel, 1997) and thus more negative  $\psi_{\text{soil\_root}}$ .

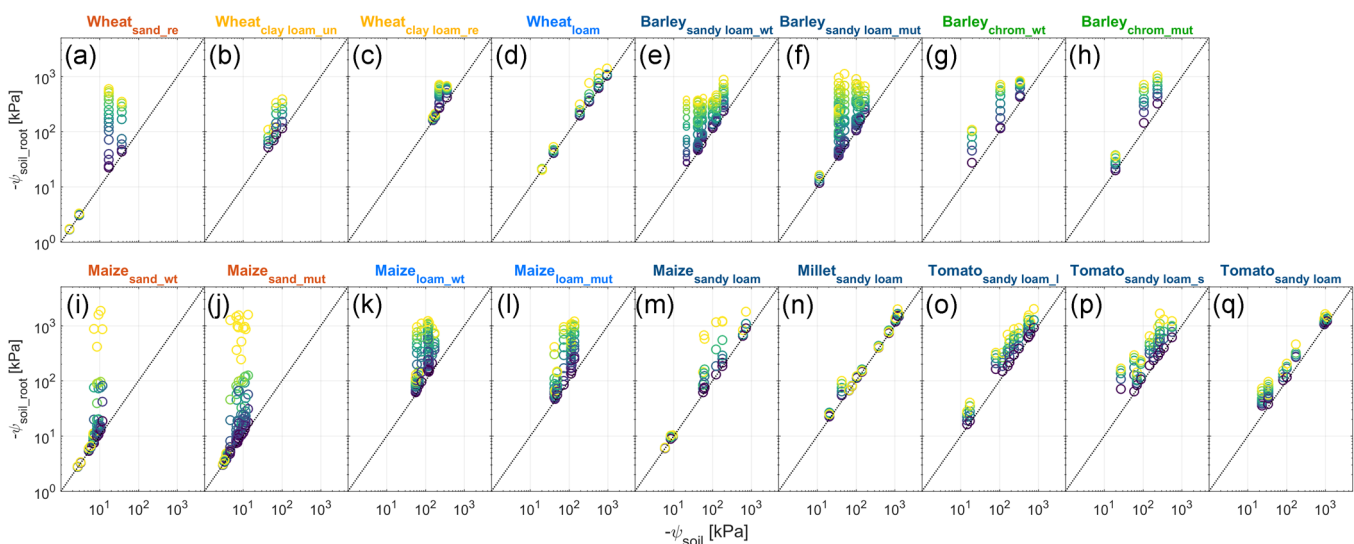
### 3.2.2 | Root length

Concerning root length (and assuming a similar root diameter), longer roots (thus larger root surface) result in a lower water flow velocity in soil and allow for sustaining higher transpiration rates in drying soils (Faiz & Weatherley, 1982; Taylor & Willatt, 1983), which lead to a

less negative  $\psi_{\text{soil\_root}}$  (Xu, 2001). This could explain the nearly 1:1 relation between  $\psi_{\text{soil\_root}}$  and  $\psi_{\text{soil}}$  in millet (Figure 4n) and tomato (Figure 4q), since their root length was much longer than other varieties (Figure 4e,f,o,p). Note that these plants (Figure 4n–q) were grown in the same soil. It also explains the slightly less negative  $\psi_{\text{soil\_root}}$  of the tomato with long roots (Figure 4o) than with short roots (Figure 4p). However, it is neither the case for barley (Figure 4e,f) and tomato with contrasting root systems (Figure 4o,p) as these plants had a similar root length (Table 1), and nor the case for wheat (Figure 4d) and maize in loam (Figure 4i–k) as maize even had longer roots. This apparent contradiction could be explained by the differences in transpiration rate and root hydraulic conductance.

### 3.2.3 | Species

The transpiration rate of maize in loam was around two to four times higher than that of wheat at similar soil matric potential and leaf water potential (Cai et al., 2021; Passioura, 1980). Additionally, the difference in soil hydraulic conductivity curves of the two soils is very close (Figure 2). Therefore, maize in loam showed more negative  $\psi_{\text{soil\_root}}$ . For tomato and barley in sandy loam, it is not the case. Although  $K_{\text{root}}$  and transpiration of tomato were both higher than that of barley (Figure 3; Abdalla, Carminati, et al., 2021; Carminati et al., 2017), the decline in soil hydraulic conductivity of the loamy soil for tomato was steeper compared to that for barley (Figure 2). Lower  $\psi_{\text{soil\_root}}$  of barley in sandy loam indicates that the role of soil hydraulic conductivity is rather dominating the variations in  $\psi_{\text{soil\_root}}$ . The results also indicate that, during soil drying, variations in  $\psi_{\text{soil\_root}}$  with increasing transpiration are not determined by a single root



**FIGURE 4** Variation of water potential at the soil–root interface ( $\psi_{\text{soil\_root}}$ ) with increasing transpiration rates (from dark blue to yellow, note that same colour does not mean same transpiration rate) as a function of soil matric potential ( $\psi_{\text{soil}}$ ). (a,b,c,d) Wheat grown in repacked sand, undisturbed clay loam, repacked clay loam, and loam. (e,f) Wildtype and mutant (*roothairless brb*) barley grown in sandy loam. (g,h) Same barley genotypes as (e,f) grown in chromosol soil. (i,j) Wildtype and mutant (*roothairless 3*) grown in sand. (k,l) Same maize genotypes as (i,j) grown in loam. (m) Maize grown in sandy loam. (n) Millet grown in sandy loam. (o,p) Tomato with long and short root length grown in sandy loam, respectively. (q) Tomato grown in sandy loam. chrom, red chromosol soil; l, long root system; mut, mutant; re, repacked; s, short root system; un, undisturbed; wt, wild type. The same colour is used for similar soil textures [Color figure can be viewed at [wileyonlinelibrary.com](https://onlinelibrary.wiley.com)]

hydraulic parameter but are impacted by both soil and root hydraulic properties.

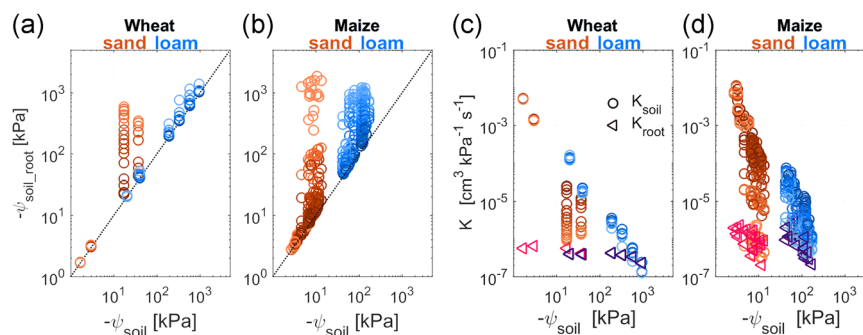
### 3.2.4 | Root hairs

Root hairs were hypothesized to soften the drop in soil matric potential around the roots (Ahmed, Passioura, et al., 2018; Cai et al., 2021; Carminati et al., 2017). Less negative  $\psi_{\text{soil\_root}}$  manifested in wildtype barley (Figure 4e) compared with mutant in sandy loam, which was, however, not observed in the wild type of maize either in sand or in loam (Figure 4i-l). This might be due to differences in root hair length between barley and maize (Cai et al., 2021). Indeed, recent studies showed that

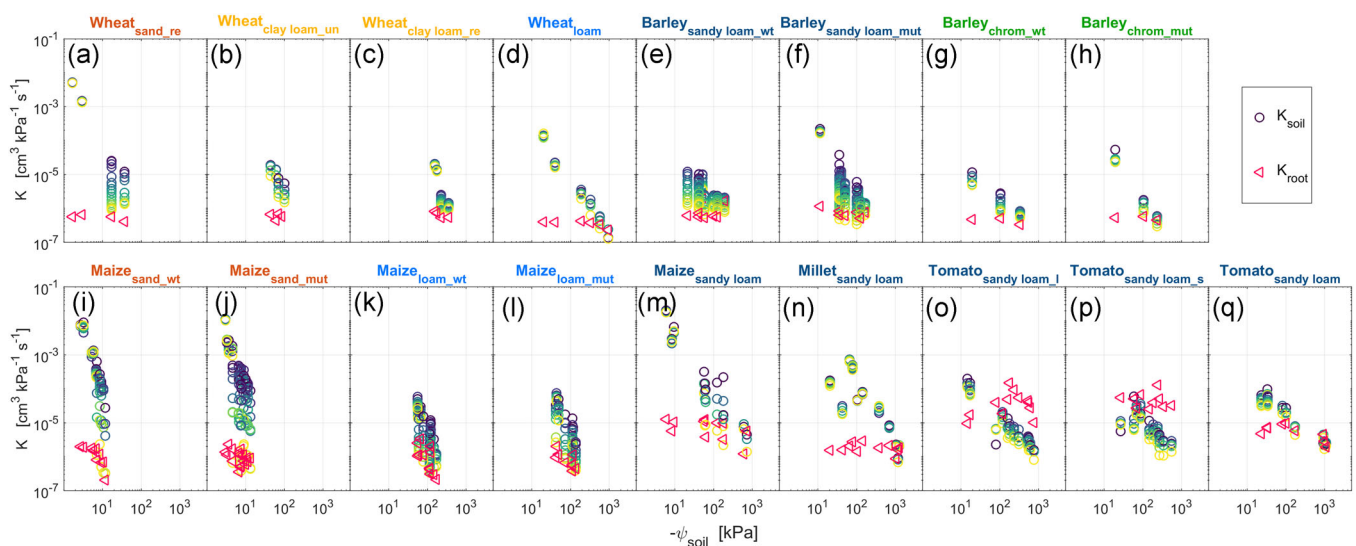
root hairs of barley were denser and around as two times long as that of maize both in lab and field measurements (Burak et al., 2021; Cai et al., 2021; Carminati et al., 2017; Marin et al., 2021). Furthermore, Burak et al. (2021) showed that, in barley, denser and longer root hairs formed thicker rhizosheath than that of maize, which probably amplified intimate contact between root and surrounding soils.

### 3.3 | Comparison between root and soil hydraulic conductances

Figures 5c,d and 6 show the comparison between  $K_{\text{soil}}$  and  $K_{\text{root}}$  across species and textures during soil drying. In wet conditions,  $K_{\text{soil}}$



**FIGURE 5** Comparison of water potential at the soil–root interface and of soil and root hydraulic conductance between sand and loam. (a,b) Water potential gradient at the soil–root interface ( $\psi_{\text{soil\_root}}$ ) with increasing transpiration rate and soil ( $\psi_{\text{soil}}$ ) in sand and loam for both wheat and maize (from Figure 4a,d,i-l). (c,d) Difference between soil hydraulic conductance ( $K_{\text{soil}}$ ) with increasing transpiration rate and root hydraulic conductance ( $K_{\text{root}}$ ) in sand and loam for both wheat and maize [Color figure can be viewed at [wileyonlinelibrary.com](https://onlinelibrary.wiley.com)]



**FIGURE 6** Variation of soil hydraulic conductance ( $K_{\text{soil}}$ ) and root hydraulic conductance ( $K_{\text{root}}$ ) for varying transpiration rates (from dark blue to yellow, note that same colour does not mean same transpiration rate) as a function of soil matric potential  $\psi_{\text{soil}}$ : soil matric potential. (a,b,c,d) Wheat grown in repacked sand, undisturbed clay loam, repacked clay loam, and loam. (e,f) Wildtype and mutant (*root-hairless brb*) barley grown in sandy loam. (g,h) Same barley genotypes as (e,f) grown in chromosol soil. (i,j) Wildtype and mutant (*root-hairless 3*) grown in sand. (k,l) Same maize genotypes as (i,j) grown in loam. (m) Maize grown in sandy loam. (n) Millet grown in sandy loam. (o,p) Tomato with long and short root length grown in sandy loam, respectively. (q) Tomato grown in sandy loam. The same colour is used for the titles of similar soil textures [Color figure can be viewed at [wileyonlinelibrary.com](https://onlinelibrary.wiley.com)]

was maintained at a high level regardless of the variations of transpiration rate.  $K_{soil}$  was around four orders of magnitude higher than  $K_{root}$  in sand (Figure 6a,i,j) and one order of magnitude higher than or even close  $K_{root}$  in sandy loam for tomato (Figure 6o-q). As the soil dried,  $K_{soil}$  decreased more rapidly than  $K_{root}$ . The decline of  $K_{soil}$  was sharper in sand than in other soils, and more marked at higher transpiration rates (see the comparison in Figure 5c,d). As transpiration rate increased,  $K_{soil}$  decreased by one to two orders of magnitude at a specific  $\psi_{soil}$  (e.g., Figure 6a,b,e,f,i,l) due to a remarkable decline in  $\psi_{soil\_root}$  (Figure 4). Though with different reduction patterns, decreasing  $K_{soil}$  with increasing transpiration approached dropping  $K_{root}$  in all drying soils. Afterwards,  $K_{soil}$  became even much lower than  $K_{root}$ , especially in tomato (Figure 6o,p). The effect of root length on  $\psi_{soil\_root}$  (Figure 4) was also reflected on  $K_{soil}$ , for example, the drop in  $K_{soil}$  with increasing transpiration was less abrupt in plants with longer roots (Figure 6k,l,n,q).

### 3.4 | Soil water limitation as a function of soil and plant hydraulic traits

We define  $\psi_{soil}^c$  as the point at which soil hydraulic conductance drops to lower values than the root hydraulic conductance. This is the point when the soil starts to limit the water fluxes in the soil-plant continuum relative to other plant tissues. This point corresponds to the onset of nonlinearity in the  $E(\psi_{leaf\_x})$  relation (Figure 7), similarly to the definition given in Carminati and Javaux (2020). In the next sections, we explored how soil and root hydraulic traits impact  $\psi_{soil}^c$ .

We fitted the decreases of  $K_{root}$  and  $K_{soil}$  at the highest transpiration rate and obtained the soil matrix potential at the intersection ( $\psi_{soil}^c$ ) of the two fittings (Figure 7). There was no unique  $\psi_{soil}^c$  at which transpiration of different species was limited by soil.  $\psi_{soil}^c$  was affected not only by soil hydraulic but also by root hydraulic parameters (Figure 7b). Compared with other soils, transpiration in sand was constrained earlier, with  $\psi_{soil}^c$  around -10 to -20 kPa, when the

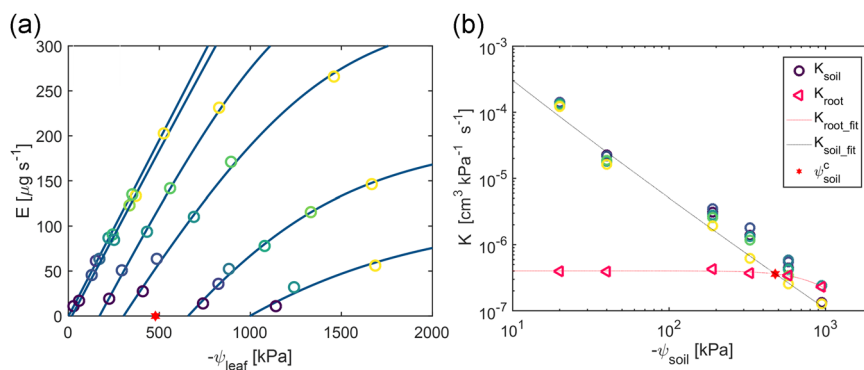
soil hydraulic conductivity was still high (Figure 2). In loamy soils,  $\psi_{soil}^c$  was around -100 to -500 kPa.

#### 3.4.1 | Soil texture

To evaluate the impact of soil properties on  $\psi_{soil}^c$ , we plotted  $\psi_{soil}^c$  versus  $\tau$ , which is the parameter that determines the steepness of the hydraulic conductivity. Lower  $\tau$  meant a less steep decline in soil hydraulic conductivity with decreasing  $\psi_{soil}$ , thereby the soil-root interface was more conductive and water flow towards root was less constrained. The larger is  $\tau$ , the steeper is the decrease in conductivity (Figure 8a). It shows that the transpiration rate is constrained at more negative  $\psi_{soil}$  by the soil with a lower  $\tau$ .  $\psi_{soil}^c$  was around -10 kPa and  $\tau$  was over 0.4 in sand, whereas these values were around -100 kPa and 0.2 and 0.3 in sandy loam and loam, respectively.

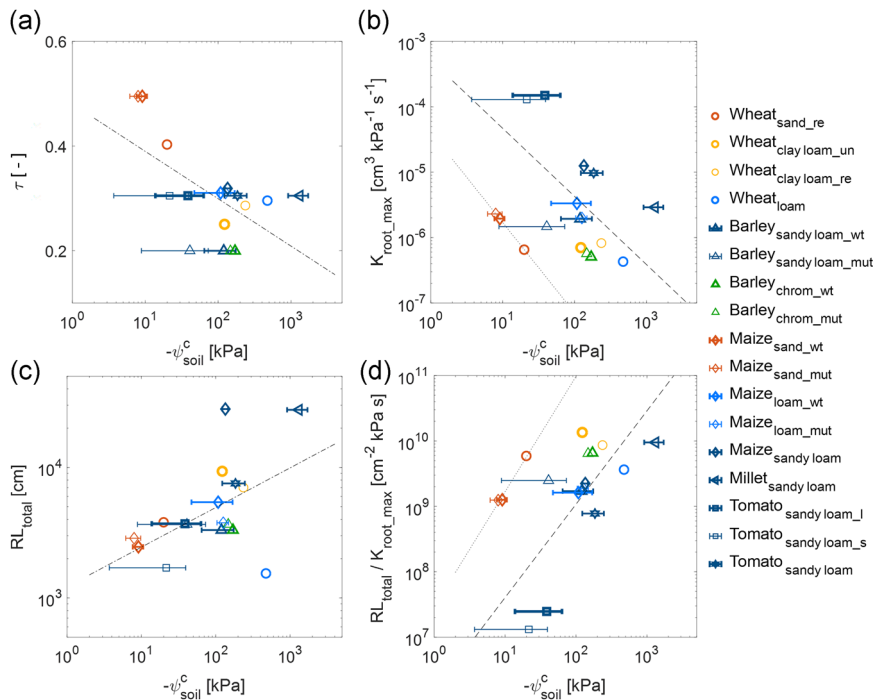
#### 3.4.2 | Root hydraulic conductance

The effect of root phenotype on  $\psi_{soil}^c$  is investigated by plotting  $\psi_{soil}^c$  with key root hydraulic properties. Figure 8b shows a general negative relation between  $K_{root\_max}$  and  $-\psi_{soil}^c$ . Plants with lower  $K_{root\_max}$  showed more negative  $\psi_{soil}^c$ . Although the relation in sandy soil seems off, the decreasing trend is similar (Figure 8b). Tomato had the highest  $K_{root\_max}$  and a high  $\psi_{soil}^c$  of ca. -30 kPa. Maize, barley, and wheat exhibited much lower  $K_{root\_max}$  and the corresponding  $\psi_{soil}^c$  was between -100 and -500 kPa. High  $K_{root\_max}$  helps plants to sustain high transpiration demand. However, it makes the  $E(\psi_{leaf\_x})$  relation more sensitive to the declines in soil hydraulic conductance (Figure 6). The total soil-root hydraulic conductance is the harmonic mean of the root and soil hydraulic conductance. As the harmonic mean is controlled by the smallest values, plants with high  $K_{root\_max}$  would be more responsive to soil drying.



**FIGURE 7** Definition of critical soil matrix potential ( $\psi_{soil}^c$ ). It is indicated as the pink star in both subplots. (a) Relation between transpiration rate ( $E$ ) and leaf xylem water potential ( $\psi_{leaf\_x}$ ) in wheat (Passioura, 1980). (b) Soil and root hydraulic conductance ( $K_{soil}$  and  $K_{root}$ ) as a function of soil matrix potential  $\psi_{soil}$ . Colour from dark blue to yellow in (a) stands for increasing  $E$ . The critical soil matrix potential  $\psi_{soil}^c$  is defined as the point where  $K_{root}$  equals  $K_{soil}$  was obtained by finding the intersection of fitted curves across  $K_{root}$  ( $K_{root\_fit}$ ) and  $K_{soil}$  ( $K_{soil\_fit}$ ) at the highest transpiration rate [Color figure can be viewed at [wileyonlinelibrary.com](http://wileyonlinelibrary.com)]





**FIGURE 8** Relation between critical soil matric potential ( $\psi_{\text{soil}}^c$ ) and soil and root hydraulic parameters. Relation between  $\psi_{\text{soil}}^c$  and (a)  $\tau$ , which determines the slope of soil hydraulic conductivity, (b) root hydraulic conductance in wet conditions ( $K_{\text{root\_max}}$ ), (c) root length ( $RL_{\text{total}}$ ) and (d) ratio of  $RL_{\text{total}}$  to  $K_{\text{root\_max}}$ . Fittings were justified using analysis of covariance (ANCOVA) for sand and other soils, and only one regression line was justified in (c).  $r^2$  of the regression is 0.35 in (a), 0.31 (without sand) and 0.999 (sand only) in (b), 0.25 in (c) and 0.50 (without sand) and 0.98 (sand only) in (d) [Color figure can be viewed at [wileyonlinelibrary.com](http://wileyonlinelibrary.com)]

### 3.4.3 | Root length

The effect of root length is shown in Figure 8c. It shows a positive relationship between root length and  $-\psi_{\text{soil}}^c$  in all soils. The effect of root length was distinct for millet and tomato grown in the same soil (sandy loam) and for maize grown in sand and loam. Maize in sandy loam and wheat in loam were off this relation but followed better the relation in Figure 8b. The mechanism by which root length impacts  $\psi_{\text{soil}}^c$  is that a longer root system reduces the local fluxes around roots and thus attenuates the gradient in soil water potential and delays the onset of nonlinearity of the  $E(\psi_{\text{leaf-x}})$  relation. Thus, plants with longer roots are expected to have a greater ability to extract water from drying soils. However, it has to be kept in mind that root length is not akin to root length active in water uptake (Ahmed, Zarebanadkouki, et al., 2018, 2016), which depends on the radial and axial hydraulic conductivities of roots. Indeed, in this review, the estimated fraction of roots active in water uptake across species and soils ranged from 4% to 30%.

### 3.4.4 | Root hairs

The relation between root hairs and soil-plant hydraulics remains inconclusive. In barley, the wild type with similar  $K_{\text{root\_max}}$  showed  $-10$  to  $-50$  kPa lower  $\psi_{\text{soil}}^c$  than that of the mutant in sandy loam but not in red chromosol (probably due to fewer measurements) (Figure 8b). Maize with similar  $K_{\text{root\_max}}$  did not show a significant ( $p > 0.05$ ) effect of root hairs on  $\psi_{\text{soil}}^c$  between wild type and mutant either in sand or in loam (Cai et al., 2021). Root hair length was two to three times longer in barley than in maize (Burak et al., 2021; Carminati et al., 2017; Marin et al., 2021). Longer root hairs were

probably more effective in maintaining the hydraulic continuity across the rhizosphere than varieties with shorter hairs. The underlying mechanism is that root hairs facilitate water uptake by increasing root absorption radius (Segal et al., 2008) or buffering the drop in matric potential at the soil-root interface (Ahmed et al., 2020; Ahmed, Passioura, et al., 2018; Carminati et al., 2017). Up to now, studies of the role of root hairs in water uptake are still limited to a few species. Further investigations should be performed on more species under different soil and water conditions to explore their contributions to root hydraulic properties and water uptake, especially given that the development (length and density) of root hair may be a plastic trait under nutrient deficiency conditions (Bahmani et al., 2016; Zhu et al., 2010).

The effect of root hydraulic traits on  $\psi_{\text{soil}}^c$  shown in Figure 8b,c was integrated into Figure 8d. Figure 8d considered root water uptake ability ( $K_{\text{root\_max}}$ ), root length (including root hairs), and uptake circumstances (soil texture). Due to the particular hydraulic property of sand (Figure 2), plants in sand did not follow the same relation between root hydraulic traits and  $\psi_{\text{soil}}^c$  as those in other soils but showed a similar trend. This relation indicates that plants with a lower  $K_{\text{root\_max}}$ , a longer root system and longer root hairs remain in the linear part of the relationship between transpiration rate and water potential for a broader range of soil water potentials.

## 3.5 | Implications

Among root hydraulic traits that impacted soil water availability, we investigated the effects of  $K_{\text{root}}$ , root length and root hairs. We introduced the concept of critical soil water potential,  $\psi_{\text{soil}}^c$ , as the point when the soil starts to limit root water uptake relative to root

conductivity. Here, we discuss the meaning and importance of these traits as well as the critical soil water potential for water use.

### 3.5.1 | Root hydraulic conductance

$K_{\text{root}}$  comprises the effects of root length, root architecture and the distribution of axial and radial conductivities (Doussan et al., 2006). The axial conductivity depends on the xylem vessel radius and number (Frensch & Steudle, 1989). The radial conductivity depends on the formation of Casparian band in exodermis (Zimmermann & Steudle, 1998), suberization (Kreszies et al., 2019) of the endodermis (Kreszies et al., 2020; Ranathunge et al., 2017), root cortical senescence (Schneider et al., 2017) and AQP expression (Martre et al., 2001). Variation of these elements impacts root hydraulic properties and thus root water uptake pattern (Ahmed, Passioura, et al., 2018; Ahmed, Zarebanadkouki, et al., 2018; Bramley et al., 2009; Javaux et al., 2008; Vadez, 2014).

Besides the absolute value of  $K_{\text{root}}$  in wet conditions ( $K_{\text{root,max}}$ ), the decline of  $K_{\text{root}}$  for decreasing soil water potentials is central to predict water use by crops. A loss of  $K_{\text{root}}$  would induce nonlinearities in the  $E(\psi_{\text{leaf,x}})$  curve. The decline in  $K_{\text{root}}$  during soil drying has several potential causes. Daily downregulation of AQP activity has been shown to occur particularly in water-stressed plants and has been suggested as a mechanism to trigger stomatal regulation at the peak of transpiration demand (Caldeira et al., 2014). A passive mechanism leading to a decrease in  $K_{\text{root}}$  is the formation of lacunae in the root cortex (Cuneo et al., 2016). Besides these mechanisms taking place at cellular and root cross-section scales, a root-system mechanism that affects  $K_{\text{root}}$  is that during soil drying the drying front moves deeper in the soil, and fewer roots have access to soil water (Hayat et al., 2019). The relevance of this mechanism depends on the root architecture and its temporal changes (Ahmed, Passioura, et al., 2018).

Our analysis suggests that root water uptake of plants with higher  $K_{\text{root,max}}$  was limited by the soil at less negative soil matric potential. The meaning of this result is not straightforward. According to the concept introduced by Carminati and Javaux (2020), the point at which the soil starts to limit water fluxes,  $\psi_{\text{soil}}^c$ , is reflected into a nonlinearity in the  $E(\psi_{\text{leaf,x}})$  relationship, and this triggers stomatal closure. The idea behind this model is that stomatal regulation prevents plants from nonlinearities in the  $E(\psi_{\text{leaf,x}})$  relationship. A similar concept was proposed by Sperry and Love (2015) and Sperry et al. (2016). A mechanism that allows stomata to accomplish this function has been recently proposed by Wankmüller and Carminati (2021). The idea is that stomatal conductance is regulated by the abscisic acid (ABA) level, which depends on the dynamic equilibrium between ABA production (assumed to increase with decreasing leaf water potential) and catabolism (assumed to increase with assimilation). The result of such a model is that stomatal conductance depends on the ratio between transpiration rate and leaf water potential. This concept still needs to be proven, but some preliminary evidence has been reported in Carminati and Javaux (2020) and Abdalla, Carminati, et al.

(2021). If we apply the concept of Carminati and Javaux (2020) and Sperry and Love (2015), we would reach the conclusion that plants with high conductivity would show an earlier decrease in transpiration during soil drying. In other words, their stomata would start to downregulate transpiration at a less negative  $\psi_{\text{soil}}^c$ . This conclusion is counterintuitive and has to be taken cautiously. Indeed, if stomatal conductance was simply a function of leaf water potential, a high plant hydraulic conductance would have the opposite effect on transpiration response to soil drying: the plant would close the stomata at a more negative leaf water potential. The latter would fit with the results of a breeding programme initiated for wheat around 30 years ago (Richards & Passioura, 1989). The authors showed that smaller xylem vessels with lower hydraulic conductance would decelerate root water uptake and extend the plant water use. It is not necessarily that these two concepts are in contradiction, as the time scales of these observations are different, that is, daily regulation of transpiration in our concept compared to seasonal water use in the work of Richards and Passioura (1989). Possibly, a high plant hydraulic conductance would maintain transpiration in wet soils at a high vapour pressure deficit (VPD), while it would trigger stomatal closure during soil drying. The underlying mechanism of stomatal closure is still not fully explored (Brodribb & Holbrook, 2003; Buckley, 2005; Tardieu, 2016; Tombesi et al., 2015). Stomatal closure is proposed to be driven by ABA (Buckley, 2019), and ABA biosynthesis increases with decreasing leaf water potential (McAdam & Brodribb, 2016). But transport (Tardieu & Davies, 1993) and catabolism of ABA are also relevant for ABA levels. So, it is possible that stomatal conductance is both function (1) of leaf water potential, and (2) of a ratio between transpiration (or assimilation) and leaf water potential, with the latter component being important for the fine regulation of stomatal opening. In conclusion, the fact that a high plant conductance causes an earlier nonlinearity in the relationship between transpiration and leaf water potential does not translate in a straightforward way into an earlier decrease in transpiration rate.

### 3.5.2 | Root length

We have shown that root length is an important root trait impacting the easiness of water uptake. First of all,  $K_{\text{root}}$  increases with root length (Figure 3a). Additionally, a long root length means a slow rate of water flow in soils. Thus, it is beneficial to maintain water uptake under moderate drought. Opposite to the discussion on  $K_{\text{root,max}}$ , independently from the stomatal model, a longer root length is expected to sustain transpiration during soil drying. Deeper rooting was shown to be important under dry conditions (Fan et al., 2017; Lynch, 2013; Wasson et al., 2012; White & Kirkegaard, 2010). However, the effects of a profuse root system on soil water availability and water use are not unique and their effects on water uptake are context-dependent (see the discussion in Ahmed, Zarebanadkouki, et al., 2018, and Ahmed et al., 2020). Vadez (2014) showed that plants with a profuse root system (high root length densities) contributed positively to water uptake under drought in

some studies whereas in others did not. A profuse root system may not be conducive to water use if most roots distribute at surface layers.

### 3.5.3 | Root hairs

Another root hydraulic phenotype that needs more attention is root hairs. Root hairs have been shown to play an important role in nutrient absorption; however, their role in water uptake is still unclear. Root hairs are expected to increase the effective root radius and reduce the gradients in soil water potential needed to sustain high transpiration rates in drying soils (Segal et al., 2008). Experiments with barley with and without hairs supported this concept, both in the lab (Carminati et al., 2017) and in the field (Marin et al., 2021). However, there are measurements with similar genotypes that did not show significant differences (Dodd & Diatloff, 2016). For maize, Cai et al. (2021) found no difference between a wild type and its hairless mutant. The reason might be related to the shorter length of root hairs of maize compared to barley. Additionally, the effects of root hairs might be soil-dependent. Therefore, it seems that there is no unique answer about the positive role of hairs on soil–plant hydraulics.

In summary, we proposed a concept to link root hydraulic phenotypes to water use by crops. Root hydraulic traits that attenuate declines in soil limitation (for instance, root hairs or long roots) would help to sustain the use of water during soil drying. Traits that cause an earlier drop in conductance, such as a high  $K_{\text{root,max}}$  or a decline in  $K_{\text{root}}$  with soil drying, could lead to a more conservative use of water. The ‘conservation’ here only links with plant water budget or plant hydraulics, not with drought tolerance (Ahmed, Passioura et al., 2018; Vadez, 2014).

### 3.6 | Limitations

Besides the factors discussed above, root hydraulic conductance can be impacted by other factors, such as root architecture (Doussan et al., 2006; Javaux et al., 2008), root types (Ahmed, Zarebanadkouki, et al., 2016, 2018) and rhizosphere hydraulics changing over time (Ahmed et al., 2014; Ahmed, Kroener, et al., 2016), which further affect water uptake. In this study,  $K_{\text{root}}$  was a ‘effective conductance’, and was an integration of root hydraulic properties of the whole root system during the drying process. Except in wet conditions, when  $E$  was not limited by soil,  $K_{\text{root}}$  was an optimized variable based on the measurements of  $E$ , water potential in the soil and leaves. More complex models that consider dynamic root hydraulic properties, root architecture and nonuniform soil moisture distribution may be better in estimating root water uptake. However, estimation from those models depends on the accuracy of detailed root distribution and hydraulic properties of individual roots. These parameters are not available and cannot be easily measured, especially in roots growing in soils.

## 4 | CONCLUSIONS

This review highlights the role of soil and key root hydraulic properties to predict water uptake by crops under limiting soil water conditions. We showed at what soil water potential the soil starts to limit the water fluxes through the soil–plant continuum across species and soils. Coarse textured soils exhibit an earlier limitation to root water uptake due to their sudden loss in conductivity for decreasing soil water potentials. We have introduced the concept of critical soil water potential  $\psi_{\text{soil}}^c$  and showed how it is affected by key hydraulic traits. The critical soil matric potential was not unique and varied with soil textures and root hydraulic phenotypes. Lower  $K_{\text{root,max}}$ , longer root length and longer root hairs have a more negative  $\psi_{\text{soil}}^c$ .

According to the concept proposed by Carminati and Javaux (2020) and Sperry and Love (2015),  $\psi_{\text{soil}}^c$  corresponds to the point at which plants downregulate transpiration. This would imply that plants with a high root hydraulic conductivity would be more responsive in reducing transpiration during soil drying. However, a high hydraulic conductance would result in a less negative leaf water potential, which might have the opposite effect of maintaining transpiration. Despite this controversy on the effect of plant hydraulic conductance on stomatal regulation, this review of hydraulic traits and their impact on soil–plant conductance could help us to reveal the role of root hydraulics in crop water use. The effect of different root hydraulic phenotypes on drought tolerance will depend on environmental conditions and precipitation/irrigation patterns (Ahmed et al., 2020; Cai et al., 2021; Comas et al., 2013). Root hydraulic phenotypes that trigger an earlier stomatal closure would be convenient to save water transiently. On the other hand, root hydraulic phenotypes that maintain transpiration during soil drying would be favourable to maintain growth (at the cost of faster soil water depletion).

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### CONFLICT OF INTERESTS

The authors declare that there are no conflict of interests.

### DATA AVAILABILITY STATEMENT

The raw data are from published studies (described in Table 1). The analysed data that support the findings of this study are available from the first and corresponding authors upon reasonable request.

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## SUPPORTING INFORMATION

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## APPENDIX

Parameters and units used in the figures and table:

$E$ : transpiration rate (measured),  $\text{cm}^3 \text{s}^{-1}$

$\psi_{\text{soil}}$ : soil matric potential (measured), kPa

$\psi_{\text{soil\_root}}$ : water potential at the soil–root interface (modelled), kPa

$\psi_{\text{soil}}^c$ : critical soil matric potential (modelled), kPa

$\psi_{\text{leaf}}$ : leaf xylem water potential (measured), kPa

$K_{\text{root}}$ : root hydraulic conductance (modelled),  $\text{cm}^3 \text{kPa}^{-1} \text{s}^{-1}$

$K_{\text{root\_max}}$ : maximum root hydraulic conductance (modelled),  $\text{cm}^3 \text{kPa}^{-1} \text{s}^{-1}$

$k_{\text{root}}$ : root hydraulic conductivity (modelled),  $\text{cm kPa}^{-1} \text{s}^{-1}$

$k_{\text{root\_max}}$ : maximum root hydraulic conductivity (modelled),  $\text{cm kPa}^{-1} \text{s}^{-1}$

$K_{\text{soil}}$ : soil hydraulic conductance (modelled),  $\text{cm}^3 \text{kPa}^{-1} \text{s}^{-1}$

$k_s$ : soil hydraulic conductivity (measured),  $\text{cm s}^{-1}$

$RL_{\text{total}}$ : measured total root length (measured), cm