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Tree rings reveal the transient risk of extinction hidden inside climate envelope forecasts

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Supporting Information Text

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Biodiversity Forecasting with Occurrence Data

 Predicting the impact of climate change on biodiversity has been dominated for the last twenty years by an approach that associates a species' occurrence (either presence-absence or presence-only data) with climate at those locations. The principle is illustrated in Fig. S1 with respect to a single temperature variable. In practice, this statistical model, known as a *species distribution model (SDM)*, *climate envelope model*, or *ecological niche model*, is created using many climate (or other predictor) variables at once. Excellent reviews of the methods are found in [\(1–3\).](https://www.zotero.org/google-docs/?8igUoc) The statistical model is then used to predict or forecast change in the species' geographic distribution with changing climate. In the single-variable example of temperature, with warming, the species' probability or rate of occurrence is predicted to increase at relatively cool locations (the *leading edge*) and decline at warmer locations (the *trailing edge*; Fig. S1B).

Climate responses across scales

 In order for climate envelope or SDM predictions of range dynamics to actually come to pass, lower-level demographic variables, such as population size or density (abundance), population- level growth rate, or the net effect of individual-level vital rates (including growth, survival, and birth rates) should respond similarly to changing climate, either matching the entire response, including breadth (solid lines, Fig. S2), or, if more limited in breadth, they should reflect the local slope of the species-scale climate response curve (dashed lines, Fig. S2). Ecological theory supports this assumption: the *center-periphery hypothesis* predicts that individual-level performance and population-level growth rate decline from the center to the edge of a species' geographic and environmental distribution (Fig. S2). Under the *abundant-center hypothesis* and *Whittaker's continuum concept*, it is expected that a species' abundance is greatest at the center of its ecological niche (*i.e.*, climatic tolerances; Fig. S2). Though we note – these theoretical expectations concern time-averaged responses to spatial variation in climate (not responses to time-varying climate). Further, range dynamics in the real world are complicated by the lack of correspondence between position in geographic and environmental space – climatically average (or extreme) conditions are found throughout species' geographic distributions [\(4–6\).](https://www.zotero.org/google-docs/?g4n8VI) Climate response curves at different biological scales (individual, population, and species) and in response to spatial vs. temporal climate variation are influenced by a variety of different ecological processes. Further, they can take on a variety of shapes beyond a Gaussian, bell-shaped curve (unimodal, symmetric; Fig. S10A).

 Individual Scale. Focusing on the scale of an individual organism, the phenotypic response of a single genotype to different environmental conditions is termed by population geneticists its *reaction norm*. A time series of an individual's growth rate is an example of a single genotype expressing different phenotypes (growth rates) when exposed to different conditions (interannual climate variation and any other differences between years). This interannual variation in growth is the result of *phenotypic plasticity*. *Thermal performance curves*, plastic responses to different temperature conditions, have been the subject of heightened interest in the context of anthropogenic warmin[g \(7\).](https://www.zotero.org/google-docs/?4tAvwu) They are understood to be bounded at both a minimum and maximum value – organismal functionality goes to zero below a minimum temperature and above a maximum temperature (the *critical thermal maximum*). Between the minimum and maximum 91 lies an optimum – a temperature at which individual-level physiological functioning reaches its peak. This thermal response curve may be symmetric (Gaussian) or left- or right-skewed (Fig. S10A). If skewed, the response may be approximated by a linear function across much of the range of temperatures to which the organism is regularly exposed and hence adapted (Fig. S10B). Other climate response curves at the individual scale may be saturating. For example, performance in plants may be a saturating function of soil moisture – increasing to a plateau, above which additional soil moisture leads to neither increased nor decreased performance (Fig. S10A).

 Population scale. Climate response curves at the population scale, involving population size, density, or growth rate as responses, emerge as the net result of five demographic processes or *vital rates*: birth, death, immigration, emigration, and (in structured populations) changes in age, size, or developmental stage. Different vital rates may be affected by climate differently, even in an opposite manner – known as *demographic compensation*. The net effect of a climate variable on population growth rate is a function of its effect on each vital rate and the sensitivity of population growth rate to changes in each vital rate. Additionally, population-scale climate response curves reflect community dynamics, including competition between species, facilitative or mutualistic interactions, food web (consumer-resource) relationships, disturbance processes, source-sink dynamics, and community sorting (*e.g.*, priority effects).

 Species Scale. At the scale of a species' geographic distribution, climate response curves reflect all the above processes. In addition, it is widely recognized that climatically suitable habitat may not be occupied because the species' rate of dispersal may prevent (limit) it from colonizing a given location (*dispersal limitation*). In fact, all these processes (plasticity, demography, competition and other interspecific interactions, disturbance processes, evolution, dispersal, *etc.*) operate at all scales all the time. Different sources of data, with different temporal and spatial

 extent and resolution, offer more or less insight into direct cause-effect relationships and which processes are important influences on pattern at each scale.

Study species

 Pinus edulis Engelm. (Colorado, common, or two-needle piñon) is a small-statured, slow-growing, stress-tolerant pine that can survive more than 500 years. It is endemic to the Colorado Plateau of the southwestern U.S. states of Arizona, Colorado, New Mexico, and Utah, where potential evapotranspiration exceeds precipitation in most months [\(8\).](https://www.zotero.org/google-docs/?pfjIN5) Because it occurs across a wide range of elevation (1400-2700 m), *P. edulis* grows under a wide range of temperature conditions (mean annual temperature [MAT] of 4-17° C; Fig. S3). There is a gradient of monsoon moisture (in July and August, Fig. S11) increasing from north to south across *P. edulis*' distributio[n \(8\).](https://www.zotero.org/google-docs/?t6OSIV)

Tree-Ring Data

 Tree-ring data were derived from increment cores sampled in the U.S. Forest Service Forest Inventory and Analysis (FIA) Program's spatial network of permanent sampling plots, during forest inventories between 1995 and 2013 (especially 1995-1997; [\(9, 10\)](https://www.zotero.org/google-docs/?x8TBM0). Samples were processed to generate annually resolved time series at the University of Arizona Laboratory of Tree-Ring Research and at Utah State University, following standard dendrochronological protocols [\(9, 11\)](https://www.zotero.org/google-docs/?0Bs5kT): mounted on grooved boards, sanded to a fine polish, crossdated to assign a year of formation to each growth ring, and measured on a calibrated sliding stage micrometer. Year assignments were verified using COFECHA [\(12\)](https://www.zotero.org/google-docs/?jwAm4S) before any analyses of the ring width time series. Start dates of the time series ranged from 1530 to 1983, but data analysis was limited to the period for which gridded climate data products are available (1895-1995).

Regression modeling of tree-ring width

 We used a Bayesian hierarchical regression model to evaluate the predictions of Hypotheses 1 vs. 2 about patterns of variation in *P. edulis*' performance with variation in climate across space and time. This model predicted the log-transformed width of growth rings as a function of tree size and climate variables. Tree size (stem diameter at root collar, DRC) was included as a predictor because ring widths are known to change with tree size – radial growth increments are wide when a tree is small and become narrower as a tree becomes larger [\(13\).](https://www.zotero.org/google-docs/?1Xyj4D) Starting with the DRC measurement at the time that an increment core was sampled (available from the FIA database), DRC was back-calculated based on the tree-ring time series, generating an inferred DRC for every tree in every year. In addition to these fixed effects, we specified random tree effects – *i.e.*,

 a modification of the model intercept specific to each tree – capturing heterogeneity among trees in average growth rate not explained by site-specific mean annual temperature (MAT) and mean annual precipitation (MAP) data.

 Both spatially varying and spatio-temporally varying climate predictors were included in models of *P. edulis* growth, the first to capture how growth varies across spatial gradients of climate and the second to capture plastic responses to interannual climate variability (*i.e.*, reaction norms). Both types of climate predictors were created from time series of monthly, 4-km resolution climate data downloaded from the PRISM Climate Grou[p \(14\)](https://www.zotero.org/google-docs/?iFDR4u) for the period 1895- 2018. Climate normals, which vary strictly across space, were created by averaging monthly mean temperature and summing monthly precipitation of each year (January-December), then averaging across years to calculate mean annual temperature (MAT) and mean annual precipitation (MAP). To make time-varying climate predictors, the time series of monthly climate data were aggregated across a twelve-month time frame (previous September to current August) and four biologically relevant seasons: fall (September through October of the previous year), winter (previous November through the current year's March), spring (current year April through June) and monsoon (July through August). These four seasonal climate variables represent two wet seasons, one warm (monsoon) and one cold (winter), along with two warm, dry seasons (spring, fall; Fig. S11). Responses to time-varying climate were modeled as linear, which is traditional in dendrochronology [\(11, 13\).](https://www.zotero.org/google-docs/?L6mdon) While a few recent analyses of tree-ring data have shown nonlinearity in responses to climate [\(15, 16\),](https://www.zotero.org/google-docs/?cFbGHR) and some reaction norms (*e.g.*, thermal response curves) are theoretically expected to peak between minimum and maximum climate values [\(7\),](https://www.zotero.org/google-docs/?99VV9x) field-observed values of performance vs. climate may be adequately predicted with a linear response across the historic, field-observed range of variability of climate, as described above, in "*Climate Responses across Scales*" (see Fig. S10B).

 Though we had specific hypotheses to test (Fig. 1), it is important to verify that the model structure corresponding to those hypotheses adequately describes variation in *P. edulis* performance across space and time, *i.e.*, is not outcompeted by alternative models. We report here the fit to data of ten alternative Bayesian hierarchical regression models (Table S2), which included different combinations of climate predictors (normals, time-varying) and different methods of scaling during standardization (*i.e.*, centering and scaling of predictors to a mean of zero and standard deviation of one). Specifically, time-varying climate predictors were standardized using either global or local scaling. Global scaling compares temperature or precipitation at a given site each year to the global average for that variable across both space and time, whereas local scaling compares each year's climate to the local site average for that variable. Locally scaled variables thus reflect climate anomalies relative to average conditions at each site. The hypothesis under global scaling is that a temperature of 10° C and temperature variation of 1° C has the same effect on all trees throughout the geographic distribution, whereas

191 the hypothesis under local scaling is that a temperature of 10° C may be relatively cool at the warm edge of the distribution and relatively warm at the cool edge of the distribution, and that the effect of climate variability is best described relative to what trees are adapted to at a given location. Hence, the fit of data to models with globally vs. locally scaled climate predictors represents a secondary test comparing the hypotheses that climate responses are best described at the species scale (Hypothesis 1) vs. population or individual scale (Hypothesis 2).

 The ten models are generally ordered from simpler to more complex. Model 1 tested the hypothesis that the climate normals MAP and MAT alone are sufficient to explain variation in *P. edulis* growth. Model 2 repeated Model 1 with the addition of tree size. Model 3 tested the hypothesis that time-varying climate, in addition to climate normals and tree size, explain growth variation, using globally scaled, 12-month cumulative precipitation and average temperature. Models 4-6 tested alternative seasonal time-varying climate predictors (both precipitation and temperature), which were globally scaled (Table S2). Models 7 and 8 test the hypothesis that the climate normals MAP and MAT do not significantly influence *P. edulis* growth (by removing them), with Model 7 using global scaling and Model 8 local scaling of the time-varying climate predictors. Model 9 tested the same seasonal climate predictors as Model 6, using local scaling for the time- varying predictors. Model 10 was unique in testing the use of both locally and globally scaled versions of the time-varying climate predictors of Model 6. All models included all two-way interactions between predictors, and all but Model 1 included the effect of tree size (Table S2).

 Regression modeling was implemented using R, RStudio, and STAN [\(17\),](https://www.zotero.org/google-docs/?ec8lFj) in the computational environment of VICE and the Discovery Environment of CyVerse (www.cyverse.org). Markov chain Monte Carlo simulations were run with 3 chains of 5000 iterations each, discarding the first 1000 iterations as warmup, resulting in 12,000 posterior samples. Convergence was assessed using visual inspection of traceplots and Gelman-Rubin diagnostics [\(18\).](https://www.zotero.org/google-docs/?mfk8Cd) Models were fit with a randomly selected 80% of the growth ring measurements and were validated with the remaining 20% of these held-out data using the root mean squared error (MSE) of model-predicted tree-ring width under a 5-fold cross-validation procedure. In addition, a full data fit using the approximate leave-one-out (elpd) information criterion [\(19–21\);](https://www.zotero.org/google-docs/?4KduZc) Fig. S9) was used to evaluate model performance. Model fit was also assessed visually using posterior predictive checks [\(18\).](https://www.zotero.org/google-docs/?eiqx2L) The model fit statistics showed that seasonal time-varying climate variables predict *P. edulis*'s growth variability better than 12-month time-varying climate variables (Models 5 and higher fit better than Model 4), and that local scaling outperforms global scaling of the time-varying climate variables (Model 9 compared to Model 6, and Model 8 compared to Model 7; Fig. S9). Further, the inclusion of both climate normals and time-varying climate variables is supported (Model 9 fit the data better than Models 8 and 7; Fig. S9). Among the ten models fit to the ring-width data, two rose to the top in terms of low MSE and elpd: Models 9 and 10 (Fig. S9). We selected Model 9 as the preferred best-fit model, which included MAP and MAT, four seasonal time-varying climate predictors - spring and fall temperatures and winter and monsoon precipitation - and all 2-way interactions between fixed effects.

 The posterior predictive distribution diagnostics for Model 9 (pointwise posterior predictive credible intervals, probability integral transform overlay, and probability integral transform q-q plots; Fig. S12) illustrate a good fit to the bulk of the data but show some evidence of misfit for very small and very large tree-ring width observations. We plotted model residuals as a function of each of the time-varying climate predictors to check for systematic misfit that might result from modeling log-transformed ring widths as a linear response of interannual climate variation. The residuals showed no trend with respect to spring or fall temperature variation or monsoon precipitation variation among years, but there was some trend in the residuals with respect to variation in winter precipitation (blue trend lines, Fig. S13). There is a modest pattern of more extreme negative residuals at both the lowest and highest observed values of winter precipitation, indicating that the model tends to over-predict growth ring widths in the driest and wettest winters, respectively (Fig. S13B). Noticeable in all four panels of Fig. S13 are data points with strong negative residuals (the cloud of points below most of the rest of the residuals, in each panel) – these are cases where observed growth ring widths are much smaller than the predictions of the model. These are the year and tree combinations in which a growth ring was not produced, which we replaced with 0.001 mm (the smallest non-zero observation in our data), since the log of zero is undefined. Part of the trend in the residuals in Fig. S13B seems to arise from the association of the "missing ring" years with low values of winter precipitation. In other words, missing rings tend to be produced when winter precipitation is low. The trend of negative residuals at high values of winter precipitation suggests that there may be some leveling off (saturation) of the response of tree growth to the highest values of winter precipitation, as discussed above in "*Climate Responses across Scales*" (see Fig. S10A). Overall, this trend in residuals is modest and we considered linear responses to climate to be adequate to describe the variation in the data. A future improvement to modeling tree ring widths would be to use a Tobit or other mixture model to better capture the conditions that drive the formation of missing rings. Because Model 9 performed well across predictive metrics and showed good evidence of

 fit from the posterior predictive diagnostic plots, the output of Model 9 was used for testing Hypotheses 1 vs. 2 and to create Fig. 2-3 and Fig. S4-S7. We note that across all ten models, the effects of climate were consistent in sign and magnitude, lending confidence in the robustness of these effects to alternative model structures.

Climate responses

 All estimated fixed effects (including interactions) from Model 9 are shown in Fig. S4. In a soil moisture-limited species (*P. edulis*), we expected to see higher-than-average growth in both

 cooler-than-average years and locations (Fig. 1D) because lower rates of evaporation with lower temperatures should weaken soil moisture-limitation of growth. Instead, the negative effect on *P. edulis'* growth of interannual variability in temperature $(\beta_{\text{springtemp}} = -0.0965, 95\% \text{Cl} = -0.1076 - -$ 268 0.0857 and β falltemp = -0.0723, 95%CI = -0.826 - -0.0618) is contrasted by a positive effect of 269 spatial variation in mean annual temperature (β_{MAT} = 0.2083, 95% CI = 0.1661 – 0.2506). Further, 270 there is a significant positive interaction between the climate normals (fig. S4, $\beta_{\text{MAT}^*MAP} = 0.0438$, 95% CI = 0.0097 – 0.0775), indicating that as mean annual temperature increases, the positive 272 influence of mean annual precipitation $(\beta_{MAP} = 0.2355, 95\% \text{ Cl} = 0.1940 - 0.2771)$ on growth increases (and *vice versa*). Hence, growth rates of *P. edulis* are greatest at warmer-than-average sites, but especially those that are also wetter-than-average (blue lines, Fig. 2C-D). This may best be interpreted in terms of Liebig's law of the minimum. Among wetter-than-average sites, where soil moisture is less limiting, the influence of another limiting factor becomes evident: trees at cold but relatively wet sites are still unable to grow rapidly (blue lines, Fig. 2A-B), whereas those at warm but relatively wet sites can. We summarize *P. edulis*' observed climate-growth relationships in abstract form in Figure S8B, in contrast to the predicted climate-growth relationships in fig. S8A.

 Finally, negative model-predicted responses to greater-than-average winter precipitation are seen at a very small number of cold, wet locations (*e.g.*, blue lines, Fig. 3A; Fig. S6C), which may represent high-elevation sites where a year of deep snowpack can limit the growing season and hence tree growth.

375 **Fig. S1.** Conceptual illustration of (A) a species' climate envelope, inferred from presence-
376 absence or presence-only data, with respect to a single dimension of climate (a temperatur

376 absence or presence-only data, with respect to a single dimension of climate (a temperature
377 variable). This species-level climate response is then used to project (B) the effect of climate 377 variable). This species-level climate response is then used to project (B) the effect of climate
378 change (warming) on the species' rate or probability of occurrence.

378 change (warming) on the species' rate or probability of occurrence.

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383 **Fig. S2.** Responses to temperature variability at the (A) individual, (B-C) population, and (D) 384 species scale. The use of climate envelope models to forecast change in species' distribution 384 species scale. The use of climate envelope models to forecast change in species' distributions
385 effectively assumes that climate responses are the same across scales, as illustrated here. Sol 385 effectively assumes that climate responses are the same across scales, as illustrated here. Solid
386 lines indicate a scenario in which individual- and population-scale climate responses are equal in lines indicate a scenario in which individual- and population-scale climate responses are equal in 387 breadth to the species-scale response; dashed lines indicate the case where individual- and
388 population-scale climate tolerances are narrower than species-scale climate tolerances, but to 388 population-scale climate tolerances are narrower than species-scale climate tolerances, but they
389 match in local slope (the sign of the response). Ecological theories that support the assumption of match in local slope (the sign of the response). Ecological theories that support the assumption of 390 scale-invariant climate responses are placed above the biological scale axis.

Fig. S3. Location in (A) geographic and (B) climate space of the 977 U. S. Forest Service Forest

 Inventory and Analysis (FIA) plots in the U. S. states of Arizona, New Mexico, Colorado, and Utah where *Pinus edulis* increment cores were collected. Green in panel A is the geographic

distribution of *Pinus edulis*, as defined in the atlas of U. S. trees [\(22\).](https://www.zotero.org/google-docs/?ajmy7G)

401 **Fig. S4.** Estimates of effects from a multiple regression model of *P. edulis* growth (model 9 in 402 Table S2). Each dot-whisker shows the Bayesian posterior distribution of a model parameter
403 (posterior mean +/- 95% central density). Predictors were centered and scaled, with the exce

403 (posterior mean +/- 95% central density). Predictors were centered and scaled, with the exception
404 of tree size. Main effects are color-coded into three groups: climate normals (green), time-varying of tree size. Main effects are color-coded into three groups: climate normals (green), time-varying

405 climate variables (blue), and tree size (pink). Two-way interaction effects are color-coded

accordingly both in the y-axis label and the dot-whisker.

409 **Fig. S5.** Model-predicted responses to fall temperature variability of all 1,558 common pinon trees 410 in the dataset, at locations that vary from cool to warm (quantiles of mean annual temperature
411 [MAT], averaged over the period 1895-2018), with each response colored by the mean annual 411 [MAT], averaged over the period 1895-2018), with each response colored by the mean annual
412 precipitation (MAP) at that location, from dry (red) to wet (blue). Fall temperature is the average

412 precipitation (MAP) at that location, from dry (red) to wet (blue). Fall temperature is the average of 413 monthly average temperatures September-October of the previous calendar year, locally scaled,

413 monthly average temperatures September-October of the previous calendar year, locally scaled,
414 *i.e.*, anomalies relative to site-specific average fall temperature. Responses are plotted for a i.e., anomalies relative to site-specific average fall temperature. Responses are plotted for a

- 415 constant tree size of 20 cm.
- 416

418 **Fig. S6.** Histograms of the sensitivity of the growth of 1,558 *Pinus edulis* trees to four time-

419 varying climate predictors: (A) spring and (B) previous fall temperature, and (C) winter and (D)
420 monsoon precipitation. These sensitivities correspond to the slopes of the lines in Fig. 2, Fig. S

420 monsoon precipitation. These sensitivities correspond to the slopes of the lines in Fig. 2, Fig. S5,
421 Fig. 3, and Fig. S7, respectively, although here the slope is change in *log-transformed* growth-ring

421 Fig. 3, and Fig. S7, respectively, although here the slope is change in *log-transformed* growth-ring

width in response to variation in each climate predictor.

425 **Fig. S7.** Model-predicted responses to monsoon precipitation variability of all 1,558 common 426 pinon trees in the dataset, at locations that vary from cool to warm (quantiles of mean annual
427 temperature [MAT], averaged over the period 1895-2018), with each response colored by the 427 temperature [MAT], averaged over the period 1895-2018), with each response colored by the 428 mean annual precipitation (MAP) at that location, from dry (red) to wet (blue). Monsoon 428 mean annual precipitation (MAP) at that location, from dry (red) to wet (blue). Monsoon
429 precipitation is the sum of precipitation in July and August of the current calendar year, k

429 precipitation is the sum of precipitation in July and August of the current calendar year, locally
430 scaled, *i.e.*, anomalies relative to site-specific average monsoon precipitation. Responses are

scaled, *i.e.*, anomalies relative to site-specific average monsoon precipitation. Responses are

431 plotted for a constant tree size of 20 cm.

- 434 **Fig. S8.** Predicted vs. observed relationships between tree growth and interannual temperature
435 variability across the distribution of a (A) hypothetical soil moisture-limited species (as in Fig. 1E)
- variability across the distribution of a (A) hypothetical soil moisture-limited species (as in Fig. 1E),
- and (B) *Pinus edulis*.

- **Fig. S9.** Fit of ten alternative regression models to *Pinus edulis* tree-ring width data, with respect
- to (A) root mean squared error (MSE) and (B) approximate leave-one-out information criterion (elpd). All ten models are detailed in Table S2.

Fig. S10. Climate responses (A) may take a variety of shapes, from symmetric to skewed to 446 saturating. (B) A skewed response may be fit to field-collected data with a line, *i.e.*, across the

saturating. (B) A skewed response may be fit to field-collected data with a line, *i.e.*, across the

Faranting C₂, A shown to perform they see the to here that sense that then a may, here see the range of climate variability to which an organism is regularly exposed and therefore adapted.

450 **Fig. S11.** Average monthly temperature and precipitation at U. S. Forest Service Forest Inventory
451 and Analysis plot locations where *Pinus edulis* is present, based on PRISM 4-km resolution

451 and Analysis plot locations where *Pinus edulis* is present, based on PRISM 4-km resolution climate data.

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455 **Fig. S12.** Posterior predictive diagnostics for Model 9. (A) The 50% and 90% posterior predictive 456 intervals for a subset of log tree-ring widths, ordered by posterior predictive mean (y_{exp}) . Dark
457 points are the observed log ring widths (y), which mostly fall within the 50% or 90% posterior 457 points are the observed log ring widths (y), which mostly fall within the 50% or 90% posterior
458 predictive intervals. (B) Smoothed density of posterior predictive samples of log ring width in I 458 predictive intervals. (B) Smoothed density of posterior predictive samples of log ring width in light
459 blue and a smoothed density of observed log ring widths in dark blue. (C) Leave-one-out quartile-459 blue and a smoothed density of observed log ring widths in dark blue. (C) Leave-one-out quartile-
460 quartile plot of posterior predictive mean vs. observed log ring width. quartile plot of posterior predictive mean vs. observed log ring width.

Fig. S13. Residuals of model 9 plotted against the four time-varying climate predictors: (A)
464 monsoon precipitation, (B) winter precipitation, (C) spring temperature, and (D) previous fall monsoon precipitation, (B) winter precipitation, (C) spring temperature, and (D) previous fall temperature. A generalized additive model fit to the residuals in each panel, shown with a blue line, evaluates trend in model residuals.

467 **Table S1.** Corroborating evidence of scale-dependent climate responses: mismatches between 468 individual-level vital rates (*e.g.*, growth, survival, fertility, recruitment), population-level growth rate 469 or abundance, and species-level occurrence, probability of occurrence, or climatic suitability
470 derived from occurrence data in the recent literature. The citation is followed by a description derived from occurrence data in the recent literature. The citation is followed by a description of 471 the scope of the data (organism, spatial and temporal replication, location), then the type of data
472 analyzed, or variables estimated from data, then a description of the mismatch, reproduced from analyzed, or variables estimated from data, then a description of the mismatch, reproduced from 473 the title or abstract of the citation or summarized succinctly. Where possible, we list (in 474 parentheses) whether the study considered spatial vs. temporal variation in the individual- and 475 population-level variables. Species-level variables (e.g., climatic suitability) vary only across 475 population-level variables. Species-level variables (*e.g.*, climatic suitability) vary only across space (with some exceptions). Citations are grouped into four categories: those addressing 1) the 477 abundant center hypothesis or 2) center-periphery hypothesis, those 3) analyzing individual-level
478 performance variation across space and time and fitting a demographic model to estimate 478 performance variation across space and time and fitting a demographic model to estimate
479 population-level growth rate, and 4) those that explicitly consider distinct variation across s 479 population-level growth rate, and 4) those that explicitly consider distinct variation across space 480 vs. time. This list of citations is surely incomplete, and alternative placement of a given paper into
481 a different category is possible. a different category is possible.

484 **Table S2.** Ten regression models predicting growth ring width variation of *Pinus edulis* as a 485 function of tree size and climate variables. A numerical label and all main effects are listed for
486 each model, as well as the type of scaling used for time-varying predictors (G for global scaling 486 each model, as well as the type of scaling used for time-varying predictors (G for global scaling or
487 L for local scaling). Seasons were defined as monsoon (m; Jul-Aug), winter (w; Nov-Mar), spring 1997 L for local scaling). Seasons were defined as monsoon (m; Jul-Aug), winter (w; Nov-Mar), spring
188 (s; Apr-Jun), and fall (f; Sep-Oct). Values of the model fit statistics root mean squared error (MSE 488 (s; Apr-Jun), and fall (f; Sep-Oct). Values of the model fit statistics root mean squared error (MSE)
489 and leave-one-out (elpd) validation are shown in Fig. S9. and leave-one-out (elpd) validation are shown in Fig. S9.

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496 **Table S3.** Model-fit statistics (deviance and AIC) for generalized additive models (GAMs) of *Pinus* 497 *edulis*' occurrence in the U.S. Forest Inventory and Analysis plots of Arizona, Colorado, New 498 Mexico, and Utah as a function of the climate conditions at plot locations (variables listed). GAMs 499 were fit with five, four, and three knots (k).

| Climate Variable | | Deviance | | | AIC | |
|------------------------------|---------|----------|---------|--------|------------|---------|
| | $k = 5$ | $k = 4$ | $k = 3$ | $k=5$ | $k = 4$ | $k = 3$ |
| Mean Annual Temperature | 252.85 | 254.76 | 254.47 | 288.96 | 289.32 | 287.7 |
| Mean Annual Precipitation | 294.76 | 296.05 | 299.84 | 331.09 | 331.67 | 334.4 |
| Monsoon Precipitation | 289.94 | 302.46 | 302.98 | 327.88 | 337.39 | 337.68 |
| Winter Precipitation | 299.73 | 302.46 | 302.98 | 336.14 | 337.39 | 337.68 |
| Fall Temperature | 257.7 | 258.14 | 257.85 | 293.28 | 292.83 | 291.21 |
| Spring Temperature | 257.54 | 258.02 | 257.85 | 292.65 | 292.21 | 290.63 |

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