# **Supplementary information**

# **Highly variable magmatic accretion at the ultraslow-spreading Gakkel Ridge**

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# **Highly variable magmatic accretion at the ultraslow-spreading Gakkel Ridge**

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## **Supplementary information**

This file contains Supplementary Figures 1–7 and Supplementary Tables 1–4.

#### **Supplementary figures**

- Fig. 1. Layout of seismic experiments and sampled rocks superimposed on the multibeam bathymetry map.
- Fig. 2. Travel time misfit of each pick for the inversion models.
- Fig. 3. Residual mantle Bouguer anomaly (RMBA) of the JASMInE region.
- Fig. 4. Estimated crustal thickness with a half spreading rate of 5 mm  $yr^{-1}$ .
- Fig. 5. Compilations of seismic-determined oceanic crustal thickness.
- Fig. 6. Gravity-derived crustal thickness of selected regions of slow- and ultraslowspreading ridges.
- Fig. 7. Processed and interpreted profile of sonobuoy data (B0801 site as an example).



**Supplementary Fig. 1 | Layout of seismic experiments and sampled rocks superimposed on the multibeam bathymetry map. a,** Bathymetry data and rock samples are collected from this study and the AMORE cruise<sup>1</sup>. Red dots with white outlines indicate the sampled basalts from this study. Dashed lines indicate the offset range where ice-station seismic signals could be identified<sup>2</sup>. **b-c**, Along- and across-axis profiles showing bathymetry (upper panels) and topography slope (lower panels). **d-h,** Photos of the JASMInE experiment. **d,** The icebreaker "Xuelong 2". **e,** G-gun sled being deployed. **f,** An OBS used in the JASMInE experiment. **g,** Shooting in sea ice. At times, severe sea ice conditions made the shot intermittent, and the shooting tracks deviated from the designed survey lines. **h,** Recovery of an OBS in sea ice.



**Supplementary Fig. 2 | Travel time misfit of each pick for the inversion models.** Red dots show the predicted travel times. The blue dots with grey vertical bars represent the observed reflection and refraction travel times. The OBS numbers are labelled in each panel. There are relatively large misfits for refraction arrivals of OBSs 13, 38, and 33. The root-mean-square misfits of Tomo2D models along the across-axis profile and the western and the eastern sections of the along-axis profile converged at 93, 96, and 81 ms, respectively.



**Supplementary Fig. 3 | Residual mantle Bouguer anomaly (RMBA) of the JASMInE region. a,** map of RMBA. The RMBA in the areas covered by transparency were calculated from the IBCAO 4.0 and DTU 17 data. In other regions, we used multibeam bathymetry data and shipborne gravity data to calculate the RMBA. The position of 85° E volcanic centre is marked by the red line. No obvious lateral migration track of the 85° E volcanic centre is observed here. The dashed frame marks the position of panel **b**. **b,** Across-axis profiles of gravity-derived RMBA illustrate the lower RMBA is persistent in the volcanic zone covered by high-resolution data. The lower RMBA is an indication of a thick crust or a lighter mantle. **c,** RMBA profiles. The thick grey line indicates the average of the nine profiles in part **b**. Note that the effect of sediment is not considered, which may result in the low RMBA at the southern end of the profiles. **d,** Gravity effects of thermal structures for RMBA correction and in the numerical mantle upwelling models. Zero level is arbitrary. **e,** Differences in gravity effects between thermal structures for RMBA correction and the numerical active mantle upwelling (yellow), passive mantle upwelling models with constant (red) and variable (black) viscosities.



**Supplementary Fig. 4 | Estimated crustal thickness with a half spreading rate of 5**  mm yr<sup>-1</sup>. Contours represent estimated crustal thicknesses from the model of passive mantle upwelling and wet melting<sup>3</sup>. The circles indicate the estimated crustal thickness with the measured  $T_p$  and water content in the JASMInE zone. The red circle shows the estimated crustal thickness with the average  $T_p$  (1,310 °C) and water content (250 ppm) of the JASMInE zone.



**Supplementary Fig. 5 | Compilations of seismic-determined oceanic crustal thickness. a,**  Map of the global mid-ocean ridge system. Positions of the seismic measurement shown in b are also shown. Triangles mark the segment-average crustal thickness from ref. <sup>4</sup>. **b**, Values of seismic-determined crustal thickness. Dots and squares represent the crustal thickness of the data points compiled by ref. <sup>5</sup> and ref. <sup>6</sup>, respectively. For comparison, results of ref. <sup>7</sup> with various mantle viscosity are also shown. In ref.  $^7$ ,  $\eta_0$  is the reference mantle viscosity at the depth just beneath the melting zone. Note that our buoyant models contain all the sources of buoyancy, including thermal, mantle depletion, and melt retention, while only melt retentionrelated buoyancy is considered in ref.<sup>7</sup>.



**Supplementary Fig. 6 | Gravity-derived crustal thickness of selected regions of slow- and**  ultraslow-spreading ridges. a, maps of crustal thickness at the SWIR 61°–66° E<sup>8</sup>, Mid-Cayman Spreading Centre (MCSC), MAR  $24^{\circ}-27^{\circ}$  N<sup>9</sup>, MAR  $2^{\circ}-5^{\circ}$  N, and MAR 45° N. Except for SWIR  $61^{\circ}$ –66° E and MAR 24°–27° N, the others are calculated in this study with the constraints of shipborne bathymetry data obtained from the Global Multi-Resolution Topography Synthesis<sup>10</sup> and IBCAO 4.0<sup>11</sup>, as well as satellite gravity data<sup>12</sup>. White lines indicate ridge axes. Red lines indicate the positions of the profile in **b**. The bathymetric map of the Gakkel at 19° E from IBCAO 4.0 is also shown. **b,** Across-axis profiles of gravity-derived crustal thickness illustrate a higher variation at ultraslow- than at slow-spreading ridges.



**Supplementary Fig. 7 | Processed and interpreted profile of sonobuoy data (B0801 site as an example). a,** Processed profile of sonobuoy B0801. The thick black line is interpreted as the sediment basement according to abrupt changes in seismic amplitudes. Seismic reflection profile (ARC 14-07)13 near the sonobuoy B0801 for calibrating the sediment thickness. **b,** Map shows the track of sonobuoy B0801 (cyan line), location of the profile in panel **a** (red line), shooting track (yellow line), and ARC 14-07 (dashed line). At the position closest to sonobuoy B0801 track, the sediment thickness on the ARC 14-07 is approximately 0.8 km. Each sonobuoy was equipped with a GPS and transmitted real-time coordinate data to the deck.

### **Supplementary tables**

- Table 1. Sr-Nd isotopic compositions on MORB lavas from JASMInE.
- Table 2. Representative major element compositions and water content of the glasses from JASMInE (In Excel format).
- Table 3. Summary of identified seismic phases.
- Table 4. Parameters used in numerical models.



## **Supplementary Table 1 | Sr-Nd isotopic compositions on MORB lavas from JASMInE.**

Note: ε<sub>Nd</sub> is calculated as  $[(^{143}Nd^{144}Nd)$ sample/( $^{143}Nd^{144}Nd$ )chondrites -1] × 10,000, where  $^{143}Nd^{144}Nd$ chondrites=0.512638.

**Supplementary Table 2 | Representative major element compositions and water content of the glasses from JASMInE.** In Excel format.



#### **Supplementary Table 3 | Summary of identified seismic phases\* .**

\* "√" represent the seismic phase can be identified in the OBS.



## **Supplementary Table 4 | Parameters used in numerical models.**

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