

# Supporting Information

Gannoun et al. 10.1073/pnas.1017332108

## SI Materials and Methods

**Sample Selection and Dissolution.** Enstatite chondrites consist of highly reduced mineral assemblages formed under very low oxygen fugacity ( $fO_2$ ) conditions. They are classified into two subgroups on the basis of the Fe/Si bulk ratios: the high-Fe, high siderophile group (EH) and low Fe, low siderophile group (EL). Both EH and EL comprise a series of members from unequilibrated to highly equilibrated ones. The EL chondrites contain more highly equilibrated members than the EH chondrites. Thirteen enstatites chondrites belonging to both EH and EL subgroups were investigated during this study. This suite comprises eight unequilibrated samples (EH3, EL3) and five of other petrographic type (from type 4 to 6). Additionally, one ordinary chondrite (Jilin, H5) and one carbonaceous chondrite (Allende, CV3) were also analyzed during this study for the sake of comparison. Chondrite samples analyzed for Sm–Nd systematics and measured  $^{147}\text{Sm}/^{144}\text{Nd}$  ratios are presented in Table S1. Different digestion techniques have been tested in the past to dissolve the carbonaceous chondrite Allende: (i) sealed beaker on hot plate using a mixture of HF:HNO<sub>3</sub> concentrated acid, (ii) again the same reagents but in bombs (higher pressure-temperature conditions), and (iii) fusion. Although this sample contains a large amount of both refractory inclusions and presolar grains, no difference has been found. In fact we measured each time very slightly different Sm and Nd isotope composition but this can be explained by the heterogeneity of the powder itself. The isotope heterogeneity at the bulk sample size (dissolution around 100 mg to 1 g) is also illustrated in Fig. S1. The large range of  $^{143}\text{Nd}/^{144}\text{Nd}$  and  $^{147}\text{Sm}/^{144}\text{Nd}$  ratios measured for different dissolution of Allende indicates that this sample is not homogeneous, even when a large quantity of powder is used for dissolution (about 1 g). But this range is not due to incomplete dissolution of refractory phases. Because of its low blank and reliable isotope data results we opt for steel-jacketed tetrafluoroethylene (TFE) Teflon lined Parr bombs for sample digestion.

About one gram of homogenized powdered samples were dissolved in a HF–HNO<sub>3</sub> mixture in the proportions of 9:3 in steel-jacketed TFE Teflon lined Parr bombs. Samples were digested in an oven at 150 °C for not less than 72 h. The solution is dried after the add of few mL of HClO<sub>4</sub> and the residue redissolved with the mixture of HCl and HNO<sub>3</sub> ultrapure acids in the proportions 1:2, transferred in Savillex perfluoroalkoxy 60 mL beakers. The beakers were placed in an oven at 150 °C for 48 h. Once a clear solution was obtained, a 5% aliquot was taken and spiked with  $^{149}\text{Sm}$ – $^{150}\text{Nd}$  tracer to measure Sm and Nd concentrations by isotope dilution. This spike was calibrated against standard solutions prepared from AMES and JNdi-1. Sm and Nd results on Allende using this spike are concordant with literature values (Fig. S1).

**Chemical Separation and Purification of Sm and Nd.** The full chemical procedure for the extraction and purification of Sm and Nd followed closely techniques described previously (1). The Sm and Nd fractions were separated with a two stage-chemistry procedure. In the first stage rare earth elements (REEs) as a group were separated by using cation resin and HCl. The REE split were then processed two times through a cation column using 2-Methylactic acid (0.2M and pH = 4.7) with a small amount of H<sub>2</sub>O<sub>2</sub> to ensure perfect separation of Nd from Ce and Sm by reducing the effect of interferences on mass 142 (from Ce) and on mass 144, 148 and 150 from Sm. The last step consists on one pass through Ln-spec resin in weak HCl acid. Organic

residues are then completely removed and the sample is ready to be loaded on Re filament. Unspiked Sm fraction from each chondrite sample was also purified using the same technique. Total procedural blanks for Sm and Nd were 1.5 and 8.3 pg, respectively. The percentage blank contribution to each sample was always less than 0.02‰, so no blank correction has been applied. The spiked fractions have been separated using the method developed by Pin et al. (2). Ce and Sm interferences on masses 142 and 144 and expressed in ppm are indicated in Table S1. Sm interferences are negligible whereas for a few samples the Ce signal accounts for more than a few tens ppm on mass 142. The Ce signal decreases rapidly during the isotope measurement. Fig. S2 shows that there is no covariation between  $^{142}\text{Nd}/^{144}\text{Nd}$  ratios and either  $^{142}\text{Ce}/^{142}\text{Nd}$  or  $^{144}\text{Sm}/^{144}\text{Nd}$  ratios.

As for Nd, the purified fractions of Sm obtained using 2-Methylactic acid were passed through Ln-spec resin and eluted in 0.5N HCl to discard all organic residues.

**Mass Spectrometry.** All the isotopic measurements presented in this study were obtained on the Thermo-Fisher Triton thermal ionization mass spectrometer at Laboratoire Magmas et Volcans. The purified Sm and Nd cuts were loaded in 2.5N HCl on out-gassed zone-refined Re filaments. The Nd was measured in static mode by using the nine faraday cups, as positive metal ions following the configuration presented below. Nd isotope measurements using both dynamic and static modes have been compared. No significant difference has been observed for the measurement of 3.8 Ga old samples coming from the Isua supracrustal belt (SW Greenland) that present  $^{142}\text{Nd}$  excesses of about 15 ppm (3). Because enstatite chondrites present the lowest Nd contents among chondrites, we have decided to use a static procedure for a better ion-counting statistic. The Faraday cup configuration used for Nd isotope measurements (mass) as follows: L4,  $^{140}\text{Ce}$ ; L3,  $^{142}\text{Nd}$ ; L2,  $^{143}\text{Nd}$ ; L1,  $^{144}\text{Nd}$ ; C,  $^{145}\text{Nd}$ ; H1,  $^{146}\text{Nd}$ ; H2,  $^{147}\text{Sm}$ ; H3,  $^{148}\text{Nd}$ ; H4,  $^{150}\text{Nd}$ .

To test the robustness of the Nd isotope data obtained by static mode, we have analyzed several fractions of JNdi-1 standard enriched artificially with known amount of  $^{142}\text{Nd}$  spike. The results are plotted in Fig. S3 and show that even for low amount of Nd (100 ng), artificial excesses of  $^{142}\text{Nd}$  as low as 17 ppm can be resolved.

All measured data were corrected for instrumental mass fractionation using the exponential law and  $^{146}\text{Nd}/^{144}\text{Nd} = 0.7219$  (Table S1). Each measurement consists of 27 blocks of 20 ratios (approximately 8 s for integration time) using the amplifier rotation and taking the background before each block. Gain was taken before the beginning of each run. The  $^{142}\text{Nd}$  signals range between 0.9 and  $4.5 \times 10^{-11}$  A for all chondrite samples. The  $^{142}\text{Nd}/^{144}\text{Nd}$  data are expressed in ppm relative to the mean value obtained on the JNdi-1 standards measured during the same campaign. Our current mean value for  $^{142}\text{Nd}/^{144}\text{Nd}$  in the JNdi-1 standard is  $1.1418331 \pm 0.000066$  ( $2\sigma_m$ ). The external reproducibility on repeated standard analyses is always better than 6 ppm for the given campaign. No correlation is observed between the different Nd ratios (Fig. S4).

The Sm was measured in static mode as  $\text{Sm}^+$  ions using double Re filaments following the cup configuration presented here: L4,  $^{144}\text{Sm}$ ; L3,  $^{146}\text{Nd}$ ; L2,  $^{147}\text{Sm}$ ; L1,  $^{148}\text{Sm}$ ; C,  $^{149}\text{Sm}$ ; H1,  $^{150}\text{Sm}$ ; H2,  $^{152}\text{Sm}$ ; H3,  $^{154}\text{Sm}$ ; H4,  $^{156}\text{Gd}$ .

Each run consists of 9 blocks of 20 cycles using the amplifier rotation and taking the background before each block. The  $^{152}\text{Sm}$  signals range between 0.3 and  $1.9 \times 10^{-11}$  A. Possible isobaric

interferences from Nd and Gd were monitored by measuring the intensity on  $^{146}\text{Nd}$  and  $^{156}\text{Gd}$  masses (Table S2). Sm data were corrected for instrumental mass fractionation using the exponential law and  $^{147}\text{Sm}/^{152}\text{Sm} = 0.56081$ . Sm isotope compositions for both standards and samples are given in Table S2. Sm isotope compositions have been measured during two different periods. Internal errors are higher in the first sequence of measurement than in the second one. We have encountered a problem with the Faraday cups during the first sequence of measurements. Then these samples have been measured twice and the Sm isotope composition presented in Table S2 for sequence 1 corresponds to the second run. Because most of samples have been measured at very low intensity ( $^{152}\text{Sm} < 0.4 \text{ V}$ ),  $^{144}\text{Sm}/^{152}\text{Sm}$  ratios have large errors. This is the reason why these samples have not been represented in the diagram presented in Fig. S5.

We have also calculated the Sm isotope ratios relative to  $^{154}\text{Sm}$  (100% r-process) and all data are fractionation corrected using the exponential law and  $^{147}\text{Sm}/^{154}\text{Sm} = 0.65918$ . Using this normalization scheme (Table S2), the deviations expressed in  $\mu$ -notation (ppm deviation relative to the ratio measured in terrestrial standard) are similar to those from previous calculation within the error bar.

**Correction on  $^{142}\text{Nd}/^{144}\text{Nd}$  Ratios.** Exposure to galactic cosmic rays modifies a sample's Sm and Nd isotopic composition and its  $^{142}\text{Nd}/^{144}\text{Nd}$  ratio. A long exposure age tends to decrease the  $^{142}\text{Nd}/^{144}\text{Nd}$  ratio as observed in lunar samples (4). The second-

ary neutron capture effect is currently quantified using Sm isotope measurements because  $^{149}\text{Sm}$  has the largest thermal neutron capture cross section (e.g., ref. 5). The neutron capture effect is clearly seen in a negative correlation between  $^{150}\text{Sm}/^{152}\text{Sm}$  and  $^{149}\text{Sm}/^{152}\text{Sm}$  ratios (Fig. S5), but  $^{149}\text{Sm}$  deficits are lower than  $-1.5$  epsilon-unit, which corresponds to a correction smaller than 1 ppm on the  $^{142}\text{Nd}/^{144}\text{Nd}$  ratios following the method developed by Nyquist et al. (6) or Rankenburg et al. (7). As is clearly demonstrated in the  $^{149}\text{Sm}$ - $^{150}\text{Sm}$  isotope diagram, several samples do not plot on the correlation line and the observed scatter could reflect isotope heterogeneities in the solar nebula when chondrites were formed.

The  $^{142}\text{Nd}/^{144}\text{Nd}$  ratios presented in Fig. 1B have been corrected for a common evolution considering a constant  $^{147}\text{Sm}/^{144}\text{Nd}$  ratio [0.196 (8) from 4.568 Ga and an initial solar system  $^{146}\text{Sm}/^{144}\text{Sm}$  ratio equal to 0.0085 (9)]. The  $^{147}\text{Sm}/^{144}\text{Nd}$  ratio considered for each sample is not the ratio measured on the aliquot but this ratio has been recalculated using measured  $^{143}\text{Nd}/^{144}\text{Nd}$  ratio and assuming that these samples have evolved from the same chondritic reservoir (present day values are 0.1960 and 0.512630 for  $^{147}\text{Sm}/^{144}\text{Nd}$  ratio and  $^{143}\text{Nd}/^{144}\text{Nd}$  ratios, respectively (8)). Measured  $^{142}\text{Nd}/^{144}\text{Nd}$  ratios are compared to corrected  $^{142}\text{Nd}/^{144}\text{Nd}$  ratios in Table S3. The  $^{142}\text{Nd}/^{144}\text{Nd}$  ratios reported in Fig. 1A are those from the first column whereas in all other diagrams we have plotted corrected  $^{142}\text{Nd}/^{144}\text{Nd}$  ratios presented in the last column of the Table S3.

1. Boyet M, Carlson RW (2005)  $^{142}\text{Nd}$  evidence for early (>4.53 Ga) global differentiation of the silicate earth. *Science* 309:576–581.
2. Pin C, Santos Zalduegui JF (1997) Sequential separation of light rare-earth elements, thorium and uranium by miniaturized extraction chromatography: Application to isotopic analyses of silicate rocks. *Anal Chim Acta* 339:79–89.
3. Rizo Garza HL, Boyet MM, Blichert-Toft J, Rosing M, Gannoun A (2010) A coupled Nd and Hf isotopic study of Isua Archean rocks and the differentiation of the Hadean mantle. AGU Fall Meeting (American Geophysical Union, San Francisco).
4. Boyet M, Carlson RW (2007) Early lunar differentiation and a non magma ocean origin for the lunar crust. *Earth Planet Sci Lett* 262:505–516.
5. Lingenfelter RE, Canfield EH, Hampel VE (1972) The lunar neutron flux revisited. *Earth Planet Sci Lett* 16:335–369.
6. Nyquist LE, et al. (1995)  $^{146}\text{Sm}$ - $^{142}\text{Nd}$  formation interval for the lunar mantle material. *Geochim Cosmochim Acta* 59:2817–2837.
7. Rankenburg K, Brandon AD, Neal CR (2006) Neodymium isotope evidence for a chondritic composition of the moon. *Science* 312:1369–1372.
8. Bouvier A, Vervoort JD, Patchett PJ (2008) The Lu–Hf and Sm–Nd isotopic composition of CHUR: Constraints from unequibrated chondrites and implications for the bulk composition of terrestrial planets. *Earth Planet Sci Lett* 273:48–57.
9. Boyet M, Carlson RW, Horan M (2010) Old Sm–Nd ages for cumulate eucrites and redetermination of the solar system initial  $^{146}\text{Sm}/^{144}\text{Sm}$  ratio. *Earth Planet Sci Lett* 291:172–181.





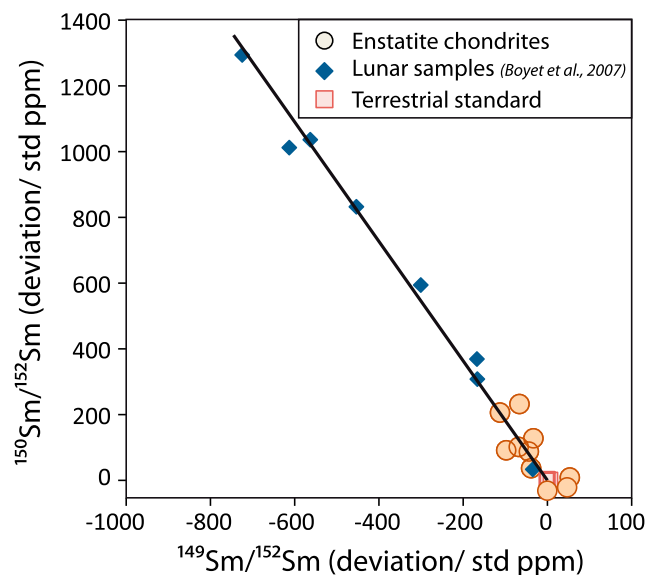


Fig. S5. The  $^{149}\text{Sm}/^{152}\text{Sm}$  vs.  $^{150}\text{Sm}/^{152}\text{Sm}$  isotopic covariation for lunar samples and enstatite chondrites.

Table S1. The  $^{147}\text{Sm}/^{144}\text{Nd}$  ratios and concentration data for enstatite chondrites analyzed in this study

Sample	Type	Source	Weight (g)	Sm (ppm)	Nd (ppm)	$^{147}\text{Sm}/^{144}\text{Nd}$
ALHA 77295	EH3	NASA	0.9135	0.1411	0.4330	0.1970
ALHA duplicate			0.9150	0.1143	0.3439	0.2009
Allende	CV3	Smith. Inst. *	1.0018	0.2406	0.7350	0.1979
Eagle 4739	EL6 Fall	AMNH †	0.9147	0.1078	0.3256	0.2001
Hvittis	EL6	dealers	0.9136	0.1199	0.3611	0.2007
Indarch	EH4	NHM ‡	0.9577	0.1013	0.3050	0.2009
Jilin	H5 Fall	NHM	1.1489	0.1732	0.5500	0.1904
Khairpur	EL6 Fall	NHM	0.9696	0.1336	0.4111	0.1965
Kota-Kota	EH3	NHM	0.9290	0.1951	0.7040	0.1675
Kota-Kota duplicate			0.9011	0.2018	0.7309	0.1669
MAC 02837	EL3	NASA	0.9370	0.1634	0.4972	0.1987
MAC 02839	EL3	NASA	1.0544	0.1871	0.5723	0.1976
MAC 88180	EL3	NASA	0.9273	0.2673	0.8187	0.1974
Sahara 97072	EH3	dealers	0.9158	0.1533	0.4942	0.1875
Sahara 97096	EH3	dealers	1.1729	0.1396	0.4305	0.1961
Sahara 97158	EH3	dealers	0.9408	0.1383	0.4404	0.1899
St. Mark's 4804	EH5 Fall	AMNH	0.9572	0.1460	0.4519	0.1953

\*Smithsonian Institution.

†American Museum of Natural History.

‡Natural History Museum of London.









Table S3. Measured Sm isotope composition of standards and samples

Sm isotope ratios normalized to $^{147}\text{Sm}/^{152}\text{Sm} = 0.56081$			$^{146}\text{Nd}/^{152}\text{Sm}$		$^{156}\text{Gd}/^{152}\text{Sm}$		$^{144}\text{Sm}/^{152}\text{Sm}$		$^{148}\text{Sm}/^{152}\text{Sm}$		$^{149}\text{Sm}/^{152}\text{Sm}$		$^{150}\text{Sm}/^{152}\text{Sm}$		$^{154}\text{Sm}/^{152}\text{Sm}$		2 s		
	Int.	152Sm	ppm		ppm		ppm	2 s	2 s	2 s	2 s	2 s	2 s	2 s	2 s	2 s	2 s	2 s	
Sequence 1	(V)																		
Standard Sm	2.96	0.48	7.35	0.114975	0.420437	0.000001	0.516856	0.000002	0.275985	0.000001	0.850761	0.000004							
Standard Sm	4.55	0.19	2.14	0.114976	0.420438	0.000001	0.516856	0.000002	0.275985	0.000001	0.850760	0.000003							
Standard Sm	4.71	0.12	1.85	0.114976	0.420436	0.000001	0.516856	0.000002	0.275984	0.000001	0.850765	0.000003							
Standard Sm	4.88	-0.90	1.45	0.114979	0.420435	0.000001	0.516854	0.000002	0.275984	0.000001	0.850770	0.000003							
Standard Sm	5.64	-0.14	0.62	0.114980	0.420434	0.000001	0.516853	0.000001	0.275983	0.000001	0.850777	0.000003							
Average (n = 5)				0.114977	0.420436	0.000004	0.516856	0.000003	0.275984	0.000002	0.850767	0.000014							
2 s ppm				35.9	7.3	0.000014	9.2	0.000003	7.6	0.000002	16.9	0.000014							
Eagle	0.34	28.60	-4.90	0.114982	0.420417	0.000014	0.516836	0.000014	0.275994	0.000011	0.850792	0.000021							
Indarch	1.84	33.30	2.79	0.114979	0.420431	0.000003	0.516833	0.000006	0.276008	0.000003	0.850781	0.000010							
Kairpur	0.63	38.30	-3.12	0.114982	0.420440	0.000011	0.516805	0.000012	0.276009	0.000010	0.850790	0.000018							
Kota-Kota	0.77	35.40	1.62	0.114977	0.420422	0.000005	0.516883	0.000008	0.275979	0.000006	0.850751	0.000015							
MAC 02839	0.31	11.90	58.60	0.114991	0.420434	0.000013	0.516820	0.000016	0.276012	0.000012	0.850833	0.000024							
St. Mark's	0.36	30.90	11.10	0.114997	0.420454	0.000014	0.516856	0.000015	0.275977	0.000012	0.850766	0.000030							
Sahara 97096	0.36	20.80	9.88	0.114980	0.420422	0.000014	0.516797	0.000012	0.276041	0.000011	0.850825	0.000020							
Sequence 2																			
Standard Sm	1.92	1.82	11.20	0.114978	0.420428	0.000001	0.516856	0.000003	0.275982	0.000002	0.850769	0.000005							
Standard Sm	2.65	0.60	1.42	0.114979	0.420427	0.000001	0.516855	0.000002	0.275984	0.000001	0.850769	0.000003							
Standard Sm	3.20	0.56	8.25	0.114978	0.420428	0.000001	0.516855	0.000002	0.275982	0.000001	0.850756	0.000003							
Standard Sm	2.70	0.74	6.36	0.114978	0.420428	0.000001	0.516855	0.000001	0.275983	0.000002	0.850765	0.000016							
Average (n = 5)				0.114978	0.420428	0.000001	0.516855	0.000001	0.275983	0.000002	0.850765	0.000013							
2 s ppm				5.5	1.6	0.000002	1.9	0.000003	7.1	0.000002	15.0	0.000005							
Hvittis	1.38	28.90	8.68	0.114982	0.420425	0.000002	0.516821	0.000003	0.276047	0.000002	0.850782	0.000005							
MAC 88180	1.70	5.91	1.88	0.114984	0.420433	0.000001	0.516883	0.000002	0.275985	0.000002	0.850757	0.000004							
Sahara 97158	0.59	26.40	51.70	0.114975	0.420444	0.000005	0.516838	0.000004	0.276018	0.000004	0.850805	0.000008							
Sm isotope ratios normalized to $^{147}\text{Sm}/^{154}\text{Sm} = 0.65918$																			
Sequence 1																			
Standard Sm	0.135142	0.000001	0.494185	0.000002	0.607518	0.000003	0.324396	0.000003	0.175415	0.000002	1.175415	0.000004							
Standard Sm	0.135143	0.000001	0.494186	0.000002	0.607522	0.000002	0.324396	0.000002	1.175416	0.000001	1.175416	0.000003							
Standard Sm	0.135143	0.000001	0.494184	0.000001	0.607517	0.000001	0.324394	0.000001	1.175411	0.000001	1.175411	0.000003							
Standard Sm	0.135147	0.000001	0.494182	0.000002	0.607514	0.000002	0.324394	0.000002	1.175406	0.000001	1.175406	0.000003							
Standard Sm	0.135149	0.000001	0.494180	0.000002	0.607511	0.000002	0.324391	0.000002	1.175399	0.000001	1.175399	0.000003							
Average (n = 5)	0.135145	0.000006	0.494183	0.000005	0.607516	0.000008	0.324394	0.000008	1.175410	0.000005	1.175410	0.000014							
2 s ppm	42.4		9.4	13.0	14.2	0.000015	14.2	0.000016	12.2	0.000005	12.2	0.000014							
Eagle	0.135155	0.000016	0.494159	0.000017	0.607483	0.000019	0.324400	0.000019	1.175384	0.000013	1.175384	0.000021							
Indarch	0.135147	0.000004	0.494176	0.000007	0.607486	0.000007	0.324419	0.000007	1.175395	0.000004	1.175395	0.000010							
Kairpur	0.135153	0.000012	0.494186	0.000013	0.607450	0.000016	0.324424	0.000016	1.175387	0.000011	1.175387	0.000018							
Kota-Kota	0.135144	0.000006	0.494169	0.000009	0.607551	0.000010	0.324391	0.000010	1.175425	0.000007	1.175425	0.000014							
MAC 02839	0.135164	0.000015	0.494176	0.000019	0.607450	0.000021	0.324414	0.000021	1.175343	0.000014	1.175343	0.000024							
St. Mark's	0.135171	0.000016	0.494202	0.000017	0.607521	0.000019	0.324393	0.000019	1.175410	0.000014	1.175410	0.000030							
Sahara 97096	0.135150	0.000016	0.494165	0.000015	0.607438	0.000016	0.324460	0.000016	1.175352	0.000014	1.175352	0.000020							
Sequence 2																			
Standard Sm	0.135145	0.000002	0.494174	0.000003	0.607517	0.000003	0.324391	0.000003	1.175407	0.000002	1.175407	0.000005							
Standard Sm	0.135146	0.000001	0.494174	0.000002	0.607516	0.000002	0.324394	0.000002	1.175408	0.000001	1.175408	0.000003							
Standard Sm	0.135146	0.000001	0.494175	0.000002	0.607518	0.000002	0.324394	0.000002	1.175421	0.000001	1.175421	0.000003							
Standard Sm	0.135146	0.000001	0.494175	0.000002	0.607517	0.000002	0.324394	0.000002	1.175421	0.000001	1.175421	0.000003							
Standard Sm	0.135146	0.000001	0.494175	0.000002	0.607518	0.000002	0.324394	0.000002	1.175421	0.000001	1.175421	0.000003							
Average (n = 5)	0.135146	0.000001	0.494175	0.000002	0.607517	0.000002	0.324393	0.000002	1.175414	0.000003	1.175414	0.000005							

Sm isotope ratios normalized to  $^{147}\text{Sm}/^{154}\text{Sm} = 0.65918$

	$^{144}\text{Sm}/^{154}\text{Sm}$	2 s	$^{148}\text{Sm}/^{154}\text{Sm}$	2 s	$^{149}\text{Sm}/^{154}\text{Sm}$	2 s	$^{150}\text{Sm}/^{154}\text{Sm}$	2 s	$^{152}\text{Sm}/^{154}\text{Sm}$	2 s
2 s ppm	6.0		3.9		3.0		9.8		13.1	
Hvittis	0.135151	0.000002	0.494170	0.000004	0.607473	0.000004	0.324463	0.000003	1.175392	0.000006
MAC 88180	0.135152	0.000001	0.494181	0.000003	0.607550	0.000003	0.324397	0.000002	1.175419	0.000004
Sahara 97158	0.135147	0.000005	0.494187	0.000006	0.607491	0.000007	0.324427	0.000005	1.175363	0.000010

Table S4. The  $^{142}\text{Nd}/^{144}\text{Nd}$  ratios are expressed in ppm deviation from the value obtained for the JNdi-1 terrestrial standard

Sample name	$^{142}\text{Nd}/^{144}\text{Nd}$ measured *	$2\sigma$ †	$^{147}\text{Sm}/^{144}\text{Nd}$ measured ‡	$^{142}\text{Nd}/^{144}\text{Nd}$ corr. §	$^{147}\text{Sm}/^{144}\text{Nd}$ cal. ¶	$^{142}\text{Nd}/^{144}\text{Nd}$ corr.	Reference
Abee	-43.6	17.1	0.1903	-35.4	0.1946	-41.5	(1)
ALHA77295	-2.3	7.4	0.1970	-3.6	0.1952	-1.2	this study
Eagle	-10.9	6.0	0.2001	-16.8	0.1982	-14.1	this study
Hvittis	-1.8	5.5	0.2007	-8.6	0.2000	-7.5	this study
Indarch	-14.6	5.0	0.1934	-10.9	0.1958	-14.3	(2)
Indarch	-17.7	7.4	0.2009	-24.8	0.1958	-17.4	this study
Khairpur	-13.2	5.1	0.1965	-13.9	0.1961	-13.4	this study
Kota-Kota	2.4	4.4	0.1675	43.3	0.1903	10.6	this study
MAC 02837	-4.4	0.3	0.1987	-8.2	0.1967	-5.4	this study
MAC 02839	-4.2	2.1	0.1976	-6.5	0.1964	-4.7	this study
MAC 88180	3.1	4.8	0.1974	1.1	0.1993	-1.7	this study
Sahara 97072	-16.1	4.3	0.1875	-3.9	0.1923	-10.8	this study
Sahara 97096	-19.9	4.3	0.1961	-20.0	0.1948	-18.2	this study
Sahara 97158	-13.9	9.9	0.1899	-5.1	0.1937	-10.7	this study
St. Mark's	-17.0	5.9	0.1953	-16.0	0.1951	-15.7	this study

\*The  $^{142}\text{Nd}/^{144}\text{Nd}$  ratios corrected for isobaric interferences and mass fractionation.

†Internal errors.

‡Measured  $^{147}\text{Sm}/^{144}\text{Nd}$  ratios.

§The  $^{142}\text{Nd}/^{144}\text{Nd}$  ratios corrected for a common evolution considering a constant  $^{147}\text{Sm}/^{144}\text{Nd}$  ratio [0.196, (3)] from 4.568 Ga and an initial solar system  $^{146}\text{Sm}/^{144}\text{Sm}$  ratio equal to 0.0085 (4). We have used the measured  $^{147}\text{Sm}/^{144}\text{Nd}$  ratios.

¶The  $^{147}\text{Sm}/^{144}\text{Nd}$  ratios calculated from measured  $^{143}\text{Nd}/^{144}\text{Nd}$  ratios assuming that all meteorites evolve from the same chondritic reservoir since 4.568 Ga.

||Corrected  $^{142}\text{Nd}/^{144}\text{Nd}$  ratios following the method explained above but using calculated  $^{147}\text{Sm}/^{144}\text{Nd}$  ratios instead of measured ratios. Note that the measured and calculated  $^{142}\text{Nd}/^{144}\text{Nd}$  deviations are similar considering the analytical errors except for Kota-Kota, which is characterized by unusual low  $^{147}\text{Sm}/^{144}\text{Nd}$  and  $^{143}\text{Nd}/^{144}\text{Nd}$  ratios.

1 Boyet M, Carlson RW (2005)  $^{142}\text{Nd}$  evidence for early (>4.53 Ga) global differentiation of the silicate earth. *Science* 309:576–581.

2 Carlson RW, Boyet M, Horan MF (2007) Chondrite barium, neodymium, and samarium isotopic heterogeneity and early earth differentiation. *Science* 316:1175–1178.

3 Bouvier A, Vervoort JD, Patchett PJ (2008) The Lu–Hf and Sm–Nd isotopic composition of CHUR: Constraints from unequilibrated chondrites and implications for the bulk composition of terrestrial planets. *Earth Planet Sci Lett* 273:48–57.

4 Boyet M, Carlson RW, Horan M (2010) Old Sm–Nd ages for cumulate eucrites and redetermination of the solar system initial  $^{146}\text{Sm}/^{144}\text{Sm}$  ratio. *Earth Planet Sci Lett* 291:172–181.