Supporting Information Appendix for

The Mechanism of Burgess Shale-type Preservation

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- Table S1. The isotopic composition of sulfur pyrite in Chengjiang and 6 other principal Burgess shale type deposits.
- Table S2. δ^{13} C vs. δ^{18} O of carbonate cements from BST-bearing claystones

Figure S1: Stratigraphic distribution of the principal Burgess Shale-type deposits across Lower (Series 2) and Middle (Series 3) Cambrian strata, modified from Conway Morris (1). Deposits analyzed in this study are highlighted in yellow. The recently- reported BST locality in the "thin" Stephen Formation of British Columbia (2) as not among the principal BST deposits identified by Conway Morris (1). Among these deposits, the Chengjiang and Burgess Shale are, by far, the most important in diversity and abundance of soft-bodied fossil biotas that they have yielded (3, 4). Mount Cap is known only from the subsurface (5), and Kinzers/Parker and Sirius Passet have experienced extensive metamorphism (e.g. (6, 7)).

Figure S2: SEM micrographs showing BST claystone microfabrics from the Chengjiang (A), Kaili Fm. (B), Spence Shale Mbr. (C), "thin" Stephen Fm. (D), Wheeler Fm. (E) and Marjum Fm. (F). All surfaces analyzed were freshly broken perpendicular to bedding. In all cases BSTbearing claystones exhibit randomly-oriented clay-microfabrics, indicative of rapid deposition from turbid suspension by bottom flowing currents (8). These images also demonstrate the predominance of clay minerals and the absence of silt or coarser-grained particles from the BST facies of all deposits. An SEM micrograph of clay fabric from the Walcott Quarry Member of the Burgess Shale is shown in Figure 2A.

Figure S3: Thin section micrographs (A, B) and X-radiographs of slab samples (C-E), illustrating the bed-capping distribution of authigenic carbonate cements in BST-bearing claystones. (A) Thin section of Marjum Formation shown in transmitted light. Carbonate cements at bed tops appear bright, clay-rich bed bases appear gray. (B) Cathodoluminescence micrograph of thin section of BST interval of Kaili Formation, arrows mark base (lower) and top (upper) of a typical BST-bearing claystone lamina, 3mm in thickness. Carbonates, which exhibit bright orange luminesce, are concentrated at bed tops and become less abundant down bed. The dark, non-luminescent portions of the image are clay-rich with subsidiary carbonate dispersed throughout the claystone fabric. (C-E) X-radiographs of slab samples of BST –bearing claystones. Carbonate-rich portions of event-deposited laminae are less easily penetrated by Xrays and appear bright in X-radiograph, whereas clay-rich portions appear dark in color, revealing concentration of carbonate cements at the tops of individual mm-scale laminae. (C) "thin" Stephen Formation, Stanley Glacier locality. (D) Wheeler Formation, Drum Mountains locality. (E) Marium Formation, Marium Pass Locality.

Figure S4. Energy Dispersive X-ray Spectroscopy (EDX) elemental maps of bed tops of BST claystones showing concentration of calcium carbonate cements in bed tops, with polished slabs at left showing the locations of the two areas represented in the elemental maps (boxes). In each elemental map, brightness corresponds to elemental abundance. (A) Burgess Shale, Walcott Quarry Member, Great *Eldoniia* Layer (GEL). Elemental maps of Ca and Fe reveal concentration of calcite and pyrite at the bed top, coincident with decreases in Al and K, which reside in clay minerals. Al and K are concentrated lower in the bed, indicating an increasing proportion of clay relative to calcium carbonate, as also shown in Figure 2. Arrow indicates vertical extent of the GEL. (B) BST-bearing interval of the Wheeler Formation, Drum Mountains locality. Elemental maps of Ca and Mg indicate concentration of magnesian calcite at bed top. Some Mg also resides in clay minerals, and as a result, Ca and Mg do not share the same precise distribution in the sample. Elemental map of Fe indicates concentration of pyrite at bed top, coincident with a decrease in Al, which resides in clay minerals, indicating an increasing proportion of clay minerals relative to calcium carbonate down the bed.

Figure S5. Plot of $\delta^{13}C$ vs. $\delta^{18}O$ of 103 analyses of carbonate cements from the Haikou core, including the Chengjiang interval, and the underlying "Upper Black Shale" and Shiyantou Formations. The two clusters of data apparent on the plot comprise fields of early vs. late diagenetic cements. Shiyantou Fm. cements contain isotopic values typical of diageneitc carbonate cements in fine-grained siliciclastics. They are characterized by more negative $\delta^{13}C$ and δ^{18} O values, indicating significant input of organodiagentically derived bicarbonate and higher temperatures of crystallization. In contrast, Chengjiang carbonate cements are characterized by δ^{13} C values that lie within the range of reported Early Cambrian seawater values (9) and display less negative $\delta^{18}O$ values, indicating early precipitation at lower temperature from a seawater source, with minor organic contribution.

Figure S6. Distribution of $\delta^{34}S_{\text{pyrite}}$ values across single event-deposited BST-bearing claystone beds in the Walcott Quarry Member of the Burgess Shale (A-B) and the Chengjiang deposit (C). Low fractionation of $\delta^{34}S_{\text{pyrite}}$ values from Cambrian seawater SO_4^2 values provides clear evidence of SO_4^2 limitation within each bed during pyrite precipitation (10) (Fig. 3, Table S1). However, $\delta^{34}S_{\text{pyrite}}$ values do not show consistent trends towards heavier values down-bed, as would be expected if SO_4^2 limitation within the porewaters was induced by reduced diffusion of SO_4^2 into the beds as a result of event-driven deposition alone (11). Therefore, an additional mechanism of SO₄² limitation during early diagenesis is required. (A) Great *Eldoniia* Layer, Walcott Quarry Member, Burgess Shale Fm. (2.8 cm thick) (B) Great *Marrella* Layer, Walcott Quarry, Burgess Shale Fm. (7.6 cm thick). (C) BST-bearing event bed from the Chengjiang deposit, Haikou core, 23.30m depth (2.1 cm thick). $\delta^{34}S_{\text{pyrite}}$ values for additional subsamples taken from cross-bed transects within single event-deposited beds of the Chengjiang deposit are reported in table S1.

Figure S7. Plot of $\delta^{34}S_{\text{pyrite}}$ vs. bed thickness for all BST deposits analyzed, showing absence of correlation.

Figure S8. Time required to precipitate pyrite contents measured in fossiliferous beds of BST deposits, calculated for a Cambrian oceanic sulfate concentration of 5 mM (12). The vertical flux is calculated as $J(x)=\Phi * Ds * (\partial C/\partial x)$ (13), where porosity (Φ) is 0.7. The diffusion coefficient (D_s) for sulfate in sediment is calculated as $D_s = D_{sw}/(1+n(1-\Phi))$, where n=2.7 (14) and D_{sw} for sulfate at 10°C = 0.68e-5 cm²/s (15). The gradient of sulfate (∂C/∂x) is set over 10 cm. Under this simple approximation, the average pyrite content of fossil-bearing beds from Chengjiang (dashed upper grey line) required no longer than 1 month to precipitate, while far less time was required to form the pyrite observed in other BST deposits (Burgess Shale, Wheeler Formation (WS), Kaili (K), Spence Shale (SS), Stephen Fm (SF), Marjum Fm (MF)).

Table S1. The isotopic composition of pyrite sulfur from 6 of the principal Burgess Shale-type deposits and from the newly-reported "thin" Stephen Formation of the Canadian Rocky Mountains. Pyrite content is generally low, implying that heavy isotopic composition resulted from restriction of sulfate, as also indicated by low fractionation from Cambrian seawater values. Reduced pyrite content of some samples may have resulted from weathering and/or metamorphism (16, 17).

Table S2. $\delta^{13}C$ vs. $\delta^{18}O$ of carbonate cements of BST-bearing claystones. Data from Chengiang are given in Figure SI5; data from all other formations analyzed are given below.

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