Supporting Information

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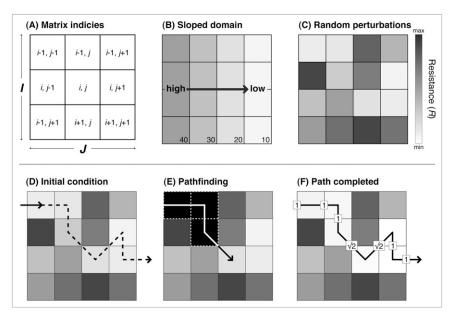


Fig. S1. Schematic of operations in our numerical flow-routing model. (*A*) Matrix indices relative to the center grid cell. (*B*) The parameter *S* determines the slope of the domain, from which (*C*) random values 0–*R* are subtracted. (*D*) From a given cell, the flow path moves to occupy whichever of its nearest neighbors has the lowest value. (*E*) As the flow path is developing, values of cells that the flow has occupied are temporarily reset to the unperturbed elevation at that cell, as in *B*. This value constitutes a local maximum that discourages the flow path from getting arbitrarily trapped by a local minimum, but does not necessarily prevent the path from recrossing itself, particularly when *R* >> *S*. (*F*) Once completed, cells occupied by the flow path are reset equal to the domain diagonal steps) divided by the length dimension of the domain. The flow path in the schematic above, excluding the steps into and out of the grid, has sinuosity $\Omega = 5.82 \div 4 = 1.5$.

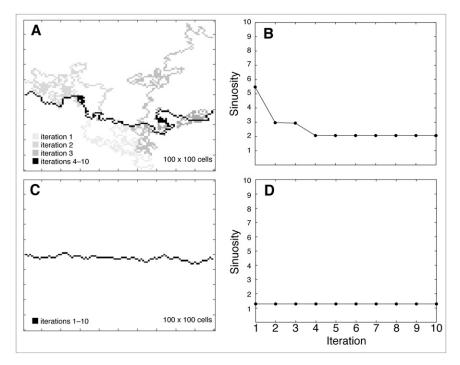


Fig. 52. Theoretically, sinuosity has no upper bound. However, where channel migration occurs in natural systems, flow dynamics tend to not only create excursive meanders but also cut them off, episodically shortening the overall channel length and reducing sinuosity. In the absence of a cutoff mechanism, resistance-dominated conditions (R/S >> 1) in our model will produce supersinuous patterns that are more numerical artifacts than useful analogs, a problem others have also encountered. We affect a cutoff-like function by updating and iterating the domain to allow the flow path to find a minimum length for a given combination of R and S that excludes numerical artifacts of supersinuous paths from the compiled results, particularly when R >> S. (A) Example of iterated flow paths and (B) corresponding sinuosities for R = 0.04 and S = 0.001. C and D show the comparatively locked planform of a slope-dominated, low-sinuosity channel (R = 0.001 and S = 0.04).

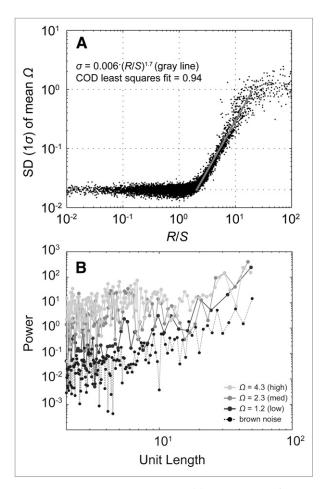


Fig. S3. (*A*) SD versus *R*/S of ensemble model runs. Variance scales with mean *R*/S. (*B*) Power spectra of representative modeled low-, medium-, and highsinuosity planforms, with the power spectrum for a brown-noise signal (integral of a white-noise signal) for comparison. The model does not produce planforms with a preferred wavelength. Both *A* and *B* are indicative of the model's fundamentally Brownian structure.

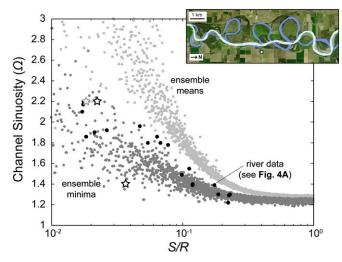


Fig. S4. Historical changes in the Sacramento River planform in terms of floodplain Froude number, plotted atop the data in Fig. 4*B*. The linked gray and black star symbols at $\Omega = 2.2$ mark the difference between assuming Manning's n = 0.15 (gray) or n = 0.10 (black) for the natural floodplain before 1874. Star symbol at $\Omega = 1.4$ reflects the effect of orchard plantations (here, n = 0.05) after 1898.

Table S1. Locations and credits for images shown in Figs. 1, 3, and 5

Fig.	Location	Source		
1 <i>A</i>	37.84°N, 126.21°E	(1)		
1 <i>B</i>	38.22°N, 109.88°W	(2)		
1C	44.57°N, 70.71°W	(1)		
1 <i>D</i>	19.17°S, 35.08°W	(3)		
1 <i>E</i>	Mars: -6.008° (centered), 153.833°E	(4) location at center of image		
1 <i>F</i>	N/A	(5)		
3A	51.51°N, 3.05°W	(1)		
3C	56.42°N, 3.12°W	(1)		
3 <i>E</i>	51.36°N, 3.70°W	(1)		
4A	64.43°N, 149.36°W	(1)		
4 <i>B</i>	64.49°N, 148.31°W	(1)		
4C	39.46°N, 121.99°W	Digitized channels based on maps in ref. 6; base image 1		

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2. National Agricultural Imagery Program (2006) Utah (natural color imagery), Tile q3335_ne_NAIP2006, NAIP Interactive Map. Available at http://mapserv.utah.gov/rasterindicies/naip2006. html. Accessed May 2012.

3. Gamboa D, Alves TM, Cartwright J (2012) A submarine channel confluence classification for topographically confined slopes. Mar Pet Geol 35(1):176-189.

4. High Resolution Imaging Science Experiment (2012) Yardangs and ridges of the edge of Aeolis Planum, Image PSP_006683_1740 (grayscale, map-projected), courtesy of NASA/ JPL/University of Arizona. Available at www.uahirise.org/PSP_006683_1740. Accessed May 2012.

5. Lunar Reconnaissance Orbiter Mission (2012) NASA LROC WAC mosaic, courtesy of NASA/JPL/University of Arizona. Available at www.nasa.gov/images/content/ 503670main_120210a.jpg. Accessed May 2012.

6. Sullivan DG (1982) Prehistoric flooding in the Sacramento Valley: Stratigraphic evidence from Little Packer Lake, Glenn County, California. MS thesis (Univ of California, Berkeley, CA).

River name	Location	Valley slope (\times 10 ⁻³)	Sinuosity	Floodplain Manning's <i>n</i>	Coordinates
Severn	United Kingdom	4.9 (1)	1.22*	0.10	52.74°N, 2.81°W
Jiu	Romania	5.0 (2)	1.29 (3)	0.10	44.17°N, 23.83°E
Republican	Nebraska	0.9 (4)	1.30 (4)	0.05	39.39°N, 97.18°W
Calamus	Nebraska	1.3 (5)	1.30 (5)	0.05	42.02°N, 99.47°W
Sacramento	California	0.8 (6)	1.39 (3)	0.05	39.44°N, 122°W
Peace	Canada	0.3 (7)	1.39 (3)	0.05	58.4°N, 114.75°W
Yampa	Colorado	0.2 (8)	1.40*	0.10	40.5°N, 107.5°W
Yellow	China	0.2 (9)	1.48 (3)	0.05	40.47°N, 109.35°E
Vyatka	Russia	0.3 (10)	1.55 (3)	0.05	58.38°N, 48.7°E
Liard	Canada	0.6 (11)	1.78 (3)	0.10	60.03°N, 128.88°W
Smoky Hill	Kansas	0.7 (4)	1.80 (4)	0.15	38.96°N, 96.91°W
Tana	Kenya	0.5 (12)	1.80 (3)	0.10	1.65°S, 40.11°E
Okavango	Botswana	0.4 (4)	1.86 (4)	0.10	18.10°W, 21.62°E
Ucayali	Peru	0.07 (13)	1.86 (3)	0.15	6.05°S, 74.85°W
Manu	Peru	0.1 (14)	1.90 (3)	0.15	12.7°S, 69.7°W
Beni	Bolivia	0.1 (15)	1.92 (3)	0.15	11.73°S, 66.78°W
Ramu	Papua New Guinea	0.5 (16)	1.96 (3)	0.15	4.13°S, 144.68°E
Jutai	Brazil	0.7 (17)	2.10 (3)	0.15	4.35°S, 67.9°W
Jurua	Brazil	0.7 (17)	2.17 (3)	0.15	4.45°S, 66.65°W
Purus	Brazil	0.7 (17)	2.20 (3)	0.15	6.88°S, 64.63°W

Table S2. Twenty meandering river reaches used in floodplain Froude number calculations

*Measured.

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