

Supporting Information

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SI Text

Geological Setting. The study area (Fig. S1) was chosen because of its rapid expansion of drilling for natural gas from the Marcellus Shale (Pennsylvania); also, it has a limited history of prior oil and gas exploration. Additionally, the study area represents portions of both the upper Susquehanna and upper Delaware watersheds that provide drinking water to >15 million people. The geological setting and methods for the work have been described previously in the works by Osborn et al. (1) and Warner et al. (2). Briefly, the sedimentary geology represents periods of deposition, burial, lithification, uplift, and subsequent erosion that form relatively simple sets of horizontal strata dipping 1° to 3° to the south and east derived from depositional environments that ranged from proposed deep to midbasin black shales to terrestrial red beds (3–5). The monocline is bounded on the north by the Precambrian Canadian Shield and Adirondack uplift (north to northeast), the west by the Algonquin and Findlay arches, and the south and east by the Appalachian fold belt (the Valley and Ridge Province) (6, 7). In general, sedimentary deposition in the northern Appalachian Basin was relatively continuous throughout the Paleozoic era. However, several unconformities erase sequence records regionally, such as the Tri-States unconformity that removed Lower Devonian strata in western New York, but complete sequences are generally found in central New York and our study region of northeastern Pennsylvania (3).

The Appalachian Basin consists primarily of sedimentary sequences of Ordovician to Pennsylvanian age that are derived from the Taconic (~450 Ma), Acadian (~410–380 Ma), and Alleghanian (~330–250 Ma) orogenic events (8). Exposed at its northern extent near Lake Ontario is the Upper Ordovician–Lower Silurian contact (Cherokee unconformity). Younger deposits (Upper Silurian, Devonian, and Mississippian) occur in successive outcrop belts to the south to the Appalachian structural front (4, 9), whereas erosion has removed most post-Pennsylvanian deposition within western-central New York and most of our study area within northeastern Pennsylvania. Bedrock thickness within the basin ranges from ~920 m along the southern shore of Lake Ontario in northern New York to ~7,600 m along the Ap-

palachian structural front to the south. A simplified stratigraphic reconstruction is presented in Fig. S2 for the study area, which constitutes a transition from the Valley and Ridge to the Plateau Province. Compared with the Valley and Ridge Province or the region near the Appalachian Structural Front, the plateau portion of the Marcellus Formation is significantly less deformed (10). Deformation began during the onset of the Alleghanian orogeny. In the plateau physiographic province, deformation is accommodated by a combination of layer parallel shortening, folding that led to low-amplitude anticline/syncline sequences, low angle thrust faulting structures, lineaments, joints, and natural fractures observable in northeastern Pennsylvania (4, 11, 12).

The Marcellus Formation is an organic-rich, hydrocarbon-producing, siliciclastic-rich black shale present beneath much of Pennsylvania, New York, West Virginia, and other northeastern states. It constitutes the stratigraphically lowest subgroup of the Middle Devonian Hamilton Group (5, 9) and was deposited in the foreland basin of the Acadian Orogeny (~385–375 Ma). The Marcellus Formation includes two distinct calcareous and iron-rich black shale members [i.e., the Union Springs (lower) and Mount Marion/Oatka Creek (upper)] interrupted by the Cherry Valley limestone].

Like the Marcellus, the upper part of the Devonian sequence is deposited in the foreland basin of the Acadian Orogeny and consists of material sourced from the Acadian orogeny as part of the Catskill Deltaic sequence. Above the Marcellus, the Hamilton Group consists of the Mahantango gray shale locally interbedded by limestones and the Tulley limestone. The Upper Devonian consists of thick synorogenic sequences of gray shales (i.e., the Brallier Formation) beneath the Lock Haven Formation sandstone and Catskill Formation clastic deltaic red sandstones. The Lock Haven and Catskill Formations constitute the two primary aquifer lithologies in northeastern Pennsylvania along with the overlying glacial and sedimentary alluvium, which is thicker in valleys than the uplands.

Additional geological information is in the work by Osborn et al. (1) and references therein and the work by Warner et al. (2) and references therein.

1. Osborn SG, Vengosh A, Warner NR, Jackson RB (2011) Methane contamination of drinking water accompanying gas-well drilling and hydraulic fracturing. *Proc Natl Acad Sci USA* 108(20):8172–8176.
2. Warner NR, et al. (2012) Geochemical evidence for possible natural migration of Marcellus Formation brine to shallow aquifers in Pennsylvania. *Proc Natl Acad Sci USA* 109(30):11961–11966.
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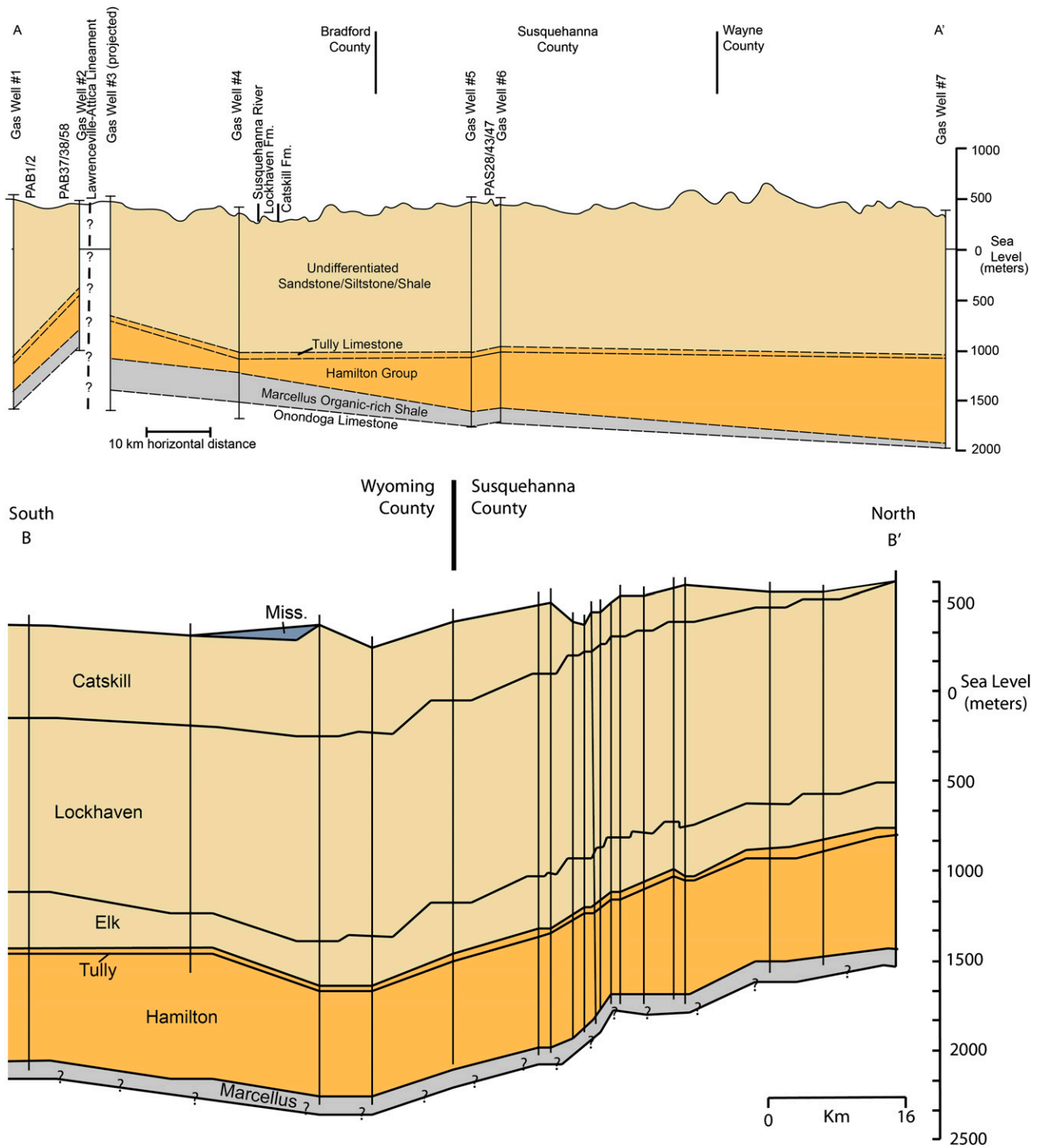


Fig. S2. Generalized stratigraphic section of the study region from the work by Osborn et al. (1), Molofsky et al. (2), and Warner et al. (3) and references therein. The cross sections shown here refer to the locations identified in Fig. S1.

1. Osborn SG, Vengosh A, Warner NR, Jackson RB (2011) Methane contamination of drinking water accompanying gas-well drilling and hydraulic fracturing. *Proc Natl Acad Sci USA* 108(20):8172–8176.
2. Molofsky LJ, Connor JA, Wylie AS, Wagner T, Farhat SK (2013) Evaluation of methane sources in groundwater in northeastern Pennsylvania. *Groundwater* 51(3):333–349.
3. Warner NR, et al. (2012) Geochemical evidence for possible natural migration of Marcellus Formation brine to shallow aquifers in Pennsylvania. *Proc Natl Acad Sci USA* 109(30): 11961–11966.

