

Supporting Information

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SI Materials and Methods

Radiocarbon Dating. Samples of charcoal were prepared at the ^{14}C laboratory, Eidgenössische Technische Hochschule Zurich, before accelerator mass spectrometry (AMS) analyses. Contamination by extraneous carbon was removed using the standard Acid Alkali Acid treatment (1). Charcoal was immersed in a weak acid bath (0.5 M HCl) at 60 °C for 12 h, removing carbonates. Following multiple washes in distilled water, the samples were treated with a weak base (0.1 M NaOH) at 60 °C for 12 h. This process removes humic acids that might also contaminate the sample. The samples were next washed at neutral pH. A final acid wash was applied to remove any atmospheric carbon that might have bonded to the charcoal during base treatment. The dried samples (1 mg of C) were converted to graphite, pressed into the targets and analyzed using AMS along with a set of blank and standards. Blanks made of ^{14}C -free coal were prepared to monitor contamination during the preparation process. Conventional ^{14}C ages (Fig. 4) were calculated following Stuiver and Polach (2). These ages include correction for fractionation based on on-line measurements of $\delta^{13}\text{C}$ on graphite. The 1- σ error includes counting uncertainty as well as the scatter of standards and blanks. ^{14}C ages were calibrated using the OxCal 4.1 program and INTCAL09 calibration dataset. We also applied the Bayesian model of OxCal 4.1 that puts constraints on the ages (prior) (3) to obtain a more precise chronology.

Modeling the Caldera Size. We modeled the size of the Samalas eruption based on pre- and posteruption topography of the Rinjani Volcanic Complex (hereafter RVC). This process of course required reconstruction of the preeruptive topography as follows (4):

- i) We located cells in the digital elevation model (DEM) likely to represent the pre-eruptive surface. They include plateaus, sectors of volcano preserved from erosion, and crests that are the uppermost interflues remaining after a significant amount of erosion. These sites provide the best approximation of paleotopography.
- ii) The preeruption RVC topography is first modeled from these located points by defining a first-order radial surface. It is a kind of surface of revolution defined by: (i) an exponential-like function (the generatrix) fitting the average concave-upwards volcano profile; and (ii) the volcano's summit location, around which the generatrix is rotated to form the surface. To obtain the best-fitting first-order surface, we introduce the following sophistication. A conventional surface of revolution with axial symmetry generates circular contour lines. Instead, as the generatrix rotates around the summit location, we stretch and contract it to obtain, in planar section, an elliptically shaped surface defined by elliptical contour lines. The optimal set of parameters that define the first-order primary volcanic surface is obtained by a least-squares method using the simplex algorithm (5). Parameters solved for include the location and elevation of the volcano summit, the eccentricity and long-axis azimuth of the contour lines, and the coefficients of the generatrix.
- iii) The first-order geometric surface is then modified according to second-order elevation variations because of local distributions of volcanic products. Residuals between elevations of the input points selected as representative of the preeruption surface and those obtained by the first-order modeled surface are calculated. Residuals are interpolated by kriging and next summed to the first-order surface elevations to obtain the

definitive preeruption surface elevation. Two independent surfaces are modeled for each of the two RVC volcanoes. The preeruption surface elevation of RVC is finally obtained by retaining for each of the two volcanoes, the area and elevations where the modeled paleo preexplosion surface lies above the postexplosion lower surface. Kriging adjustments also yield SEs for elevations of both basal and upper surfaces that provide a measure of the uncertainty in the calculated volume of material removed by the Samalas eruption.

In addition to the preeruption surface, modeled as described above, calculation of the missing volume requires knowing the elevation of the caldera base everywhere immediately after the eruption. The present caldera walls provide some constraints on this but within the crater, the floor has been mantled in products of the Samalas eruption and subsequent activity of Gunung Barujari (in the past two centuries). The basal surface is interpolated by kriging from the DEM cells located on the caldera walls but the lack of constraints on the thicknesses of accumulated material in the caldera limits its reliability.

Finally, for every DEM cell, the height difference between elevations of the preeruptive surface and caldera basal surface is multiplied by the cell area. We estimate uncertainty by similarly calculating the product of height difference SEs and cell area. The final calculation of the volume of removed material requires summation of each cell volume, and summation of errors.

Written Sources. Translation to Indonesian and to English of the historic poem Babad Lombok, written in the Old Javanese language. The original text of the historic poem *Babad Lombok* in Old Javanese language, Wacana (6):

274. *Gunung Renjani kularat, miwah gunung samalas rakrat, balabur watu gumuruh, tibeng desa Pamatan, yata kanyut bale haling parubuh, kurambangning sagara, wong ngipun halong kang mati.*

275. *Pitung dina lami nira, gentuh hiku hangebeki pretiwi, hing leneng hadampar, hanerus maring batu Dendeng kang nganyuk, wong ngipun kabeh hing paliya, saweneh munggah hing ngukir.*

276. *Hing jaringo hasingidan, saminya ngungsi salon darak sangaji, hakupul hana hing riku, weneh ngungsi samuliya, boroh Bandar papunba lawan pasulun, sarowok pili lan ranggiya, sambalun pajang lan sapit.*

277. *Yek nango lan pelameran, batu banda jejangkah tanah neki, duri hanare menyan batu, saher kalawan balas, batu lawang batu rentang batu cangku, samalih tiba hing tengah, brang bantuan gennira ngungsi.*

278. *Hana ring pundung buwak bakang, tana' gadang lembak babidas hiki, saweneh hana halarut, hing bumi kembang kekrang, pangadangan lawan puka hatin lungguh, saweneh kalah kang tiba, mara hing langko pajanggih.*

279. *Warnanen kang munggend palowan, sami larut lawan ratu hing nguni, hasangidan ya riku, hing Lombok goku medah, genep pitung dina punang gentuh, nulih hangumah desa, hing preneha siji-siji.*

The Indonesian translation of the historic poem *Babad Lombok*:

274. *Gunung Rinjani Longsor, dan Gunung Samalas runtuh, banjir batu gemuruh, menghancurkan Desa Pamatan, rumah2 rubuh dan hanyut terbawa lumpur, terapung-apung di lautan, penduduknya banyak yang mati.*

275. Tujuh hari lamanya, gempa dahsyat meruyak bumi, terdampar di Leneng (lenek), diseret oleh batu gunung yang hanyut, manusia berlari semua, sebahagian lagi naik ke bukit.

276. Bersembunyi di Jeringo, semua mengungsi sisa kerabat raja, berkumpul mereka di situ, ada yang mengungsi ke Samulia, Borok, Bandar, Pepumba, dan Pasalun, Serowok, Piling, dan Rangi, Sembalun, Panjang, dan Sapit.

277. Di Nangan dan Palemoran, batu besar dan gelundungan tanah, duri, dan batu menyan, batu apung dan pasir, batu sedimen granit, dan batu cangku, jatuh di tengah daratan, mereka mengungsi ke Brang batun.

278. Ada ke Pundung, Buak, Bakang, Tana' Bea, Lembuak, Bebidas, sebagian ada mengungsi, ke bumi Kembang, Kekrang, Pengadangan dan Puka hate-hate lungguh, sebagian ada yang sampai, datang ke Langko, Pejanggih.

279. Semua mengungsi dengan ratunya, berlindung mereka di situ, di Lombok tempatnya diam, genap tujuh hari gempa itu, lalu membangun desa, di tempatnya masing-masing.

The English translation of the historic poem *Babad Lombok*:

274. Mount Rinjani avalanched and Mount Salamas collapsed, followed by large flows of debris accompanied by the noise coming from boulders. These flows destroyed (the seat of the kingdom) Pamatan. All houses were destroyed and swept away, floating on the sea, and many people died.

275. During seven days, big earthquakes shook the Earth, stranded in Leneng (Lenek), dragged by the boulder flows, People escaped and some of them climbed the hills.

276. Hiding in Jeringo (close to Mataram), all people moved with the rest of the king's family to several places: Samulia, Borok, Bandar, Pepumba Pasalun, Serowok, Piling, and Rangi, Sembalun, Pajang, and Sapit.

277. At Nangan and Palemaron, big boulders rolled with soil, with pumices and sand, and granite sediments on the land, they evacuated to Brang Batun.

278. There were people moving to Pundung, Buak, Bakang, Tana Bea, Lembuak, Bebidas, some of them evacuated to Kembang Bumi, Kekrang, Pengadangan and Puka Puka hate-hate lungguh and also to Langko and Pejanggih.

279. Everybody took refuge together with the King, Lombok became very quiet, even seven days after the earthquakes occurred, and later they built their own houses.

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Evidence for Northern Hemisphere Winter Warming in Western Europe in 1257/1258. Arras (France).

En cest an, fut le temps si douz et si souef (chaud) que en tout l'hiver ne gela que deux jours : ou mois de janvier, trouvoit on des violettes et les fleurs de fraisiers, et estoient tous les pommiers tous blancs flouris. (7)

Translation: At that time the wheater was so mild and so hot that frost barely lasted for more than two days. In January [1258], violets could be observed, and strawberries and apple trees were in blossom.

Paris (France).

Et en l'autre année après, qui est en l'incarnation par M.CC. et LVI (1256) fist trop durement fort hyvier ou royaume de France et pluvieux esté dusqu'à la Nativité saint Jehan Baupliste (24 juin). Et en l'autre année après, fist merveilleusement chaut esté et chaut temps jusqu'à la Chandeleur ; et puis après, fist merveilleusement grant froit jusqu'à la saint Marc (25 avril). Et en cèle année meïsmes, ot par toute la France grant chierté de pain, de vin et de toutes viandes. (8)

Translation: And the following year, which is the year 1256 of the Incarnation, the winter was very harsh in the kingdom of France and the summer was very rainy until the Nativity of Saint Jean Baptist (June 24). And the following year (1257), the summer was excessively hot, and the weather was warm until Candlemass (2 February 1258), and then it was excessively cold until the St. Mark (25 April 1258). And this year also, there was throughout France a great shortage of bread, wine and any meat.

Saint Alban Abbey (England).

Annus quoque iste, cronicarum infirmitatum genitivus vix occupatum permisit aliquatenus respirare. Non enim frigus vel serenitas vel gelu saltem aliquantulum stagnorum superficiem, prout consuevit, glaciale m induravit, vel stiriam a stillicidiis coegit dependere ; sed continuae inundationes pluviarum et nebularum usque ad Purificationem beatae Virginis aera inspissarunt. (8)

Translation: This year (1257), too, generated chronic complaints, which scarcely allowed free power of breathing to any one labouring under them. Not a single frosty or fine day occurred, nor was the surface of the lakes at all hardened by the frost, as was usual; neither did icicles hang from the ledges of houses; but uninterrupted heavy falls of rain and mist obscured the sky until the Purification of the Blessed Virgin (2 February 1258) (10)

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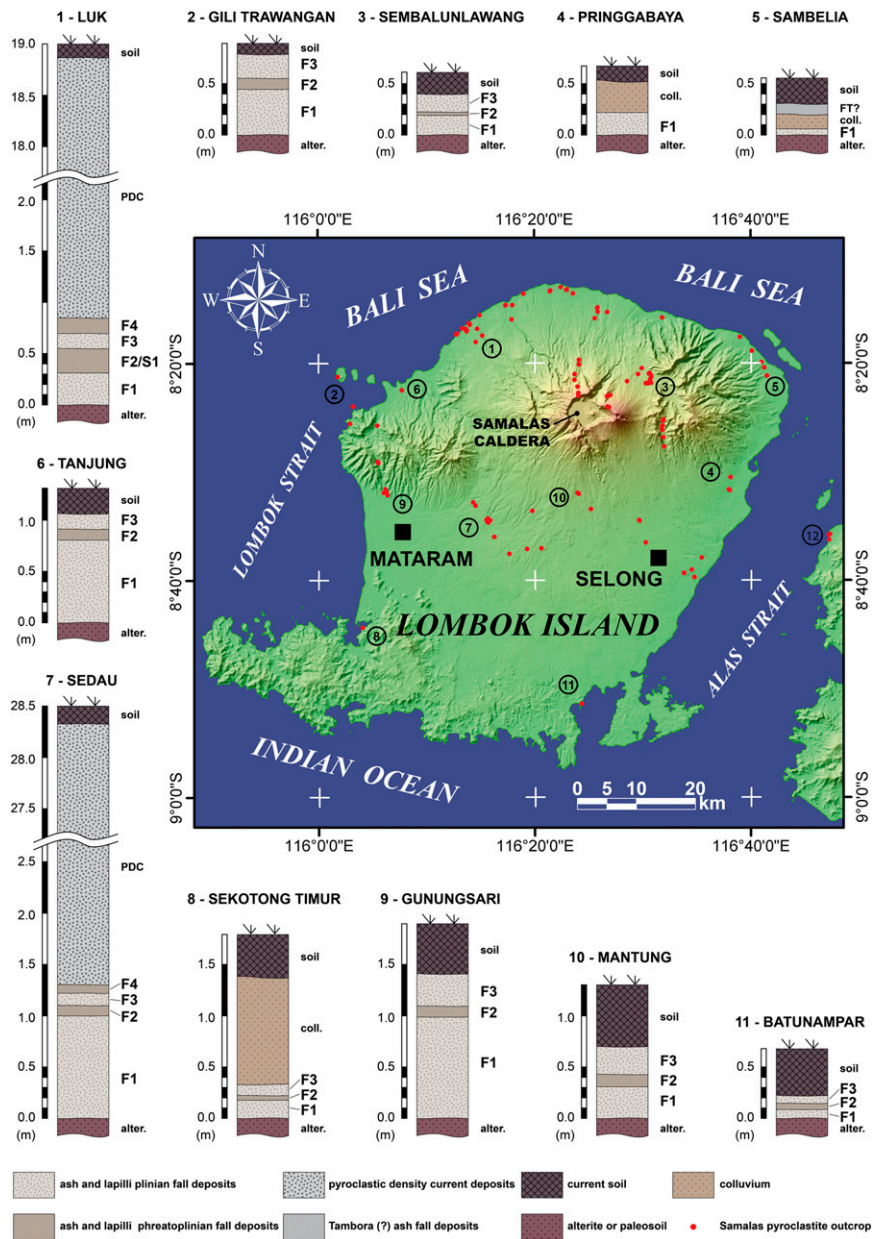


Fig. S1. Synthetic stratigraphic logs of the Samalas pyroclastic deposits (Lombok, Indonesia). The F1 Plinian fall layer is composed of pumiceous ash and lapilli with few lithics. The F2 phreatoplinian fall layer is constituted mainly by fine and coarse ash and minor small lapilli, associated in some areas with pyroclastic surge deposits. The F3 Plinian fall layer is composed of pumice and lithic ash and small lapilli, less coarse and less thick than the first Plinian fall.

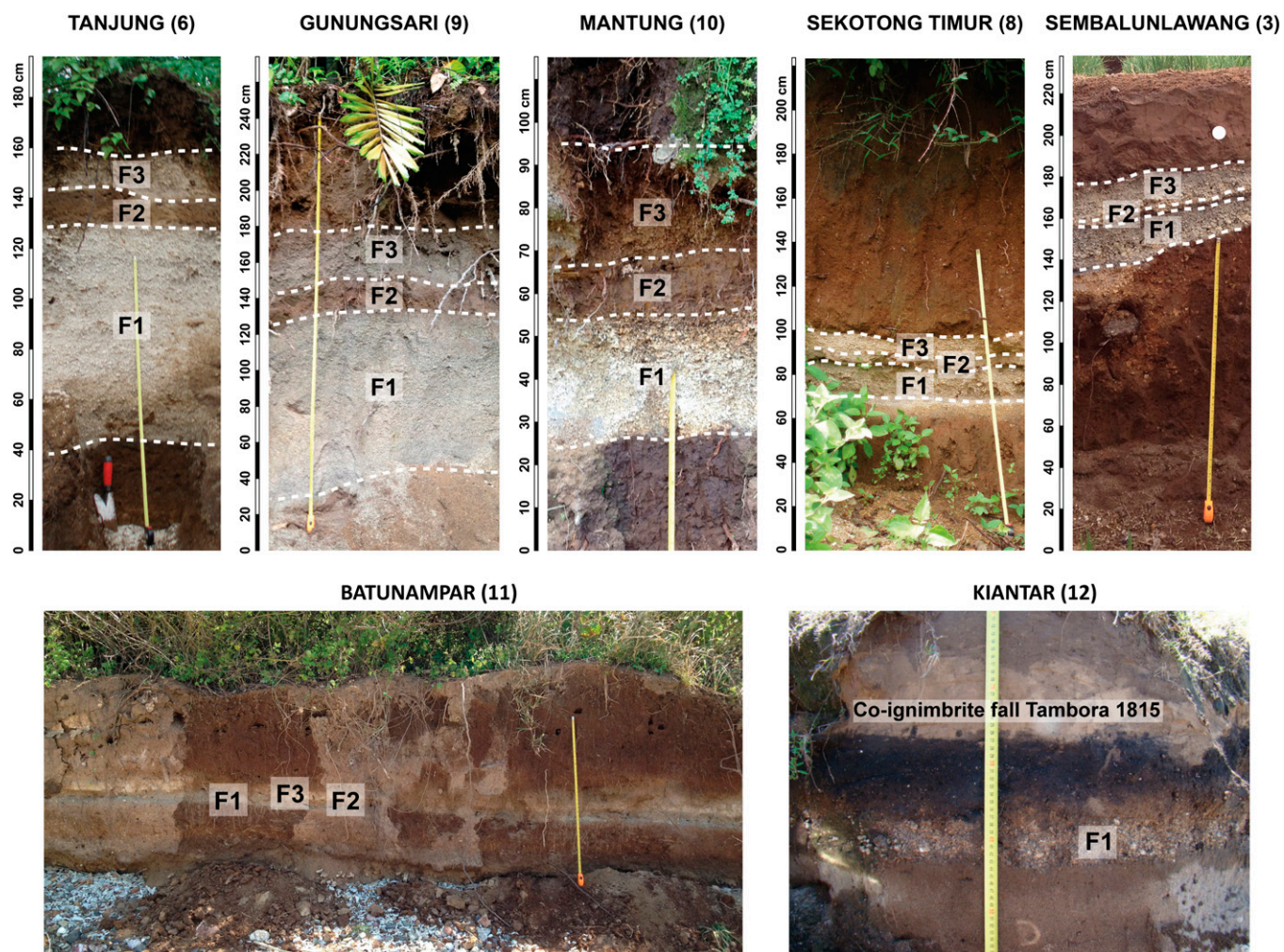


Fig. S2. Field observations on the distal pyroclastic fall layers from the Samalas (Lombok, Indonesia) (photos courtesy of J.-P.D., J.-C.K., and P.W.). Maximum thickness of the Plinian tephra fall F1 measured in the field exceeds 1 m, and cumulative fall deposit reaches up to 1.60 m (i.e., twice as much as at Tambora).

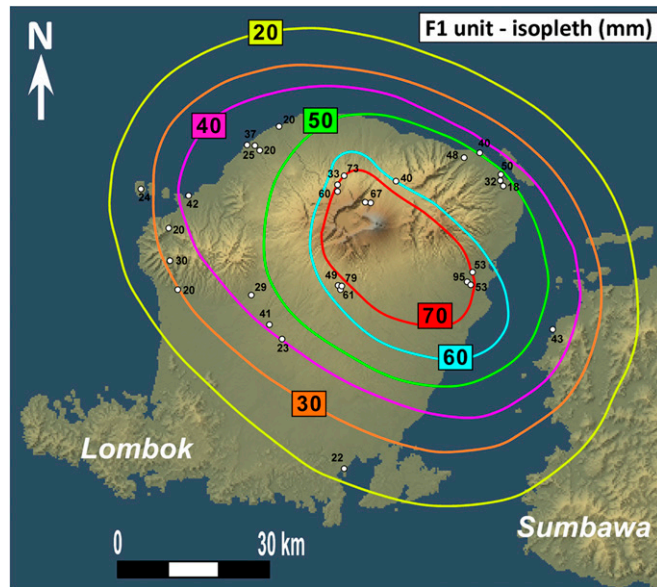


Fig. S3. Maximum pumice M_P isopleths (mm) for the F1 Plinian pumice fallout unit of the Samalas eruption. The presence of pumice clasts of up to 50 mm at 46-km distance SE from the vent on Sumbawa attests to the large magnitude of the F1 ultraplinian phase of the Samalas eruption. This tendency is confirmed by the limited M_L dataset that shows lithics of up to 35 mm at 27 km SE from the source on the eastern coastline of Lombok.

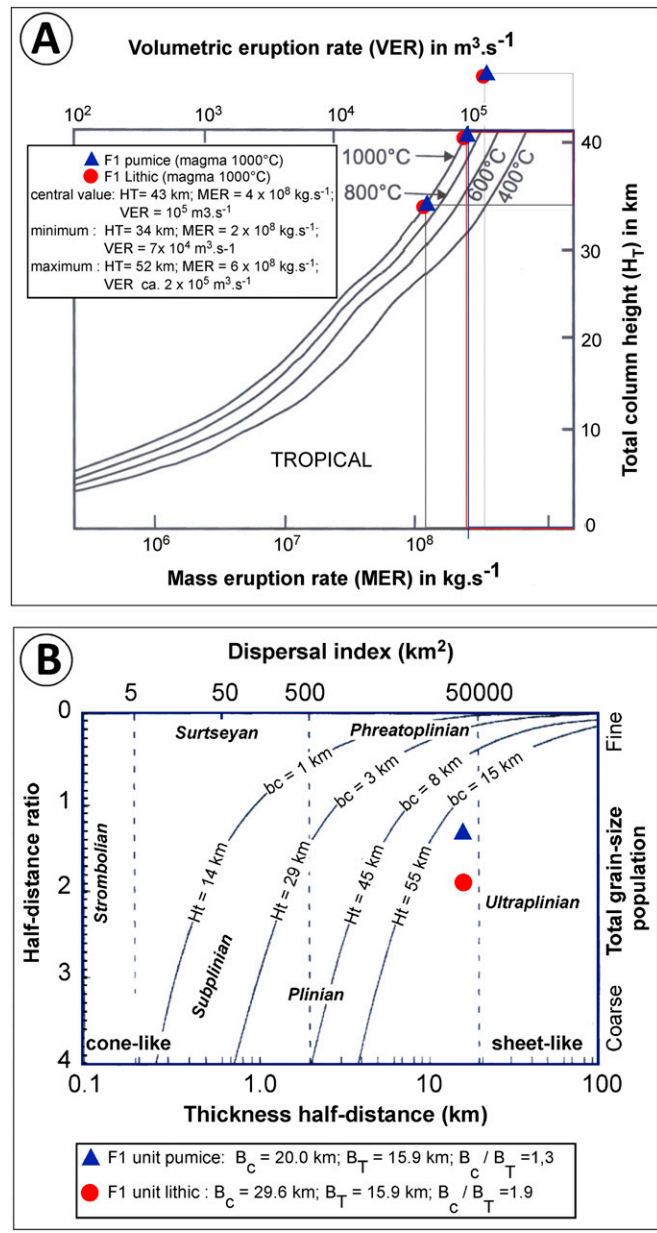


Fig. 54. Dynamic parameters of the A.D.1257 eruption of Samalas. (A) Mass discharge rate (MDR) and volumetric eruption rate (VER). The maximum height H_T of the column was estimated as $\sim 43 \text{ km} \pm 8.6 \text{ km}$ using the model of Carey and Sparks (1) and the data for the 2- and 3-cm isopleths for pumice clasts. Limited data from the maximum lithic clast 2.4-cm isopleth yield a minimum H_T value of 43 km (no lithic clasts larger than 2.4 cm were found in distal parts of the deposit). The geometry of the pumice isopleths shows, according to theoretical relationships of Carey and Sparks (1), that the F1 convective column formed in a strong crosswind with estimated velocity $> 20 \text{ m} \cdot \text{s}^{-1}$, as determined using 0.8-cm-diameter clasts with a density of $2,500 \text{ kg} \cdot \text{m}^{-3}$, which are equivalent to the F1 pumice clasts of 4-cm-diameter and a clast density of $500 \text{ kg} \cdot \text{m}^{-3}$ higher, but compatible with the density of the Samalas F1 pumice of $380 \text{ kg} \cdot \text{m}^{-3}$. (B) Thickness half-distance B_T versus clast-size half-distance – thickness half-distance ratio B_C/B_T for pumice or lithic clasts for the Samalas F1 Plinian phase. Following Pyle (2) and Bonadonna et al. (3), the Samalas F1 Plinian fall has a thickness half-distance B_T of 16.4 km and a clast-size half-distance distance of 22.6 km for pumice clasts and 29.6 km for lithic clasts, typical of some of the highest magnitude Plinian eruptions (3), as evidenced also by the very high dispersal index. The Samalas F1 plinian deposit plots in the ultraplinian field (4), with a column in excess of the theoretical limit of 55 km in still air.

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- Pyle DM (2000) Sizes of volcanic eruptions. *Encyclopedia of Volcanoes*, eds Sigurdsson H, Houghton B, Rymer H, Stix J, McNutt S (Academic, San Diego), pp 263–269.
- Bonadonna C, Ernst GGJ, Sparks RSJ (1998) Thickness variations and volume estimates of tephra fall deposits: The importance of particle Reynolds number. *J Volcanol Geotherm Res* 81(34):173–187.
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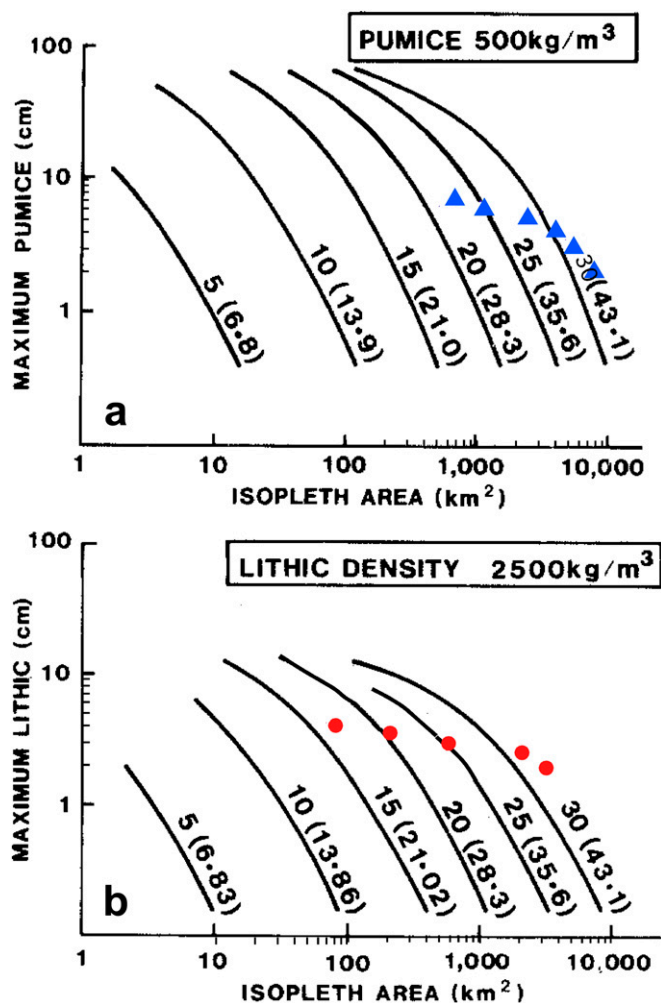


Fig. 55. Maximum height (H_t) of the column for the plinian F1 phase of the mid-13th century eruption of Samalas volcano. The maximum height H_T of the column was estimated as $\sim 43 \text{ km} \pm 8.6 \text{ km}$ using the model of Carey and Sparks (1) and the data for the 2 and 3 cm isopleths for pumice clasts (A). Limited data from the maximum lithic clast 2.4 cm isopleth (B) yield a minimum H_T value of 43 km (no lithic clasts larger than 2.4 cm were found in distal parts of the deposit). Column heights on the curves are shown in km as H_T (H_B), for example 25 (35.6), with H_T the total column height and H_B the height of neutral buoyancy of the column. These data correspond to a neutral buoyancy H_B height of 30 km derived, respectively, from the pumice and lithic isopleths for the F1 fallout unit and based on empirical relationships from Sparks (2) and Pyle (3). The geometry of the pumice isopleths shows, according to theoretical relationships of Carey and Sparks (1), that the F1 convective column formed in a strong crosswind with estimated velocity $> 20 \text{ m.s}^{-1}$ as determined using 0.8-cm diameter clasts with a density of 2500 kg.m^{-3} , which are equivalent to the F1 pumice clasts of 4 cm diameter and a clast density of 500 kg.m^{-3} higher but compatible with the density of the Samalas F1 pumice of 380 kg.m^{-3} .

- Carey S, Sparks RSJ (1986) Quantitative models of fallout and dispersal of tephra from volcanic eruption columns. *Bulletin of Volcanology* 48(109):109–125.
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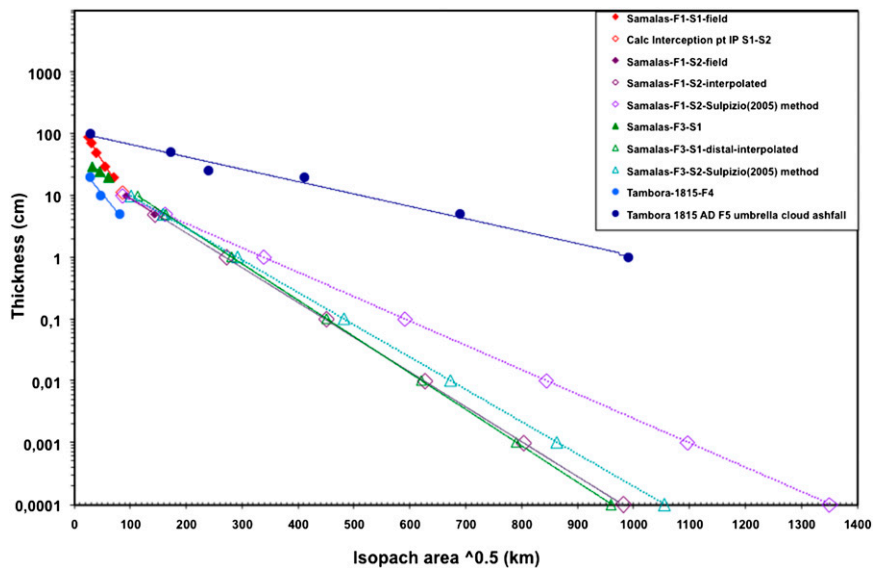


Fig. 56. Log isopach thickness versus isopach area^{0.5}. The exponential thinning of pumice fallout deposits for the Samalas F1 and F3 Plinian pumice fallout units allows a determination of the total erupted volume and its uncertainty using methodologies developed by Pyle (1), Sulpizio (2), and Fierstein and Nathanson (3). The uncertainty in the distal volume of units F1 and F3 is largely controlled by the slope of the distal segment S2 (dotted lines) of the exponential thinning law (the lower the slope, the greater the distal volume) in the volume computation following the methodology of Fierstein and Nathanson (3). The volume data is compared to the Tambora F4 Plinian fallout unit and the F5 umbrella cloud ash fallout unit, modified after Sigurdsson and Carey (4) and Self et al. (5).

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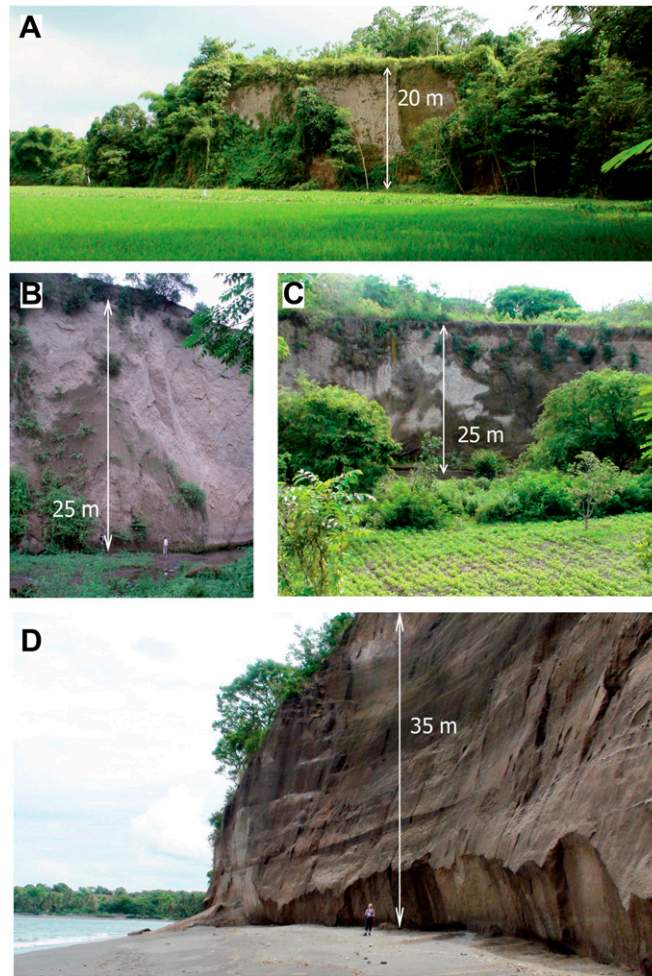


Fig. S7. Pyroclastic density current (PDC) deposits emplaced during the 13th eruption of Samalas. (A) Sedau village, southwest flank, 22 km from the caldera rim. (B and C) North flank, 20 km from the caldera rim. (D) Pumice cliff on the northwest coast of Lombok (photos: A, C, and D courtesy of F.L.; B courtesy of J.-P.D.). Between 10 and 20 km from the caldera rim, ignimbrites are deeply downcut by rivers (A–C). Successions of PDC and pyroclastic-fall deposits from the 13th century eruption form the entire valley walls up to thicknesses of 30 m. The Samalas PDCs entirely modified the precaldera topography, filling former valleys, and resulting in relief inversion following subsequent erosion. On the northwest coast, a remnant of the ignimbrite at 23 km from the caldera rim forms a 35-m-high recessive cliff (D), suggesting that a substantial part of the PDCs entered the sea.

Table S1. The largest well-documented volcanic eruptions ($M > 5$) during the Holocene

Volcano	Country	Deposit name	Bulk deposit volume (km ³)	DRE volume (km ³)	Adjusted mass (kg)	Mass eruption rate (kg/s)	Maximum magnitude*	Intensity [†]	Age	Source
Kurile Lake	Kamchatka, Russia	KO	170	80	1.92×10^{14}		7.3		6460–6414 cal B.C.	(1)
Santorini	Greece	Minoan [‡]		60	1.48×10^{14}	2.50×10^8	7.2	11.4	1627–1600 cal B.C.	(2, 3)
Mazama (Crater Lake)	Oregon, United States	Lower pumice [‡]		52	1.28×10^{14}		7.1		5677 cal B.C.	(4, 5)
Samalas	Indonesia	1257 A.D.^{‡ 5}		>40	9.90×10^{13}[¶]	1.10×10^9	7.0	12.0	Cal A.D. 1257	Present work (6, 7)
Ilopango	El Salvador	Tierra Blanca Joven	84	39	8.15×10^{13}		6.9		Cal A.D. 536	
Tambora	Indonesia	A.D. 1815 [‡]		>33	8.15×10^{13}	2.8×10^8	6.9	11.4	A.D. 1815	(8, 9)
Taupo	New Zealand	A.D. 180	105	35	8.00×10^{13}	1.10×10^9	6.9	12.0	A.D. 232 ± 5	(8, 10, 11)
Aniakchak	Alaska, United States	3430 B.P.		27	6.21×10^{13}		6.8		1645 B.C.	(12–14)
Changbaishan/Baitoushan	China/North Korea	Millenium eruption	96	24.5	5.64×10^{13}		6.8		Cal A.D. 946	(15, 16)
Quilotoa	Ecuador	800 B.P.	21.3	18.7	4.22×10^{13}	2.00×10^8	6.6	11.3	Cal A.D. 1275	(17, 18)
Katmai - Novarupta	Alaska, United States	Valley of 10 000 Smokes	17	6.8	3.00×10^{13}	1.00×10^8	6.5	11.0	A.D. 1912	(8, 19)
Krakatau	Indonesia	A.D. 1883	18–21	12.5	3.00×10^{13}	5.00×10^7	6.5	10.7	A.D. 1883	(8, 20)
Santa Maria	Guatemala	A.D. 1902	20.2	8.6	2.00×10^{13}	1.70×10^8	6.3	11.2	A.D. 1902	(8, 21)
Quizapu	Chile	A.D. 1932	9.5	4	9.72×10^{12}	1.50×10^8	6.0	11.2	A.D. 1932	(22)
Pinatubo	Philippines	A.D. 1991		5	1.10×10^{13}	4.00×10^8	6.0	11.6	A.D. 1991	(8)
Vesuvius	Italy	A.D. 79		3.25	6.00×10^{12}	1.50×10^8	5.8	11.2	A.D. 79	(8, 23)
Rungwe	Tanzania	Rungwe pumice	3.2–5.8	1.4	2.00×10^{12}	4.80×10^8	5.3	11.7	ca. 4000 B.P.	(24)
Huaynaputina	Peru	A.D. 1600 Stage I	7	2.6	1.30×10^{12}	2.40×10^8	5.1	1.4	A.D. 1600	(25)
Chichon	Mexico	Unit B 550 BP	2.8	1.1	1.05×10^{12}	1.00×10^8	5.0	11.0	Cal A.D. 1320–1433	(26)

* $M = \log_{10}(\text{erupted mass kg}) - 7$.

[†] $I = \log_{10}(\text{mass eruption rate kg/s}) + 3$.

[‡]Erupted mass is taken assuming an average of 2,470 kg-m³ for the dense-rock equivalent (DRE) density like for Samalas and Tambora.

⁵Minimum magnitude as uncertainty on distal to very distal ash bulk volume is significant.

[¶]Minimum value, that of the calculated missing caldera.

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Table S2. Geochemical composition of matrix glass from the Samalas pyroclastic fall deposits (electron microprobe analysis)

Oxide/element	F1 matrix glass		F2 matrix glass		F3 matrix glass	
	Mean	±1 σ_{SD}	Mean	±1 σ_{SD}	Mean	±1 σ_{SD}
SiO ₂ (wt.%)	66.66	1.6	67.11	1.35	67.47	1.39
TiO ₂ (wt.%)	0.47	0.11	0.45	0.1	0.43	0.1
Al ₂ O ₃ (wt.%)	16.04	0.37	15.87	0.28	15.5	0.49
FeO (wt.%)	2.77	0.19	2.57	0.21	2.56	0.2
MnO (wt.%)	0.14	0.06	0.14	0.07	0.14	0.06
MgO (wt.%)	0.74	0.07	0.67	0.05	0.61	0.09
CaO (wt.%)	2.22	0.16	2.08	0.14	1.78	0.19
Na ₂ O (wt.%)	3.99	0.33	3.84	0.37	3.74	0.3
K ₂ O (wt.%)	4.04	0.15	4.14	0.2	4.38	0.19
P ₂ O ₅ (wt.%)	0.29	0.06	0.3	0.01	0.26	0.08
S (ppm)	94	63	58	29	57	45
Cl (ppm)	1,881	477	2,241	114	2,040	787
F (ppm)	359	110	307	85	272	134
Total (wt.%)	97.57		97.43		97.09	
Na ₂ O/K ₂ O	0.99		0.93		0.85	
N major	85		28		52	
N volatile	72		17		61	
Normalized to 100% for eight oxides						
SiO ₂ (wt.%)	68.78	0.49	69.37	0.39	69.95	0.54
TiO ₂ (wt.%)	0.48	0.11	0.46	0.1	0.44	0.1
Al ₂ O ₃ (wt.%)	16.55	0.26	16.41	0.29	16.07	0.38
FeO (wt.%)	2.86	0.18	2.66	0.2	2.65	0.2
MgO (wt.%)	0.76	0.07	0.7	0.05	0.63	0.09
CaO (wt.%)	2.3	0.18	2.15	0.13	1.84	0.19
Na ₂ O (wt.%)	4.11	0.31	3.97	0.35	3.87	0.28
K ₂ O (wt.%)	4.17	0.15	4.28	0.18	4.54	0.18

Compositions was normalized to 100% for eight oxides to compare them with the available composition of the glass shard from the A.D. 1257–1259 event evidenced in the polar ice cores (1). *n* = number of analysis.

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