#### **An active atmospheric methane sink in high Arctic mineral cryosols**

#### **Supplementary Information**

This supplementary file contains detailed methodology, supplemental figures (Fig. S1- S5) and tables (Table S1-S5), and additional references.

## Methods and Materials

#### *Field flux measurements*

*In situ* CH<sub>4</sub> fluxes were measured in July 2011-2013 using a Picarro soil  $CO_2$ -CH<sub>4</sub> gas analyzer (Picarro Inc., Santa Clara, CA, USA) or Los Gatos Fast Methane Analyzer (Los Gatos Research Inc., Mountain View, CA) (Allan *et al.*, 2014; Stackhouse *et al.*, 2014 and this study).

Surface fluxes were measured in replicates using open-circuit dark chambers with continuous gas replacement from the air in 2011-2013 or closed-static chamber in 2013. For the open-circuit method, gas was continuously sampled (flow rate =  $25 \text{ cm}^3 \text{ min}^{-1}$ ) and analyzed at approximately 1 hz sample rate for CH<sub>4</sub>, CO<sub>2</sub>, H<sub>2</sub>O, and <sup>13</sup>CO<sub>2</sub>, with alternating analysis periods of atmosphere and standard calibration gas. The flux chamber interior was continuously mixed with a small fan, and sample periods ranged from 45- 600 min, with 45 min found to be the time needed to reach steady state concentration. The CH4 uptake rates were calculated from the difference between initial atmospheric CH4 concentration and the stabilized CH4 concentration, with molar volume corrected for temperature using soil surface temperature and barometric pressure. For the closed-static method, the net loss of  $CH_4$  was calculated by comparing the initial atmospheric  $CH_4$ concentration and the final CH4 concentration being measured at the end of a 4-minute duration. Subsurface CH<sub>4</sub> concentration was measured directly using gas sipper tubes installed into shallow boreholes at various depths and sealed into place with bentonite clay to prevent atmospheric contamination. Soil temperatures at corresponding depths were measured by LiCOR thermistor (Maxim Integrated Products, San Jose, CA, USA).

#### *Long-term intact core warming experiments*

The setup of this warming experiments being described briefly here is paraphrased from a submitted manuscript, (Stackhouse *et al.*, 2014), to provide basic information that is essential for understanding the metagenome and metaproteome analyses performed in this study.

Seventeen 1-m long core samples, collected from a 16x16 m rectangular, icewedge polygon at the study site, were divided into four treatments: saturated (4 cores, mimicking thermokarst-affected terrain), *in situ* (7 cores, *in situ* soil water saturation conditions), dark (2 cores), and a control group (4 cores, remain frozen below 70 cm). Initially, core samples were kept frozen in insulated 55 gallon barrels that were filled with 20% ethanol bath maintained at -3<sup>o</sup>C by recirculating thermochillers (ThermoCube 200-LT, Solid State Cooling System, USA). The cores were individually wrapped in double layers of plastic trash bag to prevent direct contact with the ethanol bath. The cores then underwent a progressively thawing, that took 13 weeks, to  $\sim$ 4 $\degree$ C in a walk-in cold room from top to bottom (except for the control group, which were thawed to 70 cm to keep the permafrost layer frozen), by adjusting the level of the ethanol bath.

Prior to thawing  $(T=0)$ , small holes were drilled using a sterilized drill bit at 5 cm, 35 cm, 65 cm and ~80 cm (below permafrost table). From one replicate core of each treatment, frozen soil samples (10-15 g) were collected with a sterilized spatula at each depth for molecular analyses. After collection, the hole was plugged with a 1.3-cm diameter butyl-rubber stopper and sealed with waterproof tape. Similarly, cryosol samples were collected after 1-week thawing at each depth (T=0.25 month) from the same replicate cores processed at T=0. At later time points (T=6, 12 and 18 months post thaw), cryosol samples were collected from different replicate cores for each treatment in order to minimize unappreciated effect on the overall porosity of the cores due to soil removal.

During the long-term warming experiment, gas samples from the headspace of all 17 cores were analyzed for  $O_2$ ,  $N_2$ ,  $CO_2$ ,  $CH_4$ ,  $H_2$  and CO. Rhizon tubes (2.5 mm diameter, 50 mm length) were installed at 5 cm, 35 cm, 65 cm and ~80 cm near the drill holes from which cryosol samples were collected. Pore water samples were collected using Rhizon tubes and were analyzed for aqueous chemistry. Operation details, results and discussion are presented in (Stackhouse *et al.*, 2014).

*Abundance of methanotrophs and methanogens in metagenomic studies*

As part of the long-term intact core warming experiments, 4 g of cryosols were collected and processed as described in (Stackhouse *et al.*, 2014). Total DNA was extracted using Fast DNA SPIN Kit (MP Biomedical, Irvine, CA) (Vishnivetskaya *et al.*, 2014). The extracted DNA samples were used to prepare metagenome shotgun libraries using the Illumina Nextera DNA library preparation kit (Illumina, Inc., San Diego, CA, USA), and sequenced (2 x 100 bp) on an Illumina HiSeq 2000 platform (Chauhan *et al.*, 2014). Results of 16 near-surface (at 5 cm depth) cryosoal samples representing different time point (T=1 week, 6 and 12 months) were analyzed.

Raw data was processed through MG-RAST pipeline (Meyer *et al.*, 2008) which first demultiplexed and removed sample identifier and then joined overlapping pair-end reads. Low quality sequences, artificial duplicate sequences were removed as part of the quality control (QC) pipeline. The number of post-QC reads per library averaged 1.7 x  $10<sup>7</sup>$  with an average library size of 4.85 Gbp. All gene features were predicted and annotated by searching against M5NR database and taxonomically classified to species level by "Best Hit Classification" using the default values (Max. e-Value Cutoff=1 $e^{-5}$ , Min. % Identity Cutoff =  $60\%$  and Min. Alignment Length Cutoff = 15). The frequency data of these 16 libraries was exported from MG-RAST into STAMP for statistical analyses (Parks & Beiko, 2010). Hit abundance data was normalized to the total sequences passing the QC pipeline. Only the methanotrophic and methanogenic genera (as reviewed by (Nazaries *et al.*, 2013)) with relative abundances  $> 0.001\%$  were considered. The relative abundance of individual genus was not statistically different at

alpha=0.05 (ANOVA in STAMP) and therefore means (and standard deviations) of all 16 samples were reported.

## *Assembly of pmo gene from metagenomes and their abundances*

*De novo* co-assembly of raw sequences from 10 libraries (five 1-week and five 6 months thawed samples at 5 cm depth) were performed using MetaVelvet (Namiki *et al.*, 2012). Functional classifications were annotated separately via IMG/ER and MG-RAST (ID: 4530050.3) using SEED Subsystem and GenBank using the default values (Max. e-Value Cutoff=1e<sup>-5</sup>, Min. % Identity Cutoff =  $60\%$  and Min. Alignment Length Cutoff = 15). Contigs identified as "methane monooxygenase" were searched for protein-coding genes because they may contain fragments of more than one gene. Phylogenetic affiliation of individual gene was queried against the NCBI non-redundant protein database using BlastX.

Raw reads were mapped to each of contigs, that are comprised of *pmo* genes (encodes for particulate methane monooxygenase), using Bowtie (Langmead *et al.*, 2009) to compute the relative abundance. Mean abundances (and standard deviations) of 16 samples were calculated by dividing the number of matched reads by the total number of mappable reads. Co-assembly of sequences from multiple libraries usually masks the genetic variations within and between populations or libraries, and the resultant contigs likely contain mixed genetic signals from the dominant population. We used prefix panto indicate that the detected genes are not derived from a single clonal population.

*De novo* assembly is more preferred when compared to reference-based assembly because the latter prevents the discovery of new genotypes variants by setting an *a priori* framework for read alignment. Nonetheless, raw reads of five 1-week thawed samples (5 cm depth) were mapped to the representative  $pmoCAB$  operon of USC $\alpha$  recovered by bacterial artificial chromosome (BAC) cloning (GenBank Acc. No. CT005232) (Ricke *et al.*, 2005) to demonstrate that a complete *pmoCAB* operon of the USC $\alpha$  genotype was successfully assembled from our data.

## *Phylogenetic analyses of pmo genes*

Phylogenetic trees were constructed from deduced amino acid (aa) sequences for *pmo*A, *pmo*B and *pmo*C genes that encodes for α, β and γ subunit, respectively. All sequences had no frame-shift errors and no curation was applied. Three datasets were created. The *pmoA* gene dataset contained (1) aa sequences of pan-*pmoA* genes from this study. One of them was too short and thus omitted; (2) aa sequences of the best three BlastX matches; (3) aa sequences of all *pmoA* gene copies in published genomes of methanotrophs; (4) aa sequences of *pmoA* genes recovered from environmental studies of permafrost and atm CH4-oxidizing sites; and (5) aa sequences of ammonia monooxygenases (*amoA* genes) that were used as the out-group. Sequences  $(2) - (5)$  were downloaded from NCBI (http://www.ncbi.nlm.nih.gov/). For cases where aa sequences were not available, nucleotide sequences were downloaded and translated. The dataset for *pmo*B and *pmoC* genes were created using the same approach.

For each dataset, sequences were aligned using MAFFT (Katoh *et al.*, 2005) included in freeware JalView package and manually edited using freeware Se-Al. Positions covered by more than half of the sequences were included while unaligned and ambiguous positions were trimmed. Alignments of 161 taxa and 171 aa (*pmo*A genes), 69 taxa and 377 aa (*pmo*B) genes) and 82 and 244 aa (*pmo*C genes) were used for phylogenetic tree construction. ProtTest (v3.3) (Darriba *et al.*, 2011) selected the best-fit amino acid evolutionary model (LG+G and LG+G for *pmo*A and *pmo*B and LG+I+G for *pmo*C gene) (Le & Gascuel, 2008) based on Bayesian Information Criterion. RAxML (v7.2.7 alpha) (Stamatakis, 2006; Stamatakis *et al.*, 2008) was used to search for the bestscoring maximum likelihood (ML) tree with the selected matrix and empirically estimated base frequencies, and to perform a rapid bootstrap analysis of 100 iterations in single run. Tree editing was done using freeware FigTree (v1.3.1).

#### *Assembly of pmo genes from metatranscriptome and their abundances*

Cryosols for metatranscriptomic analysis were collected on July 15, 2013 from an ice-wedge polygon, namely polygon interior and trough (79°24'57"N, 90°45'48"W). The samples were preserved using LifeGuard<sup>TM</sup> Soil Preservation Solution (MO BIO Laboratories Inc., Carlsbad, CA, USA) and stored at -20°C. Total RNA was extracted from 15 g of soil using the RNA PowerSoil Total RNA Isolation Kit (MO BIO Laboratories Inc., Carlsbad, CA, USA). Illumina TruSeq libraries were generated from the total RNA following manufacturer's protocols (Illumina) and sequenced on MiSeq (1 x 150 nt). After filtering and trimming using CLC Genomics Workbench (version 7.0) (CLC bio, Boston, MA, USA), the number of reads obtained for the polygon interior and trough sample were 18,390,227 and 11,577,752 respectively. Then the post-QC reads were mapped to the USCα *pmoCAB* operon (GenBank Acc. No. CT005232) (Ricke *et al.*, 2005).

The reads were also assembled using CLC Genomics Workbench (built-in assembler velvet) and uploaded to MG-RAST for annotation (ID 4548476.3 and 4548477.3 for the polygon interior and trough sample respectively). Within CLC Genomics Workbench, metatranscriptomic reads were mapped against the metatransriptome (or transcript) contigs to determine read abundances (the number of matched reads divided by the total number of mappable reads). Transcript contigs were blasted against the assembled metagenome contigs from the 5 cm soil samples (ID 4530050.3) to determine the proportion of transcript contigs that are similar to the metagenome contigs. Transcript contigs containing *pmo*B genes were translated and used as template for the alignment of the peptide sequences detected from the proteome experiment (described below).

#### *Identification of pMMO in metaproteome*

Cores showing high CH4 uptake flux in the intact core thawing experiment (Stackhouse *et al.*, 2014) were selected. Cryosols at 5 cm depth from 1-week drained cores were subsampled and kept frozen at -20°C. Three grams of cryosol was mixed with SDS-based lysis buffer and the slurry subjected to 15 min of boiling in a water bath with intermittent vortexing as described earlier (Chourey *et al.*, 2010). The slurry was briefly cooled and centrifuged at 21,000 g for 15 min and supernatant aliquoted to fresh tubes and amended with chilled 100% TCA to final concentration of 25% followed by an overnight incubation at -20°C. The TCA-precipitated proteins were collected via centrifugation at 21,000 g for 15 min and the resulting protein pellet was washed with chilled acetone (thrice), air dried and solubilized in guanidine buffer [6M Guanidine HCl, 10 mM DTT in Tris CaCl<sub>2</sub> buffer (10 mM Tris, 50 mM CaCl<sub>2</sub>, pH 7.8)] as described earlier (Chourey *et al.*, 2010). Total protein extracted from the samples was estimated using the RC/DC protein estimation kit (Bio-Rad Laboratories, Hercules, CA, USA) as per the manufacturer's protocol. The dissolved protein sample was subjected to trypsin proteolysis for 16 h at 37°C as described earlier (Chourey *et al.*, 2010; Brown *et al.*, 2006). The reaction was stopped by adding 10% formic acid to final concentration of 0.1% and kept frozen at -80°C until MS analysis.

An aliquot of digested peptides was pressure loaded onto an in-house packed SCX (Luna)-C18 (Aqua) column. The loaded sample column was subjected to an offline wash with solvent A (5% acetonitrile, 0.1% formic acid in HPLC-grade water) for 5 min followed by a gradient with 100% solvent B (70% acetonitrile, 0.1% formic acid in HPLC-grade water) over 10 min. This step was repeated 3 times for a total offline wash time of 45 min to desalt the column and get rid of any loosely attached contaminants. The sample column was then connected to an in-house C18 packed Picofrit column (New Objective, Woburn, MA) and the setup aligned on a Proxeon nanospray source in front of an LTQ-Orbitrap (Thermo Fisher Scientific Inc., San Jose, CA, USA) coupled to an

Ultimate 3000 HPLC system (Dionex<sup>TM</sup>, Thermo Fisher Scientific, Waltham, MA, USA). Peptides were chromatographically separated and analyzed via 24 h Multi-Dimensional Protein Identification Technology (MuDPIT) approach as described earlier (Thompson *et al.*, 2006; Sharma *et al.*, 2012) and the tandem mass spectra (MS/MS scans) were acquired in a data dependent mode using Xcalibur software, V2.1.0 at settings described previously (VerBerkmoes *et al.*, 2009). The raw spectra acquired by 12-step MS/MS runs were searched via SEQUEST v.27 (Eng *et al.*, 1994) against an artificially constructed pMMO database using parameters described elsewhere (Thompson *et al.*, 2006; Sharma *et al.*, 2012). The output files were sorted and filtered using DTASelect v. 1.9 (Tabb *et al.*, 2002) with Xcorr values of at least 1.8 (+1), 2.5 (+2), 3.5 (+3). Identification of at least two peptides per protein sequence was set as criteria for positive protein identifications. Three technical replicates were analyzed for each protein sample.

pMMO database included amino acid sequences translated from (1) *pmo* contigs co-assembled from DNA sequences of the 10 metagenomic libraries stated above; (2) *pmo* contigs assembled from subsets of raw metagenome reads annotated as "methane monoxygenase" by MG-RAST and JGI/IMG; (3) *pmo* contigs generated by mapping to the USCα *pmoCAB* operon (GenBank Acc. No. CT005232) (Ricke *et al.*, 2005); and (4) *pmo* genes of methane monoxygenases obtained from GenBank. Sequences of common contaminants such as trypsin and keratin were also concatenated to the database.

#### *Microcosm incubation experiments*

Two sets of microcosms were set up to study the effect of water saturation and temperature on atm CH4 oxidation rates. Prior to the experiment, 160 mL serum vials were soaked in 10% HNO<sub>3</sub> overnight to remove trace metals, rinsed using distilled water and combusted at 450°C for 8 h. New butyl rubber stoppers were boiled in 0.1 N NaOH for 45 min, soaked in distilled water for 8 hours and autoclaved.

A frozen core collected in April 2011 (Stackhouse *et al.*, 2014) was dissected into sections for every 10 cm. The peripheral rim of 5-cm thick was discarded to remove any potential contaminants from the core liner. The pristine cryosols were put into sterile Whirl-pak bags and homogenized by hand. The 0-10 cm section was used in this experiment. The original water content of the sample was determined to be  $30.4\pm2.0$  wt% by drying three subsamples of 5 g at 50°C for four days. The cryosols were visibly fully saturated, thus 10 wt%, 20 wt% and 30 wt% were regarded as equivalent to water saturation levels of 33%, 66% and 100% respectively. Cryosols were preconditioned to attain the desired water saturation by storing subsamples in a desiccator at 4°C. Eight to ten grams (wet weight) of cryosol were put into vials and sealed with treated butyl rubber stoppers and Al-crimps. Blank vials containing no soils were used to track abiotic gas exchanges and minor instrumental drift.

Manufactured air (Airgas USA LLC, PA, USA) was used to flush the headspace for 2-4 min. The gas composition of the manufactured air was analyzed by Peak Performer 1 gas chromatography systems (Peak Laboratories LLC, CA, USA), which are equipped with a thermal conductivity detector (for  $O_2$  and  $N_2$ ), a reduced compound

detector (for  $H_2$  and CO) and a flame-ionization detector (for  $CO_2$  and  $CH_4$ ). Argon (ARUHP300, Airgas USA LLC, PA, USA) was used as carrier gas at a pressure of 82 psi. Calibration curves for  $O_2$  and  $N_2$  were generated from dilutions of fresh outdoor air in Argon. Standard gas (Scotty® Analyzed Gases, Air Liquide America Specialty Gases LLC, PA, USA) containing 1% atm of CO,  $CO<sub>2</sub>$ , H<sub>2</sub>, CH<sub>4</sub> and C<sub>2</sub>H<sub>4</sub> was diluted in Argon to make gas mixtures of 0.5, 1, 2.5, 5, 7.5 and 10 ppm for all gases and 50, 100, 500 and 1000 ppm for all gases except  $H_2$  and CO. Three measurements were made for each dilution. All calibration curves have high correlation coefficient ( $R^2 > 0.99$ ).

Additional manufactured air was injected with a gas-tight glass syringe to overpressurize the vials to 1.5 atm. All treatments were run in triplicates. One set of 12 vials was incubated at 4°C while another set was incubated at 10°C. Gas was sampled from the headspace at T=0, twice for the first 2 weeks and weekly for another 2 weeks (period of incubation = 31 days). Headspace volume was maintained by replacement of respective gas. Analysis was performed on Picarro iCO2 (Model # G2101-I) using the G2101-i coordinator (Picarro Inc., Santa Clara, CA, USA). Instrumental sample dilution was accounted for by multiplying a factor of 1.302 and the values were then corrected for dilution due to replacement.

The conversion between moles and ppm followed the ideal gas law:

$$
10^{-6} mol = ppmv \times \frac{v}{RT}
$$
 where *V* is volume in L; *R* is gas constant,  
0.0821 Latm K<sup>-1</sup>mol<sup>-1</sup>; *T* is temperature in K  
Eq. 1

Atm CH4 oxidation follows first-order kinetics and hence rate constants, *k*, were calculated and used to estimate the oxidation rates at standardized CH4 concentration of 1.813 ppmv and expressed in nmol  $(g \text{ of soil})^{-1}$  day<sup>-1</sup>.

Atm CH4 oxidation rates obtained from microcosms experiments were scaled up to CH4 fluxes to compare with the *in situ* flux measurements at the field. The following formula was used:

$$
F = \frac{r \times FW}{d \times D}
$$
 where *F* is CH<sub>4</sub> flux in mg C m<sup>-2</sup> day<sup>-1</sup>; *r* is CH<sub>4</sub> oxidation rate in  
nmol g<sup>-1</sup>day<sup>-1</sup>; *FW* is formula weight of carbon, 12 g; *d* is density; *D* is depth in m

The density of  $1.8x10^6$  g m<sup>-3</sup> (Stackhouse *et al.*, 2014) was used and the assumption of methanotrophic activity within the first 5 cm of active layer was taken. This may underestimate the flux values because field flux measurements indicated that atm CH4 oxidation at AHI occurred down to 45 cm depth.

# *Temperature coefficients (Q10)*

Q10 is used to measure the rate of change of a chemical or biological reaction as a consequent of temperate increase of 10°C. It is a factor calculated from the following equation:

$$
Q_{10} = \left(\frac{R_2}{R_1}\right)^{\left(\frac{10}{T_2 - T_1}\right)}
$$
 where  $R_2$  and  $R_1$  are rates, in the same unit, measured  
respectively at  $T_2$  and  $T_1$ , in the same unit; for  $T_2 > T_1$  Eq. 3

 $Q_{10}$  values were computed for methanotrophy using (1) mean CH<sub>4</sub> oxidation rates obtained from microcosms for each treatment and corresponding incubation temperatures; and (2) mean CH4 oxidation rates reported in the literature for which temperature data was available. The criterion that the temperature difference between  $T_2$ and  $T_1$  being larger than 5 $\degree$ C was applied.

#### *Arrhenius relationship between in situ CH4 uptake flux and surface soil temperature*

Since we are cautious about quantitatively extrapolating the observed effects of temperature and water saturation under laboratory conditions to the real situation and model prediction, only CH4 fluxes and the corresponding surface soil temperatures measured during 2011-2013 expeditions were used to determine the Arrhenius relationship. Fluxes measured by open-circuit chambers were used to determine the relationship whereas those measured by closed-static chambers were excluded to eliminate the variation resulted from different collection methods (Whalen *et al.*, 1992). Natural logarithm of  $CH_4$  uptake fluxes (y-axis) was plotted against 1000/temperature (xaxis) to create an Arrhenius plot, which showed a potential change in the slope. The data was then analyzed using the "Segmented" R package (cran.rproject.org/web/packages/segmented/). Davies' test was used to determine whether the change in the slope was statistically significant. Given the result suggested a breakpoint occurred at 3.588 (equivalent to 5.6°C;  $p = 0.006$ ), function *segmented()* was used to estimate the breakpoint and the slopes (Results in Fig. S7).

Linear regression equations were fitted to data points below and above 5.6°C.

where *F* is CH<sub>4</sub> flux in mg C m<sup>-2</sup>day<sup>-1</sup>;  $E_a$  is activation energy in kJ mol<sup>-1</sup>; *R* is gas constant, 8.314 J K<sup>-1</sup>mol<sup>-1</sup>; *T* is temperature, in K; density; *A* is pre-exponential factor

 $LN(F) = \frac{-E_a}{R} \left(\frac{1000}{T}\right) + LN(A)$  Eq. 4

 $Q_{10}$  and active energy (E<sub>a</sub>) of atm CH<sub>4</sub> oxidation were derived from Eq. 3 and Eq. 4 respectively. 95% confidence intervals were calculated for each slope.

For comparative purpose, other data were overlain on the plot (Fig. 4), which include: (1) CH<sub>4</sub> fluxes estimated from our microcosm experiments at 2.0 ppmv of CH<sub>4</sub>; (2) CH4 fluxes estimated from intact core thawing experiments (Stackhouse *et al.*, 2014); and (3) atm CH4 oxidation sites at lower latitudes.

#### *Estimation of monthly and annual atm CH4 uptake fluxes*

Monthly air temperatures at AHI during 1990s and 2090s were simulated through the Climate Model Intercomparison Project (CMIP5) using 8 climate models (BCC-CSM1.1, CCSM4, CSIRO-Mk3.6.0, GFDL-ESM2M, GISS-E2-R, HadGEM2-AO, IPSL-CM5A-MR and NorESM1-M) (Taylor *et al.*, 2012). The 'high emissions scenario' assuming mitigation policies in action (RCP8.5) was used to project the climate change in 2090s. Monthly and annual atm CH4 uptake were estimated for temperatures from each model with the following assumptions:

a) Atm CH<sub>4</sub> uptake occurs at ground temperatures above  $0^{\circ}$ C. Field measurements taken during initial thaw in 2013 when soil surface temperature slowly warmed from -2°C to 10 $^{\circ}$ C, CH<sub>4</sub> uptake increased from undetectable to -0.24 mg CH<sub>4</sub>-C m<sup>-2</sup> day<sup>-1</sup>). First detectable consumptive flux corresponded roughly to the time when soil surface temperature was consistently above freezing.

b) Atm CH4 uptake fluxes increase with temperature following an Arrhenius relationship (Eq. 4) and at a faster rate below 5.6°C than that above 5.6°C.

The sum of monthly uptake fluxes multiplied by the number of days in the month equaled the mean annual uptake flux. Upper and lower 95% confidence intervals were regarded as maximum and minimum annual uptake fluxes. Multi-model means were obtained by taking average across all models.

Air temperatures at Eureka, Ellesmere Island, Nunavut, Canada (N80°00'03", W86°00'25"; 112 km NE of AHI) were available for 2010 and 2011 through Total Carbon Column Observing Network (TCCON) (Wunch *et al.*, 2011). Temperature data (T) from late March to August 2011 was used. Missing data was gap-filled by linear interpolation between the two neighboring values. The data of 2010 tracked nicely that of 2011, thus the average temperature in Sept 2010 was used to substitute the missing data of Sept 2011. Mean daily temperatures were calculated by averaging multiple measurements on the day, which were then averaged to give monthly temperatures. The monthly and annual uptake fluxes were estimated as aforementioned. Eureka is located at higher latitude where the temperature is slightly cooler than that at AHI.  $CH<sub>4</sub>$  uptake fluxes therefore were also calculated for  $T+1\textdegree C$  and  $T+6\textdegree C$  which, respectively, are more representative for our study site and also to mimic severe summer warming which was not projected by climate models.

#### Legends to supplementary figures and tables

**Fig. S1.** Phylogenetic tree of *pmo*A genes constructed from deduced amino acid sequences (161 taxa and 171 aa). Highlighted is the pan-*pmo*A gene recovered in this study. Sequences of ammonia-oxidizing monoxygenase (*amo*A) were used as the outgroup. Clusters of atmospheric  $CH_4$  oxidizers are annotated in reference to Kolb (2009) (Kolb, 2009). Bootstrap values greater than 50% are shown as branch label. The scale bar represents a substitution rate of 0.2 changes per position.

**Fig. S2.** Phylogenetic tree of *pmo*B genes constructed from deduced amino acid sequences (69 taxa and 377 aa). Highlighted are the pan-*pmo*B genes recovered in this study. Sequences of ammonia-oxidizing monoxygenase (*amo*B) were used as the outgroup. Bootstrap values greater than 50% are shown as branch label. The scale bar represents a substitution rate of 0.2 changes per position.

**Fig. S3.** Phylogenetic tree of *pmoC* genes constructed from deduced amino acid sequences (82 taxa and 244 aa). Highlighted are the pan-*pmo*C genes recovered in this study. Sequences of ammonia-oxidizing monoxygenase (*amo*C) were used as the outgroup. Bootstrap values greater than 50% are shown as branch label. The scale bar represents a substitution rate of 0.2 changes per position.

**Fig. S4.** Alignment of translated amino acids of *pmo*B contigs from metagenomic and metatranscriptomic libraries. Histidine residues (H33, H137 and H139) that coordinate the di-copper center (aka the active site of pMMO) are highlighted in blue. GenBank

sequences YP\_115247 and CAJ01562 encode for pmoB of MOB *Methylococcus capsulatus* str. Bath and atmMOB USCα.

**Fig. S5.** Predicted monthly air temperatures at Axel Heiberg Island, Canada. Monthly air temperatures at AHI during 1990s and 2090s were simulated through the Climate Model Intercomparison Project (CMIP5). T: Gap-filled monthly air temperatures in 2011 at Eureka, Ellesmere Island, Nunavut, Canada downloaded from Total Carbon Column Observing Network (TCCON).

**Table S1.** CH<sub>4</sub> field fluxes in the Northern Circumpolar permafrost region.

**Table S2.** Methanotrophic (A) and methanogenic (B) taxa identified in near-surface cryosols (at 5 cm depth) in the intact core warming experiments.

**Table S3.** Genes of methane monooxygenases and homologous enzymes identified in (a) metagenome data from near-surface cryosols (at 5 cm depth) in the intact core warming experiments and (b) metatranscriptome data from the polygon trough sample (collected on July 15, 2013).

**Table S4.** Aerobic methanotrophs detected in Arctic permafrost-affected region.

**Table S5.** Methanotrophic proteins identified in near-surface cryosols (at 5 cm depth) in the intact core warming experiments.

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Uncultured bacterium AC-A2 [CAI30616]



Nocardioidaceae bacterium Broad-1 [WP\_008359136]

# Upland Soil Cluster alpha (USCa) and Cluster 5 (CL5)

 $\mathbb{\mathsf{L}}$  Methylacidiphilum infernorum V4 [YP\_001940242]

**Methylacidiphilum fumariolicum SolV [CCG92189]** 

Methylacidiphilum fumariolicum [WP\_009060833] Methylacidiphilum infernorum V4 [YP\_001940161] Methylacidiphilum kamchatkense Kam1 [AFC75745] Methylacidiphilum fumariolicum [WP\_009059717] Methylacidiphilum kamchatkense Kam1 [AFC75742] Methylacidiphilum fumariolicum SolV [CCG92749] 99 - Methylacidiphilum infernorum V4 [YP\_001940241] Methylacidiphilum infernorum V4 [ABX56605] 100 Methylacidiphilum kamchatkense Kam1 [AFC75748] Methylacidiphilum infernorum V4 [YP\_001940158] Methylacidiphilum fumariolicum SolV [CCG92747] 100<sub>1</sub> Methylacidiphilum fumariolicum SolV [CCG92188]



Nocardioidaceae bacterium Broad-1 [WP\_008359134]

# Upland Soil Cluster alpha (USCa)





C75746] 10243] 2190] 9720]

B4 [YP\_006442845] erium rhodesiae NBB3 [YP\_005000795] d-1 [WP\_008359138]



Fig. S3. Phylogenetic tree of pmoC genes constructed from deduced amino acid sequences (82 taxa and 244 aa). Highlighted are the pan-pmoC genes recovered in this study. Sequences of ammonia-oxidizing monoxygenase (amoC) were used as the out-group. Bootstrap values greater than 50% are shown as branch label. The scale bar represents a substitution rate of 0.2 changes per position.

Fig. S4 Alignment of translated amino acids of *pmo*B contigs from metagenomic and metatranscriptomic libraries. Histidine residues (H33, H137 and H139) that coordinates the di-copper center (aka the active site of pMMO) are highlighted in blue. GenBank sequences YP\_115247 and CAJ01562 encode for pmoB of MOB *Methylococcus capsulatus* str. Bath and atmMOB USCα.



Fig. S5. Predicted monthly air temperatures at Axel Heiberg Island, Canada. Monthly air temperatures at AHI during 1990s and 2090s were simulated through the Climate Model Intercomparison Project (CMIP5). T: Monthly air temperatures in 2011 at Eureka, Ellesmere Island, Nunavut, Canada downloaded from Total Carbon Column Observing Network (TCCON)









Note:<br><sup>1</sup> Percentage by weight (wt%) unless specified<br><sup>2</sup> Ambient temperature unless specified<br><sup>3</sup> Temperature range in year 2004<br>NA: Not available

#### Table S2. Methanotrophic (A) and methanogenic (B) taxa identified in near-surface cryosols (at 5 cm depth) in the intact core warming experiments.

**(A)**



**(B)**



Notes:

1 Genera not detected: *Clonothrix*, *Crenothrix, Methylohalobius, Methylosoma, Methylothermus, Methylosphaera, Methylomarinum, Methylocaldum* and *Methylogaea.*

<sup>2</sup>The closest related isolate to atm CH<sub>4</sub>-oxidizing bacterium Upland Soil Cluster alpha (USC $\alpha$ ), as proposed by the phylogeny of the pmoCAB operon and 4 other open read frames (Ricke et al., 2005)<br><sup>3</sup>Mean abundance of

4 MMO: Methane monooxygenase; pMMO: particulate MMO; sMMO: soluble MMO

<sup>6</sup>This genus contains representatives that are capable of oxidizing CH<sub>4</sub> anaerobically. <sup>5</sup>Genera that were sporadically present in some samples with a relative abundances < 0.001% include: Methanobacterium, Methanosphaera, Methanothermus, Methanothermococcus, Methanoplanus and Methanopyrus. Table S3. Genes of methane monooxygenases and homologous enzymes identified in (A) metagenome data from near-surface cryosols (at 5 cm depth) in the intact core warming experiments and (B) metatranscriptome data from the p



**(B)**



Notes:<br>"Metagenome contigs are available on MG-RAST (ID: 4530050.3)<br>"Metan abundance of 10 cryosol samples collected at different time during the course of incubation (T=1 week and 6 months).<br>"Metatranscriptome (or transc

#### Table S4. Aerobic methanotrophs detected in Arctic permafrost-affected region



Notes.<br>Biography distances are last may be applicant support of atmospheric CH<sub>4</sub> oxidation<br>Bive highlighted species we known perchapation of a may be absolved form methanotrophs of methanotrophs in a may be a may be a spe

Verrucomicrobia

**Phylum Class Family Genus**

Proteobacteria

#### Table S5. Methanotrophic proteins identified in near-surface cryosols (at 5 cm depth) in the intact core warming experiments.



Notes:<br>"Contig sequences generated by our studies are marked with an asterisk.<br>"NSAF: Normalized Spectral abundance factor. Adjusted NSAF values are NSAF values multiplied by 1000000.<br>"Unique peptide sequences are highligh