

Supplement

Calcification response of a key phytoplankton family to
millennial-scale environmental change

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¹³ **Supplementary Discussion**

¹⁴ **Coccolith to coccospHERE relationships**

¹⁵ Figure S1 shows the relationships between coccolith dimensions and coccospHERE dimensions from
¹⁶ our culture experiments. These relationships underpin our rationale. The strong relationship be-
¹⁷ tween coccolith length (L_c) and coccospHERE diameter (D_s) in the reticulofenestrids is established-
¹⁸ Henderiks (2008):

$$D_s = 1.02 + 1.42L_c. \quad (\text{S1})$$

¹⁹ Given that,

$$A_c = \pi \left(\frac{L_c C}{2} \right)^2, \quad (\text{S2})$$

²⁰ where A_c = coccolith area, W_c and L_c = respectively the semi minor and major axes of the
²¹ elliptical coccolith and $C = \sqrt{W_c/L_c}$ = "circularity", EqS1 becomes:

$$R_s = 0.51 + \frac{1.42}{\sqrt{\pi}C} \sqrt{A_c}. \quad (\text{S3})$$

²² When a value of $C = 0.9$ is used, which is typicalHenderiks (2008), the relationship we find
²³ between the square root of coccolith area and coccospHERE radius is in very close agreement with
²⁴ EqS1. Equation S3 describes the red dashed line in Fig.1C. A more recently published relationship
²⁵ between coccolith length and coccospHERE diameter, has a different gradientMüller *et al.* (2010),
²⁶ but the measurements of length are based on coccolith volume, related to distal shield length via
²⁷ an equation from the literatureYoung & Ziveri (2000) - not from direct measurements of area
²⁸ or length. Our direct measurement of and correlation between coccolith dimensions and molar
²⁹ PIC:POC, circumvents the complications associated with allometry and with multiple layers of
³⁰ coccoliths.

³¹ Other geological coccolith time series

³² Some geological time series of coccolith mass from the literature found a decrease in coccolith
³³ mass with increasing CO₂Beaufort *et al.* (2011); Meier *et al.* (2014a), whilst others found the
³⁴ oppositeIglesias-Rodriguez *et al.* (2008); Meier *et al.* (2014b). Increasing coccolith mass in the
³⁵ absence of area changes would be expected to correspond to an increase in coccolithophore PIC:POC
³⁶ according to Eq. 3. It is possible that these contrasting conclusions may be reconciled when
³⁷ PIC:POC rather than coccolith mass alone is considered. Unfortunately the data from these studiesBeaufort
³⁸ are not appropriate for this analysis. The earlier version of SYRACO used in these studiesBeaufort
³⁹ *et al.* (2011); Meier *et al.* (2014a,b) yielded significant underestimates of coccolith area, especially
⁴⁰ for E. huxleyi, and overestimates of coccolith thickness, rendering the estimation of PIC:POC
⁴¹ somewhat insensitive to changes in coccolith area. Improvements in the version of SYRACO used
⁴² in this studyBeaufort *et al.* (2014), are an increased resolution camera, higher magnification lens
⁴³ and images taken in triplicate at different angles to remove the extinction cross. Coccolith volume
⁴⁴ aloneIglesias-Rodriguez *et al.* (2008) is inappropriate because estimates of PIC:POC using Eq.3
⁴⁵ necessitate decoupling volume into thickness and area.

⁴⁶ Supplementary figures

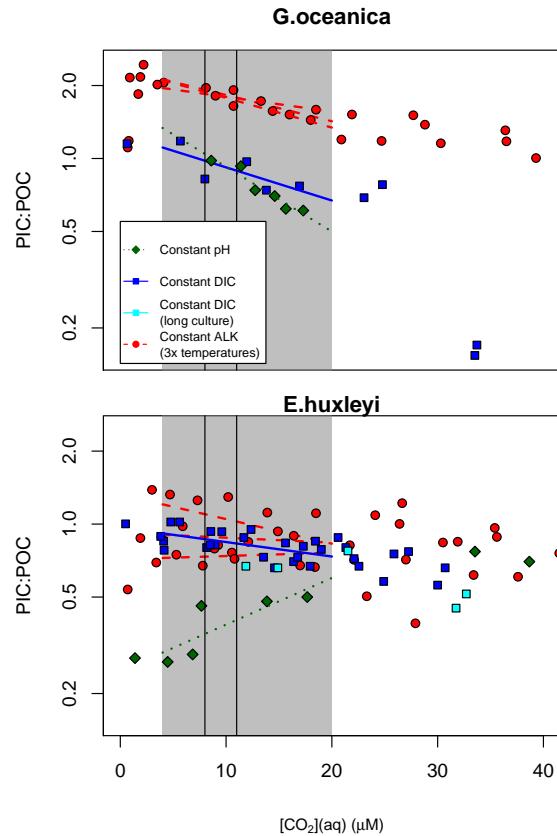


Figure 1: (S1): Plastic response of *E. huxleyi* and *G. oceanica* from the literature. Least-squares linear regressions for each method of carbon manipulation are shown, over a representative range of $[CO_2(aq)]$ (grey shaded region). The solid vertical lines delimit the range of $[CO_2(aq)]$ experienced at ODP site 1123 over the glacial terminations. Three separate regressions are given within the const.ALK experiments, for three experiments undertaken at different temperatures. Slopes are based on the $\sim 4\text{--}20 \mu M$ range of $[CO_2]$, to capture the linear part of the response, representative of the $[CO_2(aq)]$ range $\sim 8\text{--}11 \mu M$. *E. huxleyi* data from Bach *et al.* (2011); Iglesias-Rodriguez *et al.* (2008); Langer *et al.* (2009); Müller *et al.* (2010); Sett *et al.* (2014); Zondervan *et al.* (2001) and Riebesell *et al.* (2000), and *G. oceanica* data taken from Rickaby *et al.* (2010); Sett *et al.* (2014); Zondervan *et al.* (2001) and Riebesell *et al.* (2000)). Nb: The constant alkalinity *E. huxleyi* ($15^\circ C$) data from Bach *et al.* (2011) is the same as that of Sett *et al.* (2014) so has been omitted.

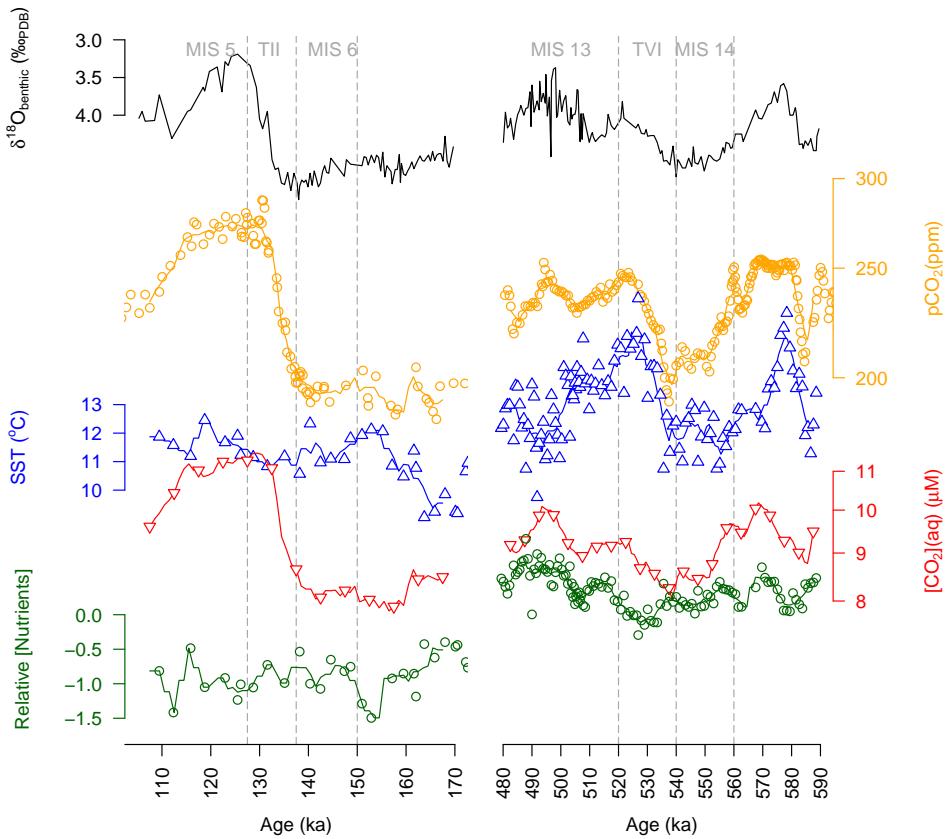


Figure 2: (S2) Time series of proxy-reconstructed climatic parameters at ODP site 1123. Benthic Oxygen isotopes from Elderfield *et al.* (2012) allow for temporal alignment with other records. Reconstructed sea surface temperature (SST) estimates are based on Mg/Ca ratios of planktic forams (see methods). $[\text{CO}_2(\text{aq})]$ is estimated from global $p\text{CO}_2$ of an assumed well mixed atmosphere from Vostok (*left*) and Dome C (*right*) Antarctic ice cores (compiled by Lüthi *et al.* (2008)), with dissolution assumed to be controlled only by SST at a constant salinity of 35. EDC3 gas age was converted to LR04 using a published conversion Parrenin *et al.* (2007). Carbon isotopic composition of planktic forams were used as a rough proxy for relative nutrient availability corrected for the effect of temperature (see methods).

⁴⁷ Regression data - fitted histograms (Fig.S3 to Fig.S10)

⁴⁸ To discount noise from the lower end of some of the mass and area size spectra, which is due to
⁴⁹ the occasional presence of coccolith fragments, and is an often unavoidable consequence of making
⁵⁰ smear slides, the values of coccolith morphometrics were found by independently fitting a gaussian
⁵¹ curve to each spectrum. The histograms responsible for the data shown in Figure 1B are presented
⁵² in FigsS4 to S11.

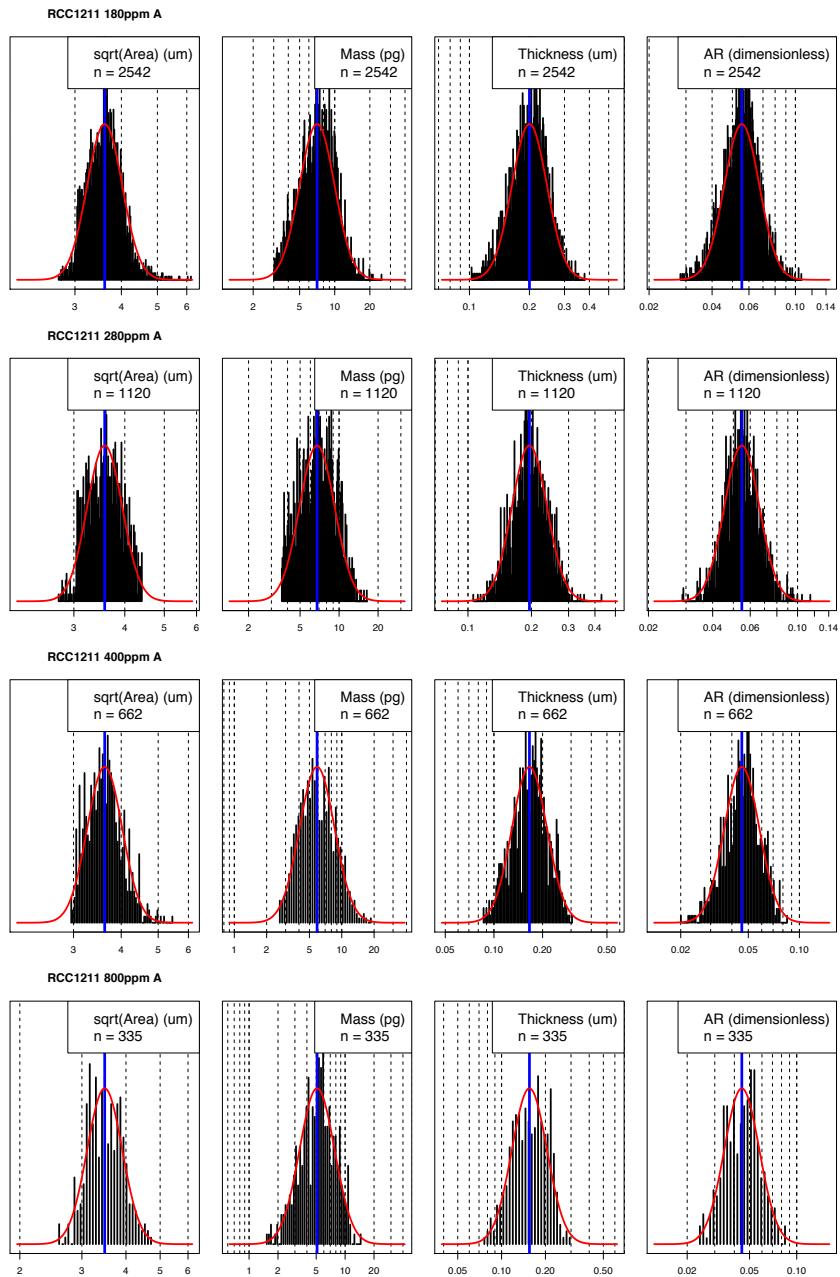


Figure 3: S3: Histograms (1)

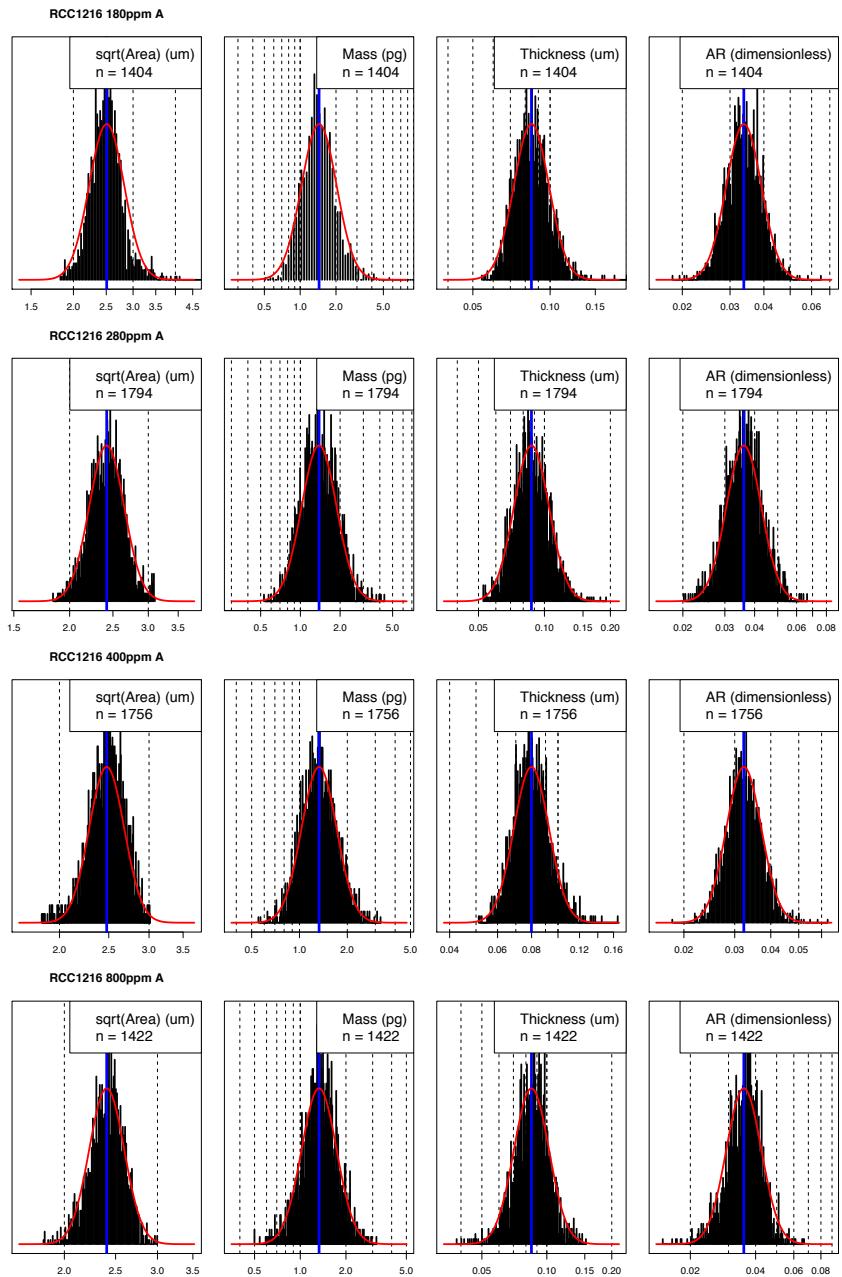


Figure 4: S4: Histograms (2)

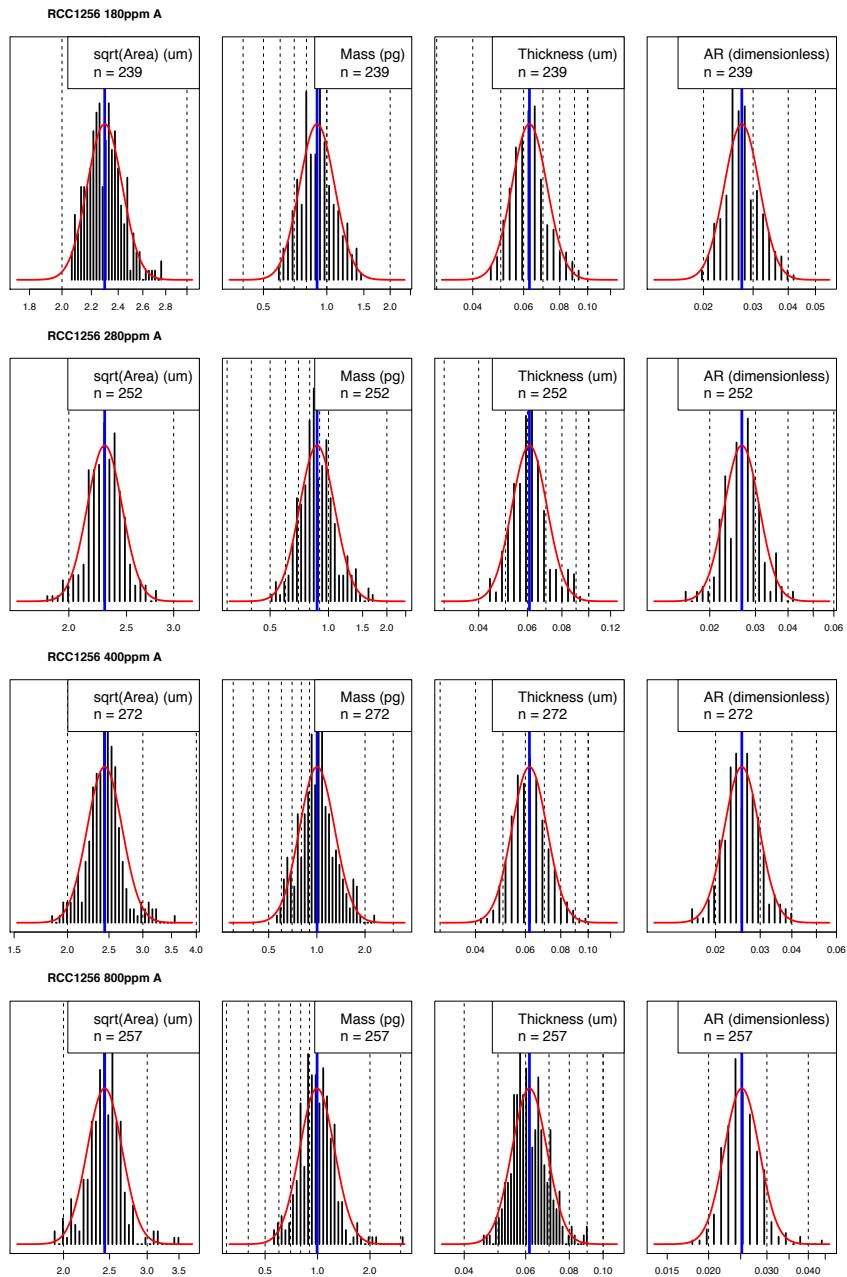


Figure 5: S5: Histograms (3)

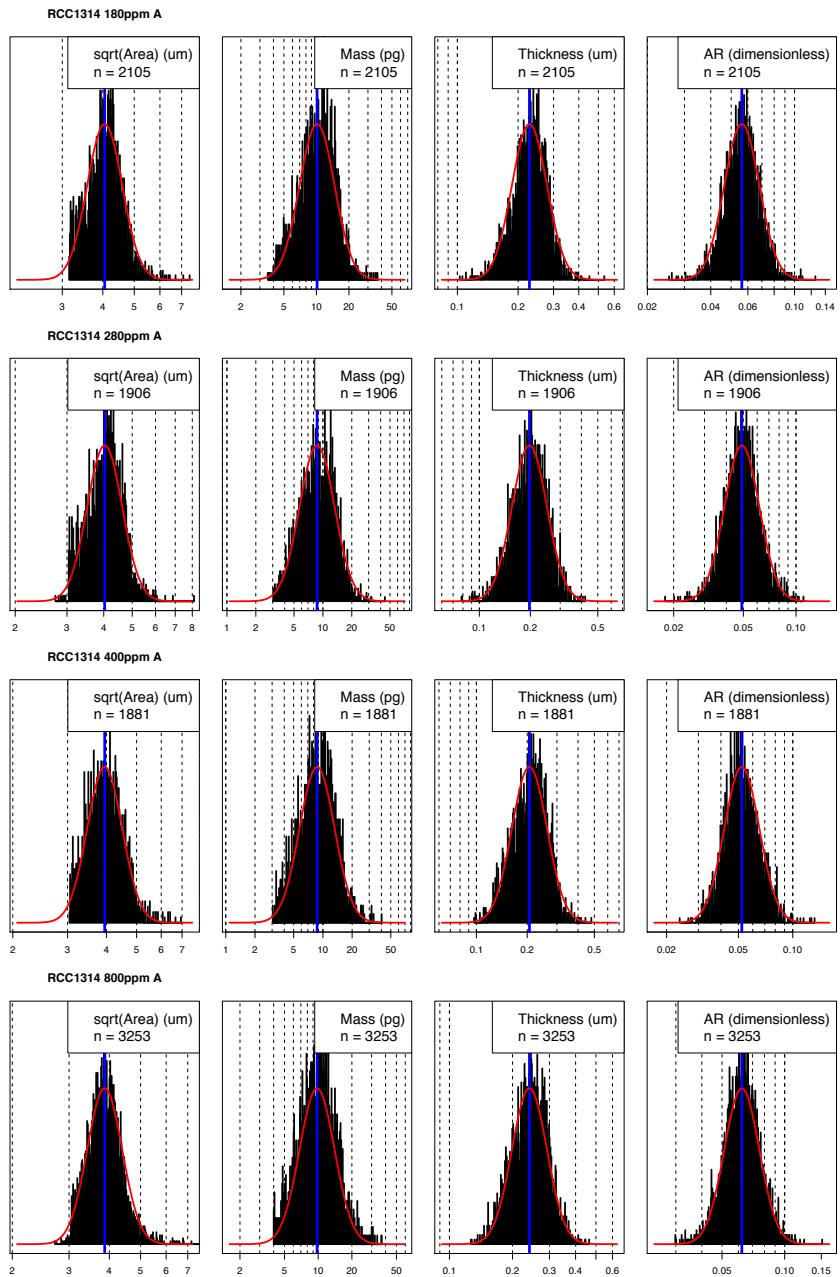


Figure 6: S6: Histograms (4)

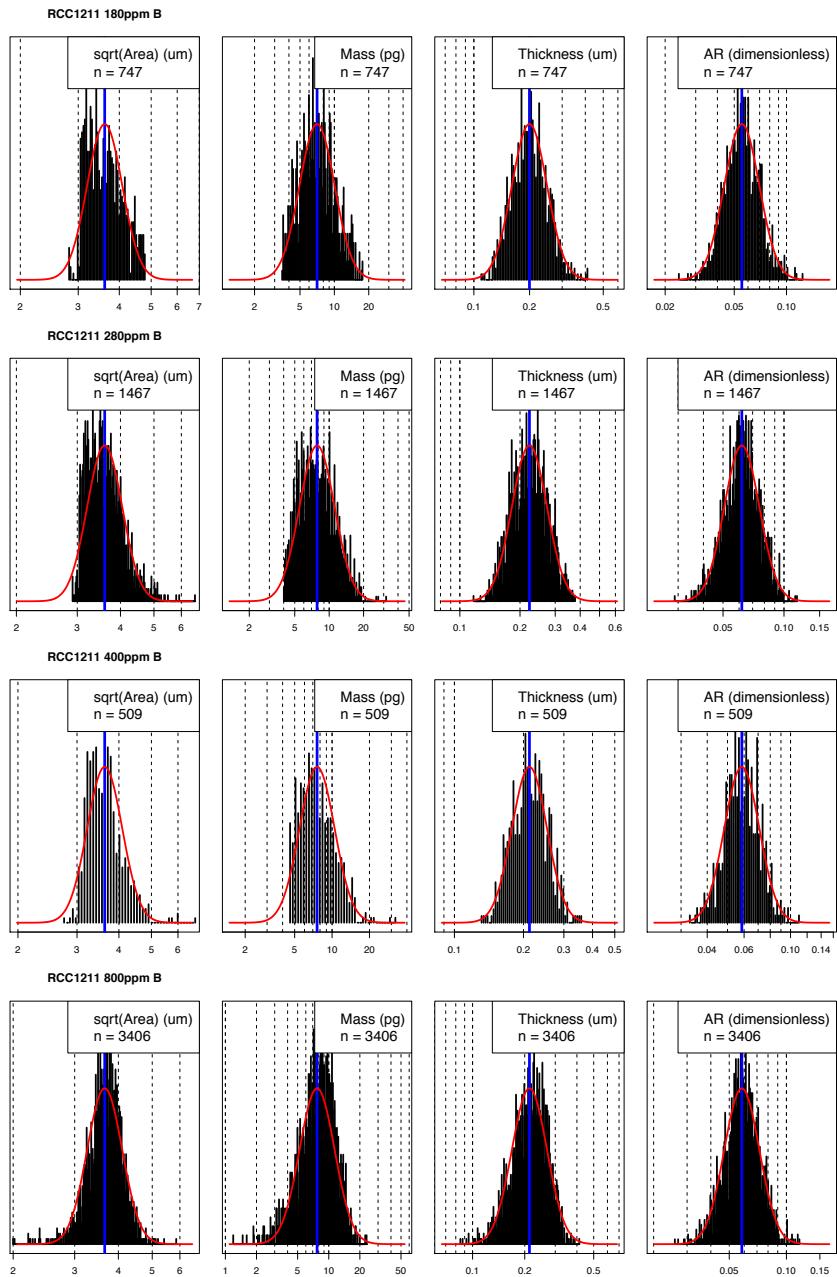


Figure 7: S7: Histograms (5)

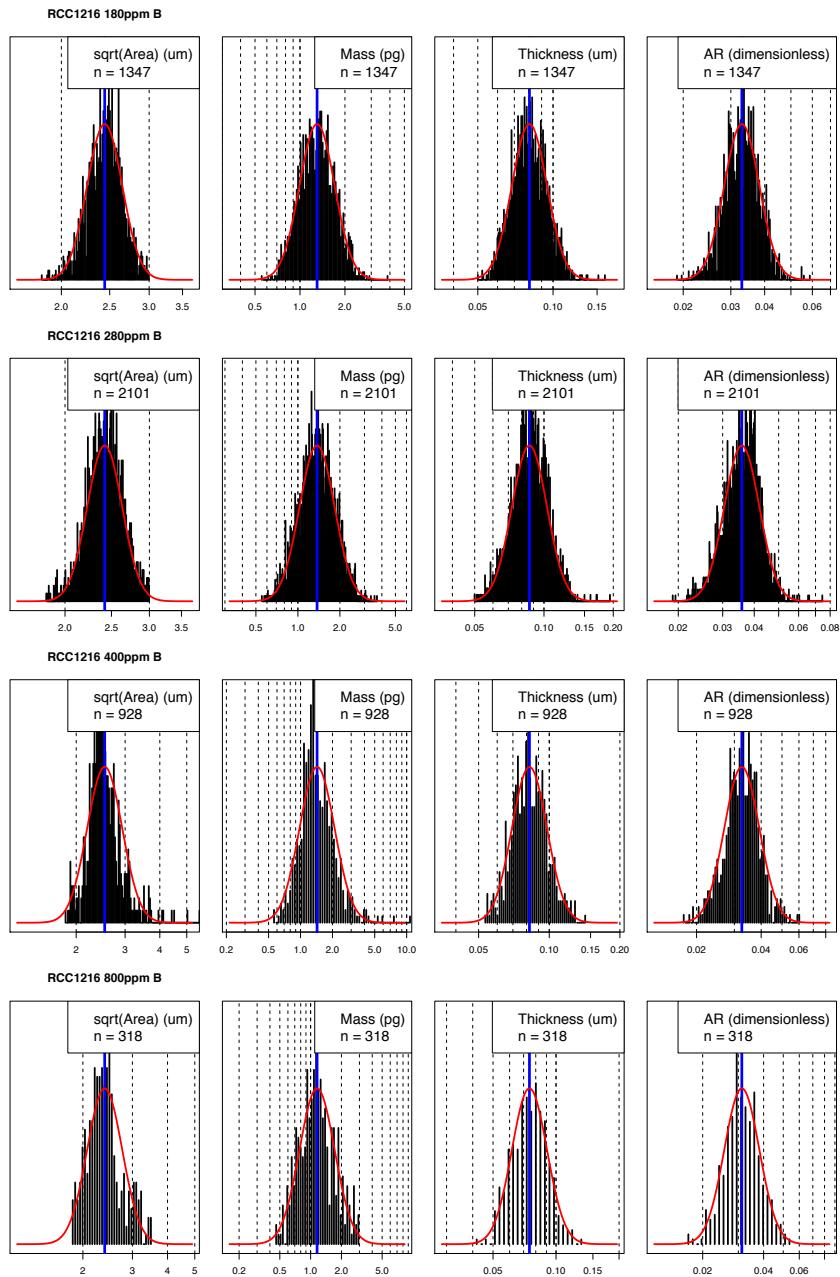


Figure 8: S8: Histograms (6)

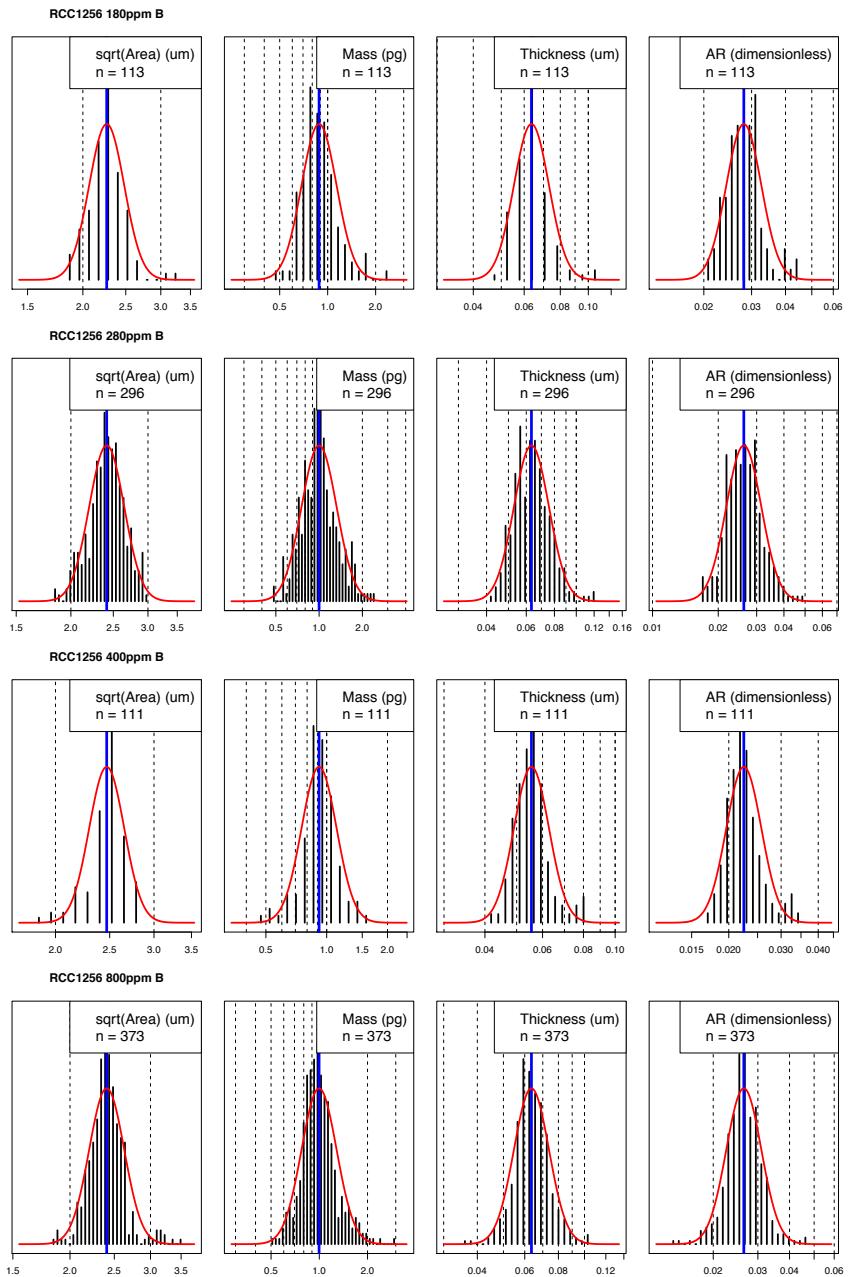


Figure 9: S9: Histograms (7)

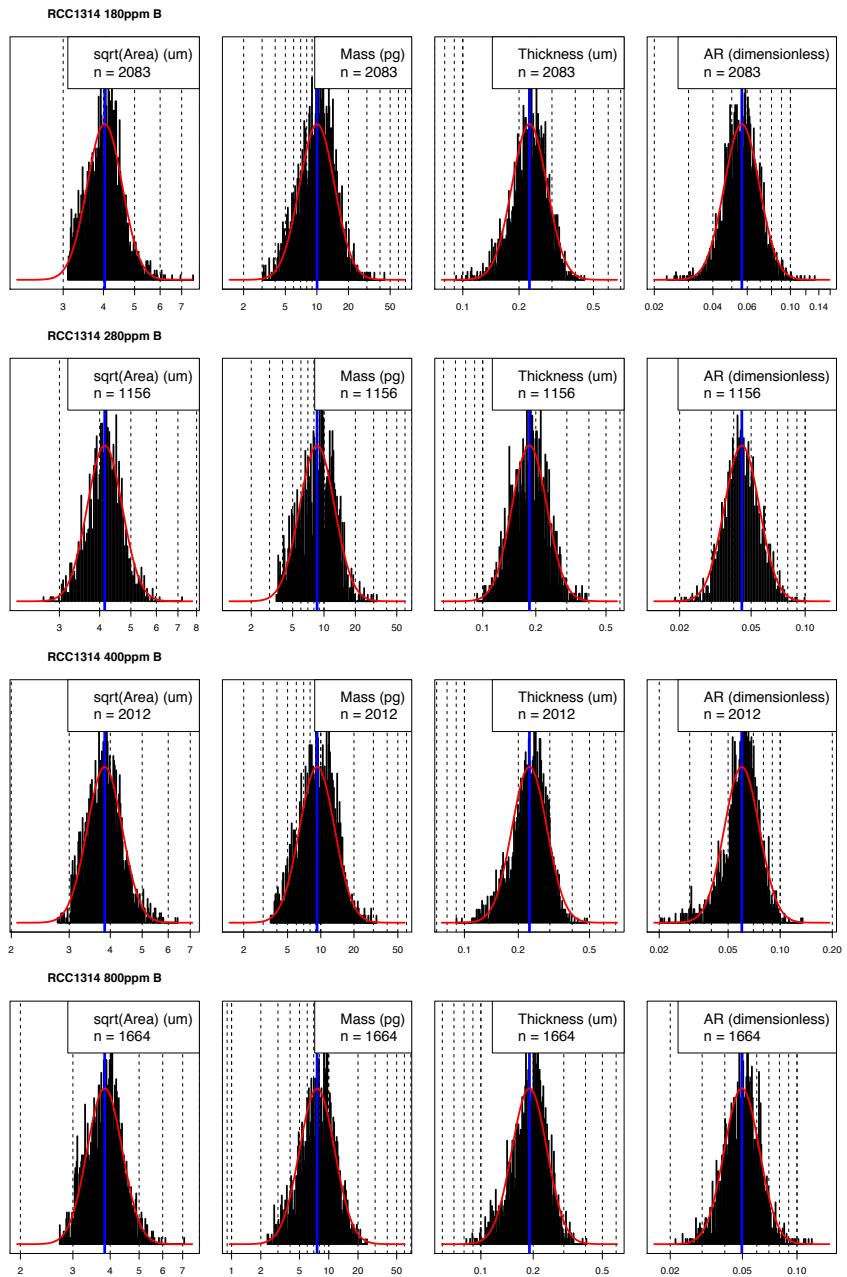


Figure 10: S10: Histograms (8)

⁵³ **Supplementary references**

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