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### Supplementary Materials for

## Reactive nitrogen chemistry in aerosol water as a source of sulfate during haze events in China

Yafang Cheng, Guangjie Zheng, Chao Wei, Qing Mu, Bo Zheng, Zhibin Wang, Meng Gao, Qiang Zhang, Kebin He, Gregory Carmichael, Ulrich Pöschl, Hang Su

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### This PDF file includes:

- fig. S1. Weakened photochemistry by aerosol dimming effects during January 2013 in Beijing.
- fig. S2. Importance of the NO<sub>2</sub> reaction pathway for sulfate production in the Beijing haze (January 2013).
- fig. S3. Influence of ionic strength (*I*) on rate of aqueous sulfate-producing reactions.
- fig. S4. Estimation of Fe<sup>3+</sup> and Mn<sup>2+</sup> concentrations as a function of aerosol water pH during Beijing hazes.
- fig. S5. Regional pollution across the NCP during January 2013.
- fig. S6. Annual precipitation pH of China in 2013.
- fig. S7. The same as Fig. 2 but with a lower limit of reaction rate constants reported by Lee and Schwartz (18).
- table S1. Previously reported concentrations of cations and anions in PM<sub>2.5</sub> during winter for cities in NCP used in Fig. 1D.
- table S2. Summary of field observation and methods in this study.
- table S3. Domain, configurations, and major dynamic and physical options used in WRF v3.5.1.
- table S4. Rate expression and rate coefficients of relevant aqueous-phase reactions.
- table S5. Constants for calculating the apparent Henry's constant  $(H^*)$ .
- table S6. Summary of suggested activity coefficient (*a*)-ionic strength (*I*) dependence.
- table S7. Influence of ionic strength (*I*) on rate of aqueous sulfate-producing reactions.

• References (108–128)

#### **Supplementary Materials**



fig. S1. Weakened photochemistry by aerosol dimming effects during January 2013 in Beijing. (A) Relationship between the observed total surface solar radiation and ground based measurement of PM<sub>2.5</sub>. (B) Diurnal pattern of observed O<sub>3</sub> and model simulated OH (7) for hours with PM<sub>2.5</sub>  $\leq 200 \ \mu g \ m^{-3}$  and hours with PM<sub>2.5</sub>  $\geq 200 \ \mu g \ m^{-3}$ . (C) The observed O<sub>3</sub> and modeled OH (7) are plotted against the observed PM<sub>2.5</sub> concentration. Only midday (10:00–15:00 local time) data when photochemistry is expected to be the strongest of the day are shown.



fig. S2. Importance of the NO<sub>2</sub> reaction pathway for sulfate production in the Beijing haze (January 2013). Sulfate production rates calculated for the aqueous phase NO<sub>2</sub> reaction pathway (red circles) and the gas phase OH reaction pathway (green crosses) compared to the missing source of sulfate (black diamonds) (geometric mean  $\pm$  SD) (5, 7). The pH of aerosol water here is calculated based on the observed size-segregated aerosol chemical composition (70) as shown in Fig. 1D (PM<sub>1.1</sub>, diamond; PM<sub>1.1-2.1</sub>, triangle).



 $I^{1/2}$  (M<sup>1/2</sup>) **fig. S3. Influence of ionic strength** (*I*) **on rate of aqueous sulfate-producing reactions.**  $k^{I=0}$  indicates the kinetic constant at I = 0 M. Detailed expressions of k and the  $\log(k/k^{I=0}) \sim I$  relationship are listed in table S7.  $O_3^{*a}$  and  $O_3^{*b}$  refer to the relationship suggested by Maahs *et al.* (53) and Lagrange *et al.* (54), respectively. More details can be found in section M8.



fig. S4. Estimation of  $Fe^{3+}$  and  $Mn^{2+}$  concentrations as a function of aerosol water pH during Beijing hazes. At higher pH, the solubility of Fe(OH)<sub>3</sub> and Mn(OH)<sub>2</sub> may become the limiting factor. The corresponding precipitation constant at 298K are 2.6 ×10<sup>-38</sup> for Fe(OH)<sub>3</sub> and 1.6×10<sup>-13</sup> for Mn(OH)<sub>2</sub> (66).



**fig. S5. Regional pollution across the NCP during January 2013.** (A) Topographic map around Beijing, and location of all cities shown below. (B) PM<sub>2.5</sub> (C) SO<sub>2</sub> and (D) NO<sub>2</sub> concentrations in Beijing and its southwest (upper panel) and southeast (lower panel) cities.



**fig. S6. Annual precipitation pH of China in 2013.** (source: Report of China's Environmental Quality in 2013, by Ministry of Environmental Protection of People's Republic of China, <u>http://jcs.mep.gov.cn/hjzl/zkgb/</u>).



fig. S7. The same as Fig. 2 but with a lower limit of reaction rate constants reported by Lee and Schwartz (18).

City	Site Type	Veen	Cations <sup>b</sup>	Anions <sup>b</sup>	Cations	Deference	
		Year	(µmol m <sup>-3</sup> )	(µmol m <sup>-3</sup> )	/ Anions	Kelerence	
	Urban	2009-2010	0.41	0.39	1.03	(108)	
	Urban	2007	0.45	0.37	1.20	(82)	
	Urban	2005-2006	0.84	0.66	1.28	(100)	
	Rural	2005-2006	0.45	0.31	1.48	(109)	
Beijing		2003	1.20	1.10	1.09	(22)	
5 0	Urban	2002-2003	0.93	1.09	0.85		
	Urban	) 2002-2003	0.96 0	0.88	1.09		
	(Industrial)					(110)	
	Urban	2002-2003	1.44	1.14	1.26		
	(Residential)						
Linon	Urban	2007~2008	1.85	1.49	1.24	(111)	
Jinan	Urban	2004~2005	1.66	1.39	1.20	(83)	
Tianjin	Urban	2003	1.78	1.63	1.10	(22)	
Qingdao	Urban	2003	1.53	1.32	1.16	(22)	

table S1. Previously reported concentrations of cations and anions in PM<sub>2.5</sub> during winter for cities in NCP used in Fig. 1D.

<sup>a</sup> Only studies which have reported the seasonal-averaged concentrations of all major water-soluble inorganic ions (i.e. Na<sup>+</sup>, K<sup>+</sup>, Ca<sup>2+</sup>, Mg<sup>2+</sup>, NH<sub>4</sub><sup>+</sup>, SO<sub>4</sub><sup>2-</sup>, NO<sub>3</sub><sup>-</sup> and Cl<sup>-</sup>) are selected. <sup>b</sup> Concentrations of cations and anions are estimated as

Cations (µmol m<sup>-3</sup>) = 
$$\frac{[Na^+]}{23} + \frac{[K^+]}{39} + \frac{[Ca^{2+}]}{20} + \frac{[Mg^{2+}]}{12} + \frac{[NH_4^+]}{18}$$
  
Anions (µmol m<sup>-3</sup>) =  $\frac{[Cl^-]}{35.5} + \frac{[SO_4^{-2-}]}{48} + \frac{[NO_3^{-1}]}{62}$ 

Species	Method	Time Resolution	Period
PM <sub>2.5</sub> and PM <sub>10</sub>	PM-712 Monitor (Kimoto Electric	1h	Jan. 1 <sup>st</sup> to Jan. 31 <sup>st</sup>
Meteorology	Milos520 Weather Station (VAISALA Inc. Finland)	1h	Jan. 1 <sup>st</sup> to Jan. 31 <sup>st</sup>
OC and EC	Sunset model 4 semi-continuous carbon analyzer (Beaverton, OR, USA)	1h	Jan. 1 <sup>st</sup> to Jan. 31 <sup>st</sup>
SO <sub>4</sub> <sup>2-</sup> and NO <sub>3</sub> <sup>-</sup>	ACSA-08 Monitor (Kimoto Electric Co., Ltd., Osaka, Japan)	1h	Jan. 1 <sup>st</sup> to Jan. 31 <sup>st</sup>
Water-Soluble Inorganic Ions <sup>a</sup>	Offline Sampling - Ion Chromatograph analysis	2h	Jan. 12 <sup>th</sup> to Jan. 24 <sup>th</sup>
Water-Soluble Inorganic Ions <sup>a</sup>	Offline Sampling - Ion Chromatograph analysis	24h	Jan. 1 <sup>st</sup> to Jan. 31 <sup>st</sup>

table S2. Summary of field observation and methods in this study.

<sup>a</sup> Including Na<sup>+</sup>, K<sup>+</sup>, Ca<sup>2+</sup>, Mg<sup>2+</sup>, NH4<sup>+</sup>, SO4<sup>2-</sup>, NO3<sup>-</sup> and Cl<sup>-</sup>.

Simulation period	Dec 2012 and Jan 2013		
	East Asia (columns: 178, rows: 133) with three extra		
Domain	grids in each boundary of Domain 1 (columns: 172,		
	rows: 127)		
Horizontal resolution	36 km		
Vartical resolution	23 sigma levels from surface to tropopause (about 100		
vertical resolution	mb)		
Mataorological IC and PC	Reanalysis data from the National Centers for		
Meteorological IC and BC	Environmental Prediction Final Analysis (NCEP-FNL)		
Shortwave radiation	New Goddard scheme (112)		
Longwave radiation	The rapid radiative transfer model (RRTM) (113)		
Land surface model	The USGS 24-category land use data		
Surface layer	Pleim–Xiu land surface scheme (114)		
Planetary boundary layer model	ACM2 PBL scheme (115)		
Cumulus parameterization	Kain–Fritsch cumulus scheme (116)		
Cloud microphysics	WSM6 (117)		
Analysis nudging	Temperature and water vapor mixing (above PBL); wind		
	(in and above PBL)		
Observational nudging	Temperature, water vapor mixing and wind (in and		
	above PBL)		
Soil nudging	Include soil moisture and temperature		
FDDA data	NCEP Automated Data Processing (ADP) surface		
TDDA data	(ds461.0) and upper (ds351.0) air data		

table S3. Domain, configurations, and major dynamic and physical options used in WRF v3.5.1. (Source: Zheng et al. (7)).

Oxidants	Rate Expression, -d[S(IV)]/dt	Reference
O <sub>3</sub>	$\begin{array}{l} (k_0[SO_2 \cdot H_2O] + k_1[HSO_3^{-1}] + k_2[SO_3^{2-1}])[O_3(aq)] \\ k_0 = 2.4 \times 10^4 \ M^{-1}s^{-1} \\ k_1 = 3.7 \times 10^5 \ M^{-1}s^{-1},  E/R = 5530 \ K^a \\ k_2 = 1.5 \times 10^9 \ M^{-1}s^{-1},  E/R = 5280 \ K \end{array}$	(12)
$H_2O_2$	$ \begin{array}{l} (k_3[H^+][HSO_3^-][H_2O_2(aq)])/(1+K[H^+]) \\ k_3 = 7.45 \times 10^7 \ M^{-1}s^{-1} \ ,  E/R = 4430 \ K \\ K = 13 \ M^{-1} \end{array} $	(12)
TMI+O <sub>2</sub> , where TMI are Fe(III) and Mn(II) <sup>b</sup>	$\begin{split} pH &\leq 4.2,  k_4 [H^+]^{-0.74} [Mn(II)] Fe(III)] [S(IV)] \\ k_4 &= 3.72 \times 10^7  M^{-1} s^{-1} \\ pH &> 4.2,  k_5 [H^+]^{0.67} [Mn(II)] Fe(III)] [S(IV)] \\ k_5 &= 2.51 \times 10^{13}  M^{-1} s^{-1} \end{split}$	(118)
NO <sub>2</sub> <sup>e</sup>	$ \begin{array}{c} k_6[NO_2(aq)][S(IV)] \\ k_{6, \ low}{}^c = (0.14 \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \ \$	(12, 18, 61)
MHP (methyl-hydrogen peroxide)	$k_7[H^+][HSO_3^-][MHP(aq)]$ $k_7 = 1.75 \times 10^7 \text{ M}^{-1}\text{s}^{-1}, \text{ E/R}= 3801 \text{ K}$	(20)
PAA (peroxyacetic acid)	$(k_8[HSO_3^-][PAA(aq)])/([H^+]+K_8)$ $k_8 = 3.64 \times 10^7 M^{-1}s^{-1}$ , E/R= 3994 K $K_8 = 1.65 \times 10^{-5} M$	(20)

table S4. Rate expression and rate coefficients of relevant aqueous-phase reactions.

<sup>a</sup> According to the Arrhenius equation, the dependence of kinetic constant *k* on temperature T could be expressed as

$$k(T) = k(T_0) \exp\left[-\frac{E}{R}\left(\frac{1}{T} - \frac{1}{T_0}\right)\right]$$
 where  $T_0 = 298$  K.

- <sup>b</sup> Only Fe(III) and Mn(II) are considered here, since other transition metal ions (TMIs), such as Sc(III), Ti(III), V(III), Cr(III), Co(II), Ni(II), Cu(II) and Zn(II), showed much less catalytic activities (*58*). In addition, it has been concluded that the decreased temperature would generally lead to a decrease in overall reaction rate, due to the effect that higher activation energy would overweigh the increased SO<sub>2</sub> solubility (*12*). Due to the lack of information on E/R, we used both kinetic and SO<sub>2</sub> solubility at 298K, which would be an overestimation in Beijing haze scenario (271K) for the sulfate production rate by TMIs reaction pathway.
- <sup>c</sup> The lower estimate of  $k_6$  ( $k_{6, low}$ ) are derived from Lee and Schwartz (*18*). They determined a  $k_6 = 1.4 \times 10^5 \,\text{M}^{-1}\text{s}^{-1}$  at pH 5.0, but only a lower limit of  $2 \times 10^6 \,\text{M}^{-1}\text{s}^{-1}$  at pH 5.8 and 6.4. Thus we used a stepwise function with pH for estimation of  $k_{6, low}$ . When pH < 5,  $k_{6, low} = 1.4 \times 10^5 \,\text{M}^{-1}\text{s}^{-1}$ , while for pH>5.8,  $k_{6, low} = 2 \times 10^6 \,\text{M}^{-1}\text{s}^{-1}$ , and for pH between 5 and 5.8 the linear interpolated values are used.
- <sup>d</sup> Clifton et al. (*61*) suggested a 10~100 times higher value than Lee and Schwartz (*18*), and their results are used as the higher estimation of  $k_6$  ( $k_{6, high}$ ). Similarly, their reported pH-dependence was taken into consideration with a stepwise function. At pH 5.3 and 8.7 (which is already beyond our interested pH range),  $k_{6, high}$  are respectively  $1.24 \times 10^7$  and  $1.67 \times 10^7$  M<sup>-1</sup>s<sup>-1</sup>, and at pH 5.3~8.7 the linear interpolated values are used.

<sup>e</sup> In Fig. 2, the average rate calculated by  $k_{6, low}$  and  $k_{6, high}$  are used.

Gas-Aqueous Equilibrium for S(IV)					
Equilibrium	Constant Symbol	H <sub>298K</sub> (M atm <sup>-1</sup> ) <sup>b</sup>	–∆H <sub>298K</sub> /R (K)		
$SO_2(g) \leftrightarrow SO_2(aq)$	H <sub>SO2</sub>	1.23	3145.3		
$O_3(g) \leftrightarrow O_3(aq)$	H <sub>O3</sub>	1.1 (-2)	2536.4		
$H_2O_2(g) \leftrightarrow H_2O_2(aq)$	H <sub>H2O2</sub>	1.0 (5)	7297.1		
$NO_2(g) \leftrightarrow NO_2(aq)$	$H_{NO2}$	1.0 (-2)	2516.2		
Aqueous-phase Ionization Equilibrium <sup>c</sup>					
Equilibrium	Constant Symbol	K298K (M) <sup>b</sup>	–∆H <sub>298K</sub> /R (K)		
$SO_2 H_2O \leftrightarrow H^+ + HSO_3^-$	K <sub>s1</sub>	1.3 (-2)	1960		
$\mathrm{HSO}_3^- \leftrightarrow \mathrm{H}^+ + \mathrm{SO}_3^{2-}$	$K_{s2}$	6.6 (-8)	1500		

table S5. Constants for calculating the apparent Henry's constant  $(H^*)^a$ .

<sup>a</sup> Read 1.0 (5) as  $1.0 \times 10^5$ . Data are extracted from Chapter 7 of Ref. (12).

<sup>b</sup> The Henry's constant (H) and ionization constant (K) are both equilibrium constant in nature, and thus have the same dependence on temperature. The H at temperature T is

$$H(T) = H(T_0) \exp\left[-\frac{\Delta H_{298K}}{R} \left(\frac{1}{T} - \frac{1}{T_0}\right)\right]$$

where  $T_0 = 298$  K. The same is true for K(T).

<sup>c</sup> The effective Henry's constant (H\*) for  $HSO_3^-$ ,  $SO_3^{2-}$  and S(IV) (which is defined as sum of  $[SO_2 \cdot H_2O]$ ,  $[HSO_3^-]$  and  $[SO_3^{2-}]$ ) are respectively

$$H^*_{HSO_3^-} = H_{so_2} \frac{K_{s1}}{[H^+]}$$
$$H^*_{SO_3^{2-}} = H_{so_2} \frac{K_{s1}K_{s2}}{[H^+]^2}$$
$$H^*_{S(IV)} = H_{so_2} [1 + \frac{K_{s1}}{[H^+]} + \frac{K_{s1}K_{s2}}{[H^+]^2}]$$

where  $H_{SO2}$ ,  $K_{s1}$  and  $K_{s2}$  are corresponding values at temperature T.

Equation Name	Equation	Targeted Species	Upper Valid Limit of I (M)	Reference
Debye-Hückel	$\log a_i = -A z_i^2 \sqrt{I}$	Ions	~ 0.01	(119)
Setschenow <sup>b</sup>	$\log a_i = bI$	Neutral species	~5 <sup>b</sup>	(120)
Simple extended Debye- Hückel	$\log a_i = -Az_i^2 \frac{\sqrt{I}}{1 + \sqrt{I}}$	Ions	~ 0.1	(121)
Davies	$\log a_i = -Az_i^{\ 2} (\frac{\sqrt{I}}{1 + \sqrt{I}} - 0.3I)$	Ions	~ 0.7	(121)
Extended Debye-Hückel ("B-dot" equation)	$\log a_i = -Az_i^2 \frac{\sqrt{I}}{1 + Ba_i^\circ \sqrt{I}} + bI$	Ions and neutral species	~1	(122)
Pitzer	$\log a_i = \log a_i^{LR} + \log a_i^{MR} + \log a_i^{SR c}$	Ions and neutral species	>1	(123, 124)

table S6. Summary of suggested activity coefficient (*a*)-ionic strength (*I*) dependence<sup>b</sup>.

<sup>a</sup> Notation of the common symbols:  $a_i$  and  $z_i$  are respectively the activity coefficient and charge number of ion *i*, *I* is the ionic strength, others are constants.

<sup>b</sup> The Setschenow equation was at first intended to describe the pattern of changed solubility with electrolyte concentration, and Debye and McAulay (*120*) provided an theoretical explanation for this equation. The valid range may change with solute and solvent; but it was generally valid when I < 5 M (e.g., (*125–127*)).

<sup>c</sup> The LR, MR and SR refer to Long-, Middle- and Short-Range effects, respectively. See the most recently studies in Rusumdar (128).

Oxidants	Rate Expression, - d[S(IV)]/dt	Relation of k/k <sup>I=0</sup> with Ionic Strength I <sup>a</sup>	<b>Condition Notes</b>	Reference
H <sub>2</sub> O <sub>2</sub>	k1[H <sup>+</sup> ][HSO3 <sup>-</sup> ][H2O2(aq)]	$\log \frac{k}{k^{I=0}} = -\frac{2A\sqrt{I}}{1+B\sqrt{I}} + 2\beta I$ A=0.509 M <sup>-0.5</sup> , B=0.17 M <sup>-0.5</sup> , β=0.18 M <sup>-1</sup>	$I_{\text{max}} = 5 \text{ M};$ <i>a-I</i> Type: Pitzer	(52) and references therein
O3 <sup>b</sup>	$k_2[S(IV)][O_3(aq)];$ $k_2 = k_{2A}+k_{2B}[OH^-]$	$\log \frac{k}{k^{I=0}} = b_2 (\frac{\sqrt{I}}{1+\sqrt{I}} - 0.3I)$ b_2=1.0 (ranges 0.7 to 1.3) <sup>b</sup>	$I_{\text{max}} = 0.4 \text{ M};$ pH = 3~6.2 <i>a-I</i> Type: Davies	(53)
03	$k_3[HSO_3^-][O_3(aq)][H^+]^{-0.5}$	$\frac{k}{k^{I=0}} = 1 + b_3 I$ b_3=1.94 M <sup>-1</sup> (ranges 1.34 to 6.13) <sup>b</sup>	$I_{\rm max} = 1.2 { m M}$	(54)
TMI+O <sub>2</sub> <sup>c</sup>	$k_4[S(IV)][TMI][H^+]^{-1};$ TMI = Fe(III) or Mn(II)	$\log \frac{k}{k^{I=0}} = b_4 \frac{\sqrt{I}}{1 + \sqrt{I}}$ b_4 = -4 (ranges -2 to -4) <sup>c</sup>	$I_{\text{max}} = 2 \text{ M}; \text{ pH} = 3.0$ <i>a-I</i> Type: Extended Debye- Hückel	(55, 56)
NO2 <sup>d</sup>	k <sub>5</sub> [S(IV)][NO <sub>2</sub> (aq)]	$\log \frac{k}{k^{I=0}} = b_5 I$ b_5=0.5 M <sup>-1 d</sup>	<i>a-I</i> Type: Setschenow	/

table S7. Influence of ionic strength (I) on rate of aqueous sulfate-producing reactions.

<sup>a</sup> All ionic strength *I* in the expressions are with the unit of M.

<sup>b</sup> Different values of *b* are reported with changing solution chemical environments, while the general form are kept (53, 54). Here we used a medium value of  $b_2=1.0$  and  $b_3=1.94$ , respectively, to show the general pattern.

 $^{c}b_{4}$  are -2 for Fe(III) and -4 for Mn(II). In fact, this expression was insufficient to fully predict the sulfate inhibition effect. As a result, here the value of -4 are used to better represent the situation (55, 56).

<sup>d</sup> Value of  $b_5$  are not reported, but is only our theoretical prediction. The value of 0.5 was assumed to show the trend. Note that a positive value is implied in its definition.