

Supplementary Note 1: Phylogeny and taxonomy

A time-calibrated informal species-level amniote supertree, including 1046 taxa, was constructed by hand based on the most up-to-date phylogenetic analyses available for Permian-Triassic amniotes. The interrelationships shown in the supertree reflect congruence between overlapping topologies of the source phylogenetic analyses; polytomies were generated when source phylogenies showed incongruent relationships (cf. the semi-strict supertree method automatized by the program TNT; ¹). This tree is given as Supplementary Data 1. This supertree includes 891 species within the interval from the Lopingian to the end of the Early Jurassic. As terrestrial biogeographic patterns were of interest, marine taxa were omitted; however, amphibious non-crocopod archosauromorphs and volant pterosaurs known from marine strata were included in the analysis.

To avoid artificially truncating branch lengths, and so the phylogenetic distances between taxa, 155 stratigraphically older taxa were also included in order to date deeper nodes. This topology was used to produce 100 time-calibrated trees, in which polytomies were randomly resolved, utilizing the ‘timePaleoPhy’ function of the paleotree package ² in R ³. Trees were dated according to first occurrence dates with a minimum branch-length of 1 Myr. The phylogenetic biogeographic analyses were performed across all of these trees, in order to account for phylogenetic uncertainty.

The informal supertree was constructed from pre-existing phylogenies, as described below.

Parareptilia

Overall parareptile phylogeny primarily follows ⁴⁻¹¹ with “nycteroleter” relationships following ¹²⁻¹⁴. Pareiasaur taxonomy and relationships primarily follow ¹⁵, with additional input from ^{8,13,14}. The two unnamed pareiasaur species known from the d’Ikakern Formation, Argana Basin, of Morocco are placed following ¹⁶. *Sauropareia* and *Colletta* are placed as

successive outgroups of the Procolophonidae following ¹⁷. Procolophonid relationships then follow ^{4,7,8,18–20}, although conflicts between analyses and incomplete taxon sampling required the collapse of many nodes. Additional input on procolophonid taxonomy was taken from ^{21,22}. *Libognathus* was considered to lie proximate to *Leptopleuron* and *Hypsognathus* following the comparisons of ^{23,24}. *Haligonia* exhibits a greatly expanded maxillary tooth otherwise only known from *Phonodus* ²⁵; consequently it was treated as the sister-taxon to *Phonodus* here based on the comparisons of ²⁵. *Orenburgia* and *Procolina* were placed in a polytomy with the other “kapoids” *Anomoiodon* and *Kapes* following the comparisons of ²⁶.

Eureptilia

Overall eureptile topology follows ^{27,28}. Captorhinid relationships follow ²⁹. *Gecatogomphius* was considered a moradosaurine after ³⁰. Early diapsid phylogeny then follows ^{27,28}. Drepanosaurids were positioned as the sister group to coelurosauravids after ³¹; internal relationships of the clade follow ^{32,33}. *Palaeagama* was placed following ³⁴ and *Saurosteon* after ³⁵.

Lepidosauromorpha – *Paliguana* is positioned basal to [Kuehneosauridae + Lepidosauria] after ³⁶. Kuehneosaurid relationships follow ^{26,37}. Stem-lepidosaur relationships follow ^{26,35,36,38,39}. *Tikiguana* is probably Cenozoic in age ⁴⁰ and most purported stem-squamates from this interval are either very poorly phylogenetically constrained, or probable procolophonians ^{21,36,41}. This leaves *Paikasaurus* and *Bharatagama* as the only named squamates recognized from the interval of interest ^{36,38,42}.

Rhynchocephalian topology follows ^{38,43–47}, although topological conflict and incomplete taxonomic sampling limited phylogenetic resolution. *Sigmala* was considered a sphenodontid due to its similarity to taxa such as *Pelecymala* ⁴⁸. “Clevosaur” taxonomy and relationships follow ⁴⁹, although taxonomic uncertainty, the inclusion of poorly-known but

biogeographically important specimens (e.g. the Dockum clevosaur⁵⁰) and the non-monophyly of “*Clevosaurus*” resulted in most of these occurrences being accommodated in a large polytomy. Other, highly uncertain, occurrences of clevosaur-grade sphenodontians e.g.^{51,52} were not included.

Archosauromorpha – Interrelations of basal archosauromorph clades follow^{28,53,54}, although uncertainty in the position of *Prolacertoides* limits resolution in this region. Tanystropheid phylogeny follows⁵⁵, with additional taxa added to this scaffold after⁵⁶. Allokotosaurian internal relationships are after^{53,57}; rhynchosaurian topology primarily follows^{53,58} with additional information from^{59–61}. Non-archosaur archosauriform relationships then also follow⁵³, with additional reference to⁶² and *Osmolskina* positioned as a euparkeriid after⁶³. Proterosuchid taxonomy and referral of specimens follows^{53,64}.

Pseudosuchia – Phytosauria is included within Pseudosuchia after⁵³. Phytosaur taxonomy primarily follows the results of the species/specimen-level analyses of^{50,65} and the summary of⁶⁶. *Rutiodon* spp. other than *R. carolinensis* are not considered valid⁶⁶; the wastebin nature of the taxon also means that *R. sp.* records could not be accommodated. *Pseudopalatus*, *Arribasuchus*, and *Redondasaurus* were considered junior synonyms of *Machaeroprotopus* after⁶⁷. Phytosaur phylogeny primarily follows^{68,69} with *Pravusuchus* considered the sister-taxon to Pseudopalatinae after⁶⁵. “*Paleorhinus*” *parvus* is positioned as an *Angistorhinus*-grade phytosaur following the comparisons of⁶⁹. The holotype of “*Angistorhinopsis*” has been generally compared to *Nicrosaurus* e.g.⁷⁰ and is hence placed close to *Nicrosaurus* as a possible synonym.

Ornithosuchidae is placed after^{53,71}; relationships within the clade then follow⁷². The position of the Erpetosuchidae was highly labile in the results of⁷³, primarily due to the taxa *Gracilisuchus* and *Turfanosuchus*, which have since been reappraised⁷¹. Consequently,

erpetosuchids are here placed according to their position upon removal of these taxa in ⁷³, as sister to [*Revueltosaurus*+Aetosauria]. Aetosaur taxonomy follows ^{74–76} and phylogeny primarily ⁷⁶ with additional input from ⁷⁷. The problematic taxa *Acompsosaurus* and *Ebrachosaurus* were excluded.

Gracilisuchid phylogeny follows ⁷¹. Pposaurid phylogeny then follows ⁷⁸. “Prestosuchid” and ravisuchian topology follows ^{78–81} with *Youngosuchus* placed after ⁵³. However, conflict between competing topologies and differences in taxon sampling mean that resolution in this region of the tree is compromised, with “preostosuchids” forming a large polytomy. *Dagasuchus* and *Decuriasuchus* are included in this polytomy based upon the comparisons of ^{82,83}.

Crocodylomorph taxonomy follows the summary of ⁸⁴ and relationships follow ^{78,85–88}, although incomplete taxon coverage and topological conflict means that the resolution of “sphenosuchian” taxa is poor. Although the material originally described as “*Dianchungosaurus*” is chimeric, it does demonstrate some mesoeucrocodylian synapomorphies ⁸⁹ and so is retained accordingly here.

Avemetatarsalia – Avemetatarsalian relationships primarily follow ⁷⁸, with additional information from ^{79,90–92}; uncertainty in the position of *Scleromochlus* means that it was placed in a polytomy with Pterosauria and more derived taxa. Taxonomy of basal forms follows ⁹³, with the “Eagle Basin lagerpetid” assigned to *Dromomeron romeri* after ⁹⁴. Pterosaur taxonomy and phylogeny follows ⁹⁵ and comparisons in ⁹⁶. *Faxinalipterus* is considered Archosauria indet. following ⁹⁷ and so was excluded from consideration.

The “Eagle Basin silesaurid” is retained as a separate OTU following ⁹⁴. The Otis Chalk and Petrified Forest silesaurids each also preserve putative apomorphies ⁹⁸ and so are also retained. *Pseudolagosuchus* was considered a junior synonym of *Lewisuchus* after ^{92,98,99} and *Agnosphitys* was included as a silesaurid after ⁹³. Silesaurid phylogeny follows ^{92,99–101},

although accommodation of poorly-known taxa and specimens means that resolution within this clade is poor.

Nyasaosaurus was placed in a polytomy with Ornithischia and Saurischia in order to accommodate the uncertainty in its phylogenetic position⁸¹. Ornithischian relationships then follow^{102–104} with heterodontosaurid taxonomy and phylogeny after^{102,103,105,106}, although accommodation of poorly-known specimens and taxa results in very limited resolution within this clade. Dubious Triassic and Early Jurassic ornithischians based on fragmentary or dental remains have been omitted^{107,108}. Although the thyreophorans *Tatisaurus* and *Bienosaurus* are of dubious validity¹⁰⁹ they can be distinguished from *Scelidosaurus*¹⁰⁹ and so are provisionally retained here as informal taxa.

Theropod relationships follow^{110–113}. Coelophysoid relationships follow¹¹⁴, although the complicated taxonomic history of *Coelophysis* and proximate taxa¹¹⁵ means that *Coelophysis bauri*, *Coeolophysis* (= *Megapnosaurus*) *rhodesiensis*, and *Campsosaurus* have been collapsed into a polytomy. Additional theropod taxa were placed on the basis of phylogenetic results or comparisons from individual studies (*Gojirasaurus*¹¹⁶, *Lophostropheus*¹¹⁷, *Dracovenator*¹¹⁸, dilophosaurid remains from the Dharmaram Formation¹¹⁹, and an unnamed neotheropod from Poland¹²⁰). *Eshanosaurus* is tentatively considered to represent the oldest coelurosaur after¹²¹.

Sauropodomorph relationships follow^{122–136}. *Asylosaurus* was positioned in a polytomy with *Thecodontosaurus* and *Pantyraco* based upon the comparisons of¹³⁷. *Euskelosaurus* was considered invalid after^{138,139}, with most material reassigned to *Plateosaurus* following^{138–143}. *Plateosaurus longiceps* is treated as a junior synonym of *P. erlenbergiensis*¹⁴⁴; *P. gracilis* is tentatively treated as a valid taxon following¹⁴⁵. Purported *Plateosaurus* sp. records from the UK cannot be reliably diagnosed above the level of Saurischia¹⁴⁶.

Doubts about the validity of “*Gyposaurus sinensis*” means that most specimens formerly referred to this taxon were omitted from analysis. However, the two specimens for which the taxon were originally named are potentially valid and have been resolved in a basal position proximate to *Sarhsaurus*¹⁴⁷. *Lufengosaurus magnus* was considered a likely junior synonym of *L. huenei* and *Yunnanosaurus robustus* a junior synonym of *Y. huangi*¹⁴¹.

The holotype of “*Gryponyx africanus*” was considered potentially valid and placed in the position resolved by¹³⁶. *Xixiposaurus* was positioned proximate to *Mussaurus* following the results of¹⁴⁸. *Coloradisaurus* was also positioned close to these taxa in¹⁴⁸. However, this is due to features which have since been demonstrated to be erroneous^{124,125} and *Coloradisaurus* is instead considered a massospondylid after^{124–126,131,133}. *Chinshakiangosaurus* was positioned on the basis of¹⁴⁹.

Synapsida

“Pelycosaur” phylogeny follows¹⁵⁰. *Raranimus* is considered the sister taxon to all other therapsids after¹⁵¹. Biarmosuchian relationships follow^{152,153}; *Biarmosuchoides* is tentatively positioned as a basal biarmosuchian, in a polytomy with *Biarmosuchus* and more derived taxa based on the comparisons of¹⁵⁴. *Ivantosaurus* and *Eotitanosuchus* are considered junior synonyms of *Biarmosuchus tener* following¹⁵⁵. Anteosaur taxonomy follows¹⁵⁶ and tapinocephalid taxonomy^{157,158}. Dinocephalian phylogeny follows¹⁵⁹, although lack of recent interest in tapinocephalid interrelationships means that resolution within this clade is poor. The affinities of *Niaftasuchus* and *Phtinosuchus* within therapsids are unknown¹⁵⁶; consequently, they were not included.

Anomodont phylogeny follows^{160,161} with emydopoid topology following¹⁶². Additional general input on taxonomy comes from^{163,164}. *Endothiodon* taxonomy follows¹⁶⁵ in the absence of formal revision, as in¹⁶⁴. Relationships within the genus follow¹⁶⁶, with *E. tolandi*

placed as the sister taxon to all other species on the basis of the comparisons of ¹⁶⁷. *Chelydontops* is considered a junior synonym of *Brachyprosopus* after ¹⁶¹ and *Diictodon tienshanensis* a junior synonym of *D. feliceps* after ¹⁶⁸. *Abajudon* was excluded as its affinities are poorly constrained ¹⁶⁹. *Kingoria* was considered synonymous with *Dicynodontoides* after ¹⁷⁰. Taxonomy of *Dicynodon* spp. and taxa formerly considered to belong to this genus follow ^{160,162,171}; the polyphyletic nature of the classical “*Dicynodon*” means that occurrences only identified to the level of *D. sp.* were omitted. “*Aulacephalodon peavoti*” has been referred to multiple species and genera, and in addition the holotype is currently lost ¹⁷². Consequentially, it was also omitted from the dataset.

Lystrosaurus spp. taxonomy follows ^{164,173}, with OUMNH TSK 2 removed from the genus after ¹⁷¹. The affinities of “*Kannemeyeria latirostris*” are contested ¹⁷⁴: given that it provides redundant biogeographic information with respect to *K. lophorhinus* it was omitted from analysis. *Rechnisaurus* was considered endemic to India, with African occurrences instead referred to *Kannemeyeria* after ^{164,175}. *Placerias gigas* was considered synonymous with *P. hesternus* ¹⁷⁶. The phylogenetic position of *Elephantosaurus* is uncertain ¹⁶⁰ and it provides redundant biogeographic information with regards to better-constrained kannemeyeriiform taxa. As a result, it was omitted here. *Sungeodon* was placed after ¹⁷⁷.

A humerus from the Upper Muschelkalk originally compared to *Placerias* ¹⁷⁸ has since been compared to multiple other kannemeyeriiform taxa ^{179,180}. Unfortunately, the specimen is now destroyed, making assessment of it even more problematical. Given these difficulties, it was omitted from analysis.

Taxonomy and phylogeny of basal gorgonopsids follow ¹⁸¹ and those of rubidgeines follow ¹⁸². The holotype of “*Gorgonops whaitsi*” is both problematic and apparently lost ¹⁸¹. Consequently, it was excluded. *Arctognathus curvimola* was considered the only species of

Arctognathus following ¹⁸³. *Arctops* is tentatively placed in a position close to *Smilesaurus* based on the comparisons of ¹⁸².

Viatkogorgon, *Suchogorgon*, and *Pravoslavlevia* are problematic, having never been included in a numerical phylogenetic analysis. *Kamagorgon* and *Dinosaurus* are also poorly phylogenetically constrained, having been compared variously with gorgonopsians ¹⁸⁴, *Phtinosuchus* and dinocephalians ¹⁸⁵. Consequently, all of these taxa were omitted.

Therocephalian relationships primarily follow ^{186,187}. Taxonomy generally follows ¹⁶³ with updates from subsequent phylogenetic analyses and descriptions ^{187–193}. The traditional “Scalposauridae” is polyphyletic ¹⁸⁹: as a result, scalposaurid taxa which have yet to be included in quantitative phylogenetic analyses cannot be placed with any precision and so were excluded. *Megawhatsia* is placed as a whaitsid after ¹⁹⁴.

Non-mammaliaform cynodont topology primarily follows ¹⁹⁵, with additional referral to ^{14,196–200} for traversodontids and ^{201–203} for prozostrodonians. Cynodont taxa known only from dental remains e.g. ^{204,205} have proven to be phylogenetically problematic ²⁰⁵ and so were omitted from analysis here. *Cyrbasiodon* is considered a junior synonym of *Procynosuchus* after ²⁰⁶. *Procynosuchus rubidgei* is considered synonymous with *P. delharpae* after ¹⁶³. *Trirachodon* species taxonomy follows ^{163,207}. *Bolotriodon* was positioned after ²⁰⁸; *Cromptodon* was placed proximate to *Bolotriodon* and Galesauridae based upon the comparisons of ²⁰⁹. *Titanogomphodon* is considered a diademodontid following ²¹⁰.

“*Pachygenelus milleri*” is undiagnostic above the level of ?Eucynodontia indet. ⁵⁰ and so was excluded, as was the problematic “madysaurid” *Madysaurus*. *Probelesodon* and *Belesodon* were treated as junior synonyms of *Chinquodon* after ²¹¹. *Abelobasileus* was placed following ²¹² and *Dinnebitrodon* was considered to be a tritylodont after ²¹³. “Dromatheriids” are poorly known, and their phylogenetic position remains uncertain ^{214,215}, preventing their inclusion.

The same is true for most “therioherpetids”, which have not been included in quantitative phylogenetic analyses. The exceptions to this are *Therioherpeton* and *Riograndia*, which have generally not been found to form a monophyletic group¹⁹⁵. Other “therioherpetids” hence could not be accommodated. *Oligokyphus* spp. referral follows²¹⁶. *Protheriodon*, *Panchetocynodon*, and *Minicynodon* are positioned based on the results and comparisons of²⁰⁹.

The interrelationships of basal mammaliaform clades follow^{199,217–222}. *Hadrocodium* was positioned after²²³ and *Trishulotherium* after²²⁴. *Dyskritodon* is considered a triconodontid²²⁵ and *Brachyzostrodon* a megazostrodonid with the informal species of²²⁶ provisionally included. The content of the Morganucodontidae follows^{227,228} and that of Haramiyidae follows^{199,229}. *Indobaatar* is considered to represent the earliest multituberculate after²³⁰. Mammaliaform tooth taxa (e.g. “symmetrodonts”) are generally phylogenetically and taxonomically problematic, and so were mostly excluded, as with non-mammaliaform cynodonts.

Supplementary Note 2: Taxon occurrences and ages

Occurrence data for these taxa was taken primarily from the Paleobiology Database²³¹, with the addition of some occurrences from the literature (see Supplementary Data 2). Taxa were dated at stage level. They were then placed in the following time bins for analysis: Lopingian, Early Triassic (Induan and Olenekian), Anisian, Ladinian, early Late Triassic, late Late Triassic, early Early Jurassic (Hettangian, Sinemurian), and late Early Triassic (Pliensbachian, Toarcian).

The Late Triassic was not split into its constituent stages due to the disproportionately long Norian e.g.^{232–235}. Similarly, recent re-dating of the Los Colorados Formation as being early–middle Norian in age²³⁵ prevents separation of time bins according to the Ischigualastian and

Coloradian land vertebrate ages. Instead, Late Triassic occurrences were divided into two time bins: Carnian–early Norian (Lacian) and middle/late Norian (Alaunian)–Rhaetian, with the separation between these bins dated at approximately 219–213 Ma ²³². Geological units were assigned to each of these time bins as summarized in Supplementary Tables 1 and 2. However, difficulties in the dating of individual geological units and/or uncertainty in dating the base of the Alaunian ²³² meant that occurrences from the Sonsela Member of the Chinle Formation, the Trujillo Formation, the Lockatong and lower Passaic (Neshanician) formations from the Newark Basin, the Leedstown Formation from the Taylorville Basin, Late Triassic occurrences from the Fundy Basin, the Bigoudine Formation, the *Riograndia* Assemblage Zone of the Santa Maria Supersequence, the La Esquina fauna of the Los Colorados Formation, the Upper Maleri Formation, the fissure-fill deposits of the southwestern UK, and most Middle Stubensandstein occurrences could not be satisfactorily restricted to a single time bin. To reflect this uncertainty, these occurrences were included in both Late Triassic time bins for analysis.

| Area | early Late Triassic | Referral |
|---------------------------|---|--|
| Western USA | Chinle Group: Mesa Redondo Member, Bluewater Creek Formation, Blue Mesa Member. Dockum Group: Colorado City Formation, Tecovas Formation and correlates of these units. | Otischalkian and Adamanian Land Vertebrate Faunas ²³⁶ were considered to be Carnian and early Norian in age, respectively, after ²³³ . This corresponds with radiometric dating of the overlying Sonsela Formation of the Chinle Group ²³⁷ . |
| Eastern USA | Newark Supergroup: Stockton Formation, Conewagian Formation, New Oxford Formation, Cumnock Formation. Taylorsville Basin: Newfound Formation, Port Royal Formation. | The chronostratigraphic framework of the Newark Supergroup and its correlates used herein follows that of ²³⁴ . |
| Morocco | Irohalene Formation | The Irohalene Formation is dated late Carnian-early Norian on the basis of biostratigraphy ²³⁸ and palynological dating of the overlying Bigoudine Formation ^{239,240} . |
| UK | Lossiemouth Sandstone | The Lossiemouth Sandstone is considered to be late Carnian in age on the basis of biostratigraphy ⁹¹ . |
| Continental Europe | Germanic Basin: Stuttgart Formation, Weser Formation, Arnstadt Formation, Lower Stubensandstein of the Lowenstein Formation and their correlates. Poland: the Krasiejow fauna. | The Lacinian is considered to be early Norian in age, following ^{232,241-244} . The Krasiejow fauna is dated after ²⁴⁵⁻²⁴⁷ . |
| Brazil | Santa Maria Supersequence: Santa Cruz Sequence, Candelaria Sequence. | Stratigraphy and dating of the Santa Maria Supersequence follows ²⁴⁸ , with these sequences being considered to range from late Ladinian to early Norian. |
| Argentina | Ischigualasto Formation, Cacheuta Formation, the La Chilca Fauna of the Los Colorados Formation. | Stratigraphy follows ²⁴⁹ , biostratigraphy ^{100,127} and chronostratigraphy ²³⁵ . The Cacheuta Formation is dated based on correlation with the Ischigualasto Formation ^{240,250} . The La Chilca fauna is correlated with the Candelaria Sequence based on the presence of <i>Jachaleria</i> ²⁴⁰ . |
| India | Lower Maleri Formation | The Lower Maleri fauna allows correlation with the <i>Scaphonyx-Exaeretodon-Herrerasaurus</i> and <i>Hyperodapedon</i> Assemblage Zones, and so is late Carnian to early Norian in age ¹³² . |

Supplementary Table 1: Referral of stratigraphic units to the early Late Triassic time bin as used in this analysis.

For units referred to the late Late Triassic time bin see Supplementary Table 2; for those which could not be delimited to either Late Triassic time bin, and so were included in both, see Supplementary Note 2.

| Area | late Late Triassic | Referral |
|---------------------------|---|--|
| Western USA | Chinle Group: upper Petrified Forest/Painted Desert Member, Owl Rock Member, Rock Point Member. Dockum Group: Redonda Formation, Sloan Canyon Formation, Upper Trassever Formation, and correlates of these units. | The Apachean Land Vertebrate Fauna ²³⁶ was considered to be late Norian-Rhaetian in age, after ²³³ . |
| Eastern USA | Newark Supergroup: Upper Paissic Formation (Cliftonian). New Haven Formation. | Newark Supergroup chronostratigraphy follows ²³⁴ . The age of the New Haven Formation is based on the radiometric dates of ²⁵¹ . |
| Greenland | Fleming Ford Formation | The Fleming Ford Formation is considered Rhaetian based on correlation with the Newark-APTS 2010 ²³⁴ and the Steinbergkogel ²⁵² . |
| Continental Europe | Germanic Basin: Upper Stubensandstein of the Lowenstein Formation, Trossingen Formation (Knollenmergel), Exter Formation, and their correlates. Italy: Zorzino Limestone, Dolomia di Forni Formation. | The Sevatian 1 is considered to be late Norian in age, and Sevatian 2 Rhaetian in age following ^{232,241–244} . The Italian limestone successions are dated on the basis of palynofloras and conodont ages ^{253–255} . |
| Argentina | Quebrada del Barro Formation | ²⁵⁶ . |
| South Africa | Lower Elliot Formation | ²⁵⁷ . |
| India | Dharmaram Formation | The absence of rhynchosaurs and therapsids in this fauna imply a late Norian-Rhaetian age ¹³² . |

Supplementary Table 2: Referral of stratigraphic units to the late Late Triassic time bin as used in this analysis.

For units referred to the early Late Triassic time bin see Supplementary Table 1; for those which could not be delimited to either Late Triassic time bin, and so were included in both, see Supplementary Note 2.

Supplementary Note 3: Geographic regions

Geographic areas for use in studies of palaeobiogeography are usually defined on the basis of modern continental configurations e.g. ^{258–261}. However, this is potentially problematic when dealing with occurrences on a supercontinent where, for example, eastern North American and northwestern African localities were much closer to each other than to localities in southwestern North America or southern Africa, respectively. Instead, the geographic areas to be treated as nodes in the network analyses were defined on the basis of multivariate *k*-means clustering of palaeocoordinate data.

Palaeolongitudes and palaeolatitudes for 2144 Lopingian-Toarcian terrestrial tetrapod occurrences were taken from the Paleobiology Database, and binned at epoch level, so as to mitigate the effects of continental drift on results (see Supplementary Data 2). *K*-means clustering was then performed on these epoch-level datasets within R ³, varying *k* from 5-15. Ten-thousand replicates were performed for each of these analyses, with ten random starts. The performance of each was measured as the proportion of the total variance explained by the resolved clusters (the ratio of the between clusters sum of squares: total sum of squares). The best performing iteration for each value of *k* was retained for further comparison.

Comparison between results of different values of *k* was principally performed on the basis of the proportion of variance explained by each, with those scoring <98% omitted from consideration. These results are given in Supplementary Table 3, with full results of the cluster analyses given as Supplementary Data 3. Further comparison between results for different values of *k* was performed on the following criteria: the consistency of the clusters through the time interval of interest and their consistency with previously recognized biogeographic provinces e.g. ^{240,260,262}. Results for fifteen clusters were unstable between different time bins, resulting in the designation of ten clusters to use in the network biogeography analyses, as

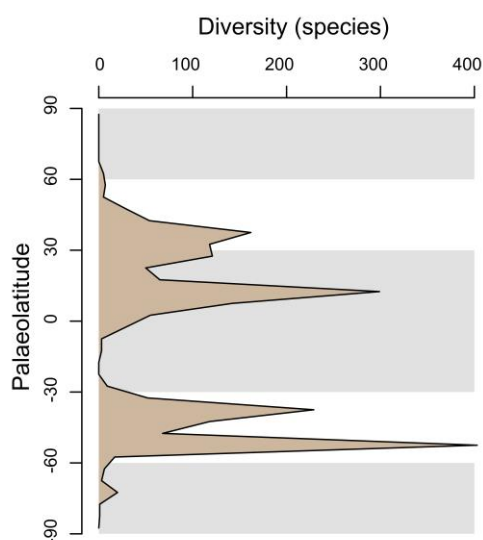
given in Fig. 1b. These then formed the basis of the bipartite taxon-locality presence/absence matrices for each time bin, given in Supplementary Data 4.

| | Number of clusters (<i>k</i>) | | | | | | | | | | |
|-----------------|---------------------------------|------|------|------|------|------|------|------|------|------|------|
| | 5 | 6 | 7 | 8 | 9 | 10 | 11 | 12 | 13 | 14 | 15 |
| late Permian | 97.3 | 97 | 97.3 | 97.4 | 98.1 | 98.7 | 98.8 | 98.8 | 98.9 | 99 | 99.1 |
| Early Triassic | 93.7 | 96.5 | 97.5 | 98 | 97.8 | 98.6 | 98.8 | 99.2 | 99.4 | 99.5 | 99.5 |
| Middle Triassic | 93 | 95.8 | 96.9 | 97.3 | 98.1 | 98.2 | 98.9 | 98.8 | 99.3 | 99.3 | 99.4 |
| Late Triassic | 91.8 | 93.6 | 94.8 | 95.6 | 96.9 | 98 | 97.6 | 97.9 | 97.7 | 97.9 | 98.7 |
| Early Jurassic | 97.6 | 98.5 | 99.2 | 99.5 | 99.5 | 99.6 | 99.6 | 99.6 | 99.7 | 99.8 | 99.8 |

Supplementary Table 3: Performance values (measured as the % total variance explained by the observed clusters) of the best performing clustering analyses for each value of *k* in each time bin. Those exceeding the 98% threshold applied within are shaded.

Supplementary Note 4: Sampling

Global sampling during this interval is uneven: in particular, there is a dearth of occurrences from low-latitude Gondwana during the late Permian, Ladinian, latest Triassic, and Early Jurassic (Supplementary Fig. 1). During the Early Triassic, at least, this paucity may have been genuine²⁶⁴, but may be masking biogeographical heterogeneity in other bins²⁶⁵. Still, the inclusion of phylogenetic information, and use of relatively broad regions, makes the analyses reported herein less vulnerable to sampling heterogeneity than similar previous studies.

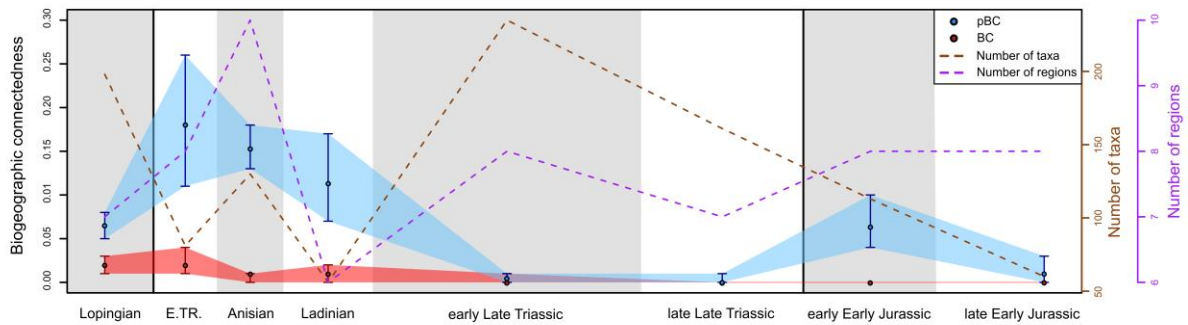


Supplementary Figure 1: Plot of the number of late Permian-Early Jurassic species sampled in this analysis against palaeolatitude.

Global pBC shows no significant correlation with either the number of taxa or the number of regions (and so, localities) sampled in each time bin (Supplementary Figs 2, 3a, b). In contrast, a significant negative correlation is observed between global pBC and the length of bins as used in the main results (Supplementary Fig. 3c). This suggests that the low values of pBC observed in the Late Triassic and Early Jurassic may be a consequence of the relative length of these bins.

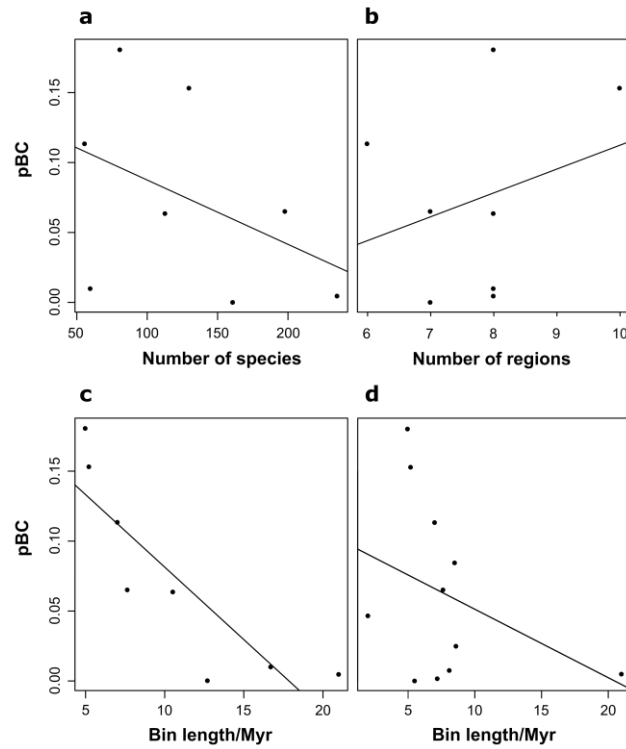
To investigate this further, the analysis was repeated with shorter-scale time bins within this interval: Carnian-early Norian, late Norian, Rhaetian, Hettangian, Sinemurian, Pliensbachian, and Toarcian (unfortunately difficulty separating late Carnian and early Norian occurrences prevented subdivision of the early Late Triassic). The same trends are resolved using these shorter time bins: low pBC during the Late Triassic, followed by a significant increase across the Triassic-Jurassic boundary after which values decline strongly in the Pliensbachian-Toarcian (Supplementary Fig. 4). Under this treatment, the relationship between pBC and bin length is no longer significant (Supplementary Fig. 3d): the late Norian and Rhaetian represent two of the shortest time bins, yet still exhibit pBC values of close to zero. Overall results

therefore appear robust to binning treatment and are not driven by bin length. As a result, the analyses herein are considered to capture valid biological signal, especially in Laurasia and temperate Gondwana.

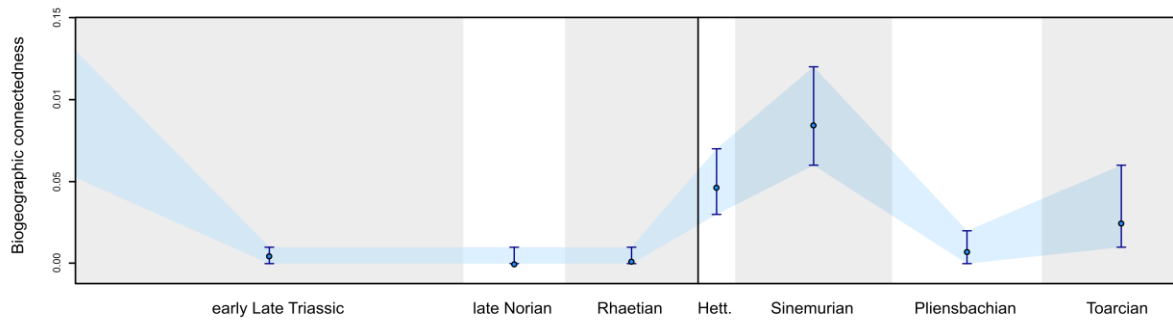


Supplementary Figure 2: Results for phylogenetic biogeographic connectedness (pBC, in blue, 95% confidence intervals calculated from jackknifing with 10,000 replicates), non-phylogenetic biogeographic connectedness (BC, in red, 95% confidence intervals calculated from jackknifing with 10,000 replicates) for the Lopingian through to the end of the Early Jurassic, plotted against the number of species and number of regions in each time bin.

However, the poor and uneven sampling of low-latitude Gondwanan localities means that results for the southern hemisphere should still be interpreted more cautiously. This is particularly true in the late Late Triassic time bin, which contains only 29 Gondwanan species, of which over 48% are known from a single area. Consequently, the potential decoupling of local pBC signals within Gondwana from global patterns identified in this time bin requires corroboration from future sampling of additional Gondwanan localities.



Supplementary Figure 3: Sampling of taxon occurrence data. a) Linear regression of phylogenetic biogeographic connectedness versus the number of species included in each time bin, showing a non-significant relationship ($p = 0.2855$, adjusted $r^2 = 0.05076$). b) Linear regression of phylogenetic biogeographic connectedness versus the number of regions in each time bin, showing a non-significant relationship ($p = 0.4924$, adjusted $r^2 = -0.07128$). c) Linear regression of phylogenetic biogeographic connectedness versus the length of each time bin, showing a significant relationship ($p = 0.0060103$, adjusted $r^2 = 0.6971$). d) Linear regression of pBC against time bin length when subdividing the Late Triassic and Early Jurassic into shorter time bins (see Supplementary Note 4, Supplementary Fig. 5), showing a non-significant relationship ($p = 0.2648$, adjusted $r^2 = 0.03979$).



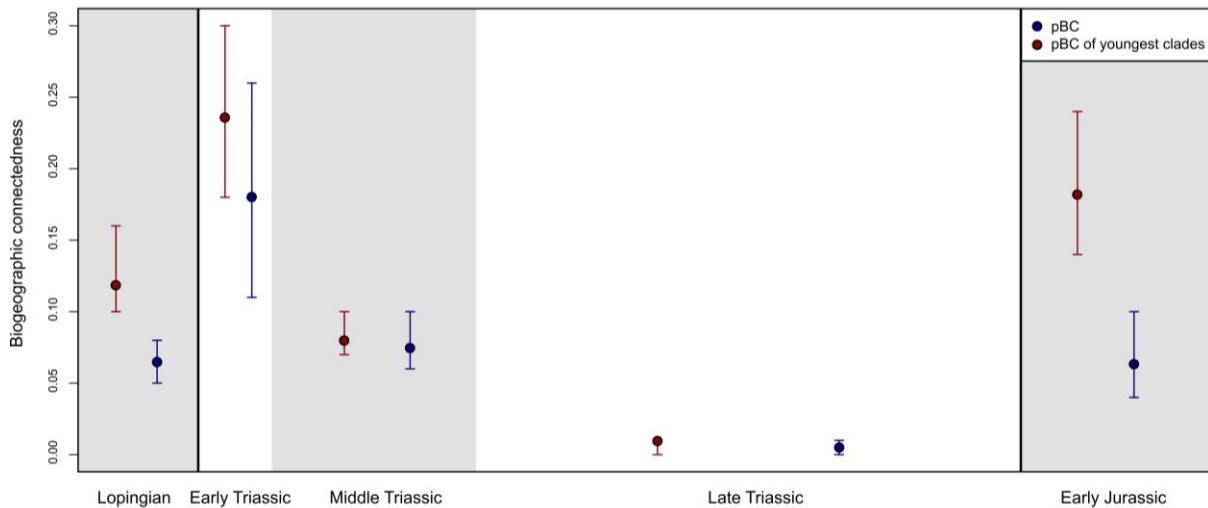
Supplementary Figure 4: Phylogenetic biogeographic connectedness results for the Late Triassic-Early Jurassic interval, employing finer time bins for the late Late Triassic and Early Jurassic. Ninety-five percent confidence intervals, calculated from jackknifing with 10,000 replicates, are given. Hett. = Hettangian.

Supplementary Note 5: Clade age sensitivity analysis

Analysis of taxonomic subsets demonstrated that the significant pBC increases observed across both the Permian-Triassic and Triassic-Jurassic boundaries were primarily driven by the radiation of novel taxa. An alternative possibility, however, is that this increase in pBC is an artefact related to a lower average clade age within these time bins. To investigate this possibility, the analysis was repeated for the entire time interval including only those clades which diverged within 2Ma above or below the lower boundary of each time bin. Unfortunately, resulting very low sample sizes in the Ladinian and latest Triassic necessitated the use of epoch-level bins for this analysis.

The results of this analysis recapitulate the overall signal seen from analysis of the complete dataset (Supplementary Fig. 5). Significant increases in pBC are still observed across both mass extinction boundaries, and there is no significant difference between pBC values derived from this subset or the total dataset in Triassic time bins. This indicates that the primary biogeographic signals observed during this interval – a decline in cosmopolitanism through the sampled interval, punctuated by increases across both the Permian-Triassic and Triassic-Jurassic boundaries – is not merely an artefact resulting from the average clade age within each time bin.

In the Lopingian time bin the pBC of these youngest clades was significantly greater than from all clades, a pattern otherwise only seen in the post-extinction time bins. The Lopingian is predated the poorly-understood end Guadalupian event, which may represent another mass extinction²⁶⁶. This offset may reflect similar biogeographic patterns accompanying turnover at this time: further analysis of a larger late Palaeozoic interval will be required to test this further.



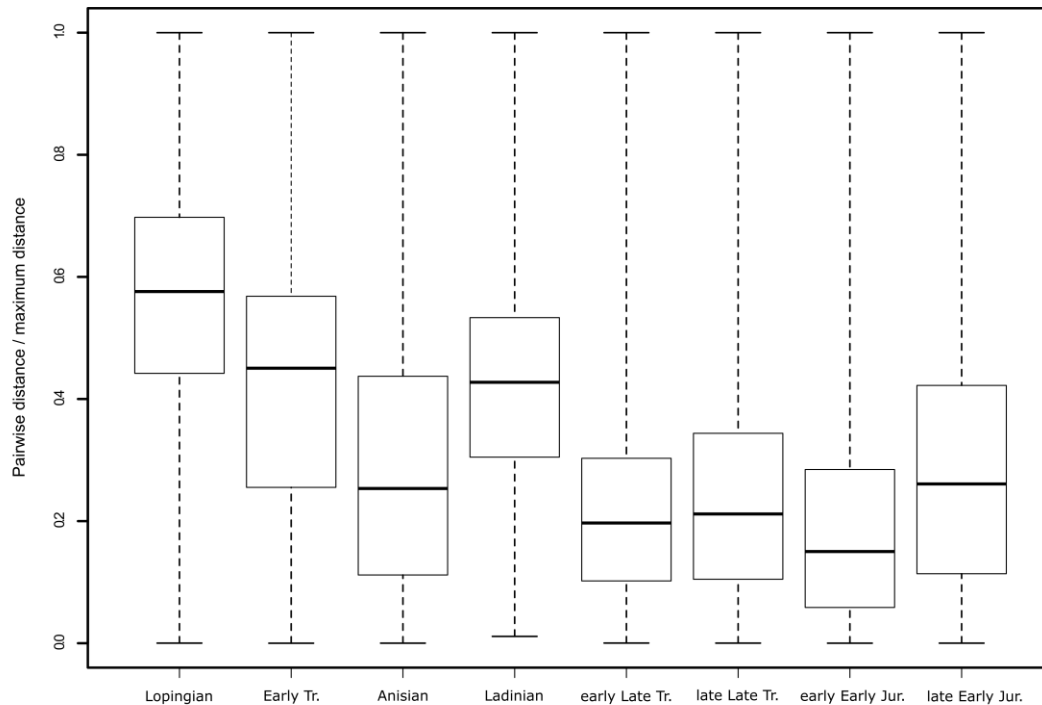
Supplementary Figure 5: Phylogenetic biogeographic connectedness results through the Lopingian-Early Jurassic, comparing results from the whole dataset (in blue) with those from a subset including only clades diverging within 2Ma either side of the lower boundary of each bin (in red). Ninety-five percent confidence intervals, calculated from jackknifing with 10,000 replicates, are given.

Supplementary Note 6: Phylogenetic distances in each time bin

It should be noted that a given value of pBC does not represent a unique solution, as the same value could be theoretically generated by a few links between closely-related taxa or more links between more moderately related taxa.

To ensure that observed pBC results are not being driven purely by differences in the average branch lengths between taxa sampled in each time bin, pairwise distances, based on branch lengths, between all tips were calculated across all input trees for each time bin, following the truncation of maximum branch lengths to the μ -value (see Supplementary Note 7). Results

indicate no significant differences between time bins (Supplementary Fig. 6), confirming that pBC results are not redundant with respect to average relatedness of sampled taxa.



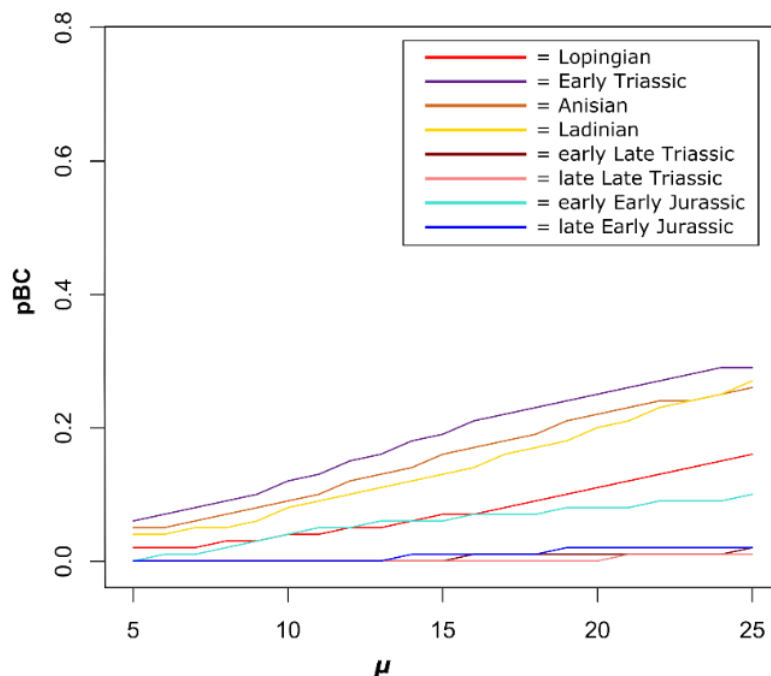
Supplementary Figure 6: Descriptive statistics of pairwise distances between all tips as measured across all 100 input trees for each time bin after truncation by a μ value of 15Ma, scaled against the maximum distance between two tips in each bin. Solid lines refer to median values, box margins to the interquartile range, and whiskers to the maximum and minimum observed values in each bin. Tr. = Triassic, Jur. = Jurassic.

Supplementary Note 7: Phylogenetic network biogeography μ sensitivity analyses

Analysis of a simulated null (stochastically generated) dataset indicated a systematic bias towards increasing phylogenetic biogeographic connectedness (pBC) through time. This is due to the increasing distance from a persistent root to the tips through time, which results in phylogenetic branch lengths between nearest relative terminal taxa becoming proportionately shorter. This was mitigated through the introduction of a constant, μ , which collapses all branches below a fixed “depth” such that root age is equal to μ million years from the tips.

The introduction of this constant also alleviates problems of temporal superimposition of biogeographic signals which may otherwise occur. It means that pBC results reported for each time bin reflect patterns generated by biogeographic processes in the preceding μ million years, preventing these recent biogeographic signals of interested from being swamped by those from deeper time intervals.

Sensitivity analyses varying the value of μ from 1–25 were performed on the Lopingian-Toarcian taxon-region matrices (Supplementary Fig. 7). The results were used to choose a μ value of 15 for further analysis, for which final results are presented in the main text. Note that changes in the value of μ make no difference to the relative pBC seen in consecutive time bins, and so the increases observed across both the Permian-Triassic and Triassic-Jurassic mass extinction boundaries are not sensitive to different values of μ . A decrease from high values of pBC in the Middle Triassic to very low values in the Late Triassic is also robust to the value of μ chosen.



Supplementary Figure 7: Results of sensitivity analyses on varying the value of μ from 5-25 in the calculation of phylogenetic biogeographic connectedness (pBC).

Once the value of μ had been chosen for further analysis, pBC was computed for each time bin. Confidence intervals for each bin were then calculated through jackknifing of occurrence records, with 10,000 replicates. This method has been made available as the “BC” function within the R package *dispeRse* (²⁶³, github.com/laurasoul/dispeRse). Additional options include bootstrapping and jackknifing of occurrence records, permutation of random trees, and measuring phylogenetic proximity by counting nodes as opposed to measuring branch lengths. An example script for pBC analyses is given in Supplementary Data 5, and a more comprehensive script covering the analyses performed in this study as Supporting Data 6.

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