# **Title: Coral calcification mechanisms facilitate adaptive responses to ocean acidification**

Running head: Coral resistance to ocean acidification

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#### **Supplementary Methods**

#### *Experimental protocol*

Carbon dioxide gas and  $CO<sub>2</sub>$ -free air were mixed using a custom system to achieve the desired  $pH<sub>T</sub>$  levels, whereas temperature was controlled using heaters and chillers  $(\pm 0.5^{\circ}C)$ . Lighting was provided via metal halide lights on a 12 hour light cycle with maximum irradiances of 400-500 µmol photons  $m^{-2}$  s<sup>-1</sup>, resulting in a daily integral light flux of 8-10 mol photons  $m^{-2} d^{-1}$ . Aquarium pumps provided water motion and mixing in each tank, resulting in flow speeds of  $\sim$ 5-15 cm s<sup>-1</sup> near the corals. Corals were fed weekly using ~1 g dry weight of newly hatched *Artemia* nauplii or Reef-Roids coral food [1]. At the end of the experiment, corals were frozen at -80°C and subsequently stored at -20°C until processing for geochemical analyses.

#### *Environmental monitoring*

Temperature was measured in each aquarium about 5-6 days per week at various times of day using a thermometer accurate to  $\pm 0.05^{\circ}$ C. Salinity was measured 1-2 times per week, while pH was measured in the morning and evening on a given day, typically 3 days per week. Salinity was measured with a YSI conductivity meter, and pH was determined spectrophotometrically using m-cresol purple [2]. Total alkalinity (TA) was measured ~weekly using a modified Gran titration, and the accuracy of titrations was verified with Certified Reference Materials obtained from Andrew Dickson [2]. TA and salinity sampling were uneven across the experiment, so means and errors were temporally weighted. To better characterise the diel cycle of chemistry variation, pH and temperature were also measured in

each aquarium every 3 hr for a 24 hr period (TA and salinity were measured every 6 hr) about midway through the experiment. This 24 hr sampling showed that diel pH variation was quasi-sinusoidal and that the morning and evening pH sampling scheme described above provided robust estimates of daily minimum, mean, and maximum pH. The daily pH range averaged 0.22±0.01 units, with the variation about evenly distributed around the mean. TA tended to increase by  $\sim$ 10-40  $\mu$ eq kg<sup>-1</sup> at night in the aquariums as compared to daytime values, likely due to lower night time calcification rates by the corals. TA averaged  $2212\pm6$ μeq kg<sup>-1</sup> in the incoming seawater. Median daytime values for pCO<sub>2</sub> and  $\Omega_{\text{drag}}$  were calculated using CO2SYS [3].

#### *Geochemical analyses*

Prior to shipping the corals from HIMB to the University of Western Australia (UWA), corals were soaked in a 1:1 solution of household bleach and deionized (DI) water for 48 hours to remove the tissue and organics within the skeleton. Skeletons were then rinsed extensively with DI water and air-dried. At UWA, the rinsing process was repeated using a sonic bath filled with DI water. The uppermost layer of the dried branch tips was then gently shaved with a diamond-tipped Dremel tool [4, 5]. Only corals that met all of the following strict criteria were sampled: (1) no partial mortality, infestation by the *Porites*-eating nudibranch *Phestilla* spp. or visible signs of bleaching occurred at any point during the experiment, (2) the stain line was clearly visible, (3) new growth beyond the stain line occurred primarily along the branch tips, as boron isotopes can differ between branch tips and other locations [6], and (4) net growth during the second temperature phase was substantially more than the material needed for boron isotope and trace element analyses (i.e., 10 mg). The only treatment that did not fulfil all four criteria was the 25.3°C/pH 7.71 treatment for *P. compressa* corals from Waimānalo Bay, which did not have net positive growth during either

the first or second temperature phase of the experiment. However, they showed extensive apical growth beyond the stain line (5-10 mm), suggesting that growth along the tips occurred throughout the experiment, while other parts of the coral experienced dissolution. These corals were therefore nevertheless sampled for geochemical analyses.

Boron isotopes and trace elements were analysed following the method of McCulloch et al. [7]. Sample weighing, chemical dissolution and boron extraction procedures were undertaken in the metal-free hepa-filtered (ISO 7) clean room complex at the Advanced Geochemical Facility for Indian Ocean Research (AGFIOR) at UWA. To ensure complete removal of organic matter, powders were cleaned with 6.25% NaClO as described previously [5]. This cleaning procedure does not influence boron isotopes or trace elements [8]. After dissolution in  $0.51$  N HNO<sub>3</sub>, an aliquot of the resulting solution diluted to 100 ppm Ca was used for analysis of B/Ca. The remaining undiluted solution was used for boron extraction using combined cation-anion exchange columns. The extracted boron was then analysed using a NU Plasma II Multi-Collector Inductively Coupled Plasma Mass Spectrometer (MC-ICP-MS; NU Instruments) at UWA, with sample measurements bracketed by an in-house standard (~19.8‰) and blank measurements. The boron isotopic composition of the skeleton  $(\delta^{11}B)$  is reported as the per mil deviation of the stable isotopes  $^{11}B$ :<sup>10</sup>B relative to NIST SRM-951 boric acid standard. All samples were analysed in duplicate. The overall precision of this method is  $\sim \pm 0.3\%$  (2sd). B/Ca was analysed on an X-Series 2 Quadrupole Inductively Coupled Plasma Mass Spectrometer (Q-ICPMS; Thermo Fisher Scientific) at UWA using the standard Xt interface and the plasma screen fitted.

### *Carbonate chemistry of the calcifying fluid*

Boron exists in seawater as two different species, boric acid  $(B(OH)_3)$  and borate  $(B(OH)\overline{4})$ . Importantly, the two species have a distinct isotopic composition and their

relative abundance in seawater is pH-dependent. Since only borate is thought to be incorporated into coral aragonite [9, 10], coral skeletal  $\delta^{11}B$  ( $\delta^{11}B_{\text{carb}}$ ) reflects the internal pH inside the calcifying fluid (pH<sub>cf</sub>) rather than seawater pH [11, 12]. pH<sub>cf</sub> was calculated using the following equation [13]:

$$
pH_{cf} = pK_B - \log \left[ \frac{(\delta^{11} B_{sw} - \delta^{11} B_{carb})}{(\alpha_{(B_3 - B_4)} \delta^{11} B_{carb} - \delta^{11} B_{sw} + 1000(\alpha_{(B_3 - B_4)} - 1))} \right]
$$
(1)

where  $pK_B$  is the dissociation constant of boric acid, adjusted to the respective treatment seawater salinity and temperature [14],  $\delta^{11}B_{sw}$  is the boron isotopic value of seawater, with a value of 39.61‰ [15], and  $\alpha_{(B3-B4)}$  is the fractionation factor, with a value of 1.0272 [16]. It was assumed that  $\delta^{11}$ B of the calcifying fluid has the same isotopic composition as seawater, since seawater is the ultimate source of boron at the site of calcification.

Recently, it has also been shown that the coral skeletal B/Ca concentration can be used to constrain the dissolved inorganic carbon (DIC) concentration of the coral calcifying fluid ( $\text{DIC}_{cf}$ ) [17, 18]. This is because the partitioning of borate versus carbonate into aragonite appears to be sensitive to  $pH_{cf}$  [12, 17, 19]. We therefore estimated the carbonate ion concentration within the calcifying fluid ( $[CO_3^{2-}]_{cf}$ ) from measurements of both pH<sub>cf</sub> and coral skeletal B/Ca using the following equation [18]:

$$
\left[\text{CO}_3^{2}\right]_{\text{cf}} = K_p \times \left[\text{B(OH)}_4\right]_{\text{cf}} / \left(\text{B/Ca}\right)_{\text{CaCO}_3}\tag{2}
$$

where

$$
K_{\rm D} = 0.00297 \exp(-0.0202 \, [\text{H}^+]_{\text{cf}}) \tag{3}
$$

Finally, the DIC<sub>cf</sub> and the aragonite saturation state within the calcifying fluid ( $\Omega_{cf}$ ) were calculated from the estimated values for pH<sub>cf</sub> and  $[CO_3^{2-}]_{cf}$ , and the percent fraction of DIC in the calcifying fluid that is present as carbonate. Biological DIC-upregulation within the calcifying fluid was calculated as the ratio of  $DIC_{cf}$  and seawater  $DIC (DIC_{cf}/DIC_{sw})$ .

For typical calcifying fluid pH<sub>cf</sub> values of  $\sim$ 8.3 to $\sim$  8.4,  $K_D \sim 0.0027$ , an order of magnitude higher than the previous estimate of Allison et al. [19]. This difference in  $K<sub>D</sub>$ values is important because our calculated  $K_D$ , based on the experimental determination of Holcomb et al. [17], is now compatible with direct substitution of  $CO_3^2$  ions with B(OH)<sub>4</sub> ions during the formation of aragonite, obviating the need for ad hoc scenarios of  $HCO<sub>3</sub><sup>2</sup>$ substitution [19]. We also assumed that the total boron and calcium concentrations of the calcifying fluid are equal to their concentrations in seawater and solely a function of seawater salinity.

## **Supplementary Tables**

## Table S1. Sample sizes per treatment for each response variable.  $cf = \text{calcifying fluid}$ ,

 $\Delta pH = biological pH-upregulation, DIC = dissolved in organic carbon, sw = seawater, [carb]$  $=$  carbonate ion concentration,  $\Omega =$  aragonite saturation state, calc. geochem corals  $=$ calcification rates of all corals used for geochemical analyses, calc. all corals = calcification rate of all corals in the experiment.



## **Table S2. Environmental conditions during the second, 9-week temperature phase of**

**the experiment.** Treatment (treatm.), salinity (sal; psu), temperature (temp;  $^{\circ}C$ ), pH<sub>T</sub>, total alkalinity (TA; µeq  $kg^{-1}$ ), carbonate ion concentration ( $CO_3^2$ ; µmol kg<sup>-1</sup>), dissolved inorganic carbon concentration (DIC;  $\mu$ mol kg<sup>-1</sup>) and aragonite saturation state ( $\Omega_{\text{arag}}$ ). Measured values reported as overall mean values ±SEM. Sample size (n) as indicated.



**Table S3. Results from generalized linear mixed models to test for the effects of species (sp.), site, pH and temperature (temp.) on boron isotopes (δ <sup>11</sup>B), calcifying fluid pH (pHcf) and biological pH-upregulation (ΔpH) of** *Montipora capitata* **(MC) and** *Porites compressa* **(PC).** Post hoc Tukey tests results are given when main effects (but no interaction terms) were significant. Effects with *p*-values  $\leq$ 0.05 are highlighted in bold. Num df = numerator degrees of freedom, den df = denominator degrees of freedom. Var. = variable, KB = Kāne'ohe Bay, Wai = Waimānalo Bay.





**Table S4. Results from generalized linear mixed models to test for the effects of species (sp.), site, pH and temperature (temp.) on B/Ca ratios, estimated calcifying fluid DIC (DICcf) and biological DIC-upregulation (DICcf/DICsw) of** *Montipora capitata* **(MC) and**  *Porites compressa* (PC). Post hoc Tukey tests results are given when main effects (but no interaction terms) were significant. Effects with  $p$ -values  $\leq 0.05$  are highlighted in bold. Num  $df$  = numerator degrees of freedom, den df = denominator degrees of freedom. Var. = variable, KB = Kāne'ohe Bay, Wai = Waimānalo Bay.





**Table S5. Results from generalized linear mixed models to test for the effects of species (sp.), site, pH and temperature (temp.) on estimated calcifying fluid carbonate ion**  concentration (carb<sub>cf</sub>), aragonite saturation state ( $\Omega_{cf}$ ) and calcification rate (calc.) of *Montipora capitata* **(MC) and** *Porites compressa* **(PC).** Post hoc Tukey tests results are given when main effects (but no interaction terms) were significant. Effects with *p*-values  $\leq$ 0.05 are highlighted in bold. Num df = numerator degrees of freedom, den df = denominator degrees of freedom. Var. = variable, KB = Kāne'ohe Bay, Wai = Waimānalo Bay.





\*Calcification rates of all corals in the experiment, not just the ones for which geochemical analyses were conducted.

#### **Supplementary Figures**



**Supplementary Figure S1. Correlation between (a) skeletal boron isotopes (δ <sup>11</sup>B) and (b) calcifying fluid pH (pHcf) and seawater pH (pHsw), respectively, for** *Montipora capitata* **(MC) and** *Porites compressa* **(PC) from Kāne'ohe Bay (KBay) and Waimānalo Bay (Waim), Hawaiʻi.** Results from linear regression analysis are indicated. In panel b, error bars ( $\pm 1$  SE) are given for pH<sub>cf</sub> averages. For seawater pH, error bars were typically  $\pm 0.01$ (SE) and are therefore not shown (see Table S2).



Supplementary Figure S2: Correlation between calcifying fluid pH (pH<sub>cf</sub>) and dissolved inorganic carbon concentration **(DIC<sub>cf</sub>)** in *Montipora capitata* **(MC)** and *Porites compressa* **(PC) from Kāne'ohe Bay (KBay) and Waimānalo Bay (Wai), Hawaiʻi.** Results from linear regression analysis are indicated. When the outliers are removed for MC Wai and MC KBay, the following two equations were obtained:  $y = -0.0001x + 8.98$ ,  $R^2 = 0.15$  (MC Wai) and  $y = -0.0002x + 9.17$ ,  $R^2 = 0.60$  (MC Kbay).

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