## **Fluctuating Radiocarbon Offsets Observed in the Southern Levant and Implications for Archaeological Chronology Debates**

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#### **Supporting Information Appendix (SI Appendix)**



#### **1. Dendrochronology**

Sections and cores of *Juniperus phoenicea* were sampled from timbers from former late Ottoman period structures of end  $19^{th}/e$ early  $20^{th}$  century AD date incorporated into the historic Hotel Taybet Zaman (site code TZM) in Taybeh Village, southern Jordan (ref. 1 pp.33-34) (for brief discussion of the Taybet Zaman site, see below). Samples were sanded to a high polish for examination and measured to the nearest 0.001 mm on a measuring table under a stereomicroscope. Analysis and cross-dating between the samples followed standard dendrochronological methods (2-4). We employed the Tellervo TRiDaS v.1.3.0 dendrochronological analysis package for measurement (5), and Corina v.1.1 (6) and COFECHA (7) for crossdating and chronology building. The set of 10 tree-ring series comprising the TZM chronology is shown with their chronological placement and summary of crossdating statistics (calculated in Corina) in Fig. S1. The TZM chronology was crossdated and placed into absolute time from AD1610-1940, against an absolutely-dated *Juniperus phoenicea* reference chronology built from living trees from south Jordan (8, 9), which is available from the International Tree Ring Databank (ITRDB):

<https://www1.ncdc.noaa.gov/pub/data/paleo/treering/chronologies/asia/jord001.crn> with the individual raw data available from:

[https://www1.ncdc.noaa.gov/pub/data/paleo/treering/measurements/asia/jord001.rwl.](https://www1.ncdc.noaa.gov/pub/data/paleo/treering/measurements/asia/jord001.rwl) A summary of the crossdating statistics comparing the TZM site chronology with the reference chronology are shown in Fig. S2. The high internal correlation and visual and statistical similarities between TZM and the juniper reference chronology indicate that TZM is a robust crossdated chronology securely placed in calendar time. The tree-ring measurements for the 10 TZM samples are listed in Table S1 and have been submitted to the International Tree Ring Databank [\(http://www.ncdc.noaa.gov/data-access/paleoclimatology-data/datasets/tree-ring\)](http://www.ncdc.noaa.gov/data-access/paleoclimatology-data/datasets/tree-ring).

The tree-ring sequences selected for  ${}^{14}$ C measurement were dissected with a steel blade under a binocular microscope. Sequences typically include 5 tree-rings (years) but in a few cases where

the tree-rings were very narrow a longer sequence (larger number of tree-rings) was dated. The sequences and their corresponding  ${}^{14}C$  data are listed in Table S2.

A second set of 14C data are also available on a *Juniperus phoenicea* sample (BADG-1) found at a different locality ~19 kilometers away in southern Jordan (the locus area was code named BADG and is about 4km NNW of Al-Bayda) ~ $30^{\circ}25'18''N$ ,  $35^{\circ}26'58''E$  (Fig.1A). This sample came from a tree stump that had been cut some years prior to our visit to the location. Two other samples from this locus offer dated tree-ring sequences AD/CE 1559-1995 and 1619-2006 (BADG-4, BADG-3 respectively). The  ${}^{14}$ C dates were run on samples cut, at that time, as approximately a time-series of consecutive 5-ring/year samples from BADG-1. In reality, the dendrochronology was not at all secure when this work was carried out, and a number of (narrow, partial and locally absent) rings were not then recognized. Thus the series is *not* knownage (contrast TZM), rather a relative ordered sequence, and, in particular, the real age span of the dated sequence is substantially longer than originally thought by, overall, ~50 years. The gaps between the mid-point of each sample dated should have been ~5 years, but in reality this is a minimum gap, and in a number of instances the gaps are in reality rather longer. The ordered sequence and the  $^{14}$ C dates on the constituent samples are listed in Table S3. We do not refer to tree-rings for the reasons just outlined for the BADG-1 samples, rather we refer to the sample bag numbers from 1 (oldest, the innermost set of rings sampled) onwards. Hence BADG Bag 1, Bag 2 … etc., below.

### **2. Radiocarbon**

*TZM*. The <sup>14</sup>C measurements on the *Juniperus phoenicea* samples were obtained via the <sup>14</sup>C Accelerator Mass Spectrometry (AMS) method at the Arizona Accelerator Mass Spectrometry Laboratory, University of Arizona (AA), and the Oxford Radiocarbon Accelerator Unit, University of Oxford (OxA). Where both laboratories ran samples of the same calendar (treering) age, this material came from identical tree-ring samples. The tree-rings were prepared via standard ABA with a bleaching step (sodium hypochlorite) to holocellulose at Arizona in a modified Jayme-Wise process as shown to achieve good pretreatment characteristics (10-12), and either as ABA plus bleach samples or as α-cellulose samples, with or without a solvent extraction step, at Oxford (the different pretreatment regimes are indicated for the Oxford samples in Table S2) (13, 14) and then dated following standard procedures (10, 11, 15). Where the identical tree-rings were pretreated by different approaches at Oxford, especially TZM-4 AD1818-1822 and TZM-3 AD1888-1892, the results are indistinguishable (Table S2) – with the exception of OxA-29735, which is an outlier  $(14)$ C age is too recent), when compared with the three other measurements on this same sample (and OxA-29735 is not employed in the analyses). The general similarities among the differently pretreated Oxford data, and between the Arizona and Oxford data (and see also the BADG samples below) conform to the generally observed pattern in previous studies that only small differences in 14C measurements usually derive from differing pretreatments ranging from  $ABA + bleach$  through to holocellulose to  $\alpha$ cellulose (16). The reported  ${}^{14}$ C ages are conventional radiocarbon ages expressed as  ${}^{14}$ C years Before Present (BP) employing the 5568 years Libby half-life and AD1950 as "present". The values have been corrected for isotopic fractionation using the  $\delta^{13}$ C values measured on the AMS. The quoted  $\delta^{13}$ C values in Table S2 were measured independently on a stable isotope mass spectrometer at each laboratory.

In the main text the Arizona and Oxford datasets are considered separately in comparisons versus IntCal13 (17): Fig. 2. We may also combine both data sets – thus pooling data for the identical samples from either or both laboratories where available into weighted average values (18). In all but one case the data are compatible with representing estimates of the same  ${}^{14}C$  age within 95% probability (18). The solitary exceptions are the data centered at 1677 which narrowly fail a  $\chi^2$ test (T=4.191 > 3.8 for df2). Fig. S3 shows the combined TZM data set versus IntCal13 (see also Table S4). Figure S4 shows the overall average offset in <sup>14</sup>C ages between the combined Arizona and Oxford Jordanian samples and IntCal13 tested with a neutral prior of  $0\pm20^{-14}$ C years using the Delta\_R function in OxCal (19):  $18.6 \pm 2.5$  <sup>14</sup>C years (for the OxCal runfile for this analysis, see Table S5). Of the 65 TZM data, 12 (18%) are not only offset from the IntCal mean value (Fig. S4A), but are incompatible as representing the same  $^{14}$ C age allowing for errors on both the TZM data and IntCal13 within 95% probability (18): Table S4. In every case the relevant significantly different TZM  $^{14}C$  age is older than the contemporary IntCal13 value. These are the data centered at 1627, 1717, 1722, 1727, 1750, 1770, 1885, 1890, 1895, 1900, 1905, 1910 (Table S4). Figure S4B shows the offset of the TZM data versus IntCal13 (linear interpolation) in terms of  $\Delta Z$  values using the equation

# $\Delta Z = \frac{T Z M 14 C sample-IntCall3 14 C value}{\sqrt{(TZM\sigma^2+IntCall3\sigma^2)}}$

Where *ΔZ* is the normalized difference (Z-Score), TZM 14Csample and IntCal13 14Cvalue are the respective  $^{14}C$  ages values of each TZM sample and the corresponding (linear interpolation) IntCal13<sup>14</sup>C mean values, and  $TZM\sigma$  and IntCal13 $\sigma$  are the respective uncertainties (see Table S4). 74% of the data are  $> 0$ , 25% < 0. One value = 0. The offset to older ages is clear.

As in Fig. 2, two approximate intervals stand out with most data involved offering offset  ${}^{14}C$ ages (with the  $TZM<sup>14</sup>C$  data from Jordanian juniper wood older in these periods than contemporary IntCal13<sup>14</sup>C data on Central and Northern European and North American wood), whereas at other times the picture is either mixed, or there is no substantive offset evident (Fig. S3). Fig. S5 shows the offset values from Fig. S3 with a range of smoothing curves through these data using 5, 7, 9 and 11 point adjacent averages, as well as 11 and 22 point LOESS (locally weighted scatter-plot smoother) curves. The two periods with mainly consistent larger offsets (evident in the raw data and in all smoothed curves), and noted in Fig. S3, are indicated. Fig. S6 shows the data on the TZM samples by individual tree sample employed for each  $^{14}C$  date. This confirms that the offsets observed, especially those in the two periods of greater offset noted in Figs. 2, S3, S5, are represented in more than one tree sample, and so are not the product of one unusual tree.

*BADG*. The <sup>14</sup>C measurements on the BADG samples were also run at Arizona and Oxford and on the same basis as the TZM samples described above. We consider an ordered sequence of samples from (oldest) Bag 1 to (most recent) Bag 33. We stopped the sequence at Bag 33 because, after Bag 33, the BADG samples start in several cases in both the Arizona (from Bag 36) and Oxford (from Bag 35) measurements either to return modern, post-bomb,  $^{14}C$  values in some cases, or they exhibit considerable variation in the <sup>14</sup>C values measured (from Bag 34 onwards in several cases) and with some of these  ${}^{14}C$  ages likely affected by modern  ${}^{14}\tilde{C}$ : see Table S3. Hence, rather than subjectively selecting what might appear to be the more acceptable results, we take the conservative view that from Bag 34 onwards there are issues with modern

(likely remaining resin or lignin-associated) contamination that affect a number of these dates and so it is best to exclude all these dates from analysis. Where available, we employ the OxA dates starting 32XXX, rather than 29XXX, as these represent re-runs of the same samples but reprocessed fully to  $\alpha$ -cellulose after the issue of (apparently) very mobile bomb-age  $^{14}C$  material became evident in a few of the more recent samples from the 29XXX series. We assume that remaining mobile resin-associated or remaining lignin material is the primary issue which has not been fully removed in all cases in the later part of the 29XXX series (20). Note: such a bomb-age/modern contamination issue, where and if relevant, would lead to incorrectly recent  $14C$  ages, and is thus the very opposite of the older  $14C$  age offset observed in the TZM and BADG series for some intervals in the  $17<sup>th</sup>$  to  $19<sup>th</sup>$  centuries AD/CE (as in Figs. 2A, S3, S5, S6). Hence, if we have missed any such contaminant issue, it would work against the finding of the offsets we report! Thus the fact we observe the  ${}^{14}C$  offsets in the BADG series in the data from both the Arizona and Oxford laboratories in the  $17<sup>th</sup>$ -19<sup>th</sup> centuries AD/CE indicates that these samples are not affected by bomb-age  ${}^{14}$ C. We also highlight that this bomb-age issue just discussed for the BADG samples is totally irrelevant to the TZM samples because the constituent TZM trees had been felled and incorporated into a building before the bomb-era, and thus no possible bomb-age  ${}^{14}C$  was ever incorporated into the cellulose of those samples.

When the BADG-1 samples were cut up, it was thought that there were  $\sim$  5 rings/years in each sample and thus the mid-points of each sample were approximately 5 years apart. As noted above, this was not in reality the case, and a number of rings were either not recognized at that time or were missing (locally absent). Subsequently, a crossdated tree-ring chronology has been constructed from a couple of the BADG samples, but these  $^{14}$ C BADG-1 samples had already been dissected and dated. (We note the obvious, for the benefit of readers: the history of the BADG-1 samples and its  $^{14}C$  date series reflects an original, mistaken, belief that we had a secure dendrochronology for this sample, and so the set of samples were cut and the  ${}^{14}C$  dates run as reported in Table S3. It was then realized that the dendrochronology was invalid and the other BADG samples were re-studied and a long secure dendrochronology achieved from BADG-3 and BADG-4 (AD/CE 1559-2006) that crossdates with both TZM (AD/CE 1610-1940: Figs.1B, S1) (t=6.7, r=0.35) and strongly with the previous Jordan juniper chronology  $(8, 9)$ AD/CE 1469-1995 ( $t=14.03$ ,  $r=0.56$ ) (6). The original aim or hope was, of course, for the BADG-1 dataset to offer a parallel dataset versus TZM. But this is impossible as the dataset we had acquired comprised an ordered series of samples and so 14C dates, but *not* a secure knownage time-series. Hence only the TZM data are employed in the main text. Nonetheless, as we discuss in the next paragraph, we can still employ the BADG-1 ordered sequence to consider the question of a  $^{14}$ C offset versus the IntCal13 calibration dataset, and hence partly 'rescue' this BADG-1 dataset and a past mistake, and make these data at least partially useful as strong independent evidence supporting the scale and approximate placement of the offsets reported from the TZM dataset.)

The BADG Bag 1 to Bag 33 sequence employed includes 29 comparisons of the same sample between Arizona (AA) and Oxford (OxA) and only in two cases are the data incompatible at the 95% level (18): Fig. S7. The BADG ordered sequence can be compared against IntCal13 (17) as a Sequence with gaps in OxCal (19). The question is the appropriate gap (and error on this estimate) to allow for between each sample, since we know there were errors and likely underestimates in a number of cases concerning rings recorded. We allow for a likely conservative

range of error estimates from 0-40%, 0-60% and 0-80% as allowing for the plausible range of error involved and likely including the correct estimate. Thus we use gaps of  $5\pm 2$  years,  $5\pm 3$ years and 5±4 years between samples in three different models. (To explain: the general issue with the initial work on the BADG-1 sample was a failure to recognize some, very narrow, treerings and locally absent rings – not usually over-counting tree rings – thus we regard the original ~5 years as the *minimum* gap between the dissected sample series, but, allowing for possible mistakes also with false rings or other errors, we in addition allow an error of  $\pm 2$  or  $\pm 3$  or  $\pm 4$ years on each of those minimum 5 year gaps to try to realistically accommodate all plausible errors). We employ the weighted average of the Arizona and Oxford dates for each sample where satisfactory (18) and for the two exceptions, samples Bag 1 and Bag 8, we consider these dates separately within a Phase (Bag 1, Bag 8) within the Bag 1 to Bag 33 OxCal Sequence. Since these two exceptions (Bag 1, Bag 8) both occur at times of rapidly changing atmospheric  ${}^{14}C$ levels (steep slopes on the  $^{14}$ C calibration curve), it may be that in one or both of these cases not exactly the same tree-rings were in fact dated, for example with a slight bias to one or a couple of older and/or more recent tree-rings between the samples and this might have exaggerated the different dates obtained. We tested for the existence, or not, of an average  ${}^{14}C$  offset between these dates on the BADG-1 Jordanian juniper ordered sequence and those from the data comprising the IntCal13 calibration dataset (17) (from wood from Central and Northern Europe and North America) employing a Delta\_R query in OxCal (19) with the IntCal13 curve resolution set at 5 years with a neutral prior of  $0\pm 20$  (we employ a 5 year curve resolution in this case, versus the 1-year in the main text, because of the much more loosely constrained data in the BADG-1 case in order to avoid a very noisy, if nonetheless overall similar/compatible, outcome with a 1-year curve resolution): Figs. S8-S13. The OxCal runfile is listed in Table S6. We note that we report examples of models which either achieved single (non-ambiguous) 68.2% probability ranges or we report the clearly most likely sub-range (e.g. >45% of total probability from the most likely 68.2% range in some cases) – since various runs offered alternative ranges (with the approximate range indicated in Figs. S8-S10 nonetheless including the most likely of the multiple ranges calculated in any such cases). Although there is more noise in the analysis as the potential error on the gap increases, from  $\pm 2$  to  $\pm 4$  years (40%-80%), the likely best fit remains very similar: Figs  $S_8-S_10$ . We therefore make further comments based on the  $5\pm 2$  years gap model – which we also believe likely best represents the real situation. The data exhibit an offset with a most likely 68.2% probability range of (rounded)  $\sim$ 26 $\pm$ 9 years (Fig. S11). If a neutral prior of 0 $\pm$ 10 is employed the offset reported (most likely 68.2% range) is ~23 $\pm$ 5<sup>14</sup>C years: Fig. S14. This is a similar finding to the TZM analysis in the main text and offers some independent confirmation of the  ${}^{14}C$  offsets reported from a different tree sample. The best placement of the BADG-1 dataset against IntCal13 (Fig. S8) shows that – as in the case of the TZM data (Fig. 2) – a number of the BADG  $^{14}C$  ages are older than the corresponding IntCal13 values. We then re-ran the BADG Bags 1-33 Sequence in Fig. S8 with the Delta\_R values of both 26 $\pm$ 9<sup>14</sup>C years and 23 $\pm$ 5<sup>14</sup>C years and with the calibration curve (IntCal13) resolution set at 5 years. The models now show a reasonably good fit, returning respective satisfactory or nearsatisfactory OxCal Amodel values of ~71.8 and ~49.1 and Aoverall values of ~62.9 and ~48.4 and the posterior densities for the calculated offsets, in dark gray, now correspond well with the assumed values, in light gray, confirming in reverse the existence of the offset at, on average, the stated level (Figs. S15, S16). We also show the  $26\pm9$  <sup>14</sup>C years case re-run with calibration curve resolution set at 1 year for comparison – it is very similar (Fig. S17). This ordered sequence of BADG data from a different tree and a different site ~19 kilometers away from TZM in southern

Jordan, although not known-age, nonetheless clearly provides valuable (if imperfect) independent confirmation of the time varying 14C offsets reported in the main text from the TZM samples (Figs. 2, S3). It is evident from Figs. S8-S10 that these offsets vary over time, and, comparing Figs. S8-S10 with Figs. 2, S3 and S5, we may observe that the periods of larger apparent offsets are similar in both the TZM and BADG series.

Quality Controls. A routine known-age sample testing program is run at the Oxford Laboratory to assess on-going dating accuracy (previous details in refs. 14, 21, 22). During the period when the main set of samples dated in this project were run at Oxford (2014-03-14 to 2016-11-04; dates OxA-29389 to OxA-34708 – four dates were run subsequently: OxA-35658, 35659, 36013, 36014), a total of 312 measurements were made on samples of known age wood and these indicate that 97.4% of dates were within  $2\sigma$  with a systematic bias of 2.6 $\pm$ 1.6 yr. As noted in the main text, we may also compare published data from Oxford on other known-age tree rings in the period AD/CE 1160 to 1660 (23, 24). These indicate a negligible offset ( $\mu \pm \sigma$ ) from 48 data points of  $3.38 \pm 4.96$  <sup>14</sup>C years versus IntCal13 (Fig. 3B, 3C), highlighting the real and substantial nature of the offset observed by the Oxford laboratory for the Jordanian juniper. A set of Oxford measurements on known age German Oak samples as employed in the IntCal13 dataset (17) offers another comparison (25). These data match the IntCal13 curve closely: Fig. S18. The mean offset testing with a neutral Delta\_R of  $0\pm 10^{-14}C$  years in OxCal (19) is effectively zero at  $0.7\pm7.2$  <sup>14</sup>C years (Fig. S19). The comparisons in Figs. 3B, 3C, S18 and S19 and other studies (26, 27) provide good grounds to regard the Oxford AMS  $^{14}C$  data as compatible with other laboratories and with the IntCal13 dataset and the modelled IntCal13 calibration curve (and hence there is no substantive laboratory offset). Arizona continuously monitors accuracy by comparison of oxalic-I and oxalic-II during every run, and uses these data to increase the quoted error where needed. As noted in the text, recent high-precision interlaboratory work demonstrates the accuracy and precision of the Arizona laboratory within stated error estimates (28, 29). Details of the calculations at Arizona are given in previous statements (30, 31).

Tel Rehov and Khirbat en-Nahas dates. The Tel Rehov  ${}^{14}C$  dates and calendar age modelling reported in Fig. 4 were taken from the  $^{14}$ C ages listed for the 11 archaeological elements shown in Fig. 2 of ref. 32. They were considered as a Sequence in OxCal  $4.3$  (19) with runs made with the <sup>14</sup>C data as published with IntCal13 (17) with curve resolution set at 5 years and then with an adjustment for a Delta\_R value of  $19\pm5$   $^{14}$ C years, or use of the approximate Jordan juniper modified calibration curve (based on the patterns of offsets in Figs. S2, S3, S5) as shown in Fig. 4A: see Table S7. The published Khirbat en-Nahas set of  ${}^{14}C$  data and Bayesian chronological model (33) were re-run comparing the calendar age ranges obtained from IntCal13 (17) with curve resolution set at 5 years versus those obtained if the <sup>14</sup>C values were adjusted by  $19\pm5$  <sup>14</sup>C years, or if the approximate Jordanian juniper (JJ) modified curve is used: see Table S8.

**3. Red Sea coral dataset and Northern Hemisphere extra-tropical temperature records** The Red Sea coral  $\delta^{18}$ O dataset (34) in Figure 2B is employed from: [https://www1.ncdc.noaa.gov/pub/data/paleo/coral/red\\_sea/felis2000\\_noaa.txt](https://www1.ncdc.noaa.gov/pub/data/paleo/coral/red_sea/felis2000_noaa.txt)

In the main text we cited one recent discussion and analysis for the reconstruction of extratropical Northern Hemisphere temperature (35), which shows, in general: (i) a warming trend

from around and in the half century or so following the peak Maunder Minimum period (the peak Maunder Minimum is dated ~AD1645-1700 (36)) – we observe a <sup>14</sup>C offset period ~1685-1762 (Fig. 2, Fig. S3) which largely corresponds to this post-peak Maunder Minimum and subsequent change period and also to an observed long, stable, wetter period (~1700-1750) in much of the Mediterranean (37); (ii) a warming trend for the period from the (post-Tambora) earlier/mid-19<sup>th</sup> century AD; and (iii) some correlation of these trends with changes in reconstructed total solar irradiance. Other recent work (38-41) for the Northern Hemisphere over the second millennium AD offers broadly consonant indications. The study of Mann (42) addressing the Middle East, while noting caveats for reconstructions for pre- $20<sup>th</sup>$  century AD period, also reconstructs generally warmer conditions for the Middle East in the mid-19<sup>th</sup> century.

#### **4. Taybet Zaman**

The Taybet Zaman village site represents an interesting, if not uncontroversial, example of attempts to re-use elements of Jordan's heritage as part of the modern tourism industry. The entire town was leased by Jordan Tourism Investments and reused/transformed into a luxury hotel and tourist attraction. The positively recounted version has the local mayor, Abu Firas, of what by the start of the 1990s was a largely abandoned village with only a handful of families and their animals, meeting with Jordanian businessmen, and then selling the idea for the villagers to rent out the old village which these investors would restore and transform, go live in a new village with modern amenities, and have work if they wanted at the new hotel-resort to be constructed in the restored old village (e.g. [http://www.independent.co.uk/travel/news-and](http://www.independent.co.uk/travel/news-and-advice/tourism-must-be-small-to-be-kind-5366036.html)[advice/tourism-must-be-small-to-be-kind-5366036.html\)](http://www.independent.co.uk/travel/news-and-advice/tourism-must-be-small-to-be-kind-5366036.html). However, while noted as an example of re-use of historic structures (ref. 43 p.305), and a winner of a 1996 tourism for tomorrow award (ref. 44 p.180), aspects of this development and economic model for the site (and other such heritage re-use developments) are reviewed and critiqued less positively in other accounts (1, 45, 46), which note that these reclamations and developments are undertaken as a specialized activity by a (remote, not associated) intellectual and economic elite, and lead often to "gentrification and alienation of the host communities rather than a "revitalization" of their social and economic life or the reaffirmation of a sense of belonging to a shared place" (ref. 46 p.13). In practical terms for our project, the site (originally and still as restored) utilizes numerous *Juniperus phoenicea* timbers in its structures. Fig. S20 shows some examples of such timbers used in doors, doorways, roof beams, etc. at the time we sampled. Our sampling happened during renovations in 2011 and 2012. The site is now the Hyatt Zaman Hotel & Resort.



**Fig. S1**. Internal crossdating of the Taybet Zaman (TZM) tree-ring chronology (Fig. 1B). Dendrochronological crossdating was evaluated using the following statistical parameters shown above: Student t-values (*t*) (47)*,* Pearson correlation coefficient (*r*), and trend coefficient (*tr*) (48, 49) were calculated using the Corina v.1.1 software package (6); the number of years overlap (*n*) between the compared tree-ring series is also shown. The set of 10 trees offer a robust tree-ring chronology following standard dendrochronological procedures (2-4).



**Fig. S2**. Crossdating of the combined (master) Taybet Zaman (TZM) *Juniperus phoenicea* chronology in Fig. S1 against the existing southern Jordan *Juniperus phoenicea* chronology (8, 9) available from the International Tree Ring Databank (ITRDB):

<https://www1.ncdc.noaa.gov/pub/data/paleo/treering/chronologies/asia/jord001.crn> with the individual raw data available from:

[https://www1.ncdc.noaa.gov/pub/data/paleo/treering/measurements/asia/jord001.rwl.](https://www1.ncdc.noaa.gov/pub/data/paleo/treering/measurements/asia/jord001.rwl) The values indicate secure crossdating and hence secure absolute dates for the TZM chronology.



**Fig. S3**. The known-age placement of the pooled Oxford and Arizona TZM  $^{14}C$  data (18) versus IntCal13 (17). Where there are two or more constituent TZM data, the symbol is blue, where there is only one  $^{14}C$  date available the symbol is red. The difference in  $^{14}C$  ages, TZM minus IntCal13, is shown below (see also Fig. S5). 1σ error bars shown. Of the 65 TZM data, 12 (18%) are not only offset from the IntCal mean value (Fig. S4), but are incompatible as representing the same  $^{14}$ C age allowing for errors on both the TZM data and IntCal13 within 95% probability (18). These are the data centered at 1627, 1717, 1722, 1727, 1750, 1770, 1885, 1890, 1895, 1900, 1905, 1910. In every case the TZM  $^{14}C$  age is older than the contemporary IntCal13 value. See Table S4.



**Fig. S4**. A: the average offset for the pooled dataset in Fig S3, calculated as a Delta\_R query in OxCal 4.3 (19) versus IntCal13 (17) with curve resolution set at 1 year with a neutral prior of 0 $\pm$ 20. This yields a posterior of 18.6 $\pm$ 2.5<sup>14</sup>C years. B: The distribution of the normalized differences ( $\Delta Z$  or  $\overline{Z}$  scores) between the <sup>14</sup>C ages for each AA+OxA pooled (where possible) value versus IntCal13 including respective errors (Fig. S3, Table S4) and compared to a normal distribution with mean 0 and standard deviation 1.



**Fig. S5**. The difference between the dataset from Fig. S3 versus IntCal13 (17), with four different adjacent average smoothed curves plotted through the data and also two locally weighted scatterplot weighting (Loess) curves. The offset fluctuates over time. The average offset in the two periods of greater largely sustained offsets, indicated by the dashed line boxes, is 24 $\pm$ 5<sup>14</sup>C years. The offset is  $\geq$  +10<sup>14</sup>C years for 63% of the comparisons.



**Fig. S6**. The TZM 14C date series, from the data in Table S2, and as shown in Fig. 2A, but indicating which data came from which of the three TZM tree samples (TZM-3, TZM-4 and TZM-15) employed in this study. 1σ error bars are shown.



**Fig. S7**. <sup>14</sup>C data on the BADG samples from Bag 1 (oldest) to Bag 33 (most recent). 1σ error bars are shown.



**Fig. S8**. The 68.2% most likely probability (highest posterior density = hpd) calendar age range fits of the BADG Bags 1 (oldest) to 33 (most recent) ordered 14C sequence against IntCal13 with a minimum 5±2 years gap between the samples and allowing for a possible offset between the BADG Jordan juniper data and the IntCal13 (17) curve via a Delta\_R query in OxCal (19) with a neutral prior of  $0\pm 20$  (see Fig. S11). All error bars indicate 68.2% probability ranges. We employed a 5 year curve resolution. For the OxCal runfile, see Table S6. All but two of the BADG 14C data lie above (older than) the corresponding IntCal13 mean value.



**Fig. S9**. The 68.2% most likely probability (highest posterior density = hpd) calendar age range fits of the BADG Bags 1 (oldest) to 33 (most recent) ordered 14C sequence against IntCal13 with a minimum 5±3 years gap between the samples and allowing for a possible offset between the BADG Jordan juniper data and the IntCal13 (17) curve via a Delta\_R query in OxCal (19) with a neutral prior of  $0\pm 20$  (see Fig. S12). All error bars indicate 68.2% probability ranges. We employed a 5 year curve resolution. For the OxCal runfile, see Table S6. All but two of the BADG 14C data lie above (older than) the corresponding IntCal13 mean value.



**Fig. S10**. The 68.2% most likely probability (highest posterior density = hpd) calendar age range fits of the BADG Bags 1 (oldest) to 33 (most recent) ordered 14C sequence against IntCal13 with a minimum 5±4 years gap between the samples and allowing for a possible offset between the BADG Jordan juniper data and the IntCal13 (17) curve via a Delta\_R query in OxCal (19) with a neutral prior of  $0\pm 20$  (see Fig. S13). All error bars indicate 68.2% probability ranges. We employed a 5 year curve resolution. For the OxCal runfile, see Table S6. All but two of the BADG 14C data lie above (older than) the corresponding IntCal13 mean value.



placed in Fig. S8 versus the IntCal13 (17) curve via a Delta\_R query in OxCal (19) with a neutral prior of  $0\pm 20^{-14}$ C years. The range of the most likely 68.2% probability region of the posterior is  $\sim$ 26 $\pm$ 9<sup>14</sup>C years.



placed in Fig. S9 versus the IntCal13 (17) curve via a Delta\_R query in OxCal (19) with a neutral prior of  $0\pm 20^{-14}$ C years. The range of the most likely 68.2% probability region of the posterior is  $\sim 20 \pm 13.75$  <sup>14</sup>C years.



Fig. S13. The overall (average) offset (posterior density) between the BADG <sup>14</sup>C ages as best placed in Fig. S10 versus the IntCal13 (17) curve via a Delta\_R query in OxCal (19) with a neutral prior of  $0\pm 20^{-14}$ C years. A split range, as shown, is typical, but with a clear most likely solution. The much less likely alternative solution tries to place the entire BADG bags 1-33 series within the  $17<sup>th</sup>$  century. This appears impossible in practical terms, considering (i) the approximate tree-ring sequence information which in fact underestimated, not massively overestimated, the ring count, so bags 1-33 contain probably no less than around 165 rings/years and likely rather more given under-counting, and (ii) samples from the post bag 33 BADG series, starting from bag 35, included samples with post-bomb  ${}^{14}C$ , so bag 33 cannot be  $17<sup>th</sup>$  century when bag 33 is only supposedly ca.10 rings plus or minus counting errors before bag 35. A placement at least in the mid-later  $19<sup>th</sup>$  century seems the earliest possible date that makes sense for bag 33. The range of the most likely sub-region of the 68.2% probability region of the posterior is  $\approx 27 \pm 7.5$  <sup>14</sup>C years.



**Fig. S14**. As Fig. S11 but with the neutral prior of  $0\pm 10^{-14}$ C years. The range of the most likely 68.2% probability region of the posterior is  $\approx 23 \pm 5^{14}$ C years.



Fig. S15. The overall (average) offset between the BADG <sup>14</sup>C age ordered sequence as best placed versus the IntCal13 (17) curve via a Delta\_R query in OxCal (19) with the prior value of  $26\pm9$  <sup>14</sup>C years (from Fig. S11) (thus the dataset in Table S6 changing the Delta<sub>R</sub> value to 26,9). This assumed prior is shown by the light gray region. The calculated posterior density (the dark gray region) shows a good overlap, confirming the average offset range proposed (contrast with the lack of overlap with the neutral priors in Figs. S13 and S14).



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Fig. S18. The D\_Sequence fit of the Oxford (OxA) data on a set of 12<sup>14</sup>C dates on 11 knownage tree-ring samples of German Oak (25) against IntCal13 (17) employing OxCal (19) (two measurements on the same rings are combined (18)). Curve resolution set at 1 year. The boxes represent the 68.2% probability ranges in terms of both  ${}^{14}C$  and calendar ranges. The IntCal13 curve is shown in blue as a 68.2% probability band. Amodel =118.8, Aoverall 122.3 well above the satisfactory threshold values of 60. All data include the known ages within the modelled most likely 68.2% hpd ranges.



Oak (25) as best placed in Fig. S18 versus the IntCal13 (17) curve via a Delta\_R query in OxCal (19) with a neutral prior of  $0\pm 10^{-14}$ C years. There is a negligible (to approximately zero) offset of  $0.7\pm7.2$  <sup>14</sup>C years.









**Fig. S20**. Photos, taken in 2011, showing some examples of *Juniperus phoenicea* timbers in use in doors, doorways, roof/ceiling beams and other structural elements at Taybet Zaman.

Table S1. Tree-ring measurements in 1/1000ths (0.001) of a mm by year in Tucson (.rwl) format for each of the ten TZM samples employed in this study and as shown crossdated in Fig. S1: TZM1 (=file TZM012 below), TZM3 (=file TZM031 below), TZM4 (=file TZM041 below), TZM9 (=file TZM091 below), TZM12 (=file TZM121 below), TZM15 (=file TZM151 below), TZM16 (=file TZM161 below), TZM17 (=file TZM171 below), TZM23 (=file TZM231 below), TZM33&36 (=file TZM330 below).









**Table S2**. Dendrochronologically dated TZM *Juniperus phoenicea* tree-ring samples (Cornell TZM numbers) and their corresponding <sup>14</sup>C dates from Oxford and Arizona. The OxA-X-2646-40 and OxA-X-2646-46 results followed non-routine chemistry; instead of employing a soxhlet extraction, test tubes and heating blocks were used. Since no significance difference in results is evident we use these dates. OxA-32832 is clearly too old and an outlier. There is no obvious explanation from laboratory data. This result is not employed in the analyses in our paper. AA-103300 is very clearly far too old and a massive outlier. There is no explanation from laboratory data. This result is not employed in the analyses in our paper. Where two or three OxA dates were run on the identical sample, a weighted average has been calculated (17) and this is used in the subsequent analysis (such as Figs. 2, 3, S3). The exception is the AD1818-1822 sample where there are four data and OxA-29735 is an outlier and is not included in the weighted average value for this sample. The codes for the Oxford pretreatment regime (14) are:  $UW =$ ABA+bleach pretreatment,  $UA = \alpha$ -cellulose pretreatment,  $* =$  solvent extraction (acetone then water for OxA-  $\leq$ 32830; 1:1 (v/v) chloroform/ethanol, then ethanol, then water for OxA- $\geq$ 34703). Quoted  $\delta^{13}$ C values are independent stable isotope mass spectrometer measurements per mil relative to VPDB.









**Table S3**. The ordered sequence of BADG *Juniperus phoenicea* tree-ring samples and their corresponding 14C dates from Oxford and Arizona. In total, the dataset comprises Bag 1, oldest, to Bag 50, most recent, but, as discussed in the SI Text, only the ordered sequence from Bag 1 to Bag 33 was employed for the analysis. The data in the gray-shaded cells were *not* employed. Where there were 32XXX series OxA dates (*OxA2* column in the table) we used these data instead (Bag 29) as these data were re-run samples using full  $\alpha$ -cellulose (and consistently returned older  $14^{\circ}$ C ages). PB = post-bomb  $14^{\circ}$ C age. Three AA dates were reported initially as post-bomb and then subsequently re-run (as α-cellulose) and returned more plausible  ${}^{14}C$  ages. Quoted  $\delta^{13}$ C values are independent stable isotope mass spectrometer measurements per mil relative to VPDB (no data are available where not listed). Where the AA and OxA data could be combined (18) a weighted average (Av) is given (and used in Figs. S8-S10).





Table S4. TZM <sup>14</sup>C data, combining (where possible) the available AA and OxA data (Fig. S3), compared to the contemporary mean IntCal13 value (17) and tested for significant (95%) difference (18). Linear interpolation is applied to the IntCal13 dataset to achieve the relevant annual year values when between the 5-year IntCal13 values.

Year AD	TZM 14C	<b>TZM</b>	IntCal13 14C	IntCal13	T value	TZM Older (71%
		<b>SD</b>		<b>SD</b>		data)
1612	380	18	357.6	7.6	1.3	X
1617	382	18	346.0	7.0	3.5	Χ
1622	306	19	333.2	7.0	1.8	
1627	396	19	330.4	7.4	10.4	X
1632	328	14	319.2	7.6	0.3	X
1637	274	19	303.2	7.0	2.1	
1642	268	19	288.6	7.0	1.0	
1647	248	19	264.6	7.0	0.7	
1652	258	19	246.8	7.0	0.3	X
1657	236	19	236.6	7.0	0.0	
1662	203	19	225.8	7.0	1.3	
1667	217	18	201.4	7.4	0.6	Χ
1672	177	17	179.0	7.6	0.0	
1677	174	18	172.4	7.4	0.0	Χ
1682	135	19	156.6	7.6	1.1	
1687	153	18	137.8	6.6	0.6	X
1692	158	18	127.0	5.6	2.7	X
1697	123	18	112.0	5.4	0.3	Χ
1702	127	19	103.8	6.0	1.4	X
1707	130	19	101.8	6.0	2.0	Χ
1712	115	17	95.4	6.4	1.2	X
1717	135	15	97.8	6.6	5.2	Χ
1722	146	16	108.2	6.0	4.9	X
1725	119	22	113.9	6.0	0.0	Χ
1727	186	17	122.6	6.8	12.1	X
1730	141	22	136.5	8.0	0.0	X
1732	185	23	143.0	8.0	3.0	Χ
1735	152	22	152.0	8.0	0.0	
1737	175	21	158.0	7.6	0.6	Χ
1740	148	17	165.9	7.0	0.9	
1742	189	20	165.4	7.0	1.2	X
1745	132	21	163.2	7.0	2.0	
1747	186	31	163.0	7.0	0.5	X
1750	228	22	162.9	7.0	8.0	Χ
1752	162	21	162.2	7.0	0.0	
1755	199	22	160.8	7.0	2.8	X
1760	177	22	156.2	7.0	0.8	X



**Table S5.** OxCal (19) runfile for the OxA Delta R plot shown in Fig. 2C. This is given as an example for the other similar plots in Fig. 2C, Fig. 3C, SI Appendix Fig. 4A which are similar *but* employ the relevant data from Table S2 and against the stated calibration dataset. IntCal13 (17) is the current default calibration dataset in OxCal. To employ SHCal13 (50), as reported in Fig. 3A, this must be specified adding the code under Options reading: Curve="ShCal13.14c";

```
Options()
{
  Resolution=1;
};
Plot()
{
 Delta_R("OxA Jordan Juniper", 0, 20);
  D_Sequence ()
 {
  R_Date("OxA 32744 1612",380,24);
  Gap(5);
  R_Date("OxA 32745 1617",404,24);
  Gap(5);
  R_Date("OxA 32746 1622",275,26);
  Gap(5);
  R_Date("OxA 32747 1627",408,23);
  Gap(5);
  R_Combine ("1632")
 {
   R_Date("OxA 32748 1632",307,23);
   R_Date("OxAX-2646-40 1632",322,23);
  };
  Gap(5);
  R_Date("OxA 32574 1637",266,26);
  Gap(5);
  R_Date("OxA 32575 1642",274,24);
  Gap(5);
  R_Date("OxA 32576 1647",260,24);
  Gap(5);
  R_Date("OxA 32577 1652",273,25);
  Gap(5);
  R_Date("OxA 32578 1657",238,24);
  Gap(5);
  R_Date("OxA 32579 1662",200,25);
  Gap(5);
  R_Date("OxA 32580 1667",234,24);
  Gap(5);
  R_Date("OxA 32581 1672",214,25);
  Gap(5);
  R_Date("OxA 32749 1677",206,24);
  Gap(5);
  R_Date("OxA 32582 1682",146,25);
  Gap(5);
  R_Date("OxA 32583 1687",157,24);
  Gap(5);
  R_Date("OxA 32584 1692",171,24);
  Gap(5);
```

```
 R_Date("OxA 32750 1697",129,23);
  Gap(5);
  R_Date("OxA 32831 1702",147,24);
  Gap(5);
  R_Date("OxA 32751 1707",120,23);
  Gap(5);
  R_Date("OxA 32752 1712",143,24);
  Gap(5);
  R_Combine ("1717")
 {
  R_Date("OxA32753 1717",141,24);
  R_Date("OxAX-2646-46 1717",126,23);
  };
  Gap(5);
  R_Date("OxA 32754 1722",122,23);
  Gap(3);
  R_Date("OxA 29718 1725",119,22);
  Gap(2);
  R_Date("OxA 32755 1727",177,23);
  Gap(3);
  R_Date("OxA 29719 1730",141,22);
  Gap(5);
  R_Date("OxA 29720 1735",152,22);
  Gap(2);
  R_Date("OxA 32756 1737",191,23);
  Gap(3);
  R_Combine ("1740")
 {
  R_Date("OxA29721 1740",122,23);
  R_Date("OxA32828 1740",175,24);
  };
  Gap(2);
  R_Date("OxA 32757 1742",184,23);
  Gap(3);
  R_Date("OxA 29722 1745",132,21);
  Gap(5);
  R_Date("OxA 29924 1750",228,22);
  Gap(2);
  R_Date("OxA 32758 1752",159,23);
  Gap(3);
  R_Date("OxA 29723 1755",199,22);
  Gap(5);
  R_Date("OxA 29724 1760",177,22);
  Gap(5);
  R_Date("OxA 29725 1765",129,22);
  Gap(5);
  R_Combine ("1770")
  {
  R_Date("OxA29726 1770",198,23);
  R_Date("OxA32829 1770",253,24);
  };
  Gap (5);
  R_Combine ("1775")
 {
  R_Date("OxA29727 1775",136,21);
   R_Date("OxA29728 1775",160,22);
```
 }; Gap (5); R\_Date("OxA 29729 1780",221,21); Gap(5); R\_Date("OxA 29730 1785",205,22); Gap(5); R\_Date("OxA 29731 1790",231,23); Gap(5); R\_Date("OxA 29732 1795",208,23); Gap(5); R\_Date("OxA 29733 1800",203,21); Gap(10); R\_Date("OxA 29734 1810",140,22); Gap(10); R\_Combine ("1820") { R\_Date("OxA32830 1820",124,23); R\_Date("OxA 35659 1820",111,25); R\_Date("OxA 36014 1820",112,23); }; Gap(5); R\_Date("OxA 29389 1825",132,25); Gap(5); R\_Date("OxA 29390 1830",119,25); Gap(5); R\_Combine ("1835") { R\_Date("OxA29445 1835",92 ,25); R\_Date("OxA32825 1835",118,26); }; Gap(5); R\_Date("OxA 29446 1840",142,25); Gap(5); R\_Date("OxA 29447 1845",163,25); Gap(5); R\_Combine ("1850") { R\_Date("OxA29717 1850",110,21); R\_Date("OxA32826 1850",161,24); }; Gap(5); R\_Date("OxA 29448 1855",126,24); Gap( 5); R\_Date("OxA 29449 1860",156,23); Gap( 5); R\_Date("OxA 29450 1865",141,24); Gap( 5); R\_Date("OxA 29867 1870",111,24); Gap( 5); R\_Date("OxA 29530 1875",150,23); Gap( 5); R\_combine ("1880") { R\_Date("OxA29531 1880",87,25); R\_Date("OxA32827 1880",103,23); };

```
 Gap(5);
  R_Combine ("1885")
  {
  R_Date("OxA29532 1885",181,24);
  R_Date("OxA34703 1885",158,23);
  };
  Gap(5);
  R_Combine ("1890")
  {
  R_Date("OxA29533 1890",131,24);
  R_Date("OxA 35658 1890",127,23);
  R_Date("OxA 36013 1890",131,24);
  };
  Gap(5);
  R_Combine ("1895")
  {
  R_Date("OxA29534 1895",162,22);
  R_Date("OxA34705 1895",114,23);
  };
  Gap(5);
  R_Combine ("1900")
  {
  R_Date("OxA29535 1900",139,23);
  R_Date("OxA34706 1900",146,22);
  };
  Gap(5);
  R_Combine ("1905")
  {
  R_Date("OxA29536 1905",133,24);
  R_Date("OxA34707 1905",98 ,24);
  };
  Gap(5);
  R_Combine ("1910")
  {
  R_Date("OxA29537 1910",105,24);
  R_Date("OxA34708 1910",148,23);
  };
 };
};
```
**Table S6**. OxCal (19) runfile for the BADG ordered sequence reported in Figs. S8, S11. Note for the data reported in Figs. S9 and S10 the lines of code below reading "Gap (5,2);" were replaced with "Gap  $(5,3)$ ;" and "Gap  $(5,4)$ ;" respectively.

```
Options()
{
 Resolution=5;
};
Plot()
{
  Delta_R ("BADG Bags 1-33",0,20);
  Sequence ("BADG Bags 1-33")
 {
  Boundary ("Start");
  Phase ("Bag 1")
 {
   R_Date("Bag 1 OxA-29538",289,23);
   R_Date("Bag 1 AA103219",213,25);
  };
  Gap (5,2);
  R_Combine ("Bag 2")
 {
   R_Date("OxA-29539",190,24);
   R_Date("AA103220",187,16);
  };
  Gap (5,2);
  R_Combine ("Bag 3")
 {
   R_Date("OxA-29540",195,23);
   R_Date("AA103221",170,21);
  };
  Gap (5,2);
  R_Combine ("Bag 4")
 {
   R_Date("OxA-29451",184,23);
   R_Date("AA103222",154,33);
  };
  Gap (5,2);
  R_Combine ("Bag 5")
 {
   R_Date("OxA-29391",162,24);
   R_Date("AA103223",166,21);
  };
  Gap (5,2);
  R_Combine ("Bag 6")
 {
   R_Date("OxA-29452",161,23);
   R_Date("AA103224",162,16);
  };
  Gap (5,2);
  R_Combine ("Bag 7")
 {
   R_Date("OxA-29392",149,24);
   R_Date("AA103225",139,22);
  };
```

```
 Gap (5,2);
  Phase ("Bag 8")
 {
 R_Date("Bag 8 OxA
-29393",197,25);
   R_Date("Bag 8 AA103226",121,16);
  };
  Gap (5,2);
  R_Combine ("Bag 9")
 {
 R_Date("OxA
-29394",133,25);
   R_Date("AA103227",144,22);
  };
  Gap (5,2);
  R_Combine ("Bag 10")
 {
 R_Date("OxA
-29395",152,25);
   R_Date("AA103228",102,22);
  };
  Gap (5,2);
  R_Combine ("Bag 11")
 {
 R_Date("OxA
-29420",163,21);
   R_Date("AA103229",121,24);
  };
  Gap (5,2);
  R_Combine ("Bag 12")
 {
 R_Date("OxA
-29421",149,22);
   R_Date("AA103230",130,16);
 R_Date("OxA
-29422",96,21);
  };
  Gap (5,2);
  R_Combine ("Bag 13")
 {
 R_Date("OxA
-29423",188,21);
   R_Date("AA103231",179,25);
  };
  Gap (5,2);
  R_Combine ("Bag 14")
 {
 R_Date("OxA
-29424",242,20);
   R_Date("AA103232",203,21);
  };
  Gap (5,2);
  R_Combine ("Bag 15")
 {
 R_Date("OxA
-29425",203,21);
   R_Date("AA103233",195,25);
  };
  Gap (5,2);
  R_Combine ("Bag 16")
 {
 R_Date("OxA
-29426",220,21);
   R_Date("AA103234",175,23);
  };
  Gap (5,2);
```

```
 R_Combine ("Bag 17")
 {
  R_Date("OxA-29427",189,20);
  R_Date("AA103235", 203,21);
  };
  Gap (5,2);
  R_Date ("Bag 18 OxA-29488",189,22);
  Gap (5,2);
  R_Combine ("Bag 19")
  {
  R_Date("OxA-29489",181,22);
  R_Date("AA103237",198,21);
  };
  Gap (5,2);
  R_Combine ("Bag 20")
  {
  R_Date("OxA-29490",176,22);
  R_Date("AA103238",195,16);
  };
  Gap (5,2);
  R_Date ("Bag 21 OxA-29453",222,25);
  Gap (5,2);
  R_Combine ("Bag 22")
 {
  R_Date("OxA-29454",202,24);
  R_Date("AA103240",222,21);
  };
  Gap (5,2);
  R_Combine ("Bag 23")
 {
  R_Date("OxA-29455",234,24);
  R_Date("AA103241", 290,22);
  };
  Gap (5,2);
  R_Combine ("Bag 24")
 {
  R_Date("OxA-29456",169,24);
  R_Date("AA103242",180,24);
  };
  Gap (5,2);
  R_Combine ("Bag 25")
 {
  R_Date("OxA-29457",153,24);
  R_Date("AA103243",187,16);
  };
  Gap (5,2);
  R_Combine ("Bag 26")
  {
  R_Date("OxA-29458",204,25);
  R_Date("AA103244",179,16);
  };
  Gap (5,2);
  R_Combine ("Bag 27")
 {
  R_Date("OxA-29823",107,22);
  R_Date("AA103245", 146,16);
```

```
 };
  Gap (5,2);
  R_Combine ("Bag 28")
  {
  R_Date("OxA-29459",125,23);
  R_Date("AA103246",128,22);
  };
  Gap (5,2);
  R_Combine ("Bag 29")
  { 
  R_Date("OxA-32071 Bag 29",89,24);
  R_Date("OxA-32119 Bag 29",69,24);
  };
  Gap (5,2);
  R_Combine ("Bag 30")
  { R_Date("OxA-32735",117,25);
  R_Date("AA103248",130,22);
  };
  Gap (5,2);
  R_Date ("Bag 31 OxA-29461",140,23);
  Gap (5,2);
  R_Combine ("Bag 32")
 {
  R_Date("OxA-29462",162,23);
  R_Date("AA103250",162,16);
  R_Date("OxA-29463",121,23);
  };
  Gap (5,2);
  R_Combine ("Bag 33")
 {
  R_Date("OxA-29464",120,23);
  R_Date("AA103251",138,16);
  };
  Boundary ("End");
 };
};
```
**Table S7**. OxCal (19) runfile for the Tel Rehov dates reported in Fig. 4B. The <sup>14</sup>C age values for the 11 stratigraphically ordered archaeological elements listed in Fig. 2 of ref. 31 are analyzed as an ordered sequence. For the data reported in Fig. 4B a first iteration is the published data with IntCal13 (17). Then this runfile is modified with the extra line added of: Delta R ("Southern Levant Average",19,5); and then a third version was run with the Jordan juniper (JJ) modified approximate calibration curve shown in Fig. 4A.

Plot()

{ Sequence ("Iteration 1: Bruins et al. 2003 Rehov with IntCal13") { Boundary ("Start"); R\_Date ("Area D Phase D6a L2836",2928,26); R\_Date ("Area D Phase D4b L1845 House", 2924,22); R\_Date ("Area D Phase D4a L1836 Street", 2885,26); R\_Date ("Area D Phase D3 Pits", 2831,18); R\_Date ("Area D Phase D2 Locus 1802", 2805,15); R\_Date ("Area C Stratum VI Locus 4426", 2768,12); R Date ("Area C Stratum V Locus 2441", 2776,9); R\_Date ("Area C Stratum V Locus 2425", 2788,14); R\_Date ("Area C Stratum V Locus 2422", 2771,8); R\_Date ("Area B Stratum V Locus 4218", 2786,22); R\_Date ("Area C Stratum IV Locus 5498", 2755,25); Boundary ("End"); };

};

**Table S8.** OxCal (19) runfile for the Khirbat en-Nahas <sup>14</sup>C data and Bayesian chronological model (32) as reported in Fig. 4B. We ran three versions of the model. Iteration 1 – *without* the yellow highlighted text in line 3 of the runfile below – was the published data and model re-run with IntCal13 (17). Iteration 2 adjusted this model by employing the approximate average Jordanian offset of  $19\pm5$  <sup>14</sup>C years (see main text) using the yellow highlighted text in line 3 of the runfile below. Iteration 3 in turn considered the situation if the model (*without* the Delta\_R line in yellow highlight) was run with the approximate Jordan juniper modified calibration curve shown in Fig. 4A.

```
Plot()
  Delta_R("Jordanian Av. Offset", 19, 5);
  Sequence ()
 {
  Sequence ()
 {
   Boundary ("Start Stratum 3");
   Sequence ("Stratum 3")
 {
    R_Date ("OxA-17646",2871,26);
    R_Date ("OxA-17644",2824,25);
    R_Date ("OxA-17647",2764,25);
    R_Date ("OxA-17643",2813,26);
    R_Date ("OxA-17703",2792,30);
    R_Date ("OxA-17640",2770,25);
    R_Date ("OxA-17642",2781,25);
    R_Date ("OxA-17641",2767,25);
    R_Date ("OxA-17645",2747,25);
   };
   Boundary ("End 3/start Strat 2 building and strat 1 slag mound");
   Phase ()
 {
    Sequence ()
\{ Boundary ("Start 2");
    Sequence ("Stratum 2 Building")
\{ R_Date ("OxA-17638",2814,25)
\{ \{ \} Outlier();
     };
     R_Date ("OxA-17633",2734,25);
     R_Date ("OxA-17632",2713,26);
     R_Date ("OxA-17636",2732,25);
     Phase ()
\{ \{ \} R_Date ("OxA-17635",2777,25);
     R_Date ("OxA-17634",2783,25);
 };
     R_Date ("OxA-17702",2740,30);
     R_Date ("OxA-17631",2676,26)
\{ \{ \} Outlier();
     };
```
{

```
 };
 Boundary ("End 2");
 };
 Sequence ()
\{ Boundary ("Start 1");
 Sequence ("Stratum 1")
\{ R_Date ("OxA-17639",2678,26);
 R_Date ("OxA-17637",2836,26)
    {
    Outlier();
 };
 R_Date ("OxA-12437",2746,35);
 R_Date ("OxA-12436",2659,32)
    {
    Outlier();
    };
    R_Date ("OxA-17630",2676,26);
    };
   Boundary ("End 1");
   };
 };
 Boundary ("End Sequence");
 };
 };
```
};

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