



Supplementary Information for  
**Growth Model Interpretation of Planet Size Distribution**

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## Materials and Methods

**Planet data** of all exoplanets used for our analysis have been downloaded from the *NASA Exoplanet Archive* (1, 2) on several occasions, with the manual addition/modification of some newly-discovered planets: in particular, on Nov.10, 2016 for initial mass-radius plot and statistical assessments, and subsequently on Feb.10, May 31, Sept.9, Oct.16, 2017, and Feb.20, Jun.15, 2018 for updates. Moreover, we have updated the radii of some planets below  $8 R_{\oplus}$  with GAIA DR1 (3) and GAIA DR2 (4). To make the mass-radius plot more precise, we select planets with density constraints better than about  $\pm 50\%$ .

Specifically, since density  $\rho \propto M/R^3$ , taking natural logarithmic gives  $\ln \rho = \ln M - 3 \ln R + \text{const}$ , and differentiation of it gives:  $\frac{d\rho}{\rho} = \frac{dM}{M} - 3 \cdot \frac{dR}{R}$ .

Error propagation assumes errors of mass and radius are independent:  $\frac{\sigma\rho}{\rho} \approx \sqrt{\left(\frac{\sigma M}{M}\right)^2 + \left(3 \cdot \frac{\sigma R}{R}\right)^2}$

Then, taking into account asymmetric error bars to avoid planets with an extremely large error bar in one direction, we chose the selection criterion as:

$$\ln\left(\frac{\rho + \sigma\rho^+}{\rho - \sigma\rho^-}\right) \approx \sqrt{\left[\ln\left(\frac{M + \sigma M^+}{M - \sigma M^-}\right)\right]^2 + \left[3 \cdot \ln\left(\frac{R + \sigma R^+}{R - \sigma R^-}\right)\right]^2} \lesssim 1$$

where the + and – signs denote the error bar in the plus and minus direction respectively. The planet data are listed in (**SI Appendix, Table S1**), with detailed references (3–80).

**Figs. 1, 2, & 4** have been plotted using this selection criterion with the built-in schemes in *Mathematica* (*Version 11*) including “ListLogLogPlot” and “ErrorListPlot”.

**Statistical analysis** of the planet data has been performed with *Mathematica*’s built-in “WeightedData”, “SmoothHistogram”, and “SmoothKernelDistribution” packages. The smoothing is done with the *Mathematica*’s default “Silverman” bandwidth-selection-method. Each planet data is weighted by the product of its positive and negative  $1-\sigma$  mass error bars. We then calculate the weighted mean and weighted standard deviation for each set of planets.

**Observations bias** of the RV method tends to detect and characterize planets of higher masses and orbiting closer to their host stars (81). Because of this bias, the RV method becomes more sensitive as mass of a planet increases at given orbit, with current detection limit of  $\sim 3 M_{\oplus}$ . When we invert the RV mass distribution to a radius distribution, we miss the planets smaller than  $\sim 1.4 R_{\oplus}$ . Because such planets are not taken into account by our simulations, our Monte Carlo histogram (gray bars in **Fig. 2**) has much less planets below  $\sim 1.4 R_{\oplus}$  than observed by the transiting method. This also results in a slightly higher radius gap compared to the population detected by the transiting method. Since this paper considers planets more massive than  $\sim 3 M_{\oplus}$ , the observational bias of the RV method has little effect on our results and conclusions.

**Multi-planet systems** which contain both super-Earths and sub-Neptunes are interesting. First of all, the goal of our paper is not to model a specific exoplanetary system but rather to use the observed statistical distributions for identifying possible nature of different planet populations. This is why we focus on the majority of exoplanets found by the *Kepler* mission. Our growth model is statistical in nature and does not intend to explain detailed architecture of a specific planetary system. Such outliers and extreme cases constitute  $\sim 5\%$  of the observed exoplanet population which our model cannot explain. Indeed, real planet-forming scenarios are more complicated and, perhaps, somewhat different around different stars. One of the critical features in different disks would be how and when the snowlines and planet-building materials migrate

during planet formation. For example, taking into account formation of a planet at a snowline followed by its inward migration, Ormel et al. 2017 (82) managed to explain the formation and architecture of compact multi-planet systems such as *Trappist-1* (70, 76, 83). *Kepler-20* (26, 84, 85) can be also explained by such scenario. We call this scenario the “snowline conveyor belt hypothesis” which is consistent with our growth model. In addition, planet-planet collisions (i.e. giant impacts), and more generally, strong dynamical interactions among planets during planet formation have long been suggested by planet formation theory, and recently have been confirmed by the observation and characterization of the *Kepler-107* system (86, 87).

**TTV planets.** The mass determination from transit-timing-variation (TTV) for multi-planet systems is known to be difficult and, often requires a later revision, because the masses determined are often degenerate with the combined eccentricities of planets in that system (88). Sometimes the N-body Markov-Chain-Monte-Carlo (MCMC) approach is used to find the best solutions. When a longer time duration of TTV observations becomes available, or data are reassessed by a different group, the planet masses can change significantly; e.g., *Trappist-1* system (70, 83, 89, 90) and *Kepler-11* system (68, 71, 72). Some TTV planets (**Fig. 2**) of 2-4  $R_{\oplus}$  seem to be puffy, i.e., of low-bulk-density. They can be explained by a core (rocky/icy, 3-10  $M_{\oplus}$ ) plus a thick, but very light (1~2% of total planet mass) gaseous envelope. Also, statistically speaking, systems with significant TTV effects are less common among all *Kepler* planetary systems; e.g., among 2599 *Kepler* Objects of Interest (KOIs) only 260 were found with significant TTVs with long-term ( $>100$  day) variations (67). TTV systems are generally characterized by tight spacing of planets, circular (low  $e$ ) and coplanar (low  $i$ ) orbits, and low obliquities of the host stars (91). The difference in planet masses between TTV and RV may be correlated with the host stellar metallicities – TTV planets prefer low-metallicity host stars (92). They can be explained by planet formation under different disk conditions, e.g., solid surface density versus gas surface density in the disk (93, 94).

**Transitional planets** (of 4-10  $R_{\oplus}$ ), in particular the ones slightly larger than 4  $R_{\oplus}$  residing to the right of the purple growth curves in **Fig. 1**, have massive cores ( $\gtrsim 20 M_{\oplus}$ ) but not much gas ( $\lesssim 20\%$  by mass). For example, K2-55 b (stellar [Fe/H]= $0.376 \pm 0.095$ ) has core mass estimated to be  $\sim 40 M_{\oplus}$  with about 10% H<sub>2</sub>/He envelope by mass, regardless of the core composition (rocky or icy) (95). Because their core masses are apparently above the critical core mass (96–98), it is interesting to understand how they could have formed with that much solid material without experiencing a runaway gas accretion. One hypothesis is that they arise as mergers of two or more less massive cores. The merging processes tend to merge the denser cores together but (partially) erode their gas envelopes (99, 100), resulting in a net mass increase accompanied by a very small radius increase, that moves them to the right on the mass-radius diagram. Disks with high metallicity may facilitate mergers. The second hypothesis is motivated by the results of micro-lensing surveys (101, 102), which are sensitive to distant planet populations orbiting their host stars typically at a few a.u. The results show that planets in the 20-80  $M_{\oplus}$  range at a few a.u. are abundant, which is contradictory to the paucity of planets of this mass range predicted by runaway gas accretion models. This suggests that the current runaway core accretion models may require significant revisions to slow down the growth of planets of 10-50  $M_{\oplus}$  and start the runaway accretion at higher masses of 50-80  $M_{\oplus}$  instead of 10  $M_{\oplus}$ . The third hypothesis assumes that part of planet formation takes place when the gas in the disk is depleted; cf. theoretical calculations of transitional disks (103) and the inferred rapid gas depletion by a few million years from the ALMA survey of the Lupus Proto-planetary disks (104, 105). However, one has to be cautious that the distinction between the total gas depletion and CO depletion cannot be inferred from observation directly, because the inferred H<sub>2</sub> depletion is calculated based on the observed CO depletion and assumed ISM-like [CO]/[H<sub>2</sub>] ratio (per. comm. Jane Huang and Sean

Andrews). The cores of these transitional planets, regardless of a hypothesis of their origin, likely contain large fractions of ices in addition to rocks, just as Uranus and Neptune.

**Mass-radius curves.** The mass-radius curves (**Figs. 1 & 2**) allow calculations of average densities and further constraining the planet bulk compositions and internal structures following the methods that we published earlier (106–109) and similar efforts from other researchers in our field (110–120). The H<sub>2</sub>-He is assumed to be a cosmic mixture of 75% H<sub>2</sub> and 25% He by mass (121–123). The Earth-like composition is assumed to be 32.5 wt.% Fe/Ni-metal plus 67.5 wt.% MgSiO<sub>3</sub>-rock. The bulk H<sub>2</sub>O in deep interior is assumed to be in solid phase along the melting curve (liquid-solid phase boundary) (124–131). The colored H<sub>2</sub>O M-R curves correspond to isothermal vapor/liquid/super-critical fluid envelope at various surface temperatures: 300K, 500 K, 700 K, and 1000 K, on top of ice VII-layer along the appropriate melting pressure (132–134). Since ices all have similar densities, and the H<sub>2</sub>O-NH<sub>3</sub>-CH<sub>4</sub> mixture is always dominated by H<sub>2</sub>O due to chemistry evaluation, the mass-radius curve of the H<sub>2</sub>O-NH<sub>3</sub>-CH<sub>4</sub> mixture is expected to be very similar to that of pure-H<sub>2</sub>O. On the other hand, the H<sub>2</sub>-He mass-radius curves are calculated along various interior adiabats at different internal specific entropies, labelled by the temperature of the corresponding specific entropy at 100-bar level in the gas envelope (400K, 1000K, 3000K). The mass-radius curves are almost parallel to one another and do not criss-cross, as in all cases at high pressure the material structure is supported mainly by the electron degeneracy pressure, which has similar functional dependence on compression. Therefore, the density of a particular material under high pressure is primarily determined by its average atomic weight, which is very different for the three major planet-building materials: rocks, ices, H<sub>2</sub>-He gas, resulting in significant density differences among them in planetary deep interiors, and secondarily by the crystal structure or temperature. The surface of a planet in our calculation is always defined at 10<sup>-3</sup> bar (1 milli-bar) level as an approximate for the level that transiting observations probe.

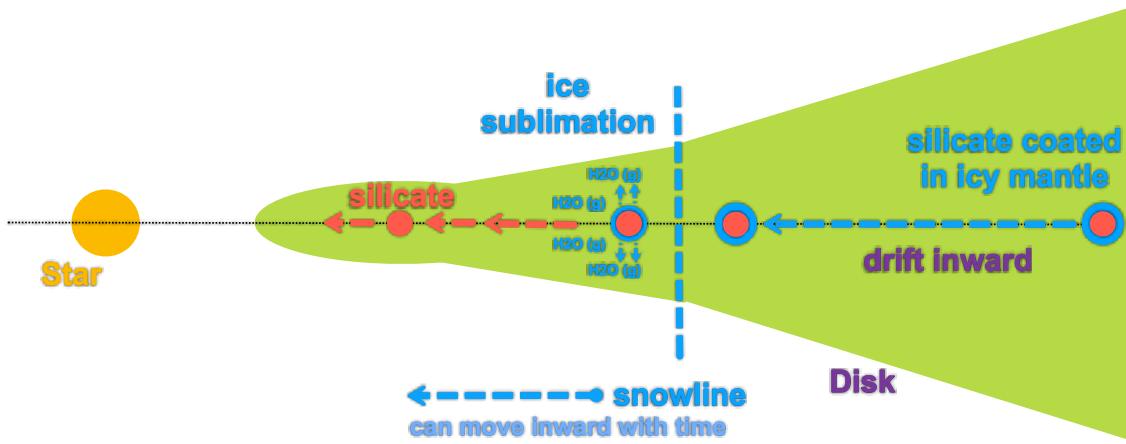
**H<sub>2</sub>O EOS.** The **equations of states (EOS)** of ice and ice-mixtures under appropriate pressure and temperature conditions are required to understand the interiors of these water worlds. Density Functional Theory calculations based on quantum mechanics have provided thermo-physical properties of water at extreme conditions, including predictions of superionic phases (127, 129). Importantly, key results from DFT simulations have been validated by experiments on Sandia's Z-machine (135) as well as by direct measurements of proton conductivity in diamond anvil cell experiments (136) and shock compression experiments (128). Various models of planet interiors utilizing those EOSs have been introduced (137). These models make predictions of the existence of intriguing phases such as the super-ionic state of hot ices and also the non-dipole magnetic field generated in the interior of such planets. They also show that H<sub>2</sub>O EOS commonly used in some previous planetary models significantly overestimate the compressibility, and consequently overestimate the density of water at a few hundred GPa (135), relevant to the interior of these intermediate-sized planets. This new water EOS (135) is adopted for our calculation.

**Growth curves.** The growth curves of either adding H<sub>2</sub>O-ice (cyan) or H<sub>2</sub>-He gas (purple) to a core are calculated (**Figs. 1 & 2**). The cyan growth curves start with 4 and 8 M<sub>⊕</sub> rocky cores and show the trajectory of a growing planet on the mass-radius diagram with increasing amount of H<sub>2</sub>O added. The purple growth curves start with 8 and 16 M<sub>⊕</sub> icy cores (half-ice and half-rock by mass) and show the trajectory of a growing planet on the mass-radius diagram with increasing amount of cold H<sub>2</sub>-He gas added. The tangential slope of the growth curve on the mass-radius diagram depends on the ratio between the density of the added material versus the density of the already-existing material on the planet. The initial slope of a growth curve on a log-log mass-radius plot can be approximated as  $d\ln R/d\ln M \sim \rho_{core}/(3\rho_{acc})$ , where  $\rho_{core}$  is the core density and

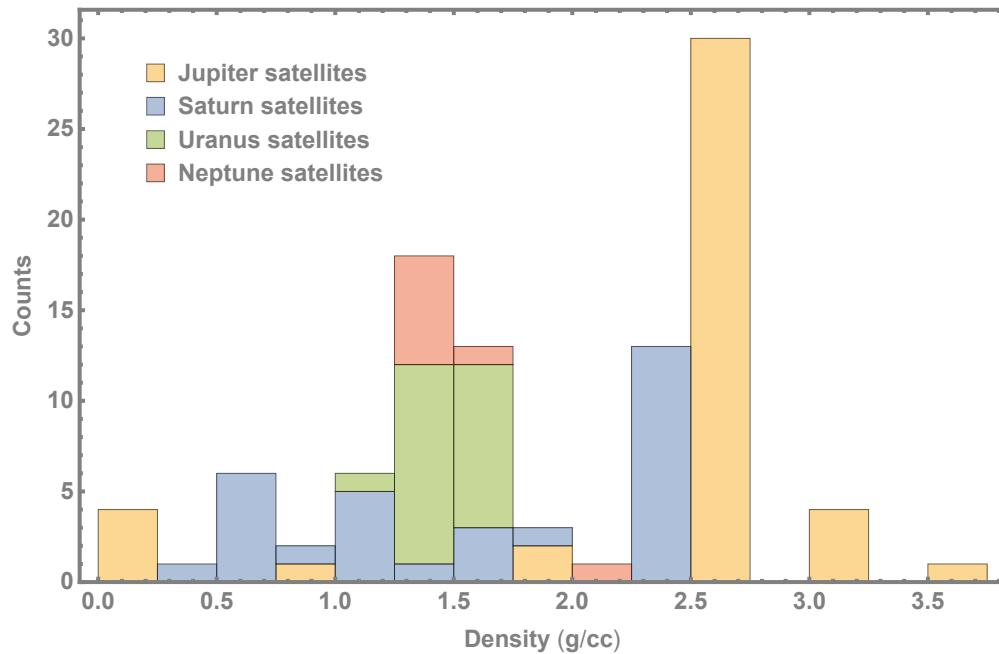
$\rho_{\text{acc}}$  is the density of added material. The tangential slope of the growth curve of H<sub>2</sub>-He is nearly vertical initially, due to the very low density of gas, regardless of the core composition---rocky or icy. The H<sub>2</sub>-He gaseous envelope of 1-2 R<sub>⊕</sub> thick is insignificant in mass---a few percent in mass% (138, 139). Until more and more H<sub>2</sub>-He is added so that the H<sub>2</sub>-He in the planetary deep interiors become self-compressed to high enough densities, so that the tangential slope of H<sub>2</sub>-He growth curve reduces. A *Mathematica* program for this calculation is available upon request. However, be cautious that a simple growth curve is not to be understood as a realistic time-evolution for a single planet growth, such that a planet core must accrete all rocks before accreting any ices. But rather, it is a statistical trend followed by comparing the overall planet populations. In fact, it is more likely that the earlier condensed dust in the nebula is entrapped into multi-layers of ices---icy pebbles (**SI Appendix, Fig. S1**). Then, a planet core, if formed beyond or near the snowline, or if accreting icy pebbles, would accrete rocks and ices simultaneously. The dichotomy of rocky versus icy bodies is also reflected in the densities of solar system giant planets' satellites (**SI Appendix, Fig. S2**).

**Monte Carlo simulation.** We conduct a series of Monte Carlo simulations, in order to test our model of rocky and intermediate-sized planetary mass-radius relation, which provide clues on the composition of the planets. Specifically, we first sample the planetary masses using the observed mass distribution, where the masses of the rocky worlds are generated following the empirical smooth kernel distribution of planets with radii between 1.4-1.9 R<sub>⊕</sub>, and the masses of the icy worlds follow the empirical smooth kernel distribution of those with radii between 2-3 R<sub>⊕</sub>. Then, we calculate the radii of the planets based on the mass-radius relation: the radii of the rocky worlds can be obtained equating  $R=M^{1/3.7}$  (108, 140), and the radii of the intermediate-sized worlds can be calculated as  $R=f\times M^{1/3.7}$ .  $f$  is an indicator of the composition of the intermediate-sized planet. For instance,  $f=1$ , if the intermediate-sized worlds still follow the rocky planetary mass-radius relation.  $f=(1+0.55 x-0.14 x^2)$ , where the ice mass fraction “ $x$ ” of the icy world is around 1/2~2/3, depending on incorporating the more volatile (C,N)-clathrate-ices or not. For a complete water world (100% water),  $f=1.41$ . In order to compare with the observed exoplanets with RV mass measurements, we assume that “ $x$ ” follows a uniform distribution in between 1/2 and 2/3 for the icy worlds, corresponding to fully incorporating H<sub>2</sub>O ices in proportion to rocks, and partially incorporating (C,N)-clathrate-ices in some cases. Then we include an additional ±7% uncertainty to account for the observational errors in planetary radii, and an additional ±30% uncertainty to account for the observational errors in planetary masses. Combining the Monte Carlo results of these two populations and adopt a number ratio of #rocky cores : #icy cores = 1 : 2, we recover the bimodality with the planet radius gap matches the observation in the mass-radius diagram (**Fig. 2**). The generated 2-dimensional mass-radius histogram and the 1-dimensional radius histogram are only meant to be compared with the observed planet populations in between 1.4-3 R<sub>⊕</sub>, since our assumption in generating this Monte Carlo simulation draws the empirical mass distribution only from within this radius range. Therefore, planets larger than about 3 R<sub>⊕</sub> generally require the presence of some gaseous envelope, and this is beyond the scope of this simulation. This gas envelope could either be primordial H<sub>2</sub>-He-dominated nebular gas or could be formed by the dissociation and vaporization of the ices near the surface and in the atmosphere of these planets. This **Monte Carlo** simulation is performed using *Mathematica*'s built-in random number generator “RandomVariate”. “DensityHistogram” is used for the 2-dimensional plot of the probability density contours from the Monte Carlo simulation results. A *Mathematica* program for generating the Monte Carlo simulation is available upon request.

**Supplementary Figures:**



**Fig. S1** Cartoon of a protoplanetary disk showing the inward drift of solid particles consisting of silicates (red) and ices (blue). When particles cross the snowline (vertical blue dashed line) they start losing ice by sublimation. Larger particles of pebble size may drift deeper into the inner zone until they completely lose ice. The snowline itself would move inward as the disk cools with time (141–143). Large icy planetary cores formed beyond snowline can later migrate inward due to planet-disk interactions (144–146) or planet-planet scattering (147).



**Fig. S2** Density variations among satellites of the outer Solar System planets. The two well-defined peaks at  $\sim 1.5$  and  $\sim 2.5$  g/cc suggest that the planet formation process involved interactions between both icy and rocky bodies.

**Table S1.**

```
# ref. (3–80).
# This file is modified from the NASA Exoplanet Archive (1, 2): http://exoplanetarchive.ipac.caltech.edu
# COLUMN pl_orbper: Orbital Period [days]
# COLUMN pl_ttvflag: TTV Flag (1=yes, 0=no)
# COLUMN st_teff: Host Stellar Effective Temperature [K]
# COLUMN st_mass: Host Stellar Mass [Solar mass = 1.9885 * 10^30 kg]
# COLUMN st_rad: Host Stellar Radius [Solar radii = 6.957 * 10^8 meter]
# COLUMN pl_name: Planet Name
# COLUMN pl_rvflag: Planet RV Flag (1=yes, 0=no)
# COLUMN pl_masse: Planet Mass [Earth mass = 5.972 * 10^24 kg]
# COLUMN pl_masseerr1: Planet Mass Upper Unc. [Earth mass]
# COLUMN pl_masseerr2: Planet Mass Lower Unc. [Earth mass]
# COLUMN pl_rade: Planet Radius [Earth radii = 6.371 * 10^6 meter]
# COLUMN pl_radeerr1: Planet Radius Upper Unc. [Earth radii]
# COLUMN pl_radeerr2: Planet Radius Lower Unc. [Earth radii]
# COLUMN met: Host Stellar Metallicity [Fe/H]
# COLUMN mass_ref: Reference to the Planet Mass adopted
# COLUMN radius_ref: Reference to the Planet Radius adopted
# COLUMN additional_ref: Additional Reference
#
```

#	pl_orbper	pl_ttvfl	st_te	st_ma	st_ra	pl_name	pl_rvfl	pl_mas	pl_masse	pl_masseerr1	pl_ra	pl_radee	pl_radeeerr1	met	mass_ref	radius_ref	add_ref
	r	ag	ff	ss	d	ag	se	rr1	rr2	de	rr1	rr2	rr2	-	-	-	-
RV planets < 1.4 Earth radii																	
1	0.584249	0	5185	0.837	0.79	K2-229 b	1	2.59	0.43	-0.43	1.164	0.066	-0.048	-	Santerne_2018	Santerne_2018	
					3								0.09				8
2	1.628928	0	3270	0.181	0.21	GJ 1132 b	1	1.66	0.23	-0.23	1.13	0.056	-0.056	null	Bonfils_2018	Dittman_2017	Diamond-Lowe_2018
	7				1												
3	0.355007	0	5058	0.76	0.74	K-78 b	1	1.87	0.27	-0.26	1.2	0.09	-0.09	-	Grunblatt_2015	Grunblatt_2015	
													0.14				5
4	3.777931	0	3216	0.179	0.21	LHS 1140 c	1	1.81	0.39	-0.39	1.282	0.024	-0.024	-	Ment_2018	Ment_2018	Dittmann_2017
					39								0.24				

RV planets 1.4-2.0 Earth radii																	
1	2.17	0	4503	0.7	0.72	K2-216 b	1	7.4	2.2	-2.2	1.8	0.2	-0.1	-	Persson_2018	Persson_2018	Mayo_2018
2	0.7365	0	5172	0.91	0.94	55 Cnc e	1	8	0.3	-0.3	1.88	0.03	-0.03	0.31	Demory_2016	Bourrier_2018	Winn_2011
3	0.8536	0	5259	0.91	0.82	CoRoT-7 b	1	5.74	0.86	-0.86	1.7	0.1	-0.1	0.05	Barros_2014	Stassun_2017	Haywood_2014
4	3.0929	0	4699	0.81	0.78	HD 219134 b	1	4.74	0.19	-0.19	1.602	0.055	-0.055	0.11	Gillon_2017	Gillon_2017	Motalebi_2015
5	6.7646	0	4699	0.81	0.78	HD 219134 c	1	4.36	0.22	-0.22	1.511	0.047	-0.047	0.11	Gillon_2017	Gillon_2017	Motalebi_2015
6	0.9596	0	5261	0.86	0.86	HD 3167 b	1	5.02	0.38	-0.38	1.7	0.18	-0.15	0.04	Christiansen_2017	Christiansen_2017	Gandolfi_2017
7	24.64	0	3896	0.6	0.56	K2-3 c	1	3.1	1.3	-1.2	1.77	0.18	-0.18	-	Damasso_2018	Damasso_2018	Sinukoff_2016
8	44.56	0	3896	0.6	0.56	K2-3 d	1	2.7	1.2	-0.8	1.65	0.17	-0.17	-	Damasso_2018	Damasso_2018	Sinukoff_2016
9	0.5713	0	5470	0.945	0.87	K2-106 b	1	8.36	0.96	-0.94	1.52	0.16	-0.16	-	Guenther_2017	Guenther_2017	Sinukoff_2017
														0.02		Adams_2017	
1	0.3693	0	5200	0.84	0.81	K2-131 b	1	6.5	1.6	-1.6	1.81	0.16	-0.12	0	Dai_2017	Dai_2017	Mayo_2018
0	0.2803	0	4599	0.71	0.68	K2-141 b	1	5.08	0.41	-0.41	1.51	0.05	-0.05	-	Malavolta_2018	Malavolta_2018	Barragan_2018
1	0.8375	0	5599	0.91	1.09	K-10 b	1	4.61	1.27	-1.46	1.489	0.07	-0.062	-	Esteves_2015	Berger_2018	Dumusque_2014
2	3.6961	0	5508	0.95	0.89	K-20 b	1	9.7	1.41	-1.44	1.707	0.08	-0.068	0.07	Buchhave_2016	Berger_2018	Gautier_2012
3	2.78578	0	6202	1.41	1.93	K-21 b	1	5.08	1.72	-1.72	1.639	0.019	-0.015	-0.1	Lopez-Morales_2016	Lopez-Morales_2016	Howell_2012
4	2.42629	0	5656	1.07	1.07	K-406 b	1	6.35	1.4	-1.4	1.558	0.148	-0.119	0.26	Marcy_2014	Berger_2018	Morton_2016
5	4.72674	0	5594	0.91	0.94	K-93 b	1	4.02	0.68	-0.68	1.642	0.069	-0.108	-0.2	Dressing_2015	Berger_2018	Ballard_2014
6	4.60358	0	4852	0.79	0.74	K-99 b	1	6.15	1.3	-1.3	1.511	0.073	-0.069	0.18	Marcy_2014	Berger_2018	Morton_2016
7	24.7371	0	3216	0.179	0.21	LHS 1140 b	1	6.98	0.98	-0.98	1.727	0.032	-0.032	-	Ment_2018	Ment_2018	Dittmann_2017
8	0.78964	0	5576	1.11	1.16	WASP-47 e	1	6.83	0.66	-0.66	1.81	0.027	-0.027	0.36	Vanderburg_2017	Vanderburg_2017	Almenara_2016
9	3.50473	1	5441	0.97	0.92	K-18 b	1	6.9	3.4	-3.4	1.765	0.185	-0.165	0.19	Cochran_2011	Berger_2018	Morton_2016
0																	

RV planets 2.0-3.0 Earth radii																	
1	41.6855	0	5766	1.67	1.08	K2-56 b	1	16.3	6	-6.1	2.6	0.1	-0.1	-	Espinoza_2016	Stassun_2017	Mayo_2018
2	1.5804	0	3026	0.15	0.22	GJ 1214 b	1	6.26	0.86	-0.86	2.85	0.2	-0.2	0.15	Harpsoe_2013	Harpsoe_2013	Charbonneau_2009
3	9.4909	0	5175	0.75	0.74	HD 97658 b	1	7.86	0.73	-0.73	2.4	0.1	-0.1	-0.3	Dragomir_2013	Stassun_2017	Van_Grootel_2014
4	9.1205	0	5089	0.78	0.72	HIP 116454 b	1	11.82	1.33	-1.33	2.5	0.1	-0.1	-	Vanderburg_2015	Stassun_2017	
5	10.0545	0	3896	0.6	0.56	K2-3 b	1	6.6	1.1	-1.1	2.29	0.23	-0.23	-	Damasso_2018	Damasso_2018	Sinukoff_2016
6	13.8638	0	5010	0.74	0.71	K2-110 b	1	16.7	3.2	-3.2	2.6	0.1	-0.1	-	Osborn_2017	Osborn_2017	Mayo_2018
7	32.9396	0	3457	0.36	0.41	K2-18 b	1	7.96	1.91	-1.91	2.38	0.22	-0.22	0.12	Cloutier_2017	Cloutier_2017	Benneke_2017
8	45.2943	1	5599	0.91	1.09	K-10 c	1	17.2	1.9	-1.9	2.35	0.1	-0.1	-	Dumusque_2014	Berger_2018	Weiss_2016
9	16.1457	0	4909	0.81	0.73	K-102 e	1	8.93	2	-2	2.43	0.1	-0.2	0.18	Marcy_2014	Berger_2018	Morton_2016
10	13.5708	0	5860	1	1.04	K-106 c	1	10.44	3.2	-3.2	2.5	0.32	-0.32	-	Marcy_2014	Marcy_2014	Morton_2016
11	43.8444	0	5860	1	1.04	K-106 e	1	11.17	5.8	-5.8	2.56	0.33	-0.33	-	Marcy_2014	Marcy_2014	Morton_2016
12	16.092	0	5787	1.02	1	K-131 b	1	16.13	3.5	-3.5	2.4	0.2	-0.2	0.12	Marcy_2014	Marcy_2014	Morton_2016
13	9.287	0	5520	0.936	0.88	K-19 b	1	8.4	1.6	-1.5	2.3	0.1	-0.1	-	Malavolta_2017	Berger_2018	Ballard_2011
14	77.6113	0	5508	0.95	0.89	K-20 d	1	10.07	3.97	-3.7	2.5	0.1	-0.1	0.07	Buchhave_2016	Berger_2018	Gautier_2012
15	6.2385	1	6285	1.19	1.36	K-25 b	1	9.6	4.2	-4.2	2.73	0.12	-0.11	-	Marcy_2014	Berger_2018	Hadden_2014
16	10.5738	0	5678	1.03	1.08	K-454 b	1	6.84	1.4	-1.4	2.06	0.1	-0.2	0.27	Gettel_2016	Berger_2018	Stassun_2017
17	9.67395	1	5235	0.88	0.85	K-48 c	1	14.61	2.3	-2.3	2.56	0.1	-0.1	0.17	Marcy_2014	Berger_2018	Hadden_2014
18	5.39875	0	5789	1.08	1.26	K-68 b	1	8.3	2.2	-2.4	2.32	0.1	-0.1	0.12	Gilliland_2013	Berger_2018	van_Eylen_2015
19	16.2385	0	5751	1	0.95	K-96 b	1	8.46	3.4	-3.4	2.64	0.1	-0.24	0.04	Marcy_2014	Berger_2018	Morton_2016
20	4.754	0	4791	0.75	0.75	K-113 b	1	11.7	4.2	-4.2	2.16	0.1	-0.2	0.05	Marcy_2014	Berger_2018	Morton_2016
21	10.8541	0	5508	0.95	0.89	K-20 c	1	12.75	2.17	-2.24	2.87	0.19	-0.13	0.07	Buchhave_2016	Berger_2018	Gautier_2012

2	3.47175	0	4975	0.83	0.78	EPIC246471	1	9.68	1.21	-1.37	2.59	0.06	-0.06	0.0	Palle_2018	Palle_2018
2	50.81895	0	5372	0.88	0.85	K2-263 b	1	14.8	3.1	-3.1	2.41	0.12	-0.12	-	Mortier_2018	Mortier_2018
3	6.2682	0	6037	1.094	1.10	$\pi$ Mensae c	1	4.82	0.84	-0.86	2.14	0.044	-0.044	0.08	Huang_2018	Huang_2018
4															Gandolfi_2018	
<b>RV planets 3.0-4.0 Earth radii</b>																
1	29.8454	0	5261	0.86	0.86	HD 3167 c	1	9.8	1.3	-1.24	3.01	0.42	-0.28	0.04	Christiansen_2017	Christiansen_2017
2	10.9542	1	5466	0.956	0.90	K-88 b	1	8.7	2.5	-2.5	3.82	0.17	-0.34	0.2	Nesvorný_2013	Berger_2018
3	2.50806	0	4728	0.81	0.74	K-94 b	1	10.84	1.4	-1.4	3.19	0.14	-0.25	0.34	Marcy_2018	Berger_2018
4	11.5231	0	5654	1.08	1.41	K-95 b	1	13	2.9	-2.9	3.14	0.14	-0.13	0.3	Marcy_2018	Berger_2018
5	7.138048	0	4975	0.83	0.78	EPIC246471	1	15.68	2.28	-2.13	3.53	0.08	-0.08	0.0	Palle_2018	Palle_2018
6	1.33735	0	5143	0.75	0.84	NGTS-4 b	1	20.6	3.0	-3.0	3.18	0.26	-0.26	-	West_2018	West_2018
														0.28		
<b>RV planets 4.0-8.0 Earth radii</b>																
1	3.33665	0	3600	0.51	0.54	GJ 3470 b	1	13.9	1.5	-1.5	4.57	0.18	-0.18	0.18	Awlphan_2016	Awlphan_2016
2	2.849272	0	4300	0.688	0.71	K2-55 b	1	43.13	5.98	-5.8	4.41	0.32	-0.28	0.37	Dressing_2018	Dressing_2018
3	3.185315	0	4985	0.85	0.81	HATS-7 b	1	38.139	3.81396	-3.81396	6.311	0.516	-0.381	0.25	Bakos_2015	Bakos_2015
4	11.81399	0	5499	1.12	1.66	HD 89345 b	1	35.7	3.3	-3.3	6.86	0.14	-0.14	0.45	Van_Eylen_2018	Van_Eylen_2018
5	4.73401	0	5474	1.12	1.75	K2-108 b	1	59.4	4.4	-4.4	5.28	0.54	-0.54	0.33	Petigura_2017	Petigura_2017
6	8.99213	0	5275	0.86	0.84	K2-32 b	1	16.5	2.7	-2.7	5.13	0.28	-0.28	-	Petigura_2017	Petigura_2017
7	10.13675	0	6120	1.07	1.31	K2-98 b	1	32.2	8.1	-8.1	4.3	0.3	-0.2	-0.2	Barragan_2016	Barragan_2016
8	3.836169	0	4910	0.84	0.76	WASP-156 b	1	40.682	3.1783	-2.86047	5.717	0.224	-0.224	0.24	Demangeon_2017	Demangeon_2017
9	4.88782	0	4708	0.88	0.77	HAT-P-11 b	1	23.4	1.5	-1.5	4.36	0.06	-0.06	0.31	Yee_2018	Yee_2018
10	3.21346	0	5857	1.22	1.49	K-4 b	1	24.472	3.814	-3.814	4.076	0.181	-0.17	0.17	Borucki_2010	Berger_2018
11	6.21229	0	5080	0.88	0.77	CoRoT-8 b	1	69.92	9.53	-9.53	6.39	0.22	-0.22	0.3	Borde_2010	Southworth_2011

1	4.23452	0	5079	1.12	0.87	HAT-P-26 b	1	22.248	6.3566	-6.3566	7.062	0.448	-0.448	-	Stassun_2017	Stassun_2017	Stevenson_2016
2							1	1						0.04			
1	7.64159	1	5441	0.97	0.91	K-18 c	1	17.3	1.9	-1.9	4.263	0.183	-0.172	0.19	Cochran_2011	Berger_2018	Hadden_2014
3				6													
1	14.85888	1	5441	0.97	0.91	K-18 d	1	16.4	1.4	-1.4	5.16	0.219	-0.215	0.19	Cochran_2011	Berger_2018	Hadden_Lithwick_2017
4				6													
1	12.7204	1	6285	1.19	1.36	K-25 c	1	24.6	5.7	-5.7	5.134	0.216	-0.414	-	Marcy_2014	Berger_2018	Hadden_2014
5													0.04				
1	2.643883	0	3416	0.47	0.45	GJ 436 b	1	21.36	0.2	-0.2	4.191	0.109	-0.109	0.02	Trifonov_2018	Turner_2016	Dressing_2018
6		12		5											Maciejewski_2014		
TTV planets <8.0 Earth radii																	
1	10.3039	1	5763	1.04	1.09	K-11 b	0	2.78	0.64	-0.66	1.916	0.095	-0.081	-	Bedell_2017	Berger_2018	
2				1									0.04				
2	13.0241	1	5763	1.04	1.09	K-11 c	0	5	1.3	-1.35	3.006	0.138	-0.124	-	Bedell_2017	Berger_2018	
3				1									0.04				
3	22.6845	1	5763	1.04	1.09	K-11 d	0	8.13	0.67	-0.66	3.338	0.158	-0.138	-	Bedell_2017	Berger_2018	
4				1									0.04				
4	31.9996	1	5763	1.04	1.09	K-11 e	0	9.48	0.86	-0.88	4.3	0.188	-0.309	-	Bedell_2017	Berger_2018	
5				1									0.04				
5	46.6888	1	5836	1.04	1.09	K-11 f	0	2.53	0.49	-0.45	2.627	0.124	-0.116	-	Bedell_2017	Berger_2018	
6				1									0.04				
6	12.28	1	3884	0.54	0.60	K-26 b	0	5.12	0.65	-0.61	3.197	0.097	-0.093	-	Jontof_Hutter_2016	Berger_2018	
7				7									0.13				
7	17.2559	1	3884	0.54	0.60	K-26 c	0	6.2	0.65	-0.65	2.982	0.278	-0.229	-	Jontof_Hutter_2016	Berger_2018	
8				7									0.13				
8	10.3384	1	5378	0.98	0.74	K-29 b	0	4.51	1.41	-1.47	2.564	0.16	-0.141	-	Jontof_Hutter_2016	Berger_2018	
9				9									0.04				
9	13.2884	1	5378	0.98	0.74	K-29 c	0	4	1.23	-1.29	2.396	0.151	-0.133	-	Jontof_Hutter_2016	Berger_2018	
1				9									0.04				
1	29.33434	1	5454	0.99	0.83	K-30 b	0	8.8	0.6	-0.5	1.837	0.124	-0.171	0.18	Hadden_Lithwick_2017	Berger_2018	
0				4													
1	13.83989	1	5979	1.07	1.65	K-36 b	0	3.9	0.2	-0.22	1.514	0.082	-0.071	-0.2	Hadden_Lithwick_2017	Berger_2018	
1				2													
1	16.23855	1	5979	1.07	1.65	K-36 c	0	7.5	0.3	-0.3	3.688	0.165	-0.153	-0.2	Hadden_Lithwick_2017	Berger_2018	
2				2													
1	45.154	1	5674	1.04	0.83	K-51 b	0	2.3	1.7	-1.6	6.559	0.316	-0.297	null	Hadden_Lithwick_2017	Berger_2018	
3				3													

1	7.1334	1	5834	1.041	1.46	K-60 b	0	3.7	0.6	-0.6	1.845	0.339	-0.117	-	Hadden_Lithwick_	Berger_2018
4				9									0.09		2017	
1	8.9187	1	5834	1.041	1.46	K-60 c	0	2	0.3	-0.5	2.062	0.378	-0.131	-	Hadden_Lithwick_	Berger_2018
5				9									0.09		2017	
1	11.8981	1	5834	1.041	1.46	K-60 d	0	3.9	0.7	-0.6	1.901	0.495	-0.17	-	Hadden_Lithwick_	Berger_2018
6				9									0.09		2017	
1	27.4029	1	6389	1.17	1.34	K-79 c	0	5.9	1.9	-2.3	3.684	0.181	-0.165	-	Jontof_Hutter_20	Berger_2018
7				5									0.07		14	
1	52.0902	1	6389	1.17	1.34	K-79 d	0	6	2.1	-1.6	7.251	0.34	-0.322	-	Jontof_Hutter_20	Berger_2018
8				5									0.07		14	
1	81.0659	1	6389	1.17	1.34	K-79 e	0	4.1	1.2	-1.1	3.054	0.76	-0.186	-	Jontof_Hutter_20	Berger_2018
9				5									0.07		14	
2	7.05246	1	4665	0.73	0.68	K-80 b	0	6.93	1.05	-0.7	2.592	0.116	-0.11	0.04	MacDonald_2016	Berger_2018
0				1									3			
2	9.52355	1	4665	0.73	0.68	K-80 c	0	6.74	1.23	-0.86	2.594	0.121	-0.109	0.04	MacDonald_2016	Berger_2018
1				1									3			
2	3.07222	1	4665	0.73	0.68	K-80 d	0	6.75	0.69	-0.51	1.561	0.081	-0.064	0.04	MacDonald_2016	Berger_2018
2				1									3			
2	4.64489	1	4665	0.73	0.68	K-80 e	0	4.13	0.81	-0.95	1.602	0.081	-0.069	0.04	MacDonald_2016	Berger_2018
3				1									3			
2	10.4208	1	5559	0.91	0.93	K-307 b	0	8.8	0.9	-0.9	2.85	0.274	-0.208	0.19	Hadden_Lithwick_	Berger_2018
4				6									2017			
2	13.0729	1	5559	0.91	0.93	K-307 c	0	3.9	0.7	-0.7	2.569	0.279	-0.211	0.19	Hadden_Lithwick_	Berger_2018
5				6									2017			
2	66.0634	1	5989	1.08	1.01	K-289 d	0	4	0.9	-0.9	2.668	0.17	-0.17	0.05	Rowe_2014	Rowe_2014
6				3									3			
2	1.510870	1	2559	0.08	0.12	TRAPPIST-1 b	0	1.22	0.15	-0.15	1.121	0.031	-0.032	0.04	Demory_2018	Grimm_2018
7	81															Delrez_2018
2	2.421823	1	2559	0.08	0.12	TRAPPIST-1 c	0	1.24	0.15	-0.15	1.095	0.03	-0.031	0.04	Demory_2018	Grimm_2018
8	3															Delrez_2018
2	4.04961	1	2559	0.08	0.12	TRAPPIST-1 d	0	0.37	0.04	-0.04	0.784	0.023	-0.023	0.04	Demory_2018	Grimm_2018
9																Delrez_2018
3	6.099615	1	2559	0.08	0.12	TRAPPIST-1 e	0	0.66	0.079	-0.075	0.91	0.026	-0.027	0.04	Demory_2018	Grimm_2018
0																Delrez_2018
3	9.20669	1	2559	0.08	0.12	TRAPPIST-1 f	0	0.97	0.08	-0.078	1.046	0.029	-0.03	0.04	Demory_2018	Grimm_2018
1																Delrez_2018
3	12.35294	1	2559	0.08	0.12	TRAPPIST-1 g	0	1.27	0.098	-0.095	1.148	0.032	-0.033	0.04	Demory_2018	Grimm_2018
2																Delrez_2018
3	18.767	1	2559	0.08	0.12	TRAPPIST-1 h	0	0.37	0.049	-0.056	0.773	0.027	-0.026	0.04	Demory_2018	Grimm_2018
3																Delrez_2018

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