1 Supporting Information 2 Unravelling metal speciation in the microenvironment surrounding 3 phytoplankton cells to improve predictions of metal bioavailability 4 Fengjie Liu^{1*}, Qiao-Guo Tan², Dominik Weiss^{1,3}, Anne Crémazy⁴, Claude Fortin⁵, Peter 5 G.C. Campbell⁵ 6 ¹Department of Earth Science and Engineering, Imperial College London, South 7 Kensington Campus, London, SW7 2AZ, United Kingdom 8 ²Key Laboratory of the Coastal and Wetland Ecosystems, Ministry of Education, 9 College of Environment and Ecology, Xiamen University, Xiamen, Fujian 361102, 10 China 11 ³Department of Civil and Environmental Engineering, Princeton University, NJ 08540, 12 United States 13 ⁴Department of Biological Sciences, University of New Brunswick, Saint John (NB), 14 E2L 4L5, Canada 15 ⁵Institut national de la Recherche scientifique, Centre Eau Terre Environnement, 490 16 de la Couronne, Québec (Québec), G1K 9A9, Canada 17

18 *Corresponding author: fengjieliu2012@gmail.com

Contents

55 **Table S1. A summary of the** *in situ* **measured pH and O2 concentrations in the phycosphere of marine and freshwater**

56 **phytoplankton cells.**

57

58 Note, the Δ pH values in the Figure 1 were calculated as the differences between the bulk waters and the phycosphere, and we

59 didn't convert the numbers into Δ [H⁺].

60 **Table S2. Basic chemical composition of freshwater and seawater used for the**

61 **calculation of metal speciation.**

62 The concentrations of major cations/anions, trace metals, and DOC were taken from

63 several publications.¹¹⁻¹⁴ *The dissolved Sm concentrations refer to two studies.^{15, 16}

65 Note, data are also shown for bulk seawater at pH 8.0. Here, the average phycosphere

66 pH change (i.e., a 0.3-unit increase in the light, and a 0.1-unit decrease in the dark) was

67 used to calculate the speciation change in the presence of low and high concentrations

68 of dissolved organic carbon (DOC).

- 69 **Table S4. Equilibrium modelled relative distribution (%) of free ion, total inorganic**
- 70 **complexes (-Inorg.) and total organic complexes (-Org.) of metals in the phycosphere**
- 71 **of freshwater phytoplankton cells.**

72 Note, data are also shown for bulk freshwater at pH 7.0. Here, the average phycosphere

73 pH change (i.e., a 1.3-unit increase in the light, and an assumed 0.5-unit decrease in the

- 74 dark) was used to calculate the speciation change in the presence of low and high
- 75 concentrations of dissolved organic carbon (DOC).

Figure S1. Fold changes in the calculated oxidation rate of Fe(II), Mn(II) and Cu(I) in the

phycosphere of freshwater or marine algae in the light (red) or in the dark (grey), in

comparison to the calculated rates in bulk solutions. The data are shown on a log scale,

81 and the dashed line indicates no difference between the phycosphere and bulk water.

82 The details of the calculation are given in the Note S2.

 Figure S2. The relationships between the calculated free Pb2+ concentration in the bulk water (left panel) or in the phycosphere (right panel) and the measured Pb uptake rate

(*J***Pb) by the freshwater alga** *Chlamydomonas reinhardtii* **CPCC11 under constant light**

88 (100 µmol photons m⁻² s⁻¹). The data are replotted.¹⁷

101 102

108 uptake test was carried out under a low light intensity (< 20 μ mol photons m⁻² s⁻¹) and

109 the algae were cultured with NH_4^+ as the nitrogen source. The low light and the supply

110 of NH₄⁺ would likely result in an overall lower phycosphere pH than that in the bulk

111 water (see Figure S5 for an example).

112

113

114

115 **Figure S5. A significant pH decrease of the exposure solution by ammonium NH⁴ + -fed**

116 **algae** *Chlamydomonas reinhardtii* CPCC11 at 2 × 10⁷ cells mL⁻¹. The experimental details

117 can be found in our previous studies. $17, 18$ The present data are not shown in those

- 118 papers.
- 119

120 **Note S1. Potential influences of several physical/chemical/biological factors, other** 121 **than pH and O2, on metal behaviour in the phycosphere.**

122 In addition to pH changes, redox dynamics in the phycosphere might differ from those in 123 the bulk water. For example, phytoplankton release various reactive oxygen species 124 (ROS) such as H₂O₂ and superoxide into the extracellular environment^{20, 21}. Such solutes 125 have slower diffusive rates (0.6-1.4 \times 10⁻⁹ m² s⁻¹ for H₂O₂)²² than OH⁻/H⁺ (5.3-9.3 \times 10⁻⁹ 126 $m²$ s⁻¹), and they could hence be locally enriched. The ROS-enriched micro-environment 127 would facilitate redox transformation of redox-sensitive metals (e.g., Fe, Cu, and Mn) 128 and ligands (e.g., thiol-containing ligands) before metals are taken up.^{23, 24} At present, 129 the level and dynamic of ROS in the phycosphere remain unknown, let alone their 130 potential influence on the local metal speciation.

131

132 Nutritional status of algae might alter chemical conditions in the phycosphere as well.

133 For example, Milligan et al. 3 observed that both Fe and carbon availability influenced the

134 phycosphere pH of the marine diatom *Thalassiosira weissflogii*; the enhancement in

135 phycosphere pH of Fe-replete cells was 0.3-unit higher than for Fe-limited cells.

136 Moreover, the phycosphere pH change in high $pCO₂$ treatments was not as significant as 137 in low $pCO₂$ treatments.³

138

139 Nitrogen sources for algae might be another factor in influencing both phycosphere pH

140 and ROS levels. Lavoie et al. ²⁵ modelled the effect of different N sources on the

141 phycosphere pH of freshwater phytoplankton, and their results suggested that

142 \sim NO₃-grown algae would have a higher phycosphere pH than NH₄⁺-grown cells. On the

143 other hand, studies with marine algae found that extracellular production of ROS was

144 associated with nitrogen nutrition; nitrate could stimulate more extracellular ROS than

145 ammonium in some algae species.²⁶

146

147 Extracellular polymeric substances (EPS) secreted by certain phytoplankton species such

S16

148 as *Phaeocystis*²⁷ might alter the chemical behaviour of metals in the phycosphere. For 149 example, the residence time of metals would be longer in the phycosphere of algae 150 having an abundance of EPS than for those without EPS, since diffusion coefficients (D) 151 of metal ions and complexes are lower in solutions of higher viscosity, η, such as in the 152 phycosphere of algae that produce EPS (i.e., due to the inverse relationship between D 153 and η in the Stokes-Einstein equation).

154

155 The local chemistry in the phycosphere might also be influenced by the microbial

156 community living in this micro-environment. In nature, a diversity of smaller

157 microorganisms such as bacteria and viruses frequently cluster near, or attach to,

158 phytoplankton.^{28, 29} Thus, the metabolism of these microbes might alter phycosphere pH

159 and redox potential (e.g., $CO₂$ release and $O₂$ consumption during respiration), and

160 extracellular release of ROS from these bacteria²¹ might influence chemical reactions

- 161 between metals and ligands. However, the potential influence of the microbes on local
- 162 metal speciation has been little examined.
- 163

164 Other than these biological factors, water turbulence can increase interfacial mixing of 165 nutrients and metabolites to and from an organism (> 100 μm) between bulk waters and 166 the biota. However, the effect from turbulence on the phycosphere conditions for 167 micro-organisms of < 100 μm in diameter would be weak, because modelling work has 168 suggested that even high levels of turbulence (e.g., at a dissipation rate 10^{-6} W kg⁻¹)

169 produce a tiny increase (around 2%) of the flux in the tiny interfacial layer.³⁰

170

171 Finally, total metal concentrations as well as competing metal interactions might not be 172 the same as those in the bulk water. For example, a higher pH and $[O_2]$ in the

173 phycosphere could result in local precipitation of Fe and $Mn^{9, 31}$ and hence alter source

174 and bioavailability of the metals to algal cells. With respect to metal-metal interactions,

175 even if there was a simultaneous enhancement of free metal ions (e.g., Zn^{2+} and Cd^{2+}) in

176 the phycosphere, this would not necessarily result in an increase in uptake of both

177 metals since they might compete with each other for the same membrane transporter.³²

178

179 **Note S2. Calculation of the differences in oxidation rates of Fe(II), Cu(I) and Mn(II)** 180 **between the phycosphere and bulk water.**

181 Fe oxidation kinetics in both freshwater and marine waters follow the rate law: 33, 34

182 Fe(II) oxidation rate =
$$
\frac{-d[Fe(II)]}{dt} = k_{Fe}[Fe(II)][O_2][OH^-]^2
$$

183

184 Thus, the fold change in the Fe(II) oxidation rate can be calculated by the following

185 equation, and it can be calculated based upon the numbers in Table S5:

186 Fe(II) oxidation rate_{phycosphere}
$$
= \frac{[O_2][OH^-]^2_{\text{phycosphere}}}{[O_2][OH^-]^2_{\text{bulk water}}}
$$

187

188 Similarly, Mn(II) oxidation kinetics in seawater follow the rate law shown below³⁵ and it

189 can be calculated based upon the numbers in the Table S5:

190
$$
\text{Mn(II) oxidation rate} = \frac{-d[\text{Mn(II)}]}{dt} = k_{Mn}[\text{Mn(II)}][0_2][0_1 - 1]^2
$$

191

192
$$
\frac{Mn(II) \text{ oxidation rate}_{\text{phycosphere}}}{Mn(II) \text{ oxidation rate}_{\text{bulk water}}} = \frac{[O_2][OH^-]^2_{\text{phycosphere}}}{[O_2][OH^-]^2_{\text{bulk water}}}
$$

193

194 Thus, the fold change in the Mn(II) oxidation rate between phycosphere and bulk water

195 is the same as that of Fe(II).

196

197 Cu oxidation kinetics in seawater follow the rate law shown below³⁶ and it can be

198 calculated based upon the numbers in the Table S5:

199

200 Cu(l) oxidation rate =
$$
\frac{-d[Cu(I)]}{dt} = k^* c_u[Cu(I)][O_2]
$$

201 The k^*_{α} (apparent oxidation rate constant, M⁻¹ s⁻¹) was calculated based on the

202 measured Cu(I) oxidation kinetics at different pH levels.³⁷ The k^*_{Cu} values varied with

203 solution pH. Specifically, the k^*_{cu} values (M⁻¹ s⁻¹) used for the seawater calculation are

- 204 2.45 at pH 7.9, 2.58 at pH 8.0, and 3.02 at pH 8.32, while in freshwater they are 5.0 at
- 205 pH 6.5, 10.0 at pH 7.0, and 22.9 at pH 8.28.
- 206

207 **Table S5. Values used for the calculation of relative changes in metal oxidation rates in** 208 **the phycosphere, in comparison to bulk waters.**

209 **Note S3. Estimation of the amount of Cu(I) oxidation and Cu complexation in the** 210 **phycosphere.**

- 211 When the residence time of Cu(I) in the phycosphere is sufficiently long, all Cu(I) can be
- 212 oxidized by oxidants such as $O₂$ before reacting with membrane-bound Cu reductases.
- 213 However, since the phycosphere of micro-algae is a thin layer of micro-meter
- 214 dimensions, the local residence time might be very short and hence little Cu(I) would be
- 215 oxidized by O_2 within the phycosphere. Here, we calculated the residence time of Cu(I)
- 216 and the percentage of total Cu(I) being oxidized by O_2 in the phycosphere of two marine
- 217 diatoms *Thalassiosira pseudonana* and *Thalassiosira oceanica*.
- 218 To calculate the residence time of Cu(I) in the phycosphere, we follow:

219 *phycosphere residue*
$$
time = \frac{[Cu(I) \; concentration] \times [phycosphere \; volumn]}{[Cu(I) \; uptake \; rate] \times [cell \; surface \; area]}
$$

220

221 For *T. pseudonana*, cell radius = $2.3 \mu m$,³⁸ theoretical phycosphere thickness = cell radius

222 ³⁹ = 2.3 μm; for *T. oceanica,* cell radius = 3.3 μm,³⁸ theoretical phycosphere thickness =

223 3.3 μ m. The ratio between $\left[Cu(I) \right]$ concentration] in bulk seawater and $\left[Cu(I) \right]$ uptake rate]

224 were taken from a short-term algal Cu uptake study, 38 assuming that

$$
\frac{[Cu(I) \; concentration]}{[Cu(I) \; uptake \; rate]} = \frac{[total \; Cu \; concentration]}{[total \; Cu \; uptake \; rate]}
$$

226 since data for algal Cu(I) uptake are rare in the literature. The calculated residence time

227 of Cu ranged from 230 s to 3000 s in the phycosphere of the two species.

228

229 To calculate the percentage of total Cu(I) being oxidized by O_2 in the phycosphere, 230 we considered:

231
$$
0xidized Cu(I)_t\% = 100\% - e^{-\frac{K_{Cu}}{t}} \times 100\%
$$

- 232 *t* = phycosphere residence time, and $K_{Cu} = k^* c_u \times$ [phycosphere O₂]. Specifically,
- 233 $k \cdot c_u = 3 \text{ M}^{-1} \text{ s}^{-1}$ (at pH 8.3)³⁷ and [phycosphere O₂] = 0.4 mM (Table S5). The
- 234 percentage of total Cu(I) being oxidized by O_2 in the phycosphere of two marine diatoms

236

237 **Table S6. Calculation of Cu residence time and percentage of total Cu(I) being**

238 **oxidized by O² in the phycosphere.**

239 Note, the Cu uptake data were taken from the literature.³⁸

240

241 Note, the calculated residence time of Cu in the phycosphere above is much longer than

242 the time required for the metal diffusion from the bulk water to the cell surface in the

243 diffusive boundary layer. Specifically, the diffusive boundary layer (i.e., a concentration

244 gradient of Cu) would form when the algal metal uptake is very rapid but the ambient

245 Cu concentration is relatively low.

246 To calculate the diffusion time, we can follow

247 (https://www.physiologyweb.com/calculators/diffusion_time_calculator):

248 Diffusion time $\approx \frac{\text{thickness of diffusive boundary layer}^2}{2 \times \text{diffusion coefficient}}$ $2 \times$ dif fusion coef ficient

249 Here, the thickness of diffusive boundary layer was assumed comparable to the cell

- 250 radius (i.e., around 2 to 3 μ m for the two diatoms), and the diffusion coefficient of Cu
- 251 ion and its complexes at the order of 10^{-6} cm² s^{-1 40} Thus, the diffusion time of Cu ions in
- 252 the phycosphere was around 8 ms.

253

254 Similarly, we calculated whether Cu complexation by DOC is fast enough to allow Cu

255 speciation to reach equilibrium in the phycosphere. Specially, we follow:

256
$$
\frac{[Cu]_t}{[Cu]_0}\% = 100\% - e^{-\frac{K_{Cu-DOC}}{t}} \times 100\%
$$

257 $t =$ phycosphere residence time, $[Cu]_0 =$ the initial Cu concentration, and $K_{Cu-DOC} =$

258 k_{Cu-DOC} \times [phycosphere DOC]. Specifically, k_{Cu-DOC} = 6 \times 10³ M⁻¹ s⁻¹ and

259 [phycosphere DOC] = 80 μ M (around 1 mg L⁻¹).⁴¹ The percentage of Cu complexation in

260 the phycosphere of the two algae was 100%.

261

262 **Note S4. Definition of Root Mean Square Logarithmic Error (RMSLE).**

263 Root mean square error (RMSE) measures the difference between model fittings and 264 observed values and is a measure of goodness of model fitting. In the Figures 3 and S4, 265 where both the metal ion concentrations and the uptake rates spanned several orders 266 of magnitude, the logarithmic version of RMSE, i.e., RMSLE, was thus used instead to 267 achieve a more balanced measure of model performance.

268

$$
\text{RMSLE} = \sqrt{\frac{1}{n} \sum_{i=1}^{n} (\log y_i - \log \hat{y}_i)^2}
$$

270 where y_i is the *i*th observed value; \hat{y}_i is the *i*th value predicted from the fitting; *n* is the 271 number of data points.

272

273
$$
RMSE = \sqrt{\frac{1}{n} \sum_{i=1}^{n} (y_i - \hat{y}_i)^2}
$$

274 Note, the RMSE was used to assess the goodness of fit between algal Pb uptake and Pb²⁺ 275 concentration (i.e., Figure S2), since the uptake test was carried out in a narrow range of 276 ambient Pb^{2+} concentrations.

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