

1 **Supporting Information for**

2 **Unprecedented ambient sulphur trioxide (SO₃) detection: possible formation mechanism and**
3 **atmospheric implications**

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39 **Text S1. Sampling site**

40 Our field measurements were conducted at BUCT (Beijing University of Chemical Technology) sampling
41 site (39.94° N, 116.30° E), which was located on the west campus of BUCT¹. This site was located at the
42 roof of a teaching building, which is approximately 15 m above the ground level. Around 130 m to the north
43 and 550 m to the west are Zizhuyuan Road and West Third Ring Road, respectively. The “West Third Ring
44 Road” is one of the main “Ring” roads in Beijing. Besides the influence of traffic, this site is also affected by
45 local commercial and residential activities. Therefore, the BUCT monitoring site is representative of an urban
46 site.

47 **Text S2. The nitrate-CI-APi-LTOF mass spectrometer**

48 A nitrate-based Chemical Ionization - Atmospheric Pressure interface -Long- Time-of-Flight (CI-APi-LTOF,
49 Aerodyne Research Inc, USA and ToFwerk AG, Switzerland) mass spectrometer was deployed to detect SO₃
50 and gas-phase sulfuric acid. The CI-APi-LTOF consists of an optimized inlet for chemical ionization (CI-
51 inlet)^{2,3} and an APi-LTOF mass spectrometer with the mass resolving power of ~10000 Th/Th. Nitrate ions
52 (NO₃⁻·(HNO₃)_n, n=0,1 and 2) were used as reagent ions. The working principle of nitrate-CI-APi-LTOF has
53 been described in many previous studies^{2,4}.

54 In the charging part of CI-inlet, the nitrate ions are electrostatically pushed into ambient sample flow to react
55 with SO₃ and H₂SO₄. In the CI-inlet, the ion-molecule reaction time was ~200 ms⁴. Pure air originated from
56 a pure air generator (Aadco 737) was used as the sheath air. Ambient air was sampled into the CI-inlet through
57 a 3/4 inch stainless steel tube. A 0.8 L min⁻¹ flow from the mixed flow entered the APi-LTOF. Data of CI-
58 APi-LTOF were acquired at 5 s time resolution and analyzed with a MATLAB tofTools package⁵.

59 **Text S3. Detection of sulfuric acid with nitrate reagent ions**

60 The ion-molecule reactions of sulfuric acid with nitrate reagent ions can be described by the following
61 reaction^{2,4}:



63 where $n = 0, 1$ or 2 and $j = 0$ or 1 . Due to H₂SO₄ being a stronger acid, de-protonation occurs during its
64 collision with nitrate ions. Thus, H₂SO₄ molecules can be detected as de-protonated monomer ions and cluster
65 ions with HNO₃ in CI-APi-LTOF.

66 To take the variation in the total reagent ions into account, neutral sulfuric acid was quantified according to
67 the following equation:

68
$$[\text{H}_2\text{SO}_4] = C \times \frac{\sum_{n=0-1} (\text{HSO}_4^-) \cdot (\text{HNO}_3)_n}{\sum_{n=0-2} (\text{NO}_3^-) \cdot (\text{HNO}_3)_n} \quad \text{E (1)}$$

69 where C (in units of cm⁻³) is a calibration coefficient from in-situ calibration.

70 **Text S4. Calibration experiment for SO₃**

71 The calibration of SO₃ was implemented by introducing a known amount of gaseous SO₃ produced by the
72 reaction of SO₂ and OH radicals formed by UV photolysis of water vapour, which is similar to the method
73 for sulfuric acid calibration in the previous literature (i.e. Kürten et al., 2012)⁶. During the calibration
74 experiment, a 10 L·min⁻¹ N₂ flow, a 100 mL·min⁻¹ pure air flow, a 300 mL·min⁻¹ SO₂ flow and a set of 20 -

75 400 mL·min⁻¹ saturated water vapour flow were mixed together as the calibration sampling flow. Then, the
 76 mixed flow was exposed to 184.9 nm UV light to produce OH radicals which reacted with SO₂ to produce
 77 SO₃. The schematic of the experimental setup was shown in Figure S2. The UV lamp was turned on in an N₂
 78 environment at least one hour before the actual calibration measurement in order to achieve a stable light
 79 intensity. During the calibration, the box was flushed with a 1 - 2 L·min⁻¹ dry N₂ flow to avoid the absorption
 80 of UV light by O₂. Different levels of SO₃ were achieved by adjusting the flow of saturated water vapour,
 81 that is, adjusting OH radical concentrations. The theoretical generated SO₃ concentrations were calculated
 82 by Numerical Model which has been introduced in Kürten et al., 2012⁶. The time profiles of water content
 83 ([H₂O]) and normalized signals of sulfuric acid and SO₃ were exhibited in Figure S3. The correlation between
 84 normalized SO₃ signals measured by CI-APi-LTOF and SO₃ concentrations formed by photo-oxidation of
 85 SO₂ by OH radicals was depicted in Figure S4. After taking the diffusion loss of the sampling line into
 86 account, a calibration coefficient of 1.7 × 10¹⁰ molecule cm⁻³ was obtained. The diffusion loss was assumed
 87 as same as that of sulfuric acid.

88 Similar to the quantification of sulfuric acid², to regard the variation in the total reagent ions, SO₃ was
 89 quantified according to:

$$90 \quad [SO_3] = C \times \frac{[SO_3 \cdot NO_3^-]}{\sum_{n=0-2} (NO_3^-) \cdot (HNO_3)_n} \quad E (2)$$

91 where C (in units of cm⁻³) is a calibration coefficient of SO₃.

92 **Text S5. Calculations for collision rate coefficients**

93 Using ion-molecule collision parametrizations from two previous studies (Su & Bowers, 1973 and the Su &
 94 Chesnavich, 1982; note that these are the standard approaches used to estimate ion-molecule collision rates
 95 also in CIMS studies)^{7, 8}, with the dipole & polarisability of H₂SO₄ computed at the same level of theory
 96 (wB97xd/aug-cc-pVTZ; numerical values given below), the results for the collision rate coefficients (unit
 97 cm³ molecule⁻¹ s⁻¹) are as follows:

98 H₂SO₄ (mass 98)

99 Dipole moment 3.15093 Debye

100 Polarisability 35.9657 bohr³ = 5.330 Ånström³

101 Collision rate at 298.15 K with HNO₃·NO₃⁻ (mass 125) according to Su & Chesnavich 1982: 2.11E-09

102 Collision rate at 298.15 K with NO₃⁻ (mass 62) according to Su & Chesnavich 1982: 2.54E-09

103 Collision rate at 298.15 K with HNO₃·NO₃⁻ (mass 125) according to Su & Bowers 1973: 2.48E-009

104 Collision rate at 298.15 K with NO₃⁻ (mass 62) according to Su & Bowers 1973: 2.99E-009

105

106 SO₃ (mass 80)

107 Dipole moment 0 Debye

108 Polarisability 28.4649 bohr³ = 4.218 Ånström³

109 Collision rate at 298.15 K with HNO₃·NO₃⁻ (mass 125) according to Su & Chesnavich 1982: 6.88E-10

110 Collision rate at 298.15 K with NO₃⁻ (mass 62) according to Su & Chesnavich 1982: 8.13E-10

111 Collision rate at 298.15 K with HNO₃·NO₃⁻ (mass 125) according to Su & Bowers 1973: 1.79E-009

112 Collision rate at 298.15 K with NO₃⁻ (mass 62) according to Su & Bowers 1973: 2.11E-009

113

114 The main charging ions are $\text{HNO}_3 \cdot \text{NO}_3^-$ and NO_3^- . The corresponding ratios (collision rate of H_2SO_4 /
115 collision rate of SO_3) are 3.07, 3.12, 1.39 and 1.42, respectively. The newer ion-molecule collision rate
116 parametrization (Su & Chesnavich, 1982) thus predicts a difference of a factor of 3, whereas the older one
117 only predicts a difference of about a factor of 1.5. A non-polar molecule collides much slower with an ion
118 than a strongly polar molecule. The product ions $\text{H}_2\text{SO}_4 \cdot \text{NO}_3^-$ (binding Gibbs free energy of -32.6 kcal/mol
119 at the wB97xd/aug-cc-pVTZ level) and $\text{SO}_3 \cdot \text{NO}_3^-$ (binding Gibbs free energy -28.4 kcal/mol at the same
120 level) are very strongly bound and stable. The sensitivity of SO_3 could be less than that of H_2SO_4 by a factor
121 of 3. Thus, a factor of 3 difference would lead to an underestimation of SO_3 if the calibration factor for
122 H_2SO_4 was used to quantify SO_3 .

123 **Text S6. Quantum chemical calculations**

124 Quantum chemical calculations demonstrate that the $\text{SO}_3 \cdot (\text{NO}_3^-)$ cluster is very strongly bound compared to
125 the $\text{HNO}_3 \cdot (\text{NO}_3^-)$ cluster (Table S1). The difference in binding is over 10 kcal/mol both in electronic and free
126 energies, with the more rigorous coupled-cluster methods predicting a larger difference than the density
127 functional theory method used here. SO_3 molecules will thus be very efficiently charged by nitrate ions in a
128 nitrate-CI-API-LTOF instrument, as the charge transfer reaction $\text{HNO}_3 \cdot (\text{NO}_3^-) + \text{SO}_3 \rightarrow \text{SO}_3 \cdot (\text{NO}_3^-) + \text{HNO}_3$
129 is highly favourable. Furthermore, the thermal evaporation rate of $\text{SO}_3 \cdot (\text{NO}_3^-)$ clusters in the CI-inlet will be
130 negligible, and also the (non-thermal) fragmentation of the cluster in the ion optics of the instrument will be
131 considerably smaller than for example that of the $(\text{H}_2\text{SO}_4)_2 \cdot \text{HSO}_4^-$ cluster, which has binding energy
132 comparable to $\text{HNO}_3 \cdot (\text{NO}_3^-)$ ⁹. All of this supports the hypothesis that the instrument sensitivity toward SO_3
133 will be very high. The optimized structure of the $\text{SO}_3 \cdot (\text{NO}_3^-)$ cluster is shown in Figure S5. The strength of
134 the $\text{O}_3\text{S} \dots \text{ONO}_2^-$ interaction is reflected in the relatively short S...O distance.

135 The thermodynamics of the $\text{SO}_3 \cdot \text{H}_2\text{O} + \text{NO}_3^- \cdot (\text{HNO}_3) \rightarrow \text{SO}_3 \cdot (\text{NO}_3^-) \cdot \text{H}_2\text{O} + \text{HNO}_3$ reaction (R5) was
136 assessed at the wB97X-D/aug-cc-pVTZ level. Three different hydrogen bonding patterns (conformers) for
137 $\text{SO}_3 \cdot (\text{NO}_3^-) \cdot \text{H}_2\text{O}$ were assessed, with the H_2O molecule placed either close to the SO_3 moiety, the NO_3^- moiety,
138 or in a bridging position between the two. The latter structure, where H_2O H-bonds to both O-S and O-N
139 oxygen atoms, was found to be the lowest in free energy (at 298 K) (see Figure S6), though the differences
140 between conformers were fairly small (below 2 kcal/mol). By comparison to the results in Table S1, it is
141 likely that higher-level energy corrections (omitted here for computational reasons) would lead to an even
142 more negative (favourable) reaction free energy.

143 **Text S7. Computational details**

144 Molecular and cluster geometries were optimized, and harmonic vibrational frequencies were calculated,
145 with the $\omega\text{B97X-D}$ density functional¹⁰ and the aug-cc-pVTZ basis set¹¹, using the Gaussian 16 program
146 suite (Gaussian 16, Revision A.03)¹². Special care was taken in the optimizations to ensure that SO_3 and NO_3^-
147 had the correct D3h symmetry point group (with a rotational symmetry number of 6). Thermal and vibrational
148 zero-point contributions to enthalpies and entropies were calculated using the standard rigid rotor - harmonic
149 oscillator approximations. Single-point energy corrections were performed on top of the $\omega\text{B97X-D/aug-cc-}$
150 pVTZ structures using explicitly correlated coupled cluster theory; specifically RHF-RCCSD(T)-F12¹³⁻¹⁵
151 with the VDZ-F12 and VTZ-F12 basis sets¹⁶. The Molpro 2015.1 program was used for these calculations
152 (MOLPRO, version 2019.2, a package of ab initio programs, see <https://www.molpro.net>)¹⁷. With the smaller
153 basis set, the “F12a” variant was used, based on recommendations in the program manual. With the larger
154 basis set, both “F12a” and “F12b” energies were calculated. The differences between binding energies

155 computed with the two variants were minimal (around 0.01 kcal/mol or less), and even the difference between
156 the basis sets was less than 0.3 kcal/mol for both $\text{HNO}_3 \cdot (\text{NO}_3^-)$ and $\text{SO}_3 \cdot (\text{NO}_3^-)$. Final results are presented
157 using the RHF-RCCSD(T)-F12b/VTZ-F12 values.

158 **Text S8. PM_{2.5}, black carbon (BC), particulate sulfate, trace gases, meteorological parameters and**
159 **UVB measurements.**

160 Mass concentrations of PM_{2.5} were recorded by a TEOM (tapered element oscillating microbalances) monitor.
161 Non-haze and haze days were categorized as daily mean PM_{2.5} mass concentrations of <100 $\mu\text{g m}^{-3}$, and \geq
162 100 $\mu\text{g m}^{-3}$, respectively.

163 Mass concentrations of BC in PM_{2.5} were measured by the aethalometer (Magee AE33). PM_{2.5} particles were
164 continuously collected through the filter tape. Then the transmission of light through the filter tape containing
165 the sample was obtained at seven different wavelengths. AE33 calculates the instantaneous concentration of
166 optically absorbing aerosols from the rate of change of the attenuation of light transmitted through the
167 particle-laden filter¹⁸. The determination of black carbon concentration is based on the measurement of light
168 absorption on a filter loaded. In this study, aerosol particles are continually sampled on the filter, and the
169 optical attenuation is measured with high time resolution 1s. Optical attenuation (ATN) is measured on two
170 spots with different sample flows and the reference spot without the flow calculated as Equation 3:

171
$$ATN = -100 * \ln\left(\frac{I}{I_0}\right) \quad \text{E(3)}$$

172 where I and I_0 represent the detector intensity signal for the measurement spot and reference signal,
173 respectively. The factor 100 is there for convenience only¹⁷.

174 Airflow (F) is measured after the air passes the filter (Eq.3), lateral airflow in the optical chamber ζ has to be
175 taken into account:

176
$$F_{in} = F_{out} (1 - \zeta) \quad \text{E(4)}$$

177 where the value of ζ is determined by measuring input and output flow and was found to be in the 0.02–0.07
178 range.

179 Attenuation coefficient (b_{ATN}) and Absorption coefficient (b_{ABS}) can be calculated as Equations 5 and 6

180
$$b_{ATN} = \frac{S * \left(\frac{\Delta ATN}{100}\right)}{F * \Delta t} \quad \text{E(5)}$$

181
$$b_{ABS} = \frac{b_{ATN}}{C} \quad \text{E(6)}$$

182 where F is the volumetric flow and S is the filter surface area on which the aerosol particles are deposited, t
183 represents the time, and C is multiple scattering parameter²⁰.

184 BC concentration can be calculated using Equation 7:

185
$$BC = \frac{b_{ABS}}{\sigma_{air}} \quad \text{E(7)}$$

186 where σ_{air} is the mass absorption cross-section, here we obtain the mass absorption cross-section value is 7.19
187 $\text{m}^2 \text{g}^{-1}$ measured at 970 nm to convert the observed light attenuation to the mass concentration of BC²¹.

188 Combining Eqs. 5, 6 and 7, we obtain Eq.8 to calculate the concentration of BC:

$$BC = \frac{S * \left(\frac{\Delta ATN}{100}\right)}{F(1 - \zeta) * \sigma_{air} * C * (1 - k * ATN) * \Delta t} \quad E(8)$$

190 where k is a loading effect parameter.

191 The mass concentrations of non-refractory PM_{2.5} including sulfate concentration were measured by an online
 192 Time-of-Flight Aerosol Chemical Speciation Monitor (ToF-ACSM, Aerodyne Research Inc., USA) equipped
 193 with a cyclone to select PM_{2.5} particles. Trace gases including SO₂, O₃ and NO_x were recorded by Thermo
 194 analyzers (Model 43i, 49i and 42i, Thermo Scientific, USA). A weather station (Vaisala Inc., Finland) was
 195 deployed to measure meteorological parameters (ambient relative humidity (RH), temperature, wind speed,
 196 visibility). The UVB (280-315 nm) intensity was measured using CUV3 radiometer (Kipp & Zonen, USA).
 197 The mixing layer height (MLH) was determined with the enhanced single-lens ceilometers CL-51 (Vaisala
 198 Inc., Finland), which utilized the strobe laser lidar technique (910 nm) to measure the attenuated
 199 backscattering coefficient profiles. The Vaisala software product BL-VIEW was used to determine the
 200 mixing layer height by finding the position with the maximum negative gradient (-dβ/dx) in the attenuated
 201 backscattering coefficient profiles as the top of the mixing layer ²².

202 **Text S9. Sub-3 nm particles measurement**

203 The number size distributions of sub-3 nm clusters/particles were measured using an Airmodus A10 particle
 204 size magnifier (PSM) coupled with an Airmodus A20 condensation particle counter (CPC) ^{4,23}. The measured
 205 mobility diameter was from ~1.2 to 2.5 nm. Diethylene glycol and n-butanol were used as working fluids of
 206 the PSM and CPC, respectively. The sub-3 nm clusters/particles can grow up to ~90 nm by condensation of
 207 diethylene glycol vapour inside the PSM and then detected by the CPC⁴. To minimize the loss of
 208 clusters/particles and increase the sampling efficiency, a core sampling system and high flow rate (7.5 L
 209 min⁻¹) were utilized ¹. The time resolution of a full scan from 0.1 L min⁻¹ to 1.3 L min⁻¹ was 240 s.

210 **Text S10. Calculation of condensation sink**

211 Condensation sink (CS) describes the condensing vapour sink caused by the particle population ²⁴:

$$CS = 2\pi D \sum_{d_p} \beta_{m, d_p} d_p N_{d_p} \quad E(9)$$

213 where *D* is the diffusion coefficient of the condensing vapour (usually assumed to be sulfuric acid), and β_m,
 214 *d_p* is the transitional regime correction factor.

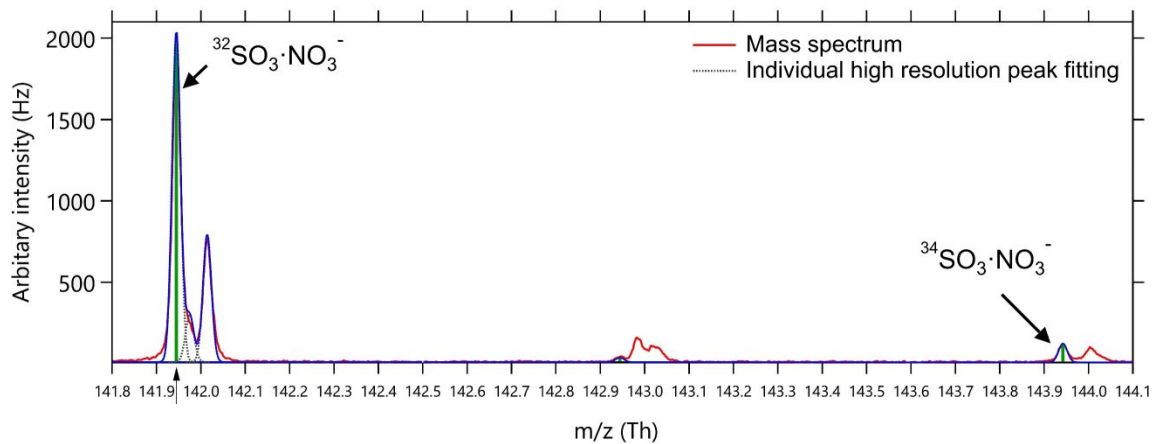
215 **Text S11. Source identification of SO₂ in winter**

216 In winter, the median concentrations of SO₂ exhibited similar diurnal trends as SO₃ (Figure 2A). A similar
 217 diurnal variation of SO₂ with an early morning peak has already been reported from another site in urban
 218 Beijing ²⁵. We also studied the evaluation of median mixing layer height (MLH) together with the diurnal
 219 trend of the median concentration of SO₂, SO₃ and UVB (Figure S8). The median MLH was merely 200-300
 220 m, and stable in the morning (~05:00 to ~08:30). Many studies have reported regional SO₂ is tightly linked
 221 with the combustion of sulfur-containing fuels and the smelting of sulfur-containing ores ²⁶⁻³². Together with
 222 stable weather conditions (Figure S8), the elevated SO₂ concentration during the early morning could mainly
 223 be attributed to local emissions (e.g. residential and industry emission) and transportation ³³.

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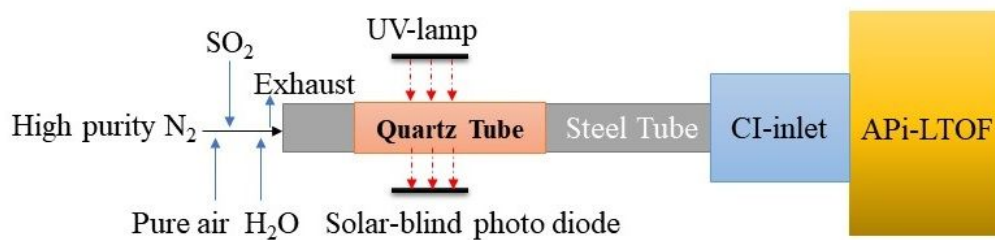
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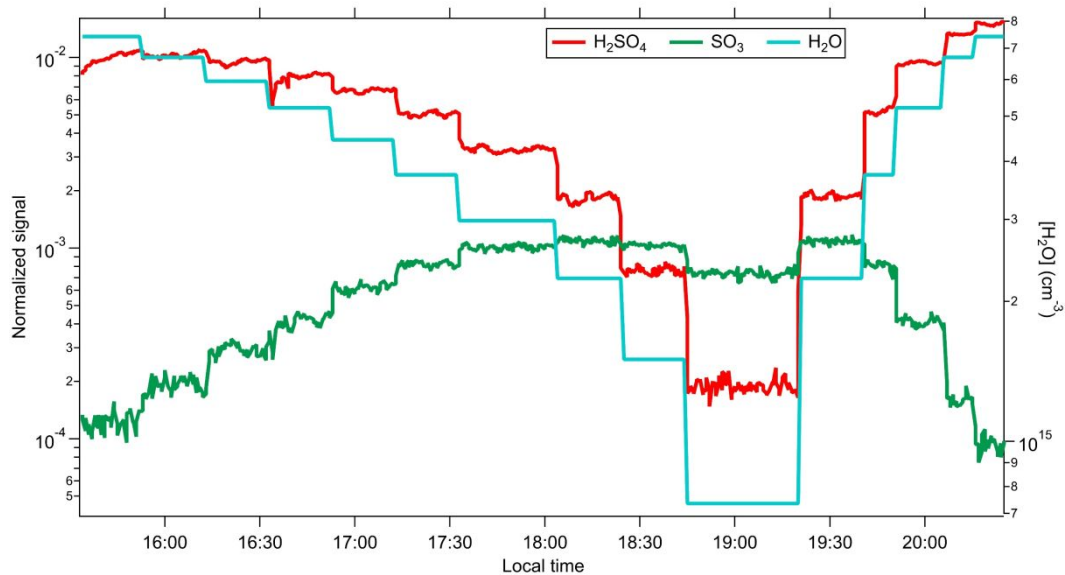
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Figure S1. High-resolution peak fitting of the peak $^{32}\text{SO}_3\cdot\text{NO}_3^-$ and its main isotope peak $^{34}\text{SO}_3\cdot\text{NO}_3^-$.



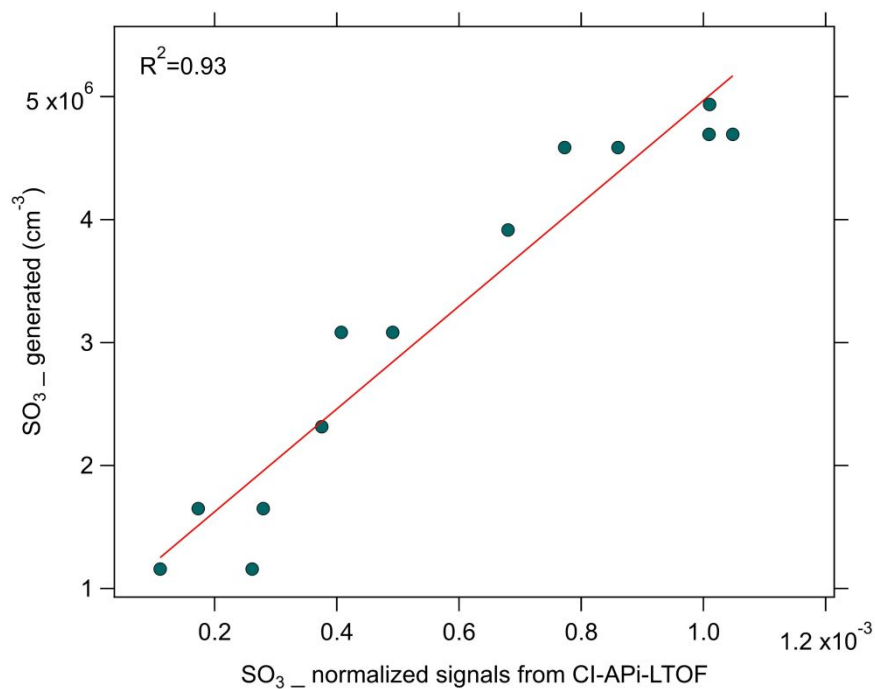
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Figure S2. The schematic of the calibration experiment setup.



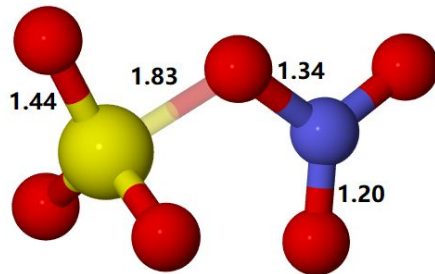
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Figure S3. Time series of normalized signals of H₂SO₄ and SO₃, and [H₂O] in the calibration experiment.



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Figure S4. The correlation between normalized SO₃ signals measured by CI-API-LTOF and SO₃ concentrations formed by the photo-oxidation of SO₂ by OH radicals. Our calibration experiment yielded a calibration coefficient of $1.7 \times 10^{10} \text{ cm}^{-3}$ for SO₃. This factor has taken diffusion loss of the sampling line of CI-API-LTOF into account.

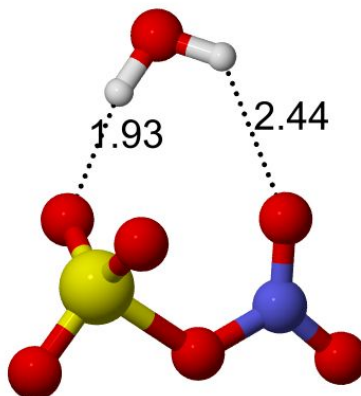


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251 **Figure S5.** Optimized structure of the $\text{SO}_3 \cdot (\text{NO}_3^-)$ cluster, at the $\omega\text{B97X-D/aug-cc-pVTZ}$ level. The S...O
 252 interaction between SO_3 and NO_3^- is shown as a covalent bond based on the relatively short interatomic
 253 distance, as well as the very strong binding energy. Key bond lengths are given in Ångström. Color coding:
 254 yellow=S, blue=N, red=O.

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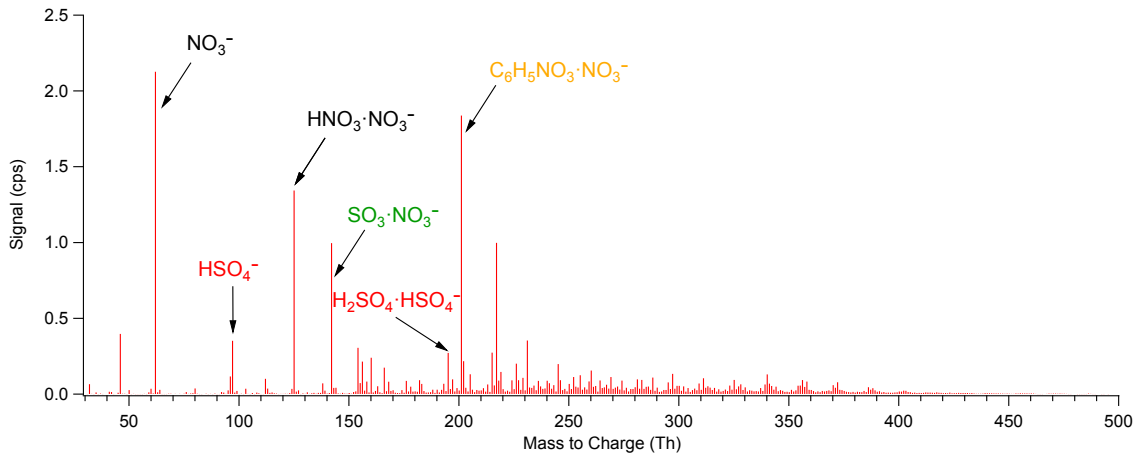
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258 **Figure S6.** Lowest free energy (at 298 K) structure found for $\text{SO}_3 \cdot (\text{NO}_3^-) \cdot \text{H}_2\text{O}$, at the wB97X-D/aug-cc-
 259 pVTZ level. Hydrogen bonds are indicated with dotted lines, and the corresponding distances are given in
 260 Ångström. Covalent bond lengths differ by less than 0.02 Å from those shown in Figure S5. Color coding:
 261 yellow=S, blue=N, red=O.

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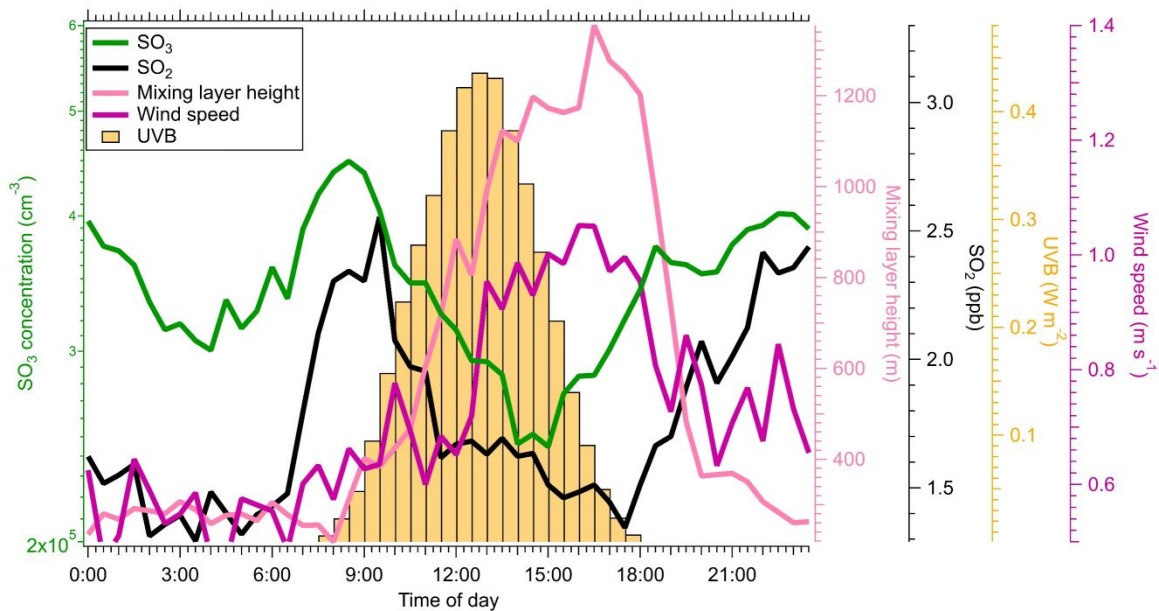
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264 **Figure S7.** The averaged mass spectra of atmospheric naturally charged ions for one whole day (10
 265 November 2018).

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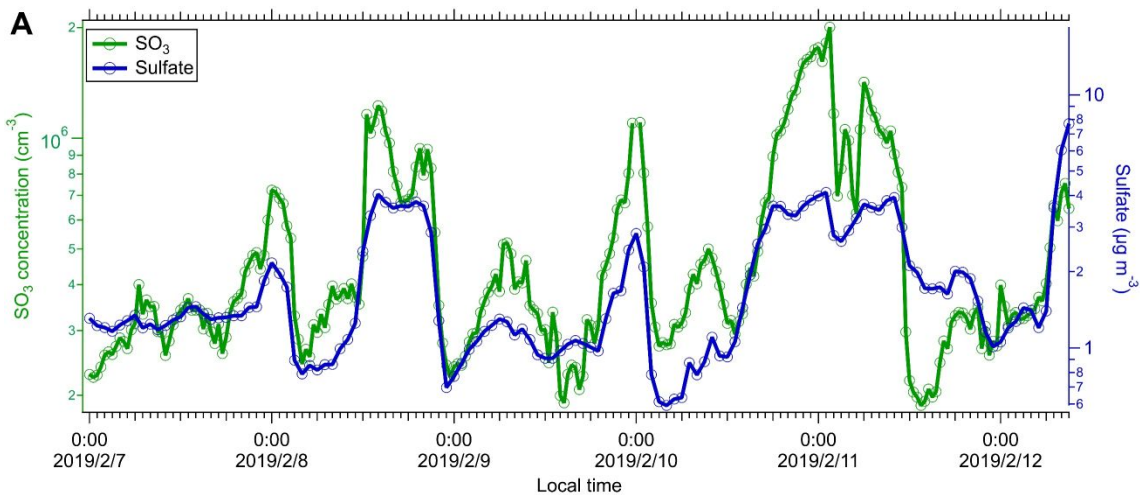
270 **Figure S8.** Median diurnal variation of concentrations of SO_3 and SO_2 , the mixing layer heights (MLH),
 271 intensities of UVB, and wind speeds during winter.

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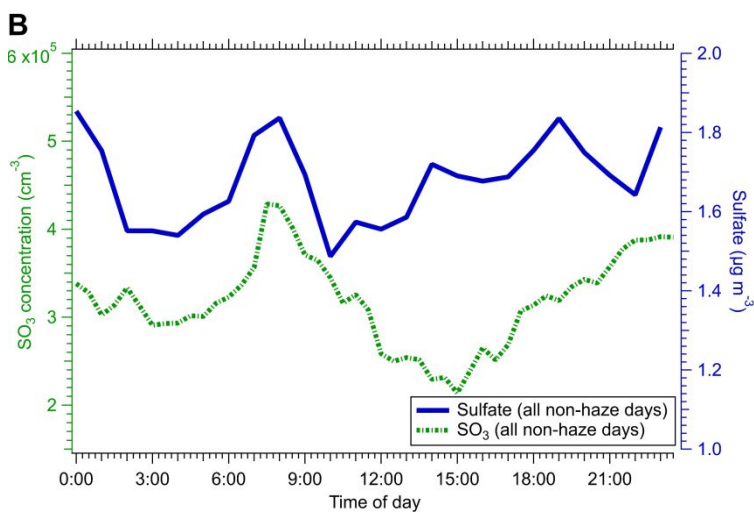
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278 **Figure S9.** Time profile of SO₃ concentration and mass concentration of sulfate in PM_{2.5} from 7 February to
 279 11 February 2019 (A) and median diel variation of SO₃ and sulfate for all non-haze days during the winter
 280 measurement period (B).

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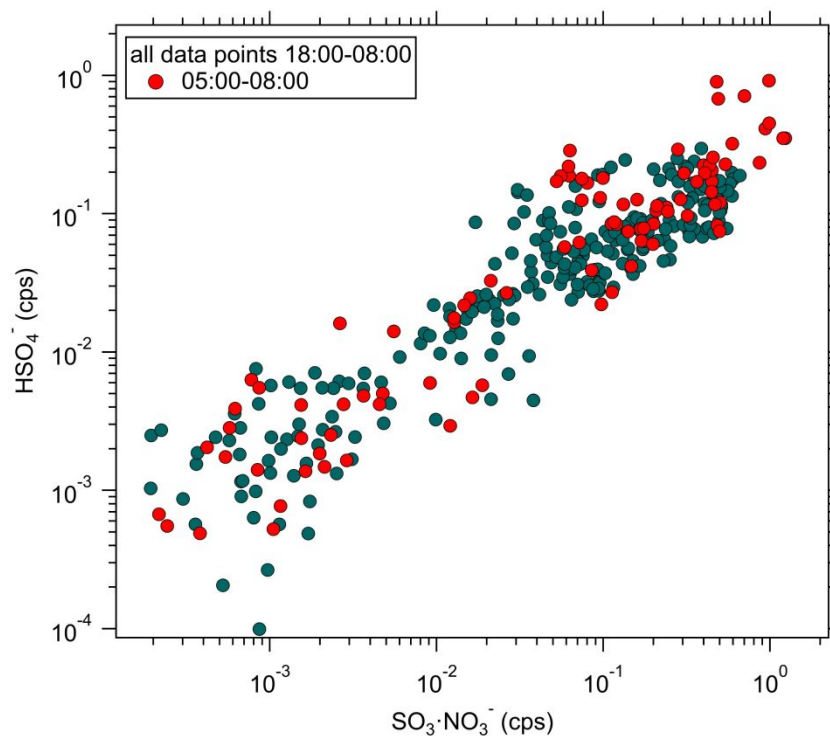
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288 **Figure S10.** The relationship between the atmospheric ion signals of HSO_4^- and $\text{SO}_3 \cdot \text{NO}_3^-$ during night time
 289 (18:00-5:00 next day) and early morning (5:00-8:00) from 9 to 22 November 2018.

290

291 **Table S1.** Comparison of the binding thermodynamics of $\text{HNO}_3 \cdot (\text{NO}_3^-)$ and $\text{SO}_3 \cdot (\text{NO}_3^-)$ ion-molecule
 292 clusters, in kcal/mol. ΔE_{DFT} and ΔE_{F12} correspond to the electronic energies (not including vibrational zero-
 293 point corrections) computed at the $\omega\text{B97X-D/aug-cc-pVTZ}$ and $\text{RHF-RCCSD(T)-F12b/VTZ-F12}$ levels,
 294 respectively, both at the optimized geometry corresponding to the former method. $\Delta G_{\text{DFT,F12}}$ corresponds to
 295 the Gibbs free energy (at 298.15 K and 1 atm reference pressure) obtained by combining the RHF-
 296 $\text{RCCSD(T)-F12b/VTZ-F12}$ electronic energy with the $\omega\text{B97X-D/aug-cc-pVTZ}$ thermal and vibrational zero-
 297 point contributions.

	ΔE_{DFT}	ΔE_{F12}	$\Delta G_{\text{DFT,F12}}$
$\text{HNO}_3(\text{NO}_3^-)$	-29.1	-29.2	-21.5
$\text{SO}_3(\text{NO}_3^-)$	-40.1	-44.4	-32.7

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306 **References:**

- 307 1. Zhou, Y.; Dada, L.; Liu, Y.; Fu, Y.; Kangasluoma, J.; Chan, T.; Yan, C.; Chu, B.; Daellenbach, K.
308 R.; Bianchi, F.; Kokkonen, T. V.; Liu, Y.; Kujansuu, J.; Kerminen, V.-M.; Petäjä, T.; Wang, L.; Jiang, J.;
309 Kulmala, M., Variation of size-segregated particle number concentrations in wintertime Beijing. *Atmos.*
310 *Chem. Phys.* **2020**, *20*, (2), 1201-1216.
- 311 2. Jokinen, T.; Sipilä, M.; Junninen, H.; Ehn, M.; Lönn, G.; Hakala, J.; Petäjä, T.; Mauldin, R. L.;
312 Kulmala, M.; Worsnop, D. R., Atmospheric sulphuric acid and neutral cluster measurements using CI-API-
313 TOF. *Atmos. Chem. Phys.* **2012**, *12*, (9), 4117-4125.
- 314 3. Eisele, F. L.; Tanner, D. J., Measurement of the Gas-Phase Concentration of H₂SO₄ and Methane
315 Sulfonic-Acid and Estimates of H₂SO₄ Production and Loss in the Atmosphere. *J. Geophys. Res-Atmos.* **1993**,
316 *98*, (D5), 9001-9010.
- 317 4. Yao, L.; Garmash, O.; Bianchi, F.; Zheng, J.; Yan, C.; Kontkanen, J.; Junninen, H.; Mazon, S. B.;
318 Ehn, M.; Paasonen, P.; Sipilä, M.; Wang, M.; Wang, X.; Xiao, S.; Chen, H.; Lu, Y.; Zhang, B.; Wang, D.;
319 Fu, Q.; Geng, F.; Li, L.; Wang, H.; Qiao, L.; Yang, X.; Chen, J.; Kerminen, V. M.; Petäjä, T.; Worsnop, D.
320 R.; Kulmala, M.; Wang, L., Atmospheric new particle formation from sulfuric acid and amines in a Chinese
321 megacity. *Science* **2018**, *361*, (6399), 278-281.
- 322 5. Junninen, H.; Ehn, M.; Petäjä, T.; Luosujärvi, L.; Kotiaho, T.; Kostianen, R.; Rohner, U.; Gonin,
323 M.; Fuhrer, K.; Kulmala, M.; Worsnop, D. R., A high-resolution mass spectrometer to measure atmospheric
324 ion composition. *Atmos. Meas. Tech.* **2010**, *3*, (4), 1039-1053.
- 325 6. Kurten, A.; Rondo, L.; Ehrhart, S.; Curtius, J., Calibration of a chemical ionization mass
326 spectrometer for the measurement of gaseous sulfuric acid. *J. Phys. Chem. A* **2012**, *116*, (24), 6375-86.
- 327 7. Su, T.; Bowers, M. T., Ion-Polar Molecule Collisions - Effect of Molecular-Size on Ion-Polar
328 Molecule Rate Constants. *J. Am. Chem. Soc.* **1973**, *95*, (23), 7609-7610.
- 329 8. Su, T.; Chesnavich, W. J., Parametrization of the Ion-Polar Molecule Collision Rate-Constant by
330 Trajectory Calculations. *J. Chem. Phys.* **1982**, *76*, (10), 5183-5185.
- 331 9. Passananti, M.; Zapadinsky, E.; Zanca, T.; Kangasluoma, J.; Mylly, N.; Rissanen, M. P.; Kurten,
332 T.; Ehn, M.; Attoui, M.; Vehkamäki, H., How well can we predict cluster fragmentation inside a mass
333 spectrometer? *Chem. Commun.* **2019**, *55*, (42), 5946-5949.
- 334 10. Chai, J. D.; Head-Gordon, M., Long-range corrected hybrid density functionals with damped atom-
335 atom dispersion corrections. *Phys. Chem. Chem. Phys.* **2008**, *10*, (44), 6615-6620.
- 336 11. Woon, D. E.; Dunning, T. H., Gaussian-Basis Sets for Use in Correlated Molecular Calculations .III.
337 The Atoms Aluminum through Argon. *J. Chem. Phys.* **1993**, *98*, (2), 1358-1371.
- 338 12. Gaussian 16, Revision A.03, Frisch, M. J.; Trucks, G. W.; Schlegel, H. B.; Scuseria, G. E.; Robb,
339 M. A.; Cheeseman, J. R.; Scalmani, G.; Barone, V.; Petersson, G. A.; Nakatsuji, H.; Li, X.; Caricato, M.;
340 Marenich, A. V.; Bloino, J.; Janesko, B. G.; Gomperts, R.; Mennucci, B.; Hratchian, H. P.; Ortiz, J. V.;
341 Izmaylov, A. F.; Sonnenberg, J. L.; Williams-Young, D.; Ding, F.; Lipparini, F.; Egidi, F.; Goings, J.; Peng,
342 B.; Petrone, A.; Henderson, T.; Ranasinghe, D.; Zakrzewski, V. G.; Gao, J.; Rega, N.; Zheng, G.; Liang, W.;
343 Hada, M.; Ehara, M.; Toyota, K.; Fukuda, R.; Hasegawa, J.; Ishida, M.; Nakajima, T.; Honda, Y.; Kitao, O.;
344 Nakai, H.; Vreven, T.; Throssell, K.; Montgomery, J. A., Jr.; Peralta, J. E.; Ogliaro, F.; Bearpark, M. J.; Heyd,
345 J. J.; Brothers, E. N.; Kudin, K. N.; Staroverov, V. N.; Keith, T. A.; Kobayashi, R.; Normand, J.;
346 Raghavachari, K.; Rendell, A. P.; Burant, J. C.; Iyengar, S. S.; Tomasi, J.; Cossi, M.; Millam, J. M.; Klene,
347 M.; Adamo, C.; Cammi, R.; Ochterski, J. W.; Martin, R. L.; Morokuma, K.; Farkas, O.; Foresman, J. B.;
348 Fox, D. J. Gaussian, Inc., Wallingford CT, **2016**.
- 349 13. Adler, T. B.; Knizia, G.; Werner, H. J., A simple and efficient CCSD(T)-F12 approximation. *J.*
350 *Chem. Phys.* **2007**, *127*, (22).
- 351 14. Knizia, G.; Adler, T. B.; Werner, H. J., Simplified CCSD(T)-F12 methods: Theory and benchmarks.
352 *J. Chem. Phys.* **2009**, *130*, (5).

- 353 15. Werner, H. J.; Knizia, G.; Manby, F. R., Explicitly correlated coupled cluster methods with pair-
354 specific geminals. *Mol. Phys.* **2011**, *109*, (3), 407-417.
- 355 16. Peterson, K. A.; Adler, T. B.; Werner, H. J., Systematically convergent basis sets for explicitly
356 correlated wavefunctions: The atoms H, He, B-Ne, and Al-Ar. *J. Chem. Phys.* **2008**, *128*, (8).
- 357 17. MOLPRO, version 2019.2, a package of ab initio programs, Werner, H.-J.; Knowles, P. J.; Knizia, G.;
358 Manby, F. R.; Schütz, M.; Celani, P.; Györfy, W.; Kats, D.; Korona, T.; Lindh, R.; Mitrushenkov, A.; Rauhut,
359 G.; Shamasundar, K. R.; Adler, T. B.; Amos, R. D.; Bennie, S. J.; Bernhardsson, A.; Berning, A.; Cooper,
360 D. L.; Deegan, M. J. O.; Dobbyn, A. J.; Eckert, F.; Goll, E.; Hampel, C.; Hesselmann, A.; Hetzer, G.; Hrenar,
361 T.; Jansen, G.; Köppl, C.; Lee, S. J. R.; Liu, Y.; Lloyd, A. W.; Ma, Q.; Mata, R. A.; May, A. J.; McNicholas,
362 S. J.; Meyer, W.; Miller III, T. F.; Mura, M. E.; Nicklass, A.; O'Neill, D. P.; Palmieri, P.; Peng, D.; Pflüger,
363 K.; Pitzer, R.; Reiher, M.; Shiozaki, T.; Stoll, H.; Stone, A. J.; Tarroni, R.; Thorsteinsson, T.; Wang, M.; and
364 Welborn, M. see <https://www.molpro.net>.
- 365 18. Drinovec, L.; Močnik, G.; Zotter, P.; Prévôt, A. S. H.; Ruckstuhl, C.; Coz, E.; Rupakheti, M.; Sciare,
366 J.; Müller, T.; Wiedensohler, A., The "dual-spot" Aethalometer: an improved measurement of aerosol black
367 carbon with real-time loading compensation. *Atmos. Meas. Tech.* **2014**, *8*, (5), 1965-1979.
- 368 19. Gundel, L. A.; Dod, R. L.; Rosen, H.; Novakov, T., The Relationship between Optical Attenuation
369 and Black Carbon Concentration for Ambient and Source Particles. *Sci. Total Environ.* **1984**, *36*, (Jun), 197-
370 202.
- 371 20. Weingartner, E., Saathoff, H., Schnaiter, M., Streit, N., Bitnar, B., and Baltensperger, U., Absorption
372 of light by soot particles: determination of the absorption coefficient by means of aethalometers. *J. Aerosol*
373 *Sci.* **2003**, *34*, 1445 – 1463.
- 374 21. Zotter, P.; Herich, H.; Gysel, M.; El-Haddad, I.; Zhang, Y.; Močnik, G.; Hüglin, C.; Baltensperger,
375 U.; Szidat, S.; Prévôt, A. S. H., Evaluation of the absorption Ångström exponents for traffic and wood burning
376 in the Aethalometer-based source apportionment using radiocarbon measurements of ambient aerosol. *Atmos.*
377 *Chem. Phys.* **2017**, *17*, (6), 4229-4249.
- 378 22. Munkel, C.; Eresmaa, N.; Rasanen, J.; Karppinen, A., Retrieval of mixing height and dust
379 concentration with lidar ceilometer. *Bound-Lay Meteorol* **2007**, *124*, (1), 117-128.
- 380 23. Vanhanen, J.; Mikkilä, J.; Lehtipalo, K.; Sipilä, M.; Manninen, H. E.; Siivola, E.; Petaja, T.;
381 Kulmala, M., Particle Size Magnifier for Nano-CN Detection. *Aerosol Sci. Technol.* **2011**, *45*, (4), 533-542.
- 382 24. Kulmala, M.; Petaja, T.; Nieminen, T.; Sipilä, M.; Manninen, H. E.; Lehtipalo, K.; Dal Maso, M.;
383 Aalto, P. P.; Junninen, H.; Paasonen, P.; Riipinen, I.; Lehtinen, K. E. J.; Laaksonen, A.; Kerminen, V. M.,
384 Measurement of the nucleation of atmospheric aerosol particles. *Nat. Protoc.* **2012**, *7*, (9), 1651-1667.
- 385 25. Xu, W. Y.; Zhao, C. S.; Ran, L.; Lin, W. L.; Yan, P.; Xu, X. B., SO₂ noontime-peak phenomenon
386 in the North China Plain. *Atmos. Chem. Phys.* **2014**, *14*, (15), 7757-7768.
- 387 26. Li, R.; Fu, H. B.; Cui, L. L.; Li, J. L.; Wu, Y.; Meng, Y.; Wang, Y. T.; Chen, J. M., The
388 spatiotemporal variation and key factors of SO₂ in 336 cities across China. *J. Clean. Prod.* **2019**, *210*, 602-
389 611.
- 390 27. Huang, Q.; Cheng, S. Y.; Perozzi, R. E.; Perozzi, E. F., Use of a MM5-CAMx-PSAT Modeling
391 System to Study SO₂ Source Apportionment in the Beijing Metropolitan Region. *Environ. Model. Assess.*
392 **2012**, *17*, (5), 527-538.
- 393 28. Kampa, M.; Castanas, E., Human health effects of air pollution. *Environ. Pollut.* **2008**, *151*, (2),
394 362-367.
- 395 29. Zhong, Q. R.; Shen, H. Z.; Yun, X.; Chen, Y. L.; Ren, Y. A.; Xu, H. R.; Shen, G. F.; Ma, J. M.; Tao,
396 S., Effects of International Fuel Trade on Global Sulfur Dioxide Emissions. *Environ. Sci. Tech. Lett.* **2019**, *6*,
397 (12), 727-731.
- 398 30. Klimont, Z.; Smith, S. J.; Cofala, J., The last decade of global anthropogenic sulfur dioxide: 2000-
399 2011 emissions. *Environ. Res. Lett.* **2013**, *8*, (1).

- 400 31. Su, S. S.; Li, B. G.; Cui, S. Y.; Tao, S., Sulfur Dioxide Emissions from Combustion in China: From
401 1990 to 2007. *Environ. Sci. Technol.* **2011**, *45*, (19), 8403-8410.
- 402 32. Zheng, H. T.; Cai, S. Y.; Wang, S. X.; Zhao, B.; Chang, X.; Hao, J. M., Development of a unit-
403 based industrial emission inventory in the Beijing-Tianjin-Hebei region and resulting improvement in air
404 quality modeling. *Atmos. Chem. Phys.* **2019**, *19*, (6), 3447-3462.
- 405 33. Liu, J.; Mauzerall, D. L.; Chen, Q.; Zhang, Q.; Song, Y.; Peng, W.; Klimont, Z.; Qiu, X. H.; Zhang,
406 S. Q.; Hu, M.; Lin, W. L.; Smith, K. R.; Zhu, T., Air pollutant emissions from Chinese households: A major
407 and underappreciated ambient pollution source. *Proc. Natl. Acad. Sci. U.S.A.* **2016**, *113*, (28), 7756-7761.
- 408
- 409