

## Supplementary Materials for

### **Groundwater production from geothermal heating on early Mars and implication for early martian habitability**

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## Supplementary Materials

### Text S1. Noachian Water Inventory

At the present, if all of the known ice were to be distributed globally, it would only make  $\sim 30$  m global equivalent layer (GEL) (62). The distribution of the present-day inventory of water ice above the predicted ice stability line of approximately +1 km during the Noachian would have resulted in regional ice sheets with an average thickness of only 700 m. The Martian water budget during late Noachian is unknown with estimates anywhere between 640m (63) to 5000 m GEL (64). The distribution of these water inventory estimates above the predicted ice stability line would have resulted in ice sheets with an average thickness of 14 - 116 km, however, it is unlikely for the entire global inventory of water in ancient Mars to have been deposited in the southern highlands alone. For our nominal model, we assume a conservative value of 2 km thick ice sheets (similar to the thickness of the Taku glacier on Earth (65, 66) in southern highlands to assess if basal melting would have been possible during the Noachian. This value of ice thickness is extremely conservative and corresponds to roughly  $\sim 14\%$  of the 640m GEL estimate and  $\sim 1\%$  of the 5000m GEL estimate.

### Text S2. The Effect of Icy Highlands on Surface Temperature

We model the thermophysical evolution of ice sheets using mean annual surface temperature of 230 K based on the 3D GCM models of Mars (11). However, the surface temperature could be considerably lower if a substantial portion of Mars was covered by ice. This is because ice is reflective and reduces the amount of energy available to warm the surface. Simulation of ice on early Mars scenarios suggest that global ice coverage could have exceeded 25 - 30 %. In such a scenario, the surface temperature may decrease by almost 25 K. Although, this effect is considerably muted in the presence of a thick CO<sub>2</sub> atmosphere with clouds (10) . We ran all the models presented here with a surface temperature of 200 K and found that no melt is produced. If Mars had thicker ice deposits in the Noachian highlands, then it is possible to generate melt even with a 200 K surface temperature. It is conceivable that thick ice deposits on Mars, if they existed, may have been covered by a thin layer of dust or volcanic ash as also observed for recently detected ice sheets (67), thus reducing the total albedo and reducing the problem associated with the icy highlands melt scenario. Both SPLD and NPLD exhibit depositions of ice with varying dust content (68), thus, it is plausible that subsequent dust deposition would have decreased the albedo of icy highlands. While speculative, the layering observed within the NPLD and SPLD suggest that this is not an unlikely scenario to have occurred on Mars.

### Text S3. Snow Accumulation Rate

No empirical evidence exists to constrain the snow accumulation rate during the Noachian. We use a nominal accumulation rate of 10 mm yr<sup>-1</sup>, which is predicted from GCM models. The snow accumulation rate is roughly exponentially dependent on temperature, so if there was a minor addition of greenhouse gases to the atmosphere or if the young Sun was slightly more luminous, then significantly faster accumulation than 10 mm yr<sup>-1</sup> would have occurred. In general, faster or increased accumulation would have led to increased basal melting. No meltwater would have been generated if the snow accumulation rate was substantially lower than considered here. Some insight about the plausibility of these values may be gathered from terrestrial estimates. The average accumulation of snow over the continent as a whole is estimated to be equivalent to about 150mm of water per year.

### Text S4. Th and K Enriched Regions of Mars

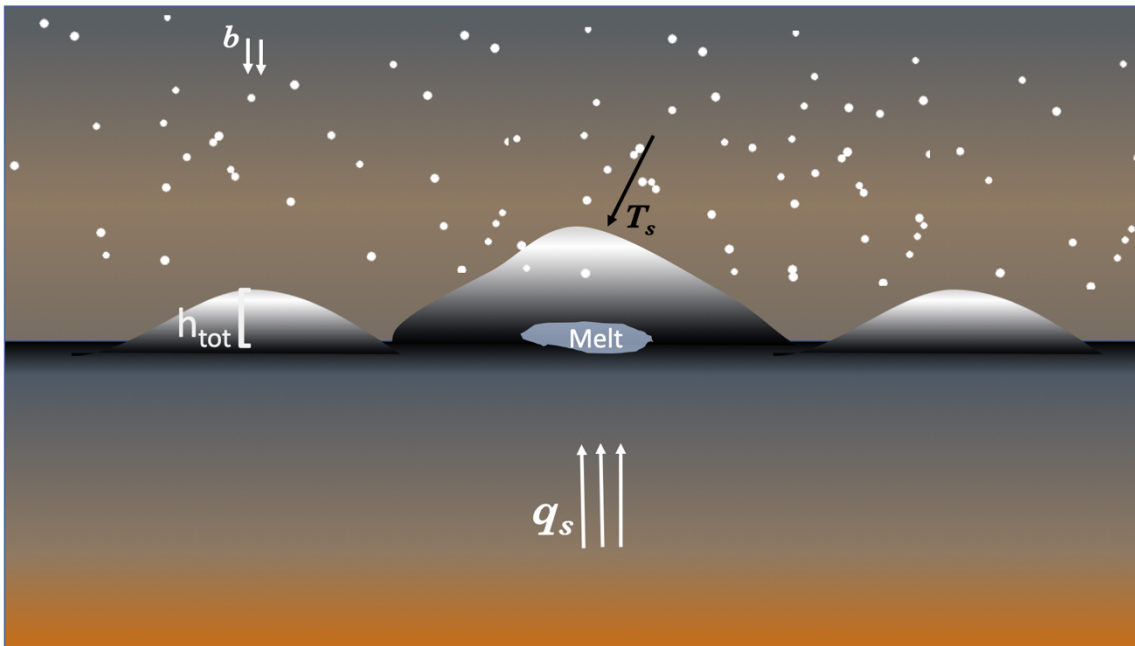
We sought to find regions on Mars that are statistically enriched in Th and K. To that end, we used an enhanced Student's t-test parameter,  $t_i$ , that measures the error-weighted deviation for each element from its bulk-average on Mars at each  $5^\circ \times 5^\circ$  chemical map grid derived from GRS data (69, 70):

$$t_i = \frac{c_i - m}{\sqrt{s_{m,i}^2 + s^2}}; \quad (1)$$

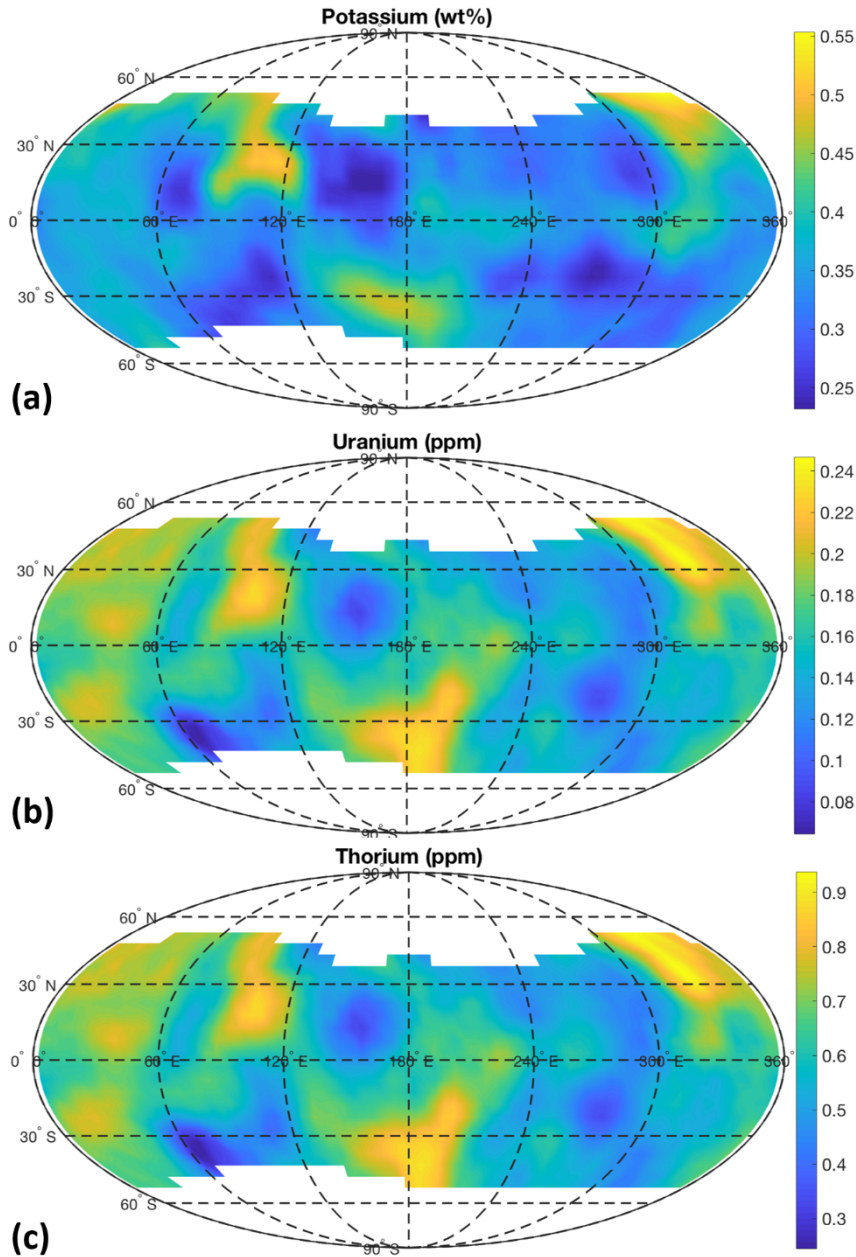
where  $c$  is the wt% of an element,  $m$  is the global arithmetic mean wt%,  $s_{m,i}$  is the numerical uncertainty of  $c$ , and  $s$  is the standard deviation of the data. The key difference between the ‘ $t$ ’ used here and the commonly used Student’s  $t$  parameter is the inclusion of ‘ $s_{m,i}$ ’ in the denominator. Additional information about the enhanced  $t$ -test can be found in our previous work (69, 70). Here we define areas of significant enrichment as those with  $t$  values greater than 1.5 (i.e. statistical confidence in directional deviation > 94 %).

### Text S5. Empirical Cumulative Distribution Functions

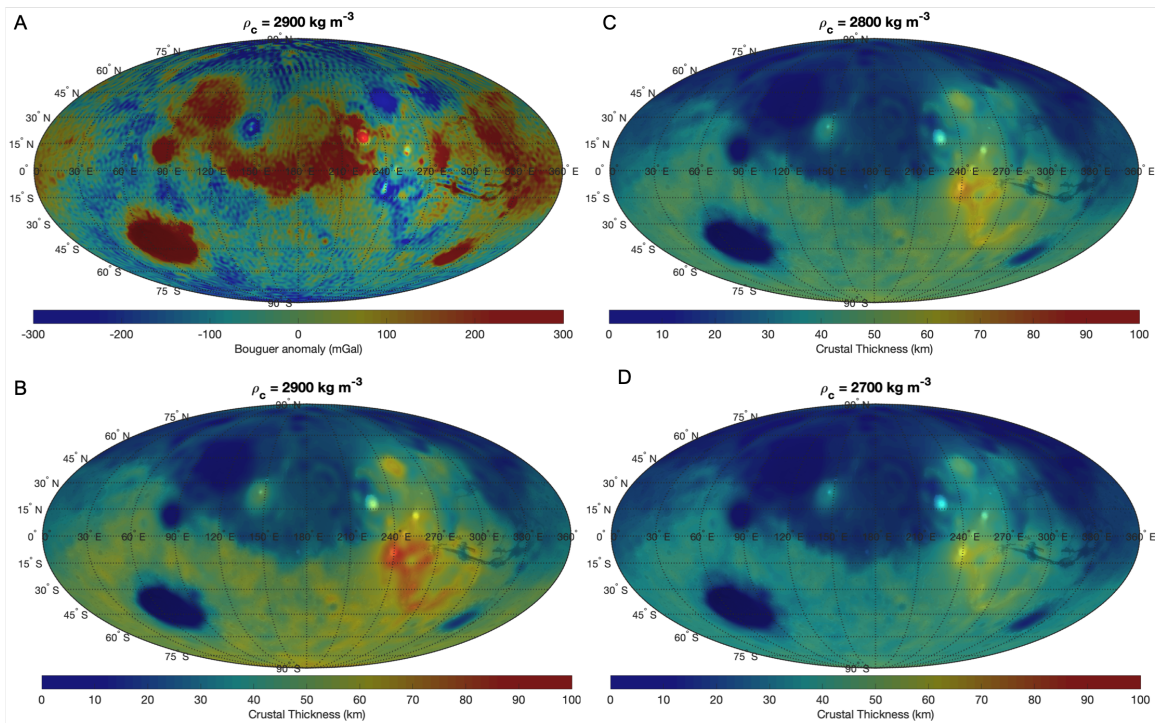
Empirical Cumulative Distribution Functions (ECDF) for all Noachian terrain reflects all pixels from the 1x1 degree resolution surface heat flow map (interpolated from the original 5x5 degree resolution map) (Fig. 3 (c)). That consists of each pixel where more than 50% of the area corresponds to Noachian age. Meanwhile, for each location where a mineral is detected – often in an outcrop via CRISM data – the surface heat flow is approximated as that of the containing heat flow map pixel. For example, two spatially distinct detections of chloride within the same pixel are assigned the same heat flow value. A value of ‘NaN’ is ascribed when the location of the hydrous minerals is outside of Noachian terrain and those sites are excluded from further statistical analysis. Matlab command ‘cdfplot’ is used to create the ECDF plots shown in Figure S6.



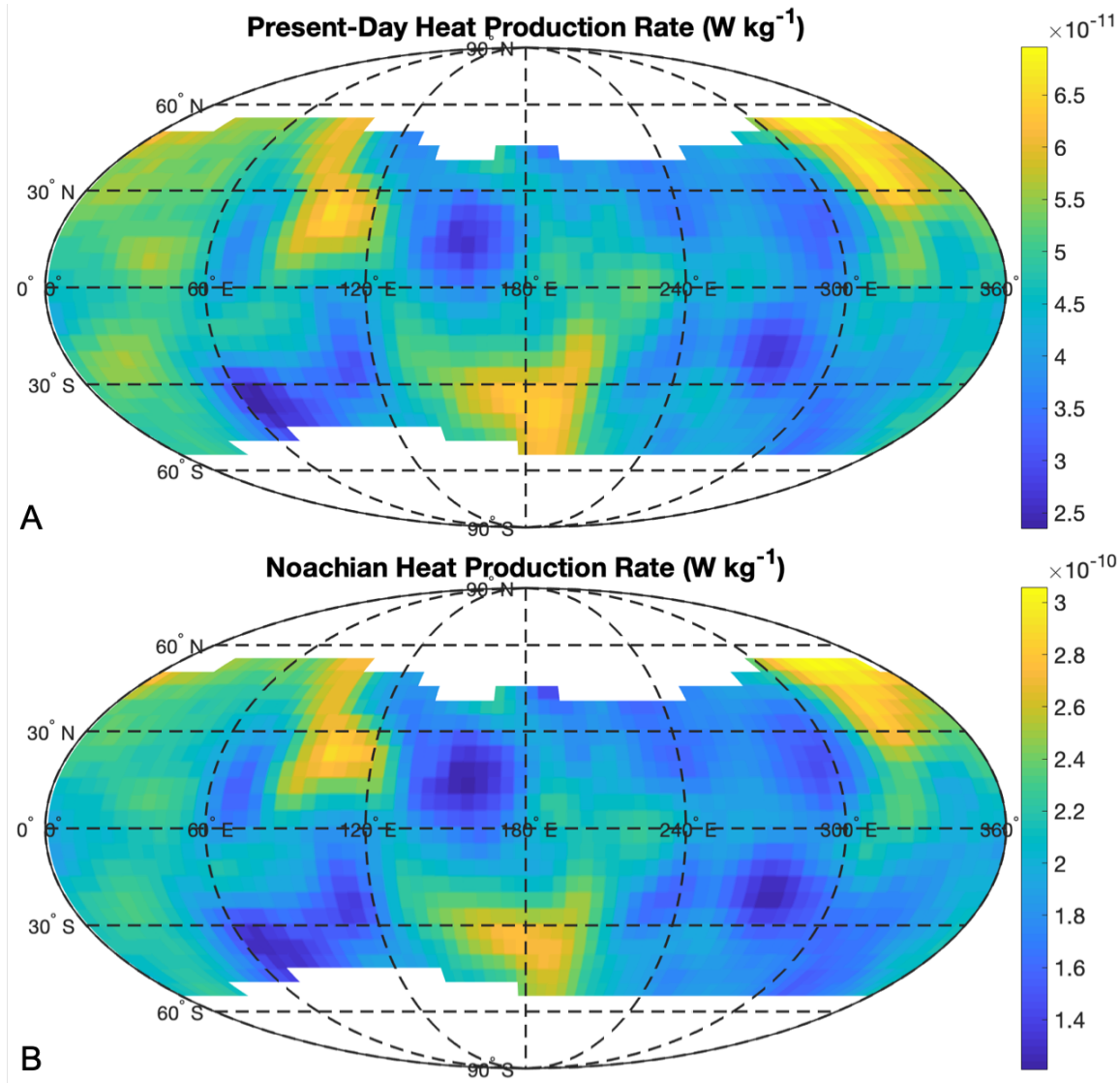
**Figure S1.** Schematic showing the setup of our model. For a given snow accumulation/precipitation rate of  $b$  ( $\text{mm yr}^{-1}$ ), total thickness of the snow/ice slab ( $h_{tot}$ ), and surface temperature ( $T_s$ ), we solve for the surface heat flow ( $q_s$ ) that can produce meltwater. Melt is only produced for the thickest of the ice slabs.



**Figure S2.** Distribution of the heat producing elements in the shallow subsurface of Mars. (a) Potassium concentration in the shallow subsurface of Mars at  $5^\circ \times 5^\circ$  in the mid-latitudes, as derived from Mars Odyssey Gamma Ray Spectrometer suite. (b) Same as (a) but showing the concentration of U calculated using the cosmochemical Th/U ratio. (c) Same as (a) and (b) but showing the concentration of Th. Rapidly increasing H abundance dilutes and increases numerical uncertainty for HPE concentrations in the polar latitudes. A mask has thus been applied to exclude such areas.

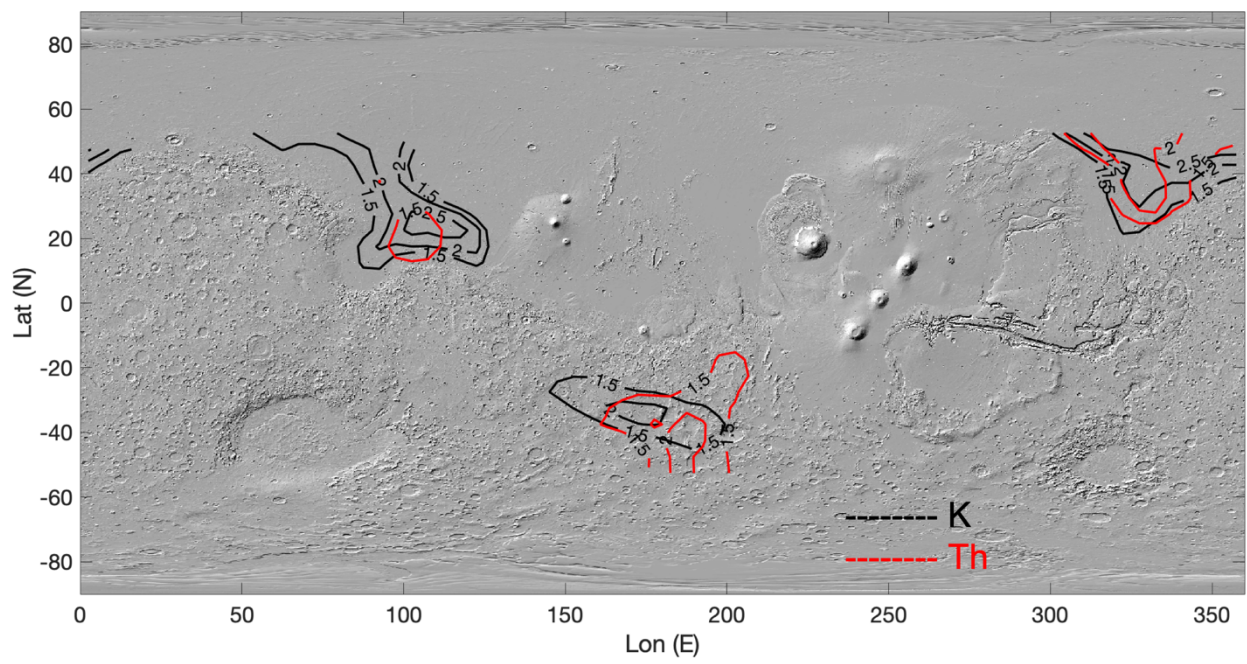


**Figure S3.** Bouguer gravity and crustal thickness maps of Mars. (A) An example of a Bouguer gravity anomaly map of Mars assuming a constant density of  $2900 \text{ kg m}^{-3}$ . (B) – (D) Crustal thickness of Mars derived from downward continuing the observed Bouguer gravity anomaly to a crust-mantle interface for crustal densities of  $2900 - 2700 \text{ kg m}^{-3}$ .

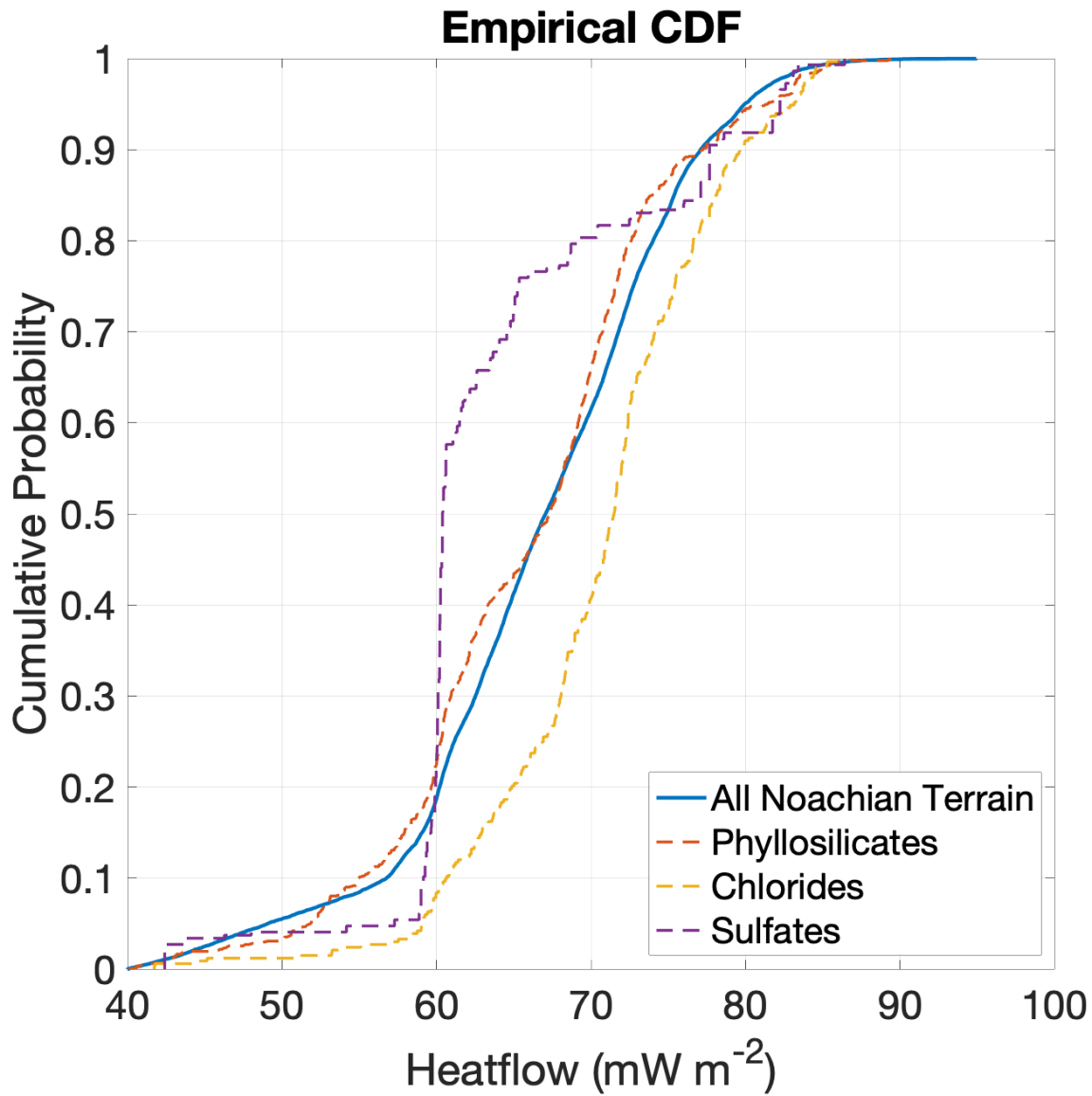


**Figure S4.** Current and Noachian crustal heat production rate for Mars. (a) Present day crustal heat production map using HPE abundance maps from Fig. S2 and crustal thickness maps from Fig. S3. (b) Crustal heat production rate during the Noachian with the current lateral HPE variability preserved.





**Figure S5.** Contour map showing regions on Mars with significant enrichment of Th and K. The contour labels correspond to a modified 't' parameter that show regions with significant enrichment of Th and K compared to the bulk-average of Mars. The background is the shaded relief map of Mars from Mars Orbiter Laser Altimeter (MOLA).



**Figure S6.** An Empirical cumulative distribution function (ECDF) plot of surface heat flow values for all Noachian terrain and regions in Noachian terrain that bear various hydrous minerals.



**Table S1: All variables and values used in the thermal simulation.**

Variable	Description	Value	Units
$\dot{b}$	Accumulation Rate	3.1688e-10 (1)	m s <sup>-1</sup> (cm yr <sup>-1</sup> )
$c_{ice}$	Specific Heat of Ice	2000	J kg <sup>-1</sup> K <sup>-1</sup>
$c_v$	Specific Heat of Voids	790	J kg <sup>-1</sup> K <sup>-1</sup>
$c_w$	Specific Heat of Melt	3985	J kg <sup>-1</sup> K <sup>-1</sup>
$f_0$	Constant Coefficient	9.5064e-16	-
$g$	Acceleration Due to Gravity	3.711	m s <sup>-2</sup>
$H$	Enthalpy	Calculated	J kg <sup>-1</sup>
$H_S$	Enthalpy of Solid Cell	Calculated	J kg <sup>-1</sup>
$H_{tot}$	Critical Ice Sheet Thickness	1300, 1400, 1500, 2000	m
$k_{ice}$	Thermal Conductivity of Ice	2	W m <sup>-1</sup> K <sup>-1</sup>
$k_v$	Thermal Conductivity of Voids	0.012	W m <sup>-1</sup> K <sup>-1</sup>
$k_w$	Thermal Conductivity of Melt	0.6	W m <sup>-1</sup> K <sup>-1</sup>
$L$	Latent Heat of Fusion	334774	J kg <sup>-1</sup>
$n$	Ice Deformation Coefficient	3	-
$\phi_v$	Void Space Volume Fraction	Calculated	-
$\phi_w$	Melt Volume Fraction	Calculated	-
$Q$	Activation Energy	45600	J mol <sup>-1</sup>
$R$	Gas Constant	8.314	J mol <sup>-1</sup> K <sup>-1</sup>
$\rho$	Density	Calculated	kg m <sup>-3</sup>
$\rho_{ice}$	Density of Ice	917	kg m <sup>-3</sup>
$\rho_v$	Density of Voids	1	kg m <sup>-3</sup>
$\rho_w$	Density of Melt	1000	kg m <sup>-3</sup>
$\rho_0$	Ice Sheet Surface Density	350	kg m <sup>-3</sup>
$t$	Time	-	s
$T$	Temperature	Calculated	K
$T_S$	Surface Temperature	230	K
$T_m$	Melting Temperature	273.15	K
$z$	Depth	-	m

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