Supplementary Information

S1 Dendrochronology

Dendrochronological dating of waterlogged oak samples from the river Main, Germany (M49 and M234)

The subfossil oak (*Quercus* sp.) samples, M49 and M234, used for this study were collected from waterlogged trunks discovered during gravel quarrying along the River Main in the 1970s and 1980s. Samples were obtained from a veneer mill near Bamberg, which obtained timber from a series of gravel pits along an 80 km stretch of the River Main between Schweinfurt (50.05N; 10.24E) and Lichtenfels (50.15N; 11.05E). The provenance of these particular timbers is not known exactly.

These samples were selected for inclusion in this study because the growth rings of the trees were > 1mm in width, and the samples were without distortion caused by growth anomalies or knots that would make the dissection of the single tree-rings for radiocarbon dating difficult.

The wet samples from the gravel pits were prepared using a scalpel blade to remove the rough wood from the sample surface to expose the tree-ring pattern. The ring-widths were then measured to a precision of 1/100 mm using a micro-computer based travelling stage.

The individual ring-width series for M234 was cross-dated against the Göttingen South German Oak Chronology¹, and was further compared with the Hohenheim Southern German Oak chronology², which over this time period was constructed separately from the Göttingen master chronology. The correlation values obtained are given in Table S1.1. The individual ring-width series for M49 were cross-dated against the Hohenheim Southern German Oak Chronology. The correlation values obtained are also given in Table S1.1. Figure S1.1 shows the ring-width series of M49 and M234 on a logarithmic scale against the Hohenheim master chronology.

Table S1.1: overlaps (years) / Gleichläufigkeit (%) / t-values (Ballie and Pilcher 1973) between M49, M234, the Göttingen

 South German Riverine Oak Chronology and the Hohenheim South German Oak Chronology.

	Hohenheim South German oak	M49	M234
Göttingen South German oak	8797 / 78 / 91.5		
Göttingen South German oak (without M234)			133 / 67 / 6.5
Hohenheim Southern German oak		118 / 69 / 10.0	133 / 69 / 8.5
M49			39 / 66 /3.0

The good visual matching of the ring-width series (Figure S1.1), along with the correlation values obtained (Table S1.1) dates the 118-ring series, M49, as spanning 7276–7159 BCE and the 133-ring series, M234, as spanning 7197–7065 BCE. The calendar dates provided are historic dates BCE without the year zero.



Figure S1.1: (a) Ring-width series of M49 and M234 on a logarithmic scale against the Hohenheim master chronology (Friedrich et al. 2004).

Timbers from the Main river gravels were originally waterlogged, but have subsequently been allowed to dry out naturally. The remaining parts of the samples are stored in the Department of Palynology and Climate Dynamics, Albrect-von-Haller Institute for Plant Sciences, Georg-August-Universität-Göttingen, Wilhelmsplatz 1, 37073 Göttingen, Germany. Ring-width data for the two tree slices sampled for radiocarbon dating are provided in Table S1.2.

Dissection for radiocarbon dating was undertaken by Alison Arnold and Robert Howard of the Nottingham Tree-ring Dating Laboratory. The samples to be used were cleaned with a razor blade and the annual growth rings marked. At this stage it became apparent that 22 rings from the centre of M232, and two from the outside, had crumbled during storage, and only the rings spanning 7175–7067 BC were available for dissection.

The annual growth rings were split from the bulk sample using a scapel blade. Whole ring samples, including both earlywood and latewood, were provided to the ETH laboratory from M49 (7276–7160 BCE) and M232 (7175–7067 BCE). Replicate whole ring samples, including both earlywood and latewood, were provided to the Bristol laboratory from M49 (7181–7160 BCE). Further replicate whole ring samples from M49 (n=29) and M234 (n=13) have been archived by Historic England.

The annual growth rings from M49 (7178–7175 BCE) were further divided using a scapel into separate samples comprising earlywood only and latewood only, which were provided to both the ETH and Bristol laboratories.

Finally, the annual growth ring for 7177 BCE from M49 was split into early and late earlywood (EW1 and EW2), and early and late latewood (LW1 and LW2) and the samples supplied to the ETH laboratory.

Table S1.2: Ring-width data for the timbers which provided single-ring samples for radiocarbon calibration (Heidelberg format)

HEADER: KeyCode=SM001D 0 DataFormat=Tree Project=S Country=S LocationCharacteristics=M SiteCode=00 TreeNo=1D CoreNo=0 SeriesType=Single curve StatisticalDataTreatment=O DataType=Ringwidth Unit=mm Dated=Dated Length=118 DateBegin=-7276 DateEnd=-7159 Species=Qusp Location=M49 PersId=HI LastRevisionDate=01.04.90 MissingRingsBefore=25 DeltaMissingRingsBefore=10 MissingRingsAfter=20

Dendrochronological dating of the bog oak samples from Derrymacfall, Co. Armagh, Northern Ireland (Q 2729 and Q 2750)

In July 1977, slices were collected from seventy-one bog oak (*Quercus* sp.) timbers from the townland of Derrymacfall, just south of Lough Neagh, Co. Armagh, Northern Ireland (54° 28'N, 6° 28'W).

The wet samples were prepared using a scalpel blade to remove the rough wood from the surface of the sample and to expose the tree-ring pattern. If the wood sample was soft or the ring pattern needed to be clarified, a razor blade was used. Finely ground chalk was rubbed onto the prepared surface to define the annual tree-ring boundaries more clearly for measurement. The ring-widths were measured to a precision of 1/50mm using a travelling stage between November 1979 and September 1980. Sometimes more than one radius was measured for an individual sample (e.g. Q2737A/B), and so some ring series are a mean of two or more measurement series on the same sample.

The ring-width series of nine of these samples as well as those of two samples from Loughmagarry, Co. Antrim (Q2527 and Q2528; 54° 52'N, 6° 17'W) were included in the Derrymacfall 2 master chronology (Table S1.3). The construction of the chronology followed the procedures described by Baillie¹. Samples with the best correlation values were combined to form sub-site masters. These were then used to date others samples. The other dated samples were incorporated into a new subsite master until a coherent master chronology was formed. The individual samples were then compared with each other giving the correlation values presented in the *t*-value matrix (Table S1.4). The correlation *t*-values are based on CROS84². At this time, the Derrymacfall 2 master chronology could not be dated by dendrochronology.

Sample	Rings	Sapwood Rings	Absolute Dating (BCE)			
Derrymacfall, Co Al	rmagh					
Q2682	159	No sapwood	BCE 5429 – 5271			
Q2717	160	No sapwood	BCE 5446 – 5287			
Q2720	100	No sapwood	BCE 5353 – 5254			
Q2729	181	24	BCE 5356 – 5176			
Q2734	91	No sapwood	BCE 5318 – 5228			
Q2737M	176	No sapwood	BCE 5457 – 5282			
Q2750	267	?h/s	BCE 5428 – 5162			
Q2779	190	No sapwood	BCE 5389 – 5200			
Q2781	139	No sapwood	BCE 5419 – 5281			
Loughmagarry, Co Antrim						
Q2527	104	No sapwood	BCE 5378 – 5275			
Q2528	133	No sapwood	BCE 5392 – 5260			

Table \$1.3: Details of the cross-matched samples from the Derrymacfall 2 master chronology (?h/s = outermost measuredring is possible heartwood/sapwood boundary)

In 1989 re-measurement of bog oak samples from a number of sites was undertaken. This work produced local site chronologies which cross-dated both with the Derrymacfall 2 master chronology and with the original Belfast Long Chronology (BLC5000). When the Derrymacfall 2 master chronology was compared with these chronologies extremely significant and consistent correlation values were found with the Tullyroan 10 and the Motorway 2 chronology.

Table S1.4: t-value matrix for ring-width series in the Derrymacfall 2 master chronology (- = no correlation value; nsm = non-significant match; nh = not highest match; * = significant match; ** = very significant match; *** = extremely significant match)

QUB ID	Q2717	Q2682	Q2750	Q2781	Q2528	Q2779	Q2527	Q2729	Q2720	Q2734
Q2737	2.33n	6.59***	2.87nh	3.55ns	2.96nm	5.32**	2.82nh	3.47nsm	4.02ns	-
Μ	h			m		*			m	
Q2717	-	4.18ns	3.67ns	3.40ns	1.86nh	6.41**	0.38nm	3.33nsm	5.38***	-
		m	m	m		*				
Q2682	-	-	5.46***	6.94***	2.62nh	7.08**	1.60nh	3.05nsm	5.09***	-
						*				
Q2750	-	-	-	3.97ns	4.51ns	5.66**	3.04nh	5.87***	7.44***	3.80ns
				m	m	*				m
Q2781	-	-	-	-	3.55ns	6.51**	3.20ns	2.48nh	4.28*	-
					m	*	m			
Q2528	-	-	-	-	-	3.92*	7.31***	2.97vw	2.45nh	-
								m		
Q2779	-	-	-	-	-	-	2.49nh	3.58nsm	6.88***	2.81nh
Q2527	-	-	-	-	-	-	-	2.68nh	1.29nh	-
Q2729	-	-	-	-	-	-	-	-	4.24*	2.98ns
										m
Q2720	-	-	-	-	-	-	-	-	-	3.29ns
										m

The Derrymacfall 2 master chronology consists of eleven samples (Figure S1.2). It is 296 years in length and dates from BCE 5457 to BCE 5162. Table S1.5 gives the correlation values of this master chronology with other site chronologies from Ireland. Only correlation values over t = 4 are included in the table. The correlation t-values are based on CROS84².

Table S1.5: Correlation values for the 296-year Derrymacfall 2 master chronology with other site and regional chronologies (*** = extremely significant match ; ** = very significant match)

Site chronology	Start and end date of chronology	Correlation value with Derrymacfall 2
Tullyroan 10, Co Armagh	5334BCE to 4680BCE	<i>t</i> = 8.39***
Motorway 2, Co Armagh	5474BCE to 5228BCE	<i>t</i> = 5.75***
Killyleagh 1, Co Down	5236BCE to 4829BCE	<i>t</i> = 4.62**

Site	Span of ring sequences							
Derrymacfall	Q2717 Q2737M Q2781 Q2682 Q2779 Q2779	Q2720 Q2734 Q2729						
Loughmagarry	Q2527 Q2528							
Calendar Years	5400 BCE	5300 BCE	5200 BCE					

Figure S1.2: Bar diagram showing the relative positions of the dated samples. White bar = heartwood; yellow bar = sapwood;?h/s = outermost measured ring is possible heartwood/sapwood boundary

Timbers from Derrymacfall were originally waterlogged, but subsequently have been allowed to dry out naturally. The remainder of the samples are stored in the Dendrochronology Laboratory, School of Natural and Built Environment, Queen's University, Belfast. Ring-width data for all the measured samples from Derrymacfall can be found at

http://www.chrono.qub.ac.uk/bennett/dendro_data/dendro.html.

Two samples from the Derrymacfall 2 site chronology, Q2729 and Q2750 (Table S1.6), have been used to supply single-year samples for radiocarbon calibration. For the selected period, the trees with the widest annual growth rings were chosen, so that extracting the individual annual growth rings is as easy as possible. The samples to be used were cleaned with a razor blade and the annual growth rings marked. The annual growth rings were split from the bulk sample using a scalpel bade and the early wood growth was removed using a scalpel. This processing means that the wood samples provided only have the latewood part of the annual growth ring.

Sample Q2729 was used to supply samples from 5300 BCE to 5211 BCE inclusive, and sample Q2750 was used to supply samples from 5210 BCE to 5175 BCE inclusive. Those single-ring samples that weighed more than 120mg were divided further using a scalpel to supply duplicate samples of 90 single rings for inter-laboratory replication.

Table S1.6: Ring-width data for the timbers which provided single-ring samples for radiocarbon calibration (Heidelberg format)

HEADER: Key Code = Q2729 Data Format = Tree Location = Derrymacfall, Co. Armagh Length = 181 Date Begin = 5356BC Date End = 5176BC Dated = Dated Comment = EDH 17-June-1980 Close to centre 24 sapwood rings not complete Key No = Q2729 Project = Derrymacfall, Co. Armagh Species = QUSP Tree No = Q2729 Pith = -WaldKante = ---Sapwood Rings = 24 Unit = 1/50 mm Data Type = Ring width Series Type = Single curve Series Start = Ring width Series End = Ring width Global Math Comment Count = 0 Image Count = 0 Comment Count = 0 Bibliography Count = 0 DATA: Single

HEADER: Key Code = Q2750 Data Format = Tree Location = Derrymacfall, Co. Armagh Length = 267 Date Begin = 5428BC Date End = 5162BC Dated = Dated Comment = EDH 9-Aug-1980 Close to centre no sapwood possible H/S boundary Key No = Q2750 Project = Derrymacfall, Co. Armagh Species = QUSP Tree No = Q2750 Pith = -WaldKante = ---Unit = 1/50 mm Data Type = Ring width Series Type = Single curve Series Start = Ring width Series End = Ring width Global Math Comment Count = 0 Image Count = 0 Comment Count = 0

Bibliography Count = 0 DATA: Single

Bristlecone Pine

Two bristlecone pine (Pinus Longaeva D.K. Bailey) samples were used in this study. For the 5259 BCE event, sample MWK80-101 was selected from the calendar dated master Methuselah Walk bristlecone pine chronology³. The sample derives from a single *Pinus longaeva* D.K.Bailey tree which grew at the Methuselah Walk site (MWK) in the White Mountains of California in the western USA. The remnant tree was preserved in a dry, high altitude (c.2900 m a.s.l.) mountain environment. The full tree sampled had 747 rings with no pith, bark or waney edge present. Standard convention for bristlecone dating is via the use of the astronomical calendar (AD/BC with 0) for which details are provided in Ferguson et al.³. Chronologies are constructed using a combination of graphical skeleton plotting⁴, and / or tree-ring measurement, and then statistical verification using COFECHA⁵. For MWK80-101, calendar dates were assigned relative to the master MWK chronology and the sample was lagged successively by 25 years using COFECHA12K version 6.06P. The sample spans -5821 to -5075 (747 years, 30 50-yr segments, 0 flags and demonstrates excellent correlation with the master MWK chronology; r = 0.770). The COFECHA outputs are provided below in AD/BC with 0. Two subsamples of MWK80-101 were marked with dated years and sampled from the cross-dated main sample. The sub-samples were cut to be c. 5 mm thick relative to the transverse section. The upper and lower surface were lightly sanded (preserving original pin-holes) but removing any surface contamination and excess resins which come up to the sample surface over time. The upper, lower and radial surfaces were clearly marked with the pin-holes and measurements re-checked prior to dissection. Dissection was carried out using the astronomical calendar (AD/BC with 0) under times 20 magnification using a binocular microscope and steel disecting blade. The same blade was used to remove any residual contamination possibilities (pinholes, any residual cells from the previously dissected year of interest) prior to sampling each seperate year of interest. Samples were dissected

working from youngest to oldest and packed in pre-marked foil packets. We note that a correction of 2 years was applied to Wes Fergusons original 1966 CE dating of this part of the MWK chronology due to an observation by Lamarche and Harlan in 1973⁶ that Ferguson had incorrectly added what he thought were two likely missing rings to his chronology in -2141 and -2680. Chris Baisan and Mathew Salzer have since demonstrated that the first of these was a clerical error and the second a false ring, placed in 1966 by Ferguson, so both these erroneous years are removed from the master chronology used in this study.



Figure S1.3: Dissected rings of sample MWK80-101 with 1mm ruler for scale. Double pinhole is the year -5250.

COFECHA12K version 6.06P output showing correlations of individual 50-year dated segments, lagged 25 years from MWK-80-101:

Flags: A = correlation under .3281 but highest as dated; B = correlation higher at other than dated position

-5700 to -5176

 2 03-876
 -6261-5548
 .63
 .49
 .79
 .81
 .64
 .65

 33 67-035
 -5553-4714
 .57
 .57
 .62
 .61
 .58
 .52
 .65
 .78
 .78
 .77
 .62
 .67
 .79
 .65
 .62

 38 69-202
 -6223-5096
 .75
 .65
 .58
 .57
 .74
 .79
 .78
 .81
 .74
 .65
 .77
 .86
 .80
 .80
 .75
 .80
 .89
 .78
 .77

 40
 70-020
 -5460-4708
 .62
 .58
 .59
 .55
 .52
 .65
 .71
 .78
 .80
 .76
 .79
 .68
 .72
 .89
 .80
 .78

 43
 70-043
 -6377-5408
 .62
 .58
 .59
 .55
 .52
 .65
 .71
 .78
 .80
 .59

 44
 70-046
 -6066-5008
 .70
 .71
 .76
 .65
 .52
 .63
 .62
 .70
 .77
 .68
 .51
 .52
 .59
 .51
 .52
 .55
 .69

48 71-059 -5922-5302 .76 .82 .84 .87 .73 .55 .60 .79 .90 .79 .69 .61 .70 .37 .44 49 71-060 -5880-5324 .62 .50 .62 .55 .41 .59 .43 .53 .71 .70 .58 .63 .83 .63 .62 50 71-061 -5419-4931 .81 .80 .82 .80 .75 .86 .90 .85 .88 53 71-161 -5340-4819 .59 .64 .77 .85 .73 .69 55 72-104 -5722-4960 .69 .74 .79 .78 .67 .65 .53 .63 .81 .70 .58 .77 .87 .57 .67 .85 .86 .78 .62 .67 56 72-175 -6059-5080 .58 .68 .76 .84 .74 .68 .80 .85 .86 .79 .79 .85 .84 .69 .75 .79 .86 .89 .82 .82 61 77-104 -5747-4487 .68 .78 .84 .87 .82 .67 .53 .63 .76 .74 .66 .74 .82 .64 .74 .85 .86 .89 .77 .73 72 77-151 -5230-4748 .67 .74 76 77-162 -5673-4821 .70 .82 .86 .80 .77 .77 .79 .80 .76 .79 .80 .84 .84 .91 .93 .88 .85 .79 .78 83 78-026 -6065-5445 .75 .57 .72 .90 .67 .63 .70 .72 .84 .81 89 78-075 -5598-4575 .73 .66 .70 .73 .80 .74 .66 .65 .80 .81 .84 .84 .70 .74 .81 .81 92 79-026 -6166-5458 .70 .64 .84 .84 .76 .57 .56 .68 .80 95 79-058 -6012-4899 .81 .78 .83 .87 .75 .64 .63 .68 .76 .70 .78 .81 .76 .75 .85 .88 .89 .87 .81 .86 104 79-085 -5377-4989 .67 .66 .74 .63 .66 .84 .52 .52 119 79-136 -6245-5448 .77 .75 .71 .75 .69 .54 .58 .70 .80 .80 125 80-101 -5821-5075 .78 .66 .70 .77 .73 .79 .69 .58 .75 .74 .68 .85 .91 .74 .69 .74 .70 .79 .75 .74 127 80-103 -5361-4491 .81 .82 .83 .77 .77 .82 .78 133 80-110 -6022-4238 .74 .65 .78 .84 .64 .60 .78 .77 .80 .80 .77 .79 .84 .83 .81 .80 .76 .72 .79 136 80-116 -5685-2822 .75 .68 .73 .81 .80 .81 .84 .79 .75 .83 .77 .81 .93 .84 .83 .72 .83 .92 .82 .76 184 93-031 -6133-5618 .62 .70 .74 234 XX-001 -6056-3998 .73 .61 .75 .84 .84 .64 .52 .60 .63 .65 .66 .69 .71 .57 .54 .67 .74 .73 .67 .68 235 XX-002 -5323-3562 .69 .80 .81 .76 .65 Av segment correlation .70 .67 .75 .78 .69 .64 .65 .70 .77 .74 .67 .72 .78 .72 .75 .74 .76 .81 .74 .73

-5200 to -5051

 33
 67-035
 -5553-4714
 .69
 .67
 .75
 .75
 .64

 38
 69-202
 -6223-5096
 .87
 .81
 .78
 .73

 40
 70-020
 -5460-4708
 .79
 .71
 .75
 .82
 .79

 44
 70-046
 -6066-5008
 .44
 .49
 .51
 .49
 .73

 46
 70-176
 -6045-5039
 .53
 .53
 .62
 .74
 .83

 50
 71-061
 -5419-4931
 .79
 .74
 .79
 .90
 .75

 53
 71-161
 -5340-4819
 .63
 .65
 .77
 .82
 .59

 55
 72-104
 -5722-4960
 .75
 .67
 .71
 .77
 .78

 56
 72-175
 -6059-5080
 .90
 .91
 .90
 .91

 61
 77-104
 -5747-4487
 .87
 .79
 .82
 .86
 .80

72 77-151 -5230-4748 .79 .71 .71 .78 .63 74 77-158 -5199-4495 .62 .65 .67 .67 .59 76 77-162 -5673-4821 .82 .85 .79 .67 .61 89 78-075 -5598-4575 .78 .81 .82 .66 .53 93 79-050 -5198-4891 .64 .81 .81 .67 .74 94 79-057 -5169-4706 .80 .80 .86 .85 95 79-058 -6012-4899 .83 .75 .81 .85 .77 98 79-075 -5095-4163 .83 101 79-082 -5077-4733 .84 104 79-085 -5377-4989 .62 .65 .76 .80 .72 107 79-091 -5102-4306 .64 .64 125 80-101 -5821-5075 .78 .83 .88 .91 .92 127 80-103 -5361-4491 .76 .70 .75 .82 .83 133 80-110 -6022-4238 .78 .81 .84 .84 .70 -5200 to -5051 (cont.) 136 80-116 -5685-2822 .80 .78 .79 .86 .85 234 XX-001 -6056-3998 .61 .57 .60 .55 .66 235 XX-002 -5323-3562 .70 .74 .77 .87 .84 Av segment correlation .73 .73 .76 .77 .74

Samples from the MWK chronology show strong cross-matching with other independently developed bristlecone pine chronologies (across 5750 years) from Sheep Mountain and Indian Garden. These chronologies are available at: <u>https://catalog.data.gov/dataset/international-tree-ring-data-bank-itrdb</u> and further details are provided in Salzer and Hughes, 2007⁷, 2010⁸, Salzer et al 2009⁹, 2014¹⁰, Kipfmueller, 2010¹¹.

Mean correlation between SHP and MWK of 115 100-year intervals overlapped every 50 years for 5750 years



Tree-ring width measurements of the dissected subsamples used in this study are provided below:

Tree ID	Astronomical calendar (AD/BC with 0)	BCE	BP	80-101B SubSample1 Ringwidth	80-101B SubSample2 Ringwidth
MWK-80-101B	-5270	5271	7220	0.408	0.599
MWK-80-101B	-5269	5270	7219	0.464	0.62
MWK-80-101B	-5268	5269	7218	0.644	0.831
MWK-80-101B	-5267	5268	7217	0.634	0.83
MWK-80-101B	-5266	5267	7216	0.526	0.656
MWK-80-101B	-5265	5266	7215	0.704	0.909
MWK-80-101B	-5264	5265	7214	0.502	0.698
MWK-80-101B	-5263	5264	7213	0.76	0.999
MWK-80-101B	-5262	5263	7212	0.628	0.811
MWK-80-101B	-5261	5262	7211	0.47	0.58
MWK-80-101B	-5260	5261	7210	0.61	0.94
MWK-80-101B	-5259	5260	7209	0.29	0.47
MWK-80-101B	-5258	5259	7208	0.342	0.501
MWK-80-101B	-5257	5258	7207	0.506	0.643
MWK-80-101B	-5256	5257	7206	0.414	0.471
MWK-80-101B	-5255	5256	7205	0.434	0.52
MWK-80-101B	-5254	5255	7204	0.644	0.856

MWK-80-101B	-5253	5254	7203	0.232	0.496
MWK-80-101B	-5252	5253	7202	0.498	0.687
MWK-80-101B	-5251	5252	7201	0.278	0.399
MWK-80-101B	-5250	5251	7200	0.338	0.489

For the 7176 BCE event, sample MWK85-051 from the non-calendar dated portion of the Methuselah Walk bristlecone pine master chronology was selected. This sample derives from a single Pinus longaeva D.K.Bailey tree which grew at the Methuselah Walk site (MWK) in the White Mountains of California in the western USA. The remnant tree was preserved in a dry, high altitude mountain environment. The full tree sampled had 474 rings with no pith, bark or waney edge present. This tree is not part of the existing MWK master chronology but was placed in time by a possible dendrochronological match backed by limited, course resolution radiocarbon dating¹² (Salzer et al. 2019). The sample was selected with the aim of testing the possible placement and using annual ¹⁴C to firmly connect this sample with the MWK master tree-ring chronology, extending this sequence securley back in time. Following dissection of the first sub-sample at annual resolution it became clear that an adjustment of around 40 years was required to this dating in order to detect the 7176 BCE event. A second sampling was carried out, with the floating astronomical calendar (AD/BC with 0) years adjusted to be 40 years older. The annual ¹⁴C analysis of the second subsample detected the 7176 BCE event, indicating a further 1-year adjustment was necessary to align the floating bristlecone pine series with the other, annual ¹⁴C measurements on calendar dated trees. This extends the continuous sequence for MWK over 10,398 years.

			8		
Dating A (AD/BC with 0) NOT CALENDAR SECURE – floating placement: course resolution radiocarbon	Dating A (BP) NOT CALENDAR SECURE – floating placement: course resolution radiocarbon	MWK-85- 051B Subsample1 Ringwidth	(AD/BC with 0) NOT CALENDAR SECURE – floating placement: higher resolution radiocarbon	Dating B (BP) NOT CALENDAR SECURE – floating placement: higher resolution radiocarbon	MWK-85- 051B SubSample2 Ringwidth
-7186	9136	402	-7160	9150	0.682
-7185	9135	260	-7159	9149	1.036
-7184	9134	490	-7158	9148	0.684
-7183	9133	602	-7157	9147	0.472
-7182	9132	656	-7156	9146	0.614
-7181	9131	596	-7155	9145	0.5
-7180	9130	720	-7154	9144	0.194

The dissected years and associated sub-sample ringwidths are provided in the table below:

Dating B

-7179	9129	760	-7153	9143	0.066
-7178	9128	640	-7152	9142	0
-7177	9127	592	-7151	9141	0.25
-7176	9126	496	-7150	9140	0.186
-7175	9125	574	-7149	9139	0.398
-7174	9124	388	-7148	9138	0.646
-7173	9123	796	-7147	9137	0.498
-7172	9122	996	-7146	9136	0.54
-7171	9121	878	-7145	9135	1.476
-7170	9120	690	-7144	9134	0.408
-7169	9119	556	-7143	9133	0.32
-7168	9118	804	-7142	9132	0.218
-7167	9117	630	-7141	9131	0.526
-7166	9116	540	-7140	9130	0.898
-7165	9115	808	-7139	9129	0.61
-7164	9114	732	-7138	9128	1
-7163	9113	924	-7137	9127	1.308
-7162	9112	338	-7136	9126	0.85
-7161	9111	446	-7135	9125	0.346
-7160	9110	884	-7134	9124	0.448
			-7133	9123	0.658
			-7132	9122	0.92
			-7131	9121	1.118
			-7130	9120	1.48
			-7129	9119	0.894
			-7128	9118	0.49
			-7127	9117	0.738
			-7126	9116	0.78
			-7125	9115	0.884
			-7124	9114	0.974
			-7123	9113	1.018

-7122	9112	0.484
-7121	9111	0.572
-7120	9110	0.482
-7119	9109	0.45
-7118	9108	0.402
-7117	9107	0.602
-7116	9106	0.336
-7115	9105	0.58
-7114	9104	0.818
-7113	9103	0.688
-7112	9102	0.632
-7111	9101	0.722
-7110	9100	0.762

The sub-samples were cut to be c. 5 mm thick relative to the transverse section. The upper and lower surface were lightly sanded, preserving original pin-holes, but removing any surface contamination and excess resins which come up to the sample surface over time. The upper, lower and radial surfaces were clearly marked with the pin-holes and measurements re-checked prior to dissection. Dissection was carried out using the two versions of the flating astronomical calendar (AD/BC with 0) scale and carried out at times 20 magnification using a binocular microscope and steel disecting blade. The same blade was used to remove any residual contamination possibilities (pinholes, any residual cells from the previously dissected year of interest) prior to sampling each seperate year of interest. Samples were dissected working from youngest to oldest and packed in pre-marked foil packets.

For both time periods covered by this study, the tree-rings were very narrow and the wood highly resinous making dissection a challenge.

Siberian Larch

Tree rings for radiocarbon analysis of Siberian Larch (*Larix sibirica* Ledeb.) were taken from the sample L4756. This sample was collected on 14 August 2019 from a subfossil stem buried in the alluvial deposits of the Tanlova River (Fig. S1.5) in the southern part of the Yamal Peninsula, which is located in the continuous permafrost zone. Most of subfossil wood found in the area grew near the recent polar tree line. Sample L4756 was found 15 km north of the northernmost currently growing trees. The coordinates of the finding location are 67.75 N 70.06 E, 30 m above sea level.

The sample L4756 contained 302 annual rings, which were dated using a 8768 years long Yamal master chronology¹³ (Fig. S1.6).

Sample L4756, like all the others subfossil samples collected in Yamal, is archived at the Institute of Plant and Animal Ecology, Ural Branch of the Russian Academy of Sciences, Ekaterinburg, Russia.



Figure S1.5. Siberian larch used in this work: a) collecting a sample L4756 (Photo by P.Fonti); b) sample L4756; c) tree rings used for analysis (5272-5247 BCE).



Figure S1.6. Ring width measurements of sample L4756

Dendrochronological dating of two larch samples from the Swiss Alps (ua-1601, tsc-182)

Two larch (*Larix decidua* Mill.) specimens sampled at Alpine glacier forefields were selected for the yearly resolved ¹⁴C analyses. The main selection criterion was comparatively wide tree rings.

tsc-182 Alpine Larch

Specimen tsc-182 was collected from a larch log which was uncovered by the retreating Tschierva glacier (46°23' N, 9°52' E), Switzerland, and subsequently displaced by the glacier creek. Field work for sampling was carried out on 15.9.2007. On the specimen tsc-182, 12 (partial) radii were measured with a resolution of 1/1000 mm and subsequently averaged into a 431-year tree-ring width series. The last 54 rings of this series are slightly compressed due to the former covering by the glacier. According to the crossdating (Tab. S1.7) to the Eastern Alpine Conifer Chronology¹⁴ (Fig. S1.7) the 431-year long series (Tab. S1.8) yielded the calendar dating 5337 – 4907 BCE (historical timescale, without year 0).

ua-1601 Alpine Larch

This larch specimen comes from a detrital log found on the forefield of the Unteraar glacier (46°34′ N, 8°12′ E), Switzerland. The log was sampled in 2016. Four (partial) radii were measured on the specimen with a resolution of 1/1000 mm and subsequently averaged to a 119-year tree-ring width series. The outermost 18 rings of specimen ua-1601 are slightly compressed. This series was crossdated (for statistics see Tab. S1.7) to the Eastern Alpine Conifer Chronology^{14,15} (Fig. S1.7) and returned the calendar dates 7242 – 7124 BCE (Tab. S1.8) (historical timescale, without year 0).

Both specimens are stored in the tree-ring sample archive at the Department of Geography, Universität Innsbruck, Austria. The dissection of the wood samples into single tree-rings for the yearly resolved ¹⁴C measurements was carried out under a microscope using razor blades.

Table S1.7 Crossdating statistics for the comparison of the two Alpine larch series tsc-182 and ua-1601 with the Eastern Alpine Conifer Chronology (EACC, Nicolussi et al. 2009, 2015). Calculations were carried out by using the program WinTSAP. Overlap: no. of years; Glk.: Gleichläufigkeit; Sign. Glk.: pointer interval Gleichläufigkeit; t-value_{BP} and t-value_H: t-values after Baillie and Pilcher as well as Hollstein; dating: historical timescale, without year 0.

Sample	Reference chronology	Overlap [n]	Glk. (%]	Sign. Glk. [%]	$t-value_{BP}$	t-value _H	Dating [year BCE]
tsc-182	EACC	431	63	66	7.5	7.4	5337 - 4907
ua-1601	EACC	119	73	76	5.8	7.0	7242 - 7124



Figure S1.7 Comparison of the tree-ring width series of the Alpine larch specimens' tsc-182 (upper) and ua-1601 (lower) with the Eastern Alpine Conifer Chronology. The series are given on a logarithmic scale, the tree rings used for single-year ¹⁴C analyses are indicated. For ua-1601, a 150-year long section of the series around the ¹⁴C-anaysed samples is shown.

Tab. S1.8 Tree-ring width series of the two Alpine larch specimen tsc-182 and ua-1601 which provided single-ring samples for radiocarbon analyses (used format: Heidelberg).

HEADER: KeyCode=tsc-182 Location=Tschierva glacier Species=Larix decidua DataFormat=Tree SeriesType=Mean curve Unit=1/1000 mm Pith=P WaldKante=---Length=431

HEADER: KeyCode=ua-1601 Location=Unteraar glacier Species=Larix decidua

Supplementary Figures





Supplementary Fig. 1 All Measurement results of the 2 new events. ¹⁴C measurements with 1- σ errors reported as Δ^{14} C of the two newly found events (7176 BCE (a), 5259 BCE (b)) in all different trees compared to the IntCal20 calibration curve¹⁶ (orange band). The Irish Oak was independently repeated by two different labs (ETH-Zurich, Bristol).



Supplementary Fig. 2 Carbon box model used to reconstruct ¹⁴C production. The carbon fluxes between boxes and their carbon contents are given in Gt/yr and Gt.



Supplementary Fig. 3 Evaluation of all known ¹⁴C events. (a) Mean data of known ¹⁴C events with 1- σ errors and result of 1000 Simulations. The fitted Gaussian shaped production spikes for all simulations are also shown. (b) Distribution of the simulated Δ^{14} C increases (blue bars) with a Gaussian fit (dashed line). (c) Distribution of excess ¹⁴C production (blue bars) with Gaussian fit (dashed line).Data Sources: 5410 BCE: Miyake et al.¹⁷, 660 BCE: Sakurai et al.¹⁸, 775 CE and 993 CE: Büntgen et al.¹⁹, 1052 CE and 1279 CE: Brehm et al.²⁰.



Supplementary Fig. 4 Production of ¹⁴C by SEP events for different values of VADM: (a) Global yield functions (Y_{p} , righthand-side axis) of ¹⁴C by protons for three values of VADM (3x, 6x and 9x10²² A m² as blue, orange and green solid lines, respectively); as well as spectral omnidirectional fluences (J(E), left-hand-side axis) of SEPs for two bounding cases, the softest and hardest-spectrum known events of 23-Feb-1956 (GLE 05, blue dashed line) and 04-Aug-1972 (GLE 24, red dashed line). (b) Calculated globally-averaged ¹⁴C production Q_{14C} during the two SEP events as a function of the geomagnetic field VADM. (c) Correction factor $Q(M_o)/Q(M)$ of the ¹⁴C production at the VADM value of M to that for a modern geomagnetic field (M_o =7.8·10²² A m²).



Supplementary Fig. 5 Comparison of two geomagnetic field reconstructions. Two geomagnetic field reconstructions by Knudsen et al.²¹ (blue) and Panovska et al.²² (orange) over the last 12000 years including $1-\sigma$ uncertainty ranges.

References

- 1 Baillie, M. G. L. *Tree-ring dating and archaeology*. (London : Croom Helm, 1982).
- 2 Munro, M. A. R. An Improved Algorithm for Crossdating Tree-Ring Series. (1984).
- 3 Ferguson, C. W. A 7104-Year Annual Tree-Ring Chronology for Bristlecone Pine, Pinus Aristata, from the White Mountains, California. *Tree-ring Bulletin* (1969).
- 4 Stokes, M. A. & Smiley, T. L. *An Introduction to Tree-ring Dating*. (University of Arizona Press, 1996).
- 5 Holmes, R. Computer-Assisted Quality Control in Tree-Ring Dating and Measurement. *Treering Bulletin* (1983).
- 6 LaMarche Jr., V. C. & Harlan, T. P. Accuracy of tree ring dating of bristlecone pine for calibration of the radiocarbon time scale. *Journal of Geophysical Research (1896-1977)* **78**, 8849-8858, doi:10.1029/JC078i036p08849 (1973).
- 7 Salzer, M. W. & Hughes, M. K. Bristlecone pine tree rings and volcanic eruptions over the last 5000 yr. *Quaternary Res* **67**, 57-68, doi:10.1016/j.yqres.2006.07.004 (2007).
- 8 Salzer, M. W. & Hughes, M. K. in *Advances in Global Change Research Advances in Global Change Research* 469-482 (Springer International Publishing, 2010).
- 9 Salzer, M., Hughes, M., Bunn, A. & Kipfmueller, K. Recent unprecedented tree-ring growth in bristlecone pine at the highest elevations and possible causes. *Proceedings of the National Academy of Sciences of the United States of America* **106**, 20348-20353, doi:10.1073/pnas.0903029106 (2009).
- 10 Salzer, M. W., Bunn, A. G., Graham, N. E. & Hughes, M. K. Five millennia of paleotemperature from tree-rings in the Great Basin, USA. *Climate Dynamics* **42**, 1517-1526, doi:10.1007/s00382-013-1911-9 (2014).

- 11 Kipfmueller, K. F. & Salzer, M. W. Linear trend and climate response of five-needle pines in the western United States related to treeline proximity. *Canadian Journal of Forest Research* **40**, 134-142, doi:10.1139/X09-187 (2010).
- 12 Salzer, M. W., Pearson, C. L. & Baisan, C. H. Dating the Methuselah Walk Bristlecone Pine Floating Chronologies. *Tree-Ring Research* **75**, 61-66, doi:10.3959/1536-1098-75.1.61 (2019).
- 13 Hantemirov, R. M. *et al.* An 8768-year Yamal Tree-ring Chronology as a Tool for Paleoecological Reconstructions. *Russian Journal of Ecology* **52**, 419-427, doi:10.1134/S1067413621050088 (2021).
- 14 Nicolussi, K. *et al.* A 9111 year long conifer tree-ring chronology for the European Alps: a base for environmental and climatic investigations. *The Holocene* **19**, 909-920, doi:10.1177/0959683609336565 (2009).
- 15 Nicolussi, K., Weber, G., Patzelt, G. & Thurner, A. A question of time: extension of the Eastern Alpine Conifer Chronology back to 10071 b2k. *GFZ Potsdam, Scientific Technical Report STR15, Potsdam*, 69-73 (2015).
- 16 Reimer, P. J. *et al.* The IntCal20 Northern Hemisphere Radiocarbon Age Calibration Curve (0– 55 cal kBP). *Radiocarbon* **62**, 725-757, doi:10.1017/RDC.2020.41 (2020).
- 17 Miyake, F. *et al.* A Single-Year Cosmic Ray Event at 5410 BCE Registered in 14C of Tree Rings. *Geophys Res Lett* **48**, e2021GL093419, doi:10.1029/2021GL093419 (2021).
- 18 Sakurai, H. *et al.* Prolonged production of 14C during the ~660 BCE solar proton event from Japanese tree rings. *Scientific Reports* **10**, 660, doi:10.1038/s41598-019-57273-2 (2020).
- 19 Büntgen, U. *et al.* Tree rings reveal globally coherent signature of cosmogenic radiocarbon events in 774 and 993 CE. *Nature Communications* **9**, doi:10.1038/s41467-018-06036-0 (2018).
- 20 Brehm, N. *et al.* Eleven-year solar cycles over the last millennium revealed by radiocarbon in tree rings. *Nature Geoscience* **14**, 10-15, doi:10.1038/s41561-020-00674-0 (2021).
- 21 Knudsen, M. F. *et al.* Variations in the geomagnetic dipole moment during the Holocene and the past 50 kyr. *Earth and Planetary Science Letters* **272**, 319-329, doi:10.1016/j.epsl.2008.04.048 (2008).
- 22 Panovska, S., Constable, C. G. & Korte, M. Extending Global Continuous Geomagnetic Field Reconstructions on Timescales Beyond Human Civilization. *Geochemistry, Geophysics, Geosystems* **19**, 4757-4772, doi:10.1029/2018GC007966 (2018).