



Divergent evolutionary trajectories of bryophytes and tracheophytes from a complex common ancestor of land plants

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Supplementary Methods

1. Selecting gene families for phylogenetic analyses

Whole genome duplications are prevalent in embryophytes⁵², leading to multiple gene copies and paralog-dominated gene families. As a result, single copy orthologous gene families used to construct supermatrices are difficult to identify – especially when sampling taxon-rich datasets. We employed OrthoFinder 2.0⁶⁷ to infer single copy orthologs from our dataset of 177 species of plants and algae. No single copy orthologous gene families were inferred. In previous studies single copy orthologous gene families have been manually curated²⁰. We chose to develop a systematic algorithm identifying the gene families closest to single copy (Supplementary Fig. 21). The filtering method employed ensured only high-quality gene families comprised the final alignment – with only 2/177 species being represented by more than 50% gaps. The algorithm is split into two scripts *prem3.py* and *super_matrix_2.py* - prior to calling *super_matrix_2.py* the gene families need to be aligned and trimmed. This can be done prior to use, however, aligning all orthogroups prior to filtering would require considerable time and computational resources.

2. Transcriptome Assembly

Raw sequencing reads were downloaded from the Sequence Reads Archive (SRA) and trimmed using trimmomatic⁹⁵ under default settings to remove adapter sequences and trim low-quality bases. Transcriptomes were de novo assembled in Trinity⁹⁶. Transcriptomes were filtered using the EvidentialGene pipeline⁹⁷ to remove isoforms and translate transcripts into amino acid sequences.

3. Root inference with Amalgamated Likelihood Estimation (ALE)

ALE models gene duplication, transfer and loss (DTL) to assess the likelihood of gene families under differing roots. Due to our phylogenetic dataset (Supplementary Table 12) containing both genomes and transcriptomes, the incomplete nature of the transcriptomes might be difficult to model accurately. In Supplementary Fig. 11, the fraction missing of each lineage is displayed (detailed in Supplementary Table 12); for many lineages, more than half of the BUSCOs are absent. As a result, we decided to curate a new dataset comprising only high-quality and more complete genomes. The dataset was comprised of 24 species, with each of the seven major lineages of embryophytes being represented by at least one genome (Supplementary Table 13). Where possible, we used “selected” proteome sets which have already been filtered to remove isoforms and splice variants. Otherwise, proteomes were filtered using CD-HIT with a cut-off of 0.99. 23 out of 24 species had more than 70% BUSCO completion, with most of the data being more than 80% complete. Moreover, the gene family clusters inferred from the high-quality dataset had greater species representation and thus might be more informative for inferring the root (Supplementary Fig. 13-14). There were also fewer large clusters with poor species representation in the high-quality dataset i.e., fewer gene families that have either undergone myriad duplications or had multiple isoforms of genes from sequencing anomalies. As a result, we only took forward the high-quality dataset for further analysis.

The high-quality gene families were analysed by ALEml – ALEml is time-consistent with the fossil-calibrated species tree and avoids time-inconsistent gene transfers (that is, gene transfers travelling back in time). We ran ALEml on 12 rooted and dated embryophyte trees, with an AU test suggesting three credible roots: hornworts, moss, and monophyletic bryophytes

(Supplementary Table 1). The credible roots were all in the same region of the tree, and the likelihood of the roots decreased with increasing distance from the root region. Moreover, analysis of the DTL signal (Supplementary Fig. 1), revealed that the signal was stable for this set of credible roots.

We also employed two additional techniques to test the root of embryophytes to ensure confidence in our topology for the reconstruction of ancestral states. Firstly, we inferred constrained maximum likelihood trees from the concatenation to be in accordance with the three credible roots selected by ALE i.e., hornworts were constrained to group with algae at the base of the tree, and thus branch from the root of embryophytes. The constrained trees were inferred in IQ-Tree⁷⁰, under the LG+C60+G4+F model. The likelihood of the three constrained trees was then assessed with an AU test⁹¹ – rejecting both moss and hornworts as earliest diverging lineages of embryophytes (Supplementary Table 2). Secondly, we ran STRIDE³⁶, an alternative rooting software which employs duplication events to determine the root. STRIDE in combination with OrthoFinder 2.0⁶⁷, reconciled 18560 orthogroups, and suggested three most-parsimonious roots: between bryophytes and tracheophytes, on the hornwort stem and on the liverwort stem (Fig. 1). These results agree in part with the ALE analysis and the outgroup-rooting analyses, but propose an additional potential root on the liverwort stem. The support from the STRIDE analysis for the hornwort-stem root was weak (0.2% probability).

4. Ancestral gene content reconstruction

In addition to reconstructing the gene content of the ancestral embryophyte, we included several algal species that had highly complete genomes; the inclusion of algae allowed us to investigate gene content evolution in the closest algal relatives of embryophytes. The inclusion of algae also enabled us to investigate further the gene loss proposed to underpin the reduced body plan of the embryophyte sister lineage, Zygnematophyceae. In addition to algae, a subspecies of *Marchantia polymorpha (ruderalis)*, was added to the dataset, allowing us to calculate values for DTL and copy number for the internal node defining liverworts. Single copy orthologs were determined using the same method as the main analysis (Supplementary Fig. 15) – 185 single copy gene families were inferred, leading to an alignment of 71855 sites. A maximum likelihood tree (Supplementary Fig. 3) was inferred from the dataset under the LG+C60+G+F model in IQ-Tree⁷⁰. The tree was rooted in accordance with the previous analysis. Again, even with this reduced dataset, we recover monophyletic bryophytes with high support (97% Bootstrap support). Due to ALEml requiring a dated tree for reconstruction of ancestral gene content, a time scale was inferred using a subset of the calibrations from the transcriptomic analysis (Supplementary Fig. 4). Divergence time analyses were performed as described for the transcriptomic analysis.

Gene family clusters were inferred using the same methods as the rooting analysis. The pipeline yielded 20822 gene family clusters with 1000 bootstrap samples for each family. The gene families were reconciled with the rooted species tree, using the ALEml algorithm with the fraction missing of each genome also incorporated. The presence and absence of each gene family across the tree was inferred using the *Ancestral_reconstruction_copy_number.py* python script – we determined that a gene family was present if it had more than 0.5 copies at the specified branch i.e., was present at the branch in more than 50% of the reconciliations. The summed number of DTL evolutionary events was determined with *branchwise_number_of_events.py*. In addition to this, we scaled the number of events by the branch length of the time scale analysis, to determine the rate of DTL for important sections of the tree. The number of DTL events and the summed gene copy number for each branch of the tree were calculated (Supplementary Table 3). The presence and absence of each gene family

cluster across the embryophyte tree, and whether an *Arabidopsis* gene is present in said cluster is displayed in Supplementary Table 4.

Clustering proteomes into gene families remains a high-throughput and potentially inaccurate process. To assess the robustness of our results to poorly-clustered or erroneous gene families, we undertook several filtering analyses. Gene families where the total trimmed alignment length was short may represent poorly designated gene families and so we progressively removed families with short alignments (>70, 60, 50 and 30 amino acids). and repeated the analysis of duplications, losses and transfers. We found that filtering by alignment length did result in fewer observed losses, but this was consistent across clades and so the losses observed in bryophytes are not the result of poorly aligned gene families. Very large gene families may also represent gene families where the clustering algorithm has failed to distinguish individual gene families, and so we progressively removed the top 1, 2.5, 5 and 10% gene families by size. We found that removing the largest gene families had a large effect on the number of inferred losses, especially among bryophytes, and that removing the largest 10% of gene families results in more inferred losses in tracheophytes than in bryophytes. This suggests that the losses in bryophytes are primarily occurring in the largest gene families. In a parsimony approach, loss in bryophytes would typically require a gene to also be present in the algal outgroups. This is not the case in a probabilistic framework with a single loss rate applied across the gene family, where high rates of loss in one part of the tree can drive losses in other regions of the tree. We repeated our analyses with a subset of gene families that contain at least one copy in one of the algal outgroups. We found that while this did decrease the number of losses along the bryophyte stem, it did so proportionally across both bryophytes and tracheophytes. Finally, high rates of duplication, transfer and loss may suggest that a gene family is poorly clustered or that its evolutionary history is poorly resolved. We repeated our analyses while progressively removing the top 1, 2.5, 5 and 10% of gene families by duplication, loss and transfer rate. We found that filtering by duplication and transfer rate had only a very small effect on the inferred number of losses, while filtering by loss rate unsurprisingly had a large effect. The number of losses decreased most in bryophytes, suggesting that the gene families with a high loss rate have disproportionately more losses in bryophytes.

To functionally annotate the gene families, the consensus sequence of each gene family was determined using HMMER3⁹². This process summarises the gene family into a single sequence most indicative of the ancestral state. The consensus sequences were functionally annotated using eggNOG-mapper 2⁹³. The gene families present at each branch were determined as described above, and the functional annotation for their corresponding consensus sequence was retrieved. The frequency of GO terms associated with the consensus sequences present at each branch was calculated using the custom python script *make_go_term_dictionary.py*. In doing so, the change in the frequency of GO terms could be tracked from node to node, allowing us to identify changes in function across the embryophyte tree. Changes in the frequency of GO terms from the embryophyte ancestor to bryophytes, and to tracheophytes were calculated – this is termed the GO term differential i.e., if there was a frequency of 10 for the GO term “Shoot Development” at the ancestral embryophyte branch, but only a frequency of 6 for the bryophyte branch, the GO term differential would be -4. The GO term differentials are presented in Extended Data Fig. 5: on the left-hand side, all GO term differentials are plotted, i.e., the change in every GO term differential between the ancestral branch and the subsequent tracheophyte and bryophyte branch is plotted. Overall, there has been a reduction in the frequency of GO terms in bryophytes in comparison to the ancestral embryophyte, suggesting there has been a wide-scale genomic reduction in this lineage. This was also suggested by the analysis of DTL rates which suggested a vast amount of loss on this branch. On the right, changes in select GO terms/functional processes are highlighted.

Reduction in the frequency of GO terms associated with reproductive shoot and vegetative shoot development are observed between the ancestral embryophyte and the bryophytes, whilst the inverse is true for tracheophytes (Supplementary Tables 5, 7 and 9).

One limitation of GO term analysis is that differences on the bryophyte and tracheophyte stems might reflect, in part, differences in the completeness of functional annotation for the extant members of the two clades. To investigate the broad-scale patterns identified in the GO term analysis in more detail, we investigated the evolutionary histories of 50 key genes involved in *Arabidopsis* stomatal function/development (Supplementary Table 10), vasculature development and symbiosis/defence (Supplementary Fig. 9-10). We were able to determine where the gene families that these genes belong to originated, were present, and were lost across the tree. Many of the orthologous gene clusters that harbour critical constituents of the stomatal development pathway such as SPCH/MUTE/FAMA, TMM, and ERECTA originated on the ancestral embryophyte branch, suggesting that stomata evolved at this point. Moreover, orthologous gene families of master-regulators of stomatal function such as HT1 and OST1 were also present. In addition to this, gene families which contain key *Arabidopsis* vasculature genes such as WOX4 and WER (MYB66) originated on the branch defining Embryophytes – highlighting a possible increase in vasculature complexity. Interestingly, our analysis suggests that these gene families were lost on the branch defining bryophytes (Supplementary Fig. 9-10), suggesting a reduction in vasculature complexity in this lineage.

To further assess the losses of important developmental genes on the branch defining bryophytes, manually curated phylogenetic trees were constructed for their respective gene families (Extended Data Fig. 6). The gene trees were inferred by first undertaking an OrthoFinder analysis of an expanded dataset, comprising further representatives of each land plant lineage. The homologues were identified within orthogroups and then aligned with MAFFT (--globalpair) and trimmed with BMGE. For each alignment, a phylogenetic tree was inferred in IQ-Tree under the best fitting model determined by BIC. The trees were then rooted in accordance with the previous analysis (Fig. 1). For the representative genes, a pattern of loss in bryophytes and concurrent expansion via duplication in tracheophytes is observed. Additionally, we see wide scale loss of gene families which harbour critical stomata and vasculature associated genes on the branches that define hornworts and liverworts (Supplementary Fig. 9-10). The loss of stomatal-associated gene families in these lineages is congruent with the findings of Harris *et al.*, (2020). Presently, the number of available genomes for non-flowering plants is limited and is a potential source of bias for estimating ancestral genome content in deeper nodes of the tree. We expect, with the increasing volume of sequence data and projects specifically targeting non-flowering plants (e.g. 10000 genome project, Sphagnum genome project), that it will be possible to develop a still more detailed reconstruction of the ancestral embryophyte.

5. Molecular clock estimates of the timing of land plant evolution

The age of land plants has become a recent controversy, with some studies favouring a far older (pre-Cambrian) origin of crown embryophytes^{23,40,41}, while others limit the origin of land plants to the late Cambrian. These differences reflect diverging approaches to the implementation of maximum age constraints derived from the fossil record. Previous studies have employed a maximum age derived from the earliest record of cryptospores. Cryptospores are spores that possess certain characteristics of embryophyte spores, but are sufficiently different to make their assignment to any extant clade difficult. However, the possibility of an embryophyte affinity cannot be disregarded. Cryptospores possess a similarly high preservation potential to the spores of extant land plants, and indeed from the Ordovician

onwards there is a good fossil record of spores that are assigned to extant land plant lineages. This provides the basis of the maximum age calibration, following the reasoning that if land plant spores did exist prior to the Cambrian, they would likely be preserved.

To address recent discussions about the value of these maxima, we implemented two additional calibration strategies. In the first (Strategy B), we relaxed the late Cambrian maximum age calibration on crown embryophytes (and all internal nodes that share the same calibration). This was achieved by setting an arbitrary ancient maxima of 1 Ga. The maximum age on tracheophytes (derived from similar reasoning regarding the presence of trilete spores), was retained. In the second strategy, we relaxed the maximum calibration on tracheophytes also to 1 Ga, providing no effective maxima on the major lineages of land plants. Analyses were performed as described in the methods of the main text and the impact on the age of crown embryophytes across calibration strategies is compared in Extended Data Fig. 4.

The absence of deep fossil calibrations in hornworts has led to greater uncertainty in the estimates of their origin. We employed a novel relative constraint based on the horizontal gene transfer of the NEOCHROME gene from hornworts into ferns to force hornworts to be older than certain fern lineages. Topological uncertainty in the NEOCHROME gene tree suggests multiple potential placements of the gene transfer, with the most probable being into the ancestor of Polypodiales. We accounted for this uncertainty by repeating the analyses with multiple relative constraints, forcing hornworts to be older than Polypodiales, Cyatheales and Gleicheniales, all of which variously contain copies of the NEOCHROME gene. Forcing hornworts to be older than the more ancient lineages (Cyatheales and Gleicheniales) resulted in more ancient estimates of the age of crown-group hornworts, including the possibility an Ordovician origin (Extended Data Fig. 3).

6. Fossil Calibrations for Molecular Clock Analyses

C_1: MRCA Viridiplantae/Chloroplastida

Fossil Taxon and Specimen: *Proterocladus antiquus* [VPIGM-4762, deposited at Virginia Polytechnic Institute Geoscience Museum], from the lower Nanfen Formation in North China [1].

Phylogenetic justification: Macrofossils from the Nanfen Formation, Northern China, have reconstructed an erect, epibenthic, multicellular alga. Tang *et al.* [1] argue that the overall morphology of *P. antiquus* is comparable to extant members of Ulvophyceae, yet no diagnostic features are described. As such, *P. antiquus* can conservatively be used to provide a minimum calibration on the origin of Viridiplantae.

Minimum age: 940.4 Ma

Maximum age: 1891 Ma

Age justification: The Nanfen Formation is positioned within the Xihe group in southern Liaoning Province, North China [2]. The Nanfeng Formation has not been dated directly. The age of the underlying Diaoyutai Formation can only be constrained by detrital zircons, the youngest population of which has been dated to 1056 Ma \pm 22 Ma [3]. A sill in the overlying Qiaotou Formation, which overlies the Nanfeng Formation, has a zircon secondary-ion mass spectrometry U-Pb age of 947.8 Ma \pm 7.4 Myr [4]. Thus, the minimum age of the Nanfeng Formation can be established on the minimum age interpretation of this sill viz. 940.4 Ma, although it is presumably significantly older. The soft maximum age is established following Morris *et al.* [5] and is based on the earliest record of eukaryotes, where despite the presence of simple eukaryotes, there is no evidence of complex or multicellular eukaryotes such as green algae.

Discussion: *Proterocladus* is determined as a crown group member of green algae based on the cellular differentiation, presence of a holdfast, inferred siphonocladous nature, and lack of an outer sheath. The absence of an outer sheath and siphonocladous nature precludes an affinity with stigonematalean cyanobacteria. A fungal affinity was also rejected based on the structure of the septae, lack of mycelium-like branching and the structure of the reproductive organs. The septation in *Proterocladus* also suggests that it is not similar to the xanthophycean alga *Vaucheria* or the rhodophyte *Griffithsia*. Nie *et al.* consider a Mesoproterozoic specimen identified as *Pterospermella* as a member of Prasinophyceae [6, 7], however the lack of clearly described structures prevents their unequivocal assignment to the green algae [5].

C_2: Chlorophyceae – Prasinophyceae

Fossil taxon and specimen: *Palaeocymopolia silurica*

Phylogenetic justification: Following Morris *et al.* [5]

Minimum age: 438.3 Ma

Maximum age: 1891 Ma

Age justification: Following Morris *et al.* [5]

C_3: Streptophyta: “Streptophytes” - Embryophytes

Fossil taxon and specimen: *Tetraedraletes cf. medinensis*

Phylogenetic justification: Following Morris *et al.*, [5]

Minimum age: 469 Ma

Maximum age: 1891 Ma

Age justification: Following Morris *et al.* [5]

C_4: Embryophyta: Bryophyta – Tracheophyta

Fossil taxon and specimen: *Cooksonia cf. pertyi*

Phylogenetic justification: Following Morris *et al.* [5]

Minimum age: 426.9 Ma

Maximum age: 515.5 Ma

Age justification: Following Morris *et al.* [5]

C_5: Marchantiophyta: Jungermanniopsida - Marchantiopsida

Fossil taxon and specimen: *Metzgeriothallus sharonae* [NYSM 17656, Paleontology Collection of the New York State Museum, Albany, USA], from the Sample CHD-4.6, Plattekill Formation of the eastern Hamilton Group, Givetian (late Middle Devonian) at Cairo Highway Department quarry, just south of New York State Route 145, 3.22 km NW of Cairo, Greene County, New York (42.32° N and 74.04°W, NAD 83).

Phylogenetic justification: Morris *et al.* [5] identified *Riccardiothallus devonicus* as the oldest liverwort but its compression-preservation precludes the identification of anatomical characters that might corroborate a liverwort affinity beyond its general form [8]. *Metzgeriothallus sharonae*, preserves anatomical detail including putative oil body cells [9, 10], bifurcating thalloid habit with an axial costa, dorso-ventral cellular differentiation, ventral ribbon-like rhizomes, and associated sporophyte that has a multivalve structure [11]. These characteristics support the interpretation of *Metzgeriothallus sharonae* as a crown group liverwort, as supported by a cladistic analysis of morphological characters [12]. Flores *et al.* placed *M. sharonae* as sister to the extant genus *Metzgeria*, albeit with very low bootstrap support (50%).

As such, we conservatively consider *M. sharonae* as a member of the Jungermanniopsida and thus capable of calibrating the divergence between Jungermanniopsida and Marchantiopsida.

Minimum age: 377.7 Ma

Maximum age: 515.5 Ma

Age justification: Attempts to resolve the age of the Plattekill Formation at Cairo Quarry through palynostratigraphy have proven unsuccessful due to poor preservation [13], however, the age of the Plattekill Formation has been established elsewhere. Willis and Bridge [14] placed the Plattekill and overlying Manorkill Formations in the mid *lemurata magnificus*–mid *optivus triangulatus* zones of Richardson and McGregor [15], indicating a latest Eifelian-early Givetian age [16]. In the absence of better stratigraphic constraint, this allows us to establish a minimum constraint on the age of the Plattekill Formation based on the Givetian-Frasnian boundary, dated to 378.9 Ma \pm 1.2 Myr [16], thus 381.7 Ma.

C_6: Anthocerotophyta: Anthocerotaceae

Fossil taxon and specimen: *Anthoceros* sp. (spore type A) [BA Pb Pm 5034, Museo Argentino de Ciencias Naturales] from the Albian Anfiteatro de Ticó Formation, Baqueró Group, Santa Cruz Province, Argentina

Phylogenetic justification: This spore was assigned to the extant genus *Anthoceros* based on the similar morphology [17, 18].

Minimum age: 116.59 Ma

Maximum age: 515 Ma

Age justification: The spores were recovered from the Anfiteatro de Ticó Formation [18], which is overlain by the Bajo Tigre Formation, constrained based on U/Pb dating to 116.85 \pm 0.26 Ma [19]. Thus, a minimum age of 116.59 Ma is proposed.

C_7: Anthocerotophyta: Notothyladaceae

Fossil taxon and specimen: *Notothylites nirulai* from the Deccan Intertrappean beds of Mohgaon-Kalan, Madhya Pradesh, India [20].

Phylogenetic justification: The overall morphology has a similar thallus size, sporophyte size and elater shape to extant *Notothylas*, as well as lacking stomata on the sporophyte [17].

Minimum age: 65.81 Ma

Maximum age: 515 Ma

Age justification: The age of the Intertrappean beds at Mohgaon-Kalan was determined to be Maastrichtian on the palynological assemblage being characterised as Maastrichtian and the magnetostratigraphy indicating the flows to be of C30N (Maastrichtian) [21]. Thus, a minimum age based on the base of the C30N polarity chron of 66.31 \pm 0.5 is proposed [22].

C_8: Anthocerotophyta: *Phaeomegaceros*

Fossil taxon and specimen: *Phaeoceros* sp. from the Uscari Formation, Costa Rica

Phylogenetic Justification: Though originally assigned to the extant genus *Phaeoceros* [23], the spore is highly similar to the extant genus *Phaeomegaceros* in size and morphology (Villareal and Renner)

Minimum age: 13.82

Maximum age: 515 Ma

Age justification: Material was deposited in the Uscari Formation in the Limon Basin of Costa Rica. Analysis of planktonic foraminifera identified an overlap of *Orbulina universa* and *Globorotalia foshi peripheroronda*, placing the the Formation in the early Middle Miocene (foraminiferal zones N9 to N10) [24]. Following Hilgen et al. [25], the base of the N10

formaniferal zone correlates with the M7 zone and the base of the Langhian and thus a minimum age of 13.82 Ma is derived.

C_9: Marchantiophyta: Pelliidae - [Jungermannidae + Metzgeriidae]

Fossil taxon and specimen: *Pellites hamiensis* [SDL-98-4-91, Institute of Paleontology and Stratigraphy of Lanzhou University, Gansu Province, China] from the Sandaoling coal mine, Hami City, Xinjiang Uygur Autonomous Region, China.

Phylogenetic justification: The gross morphology and anatomy of the thalli support a placement within Pelliaceae [26], including ribbon-like segments, thinning wings, a unistratose margin, a conspicuous costa with elongate cells and thick rectangular cells near the margin. The absence of reproductive structures led the authors to assign the samples to the fossil genus *Pellites* rather than *Pellia* [27].

Minimum age: 167.0 Ma

Maximum age: 515.5 Ma

Age justification: Material was collected from the Xishanyao Formation of the Sandaoling coal mine, which overlies the lower Jurassic Sangonghe Formation and underlies the Palaeogene Shanshan Group strata [28]. Based on biostratigraphy and plant assemblages, an Aalenian to Bajocian age is proposed for the Xishanyao formation [29]. Thus, we propose a minimum age according to the upper boundary of the Bajocian, 168.2 ± 1.2 Ma [30].

C_10: Marchantiophyta: Sphaerocarpaceae – Marchantiales

Fossil taxon and specimen: *Marchantites cyathoides*

Phylogenetic justification: Following Morris *et al.* [5]

Minimum age: 227.6 Ma

Maximum age: 515.5 Ma

Age justification: Following Morris *et al.* [5]

C_11: Marchantiophyta: Ricciaceae – Conocephalaceae

Fossil taxon and specimen: *Ricciopsis ferganica* [FG 596/X/721 in the Palaeontological collection of the Geologisches Institut, Technische Universität Bergakademie Freiberg, Germany] from the village of Madygen, southwest Kyrgyzstan, Central Asia

Phylogenetic justification: *R. ferganica* was assigned to the fossil genus *Ricciopsis* (Ricciaceae) [31] based on resemblance to the thalloid floating forms of the extant genus *Riccia*. No diagnostic features are available for the genus, however a flattened, small branching thallus and the presence of a longitudinal median furrow or groove on the dorsal side of the thallus is characteristic of Ricciaceae [32]. Further, a cladistic analysis of morphological characters supported the placement of *Ricciopsis* as sister to the extant genus *Riccia* [12].

Minimum age: 227.6

Maximum age: 515.5 Ma

Age justification: Material was collected from the lower and middle parts of the uppermost lacustrine unit at Urochishche Madygen. The Madygen Lagerstätten is deemed to be Ladinian-Carnian (Middle to Late Triassic) in age based on comparative studies of the flora [33] and insect assemblages [34]. Thus, we propose a minimum age based on the top of the Carnian of 227.3 ± 0.1 Ma [30].

C_12: Marchantiophyta: Jungermannidae – Metzgeriidae

Fossil taxon and specimen: *Cheirorhiza brittae* [515-38a, Far East Geological Institute, Far East Branch of the Russian Academy of Science] from the Bureja basin, Russia

Phylogenetic justification: This genus has complicate-bilobed, subtransversely inserted leaves, massive stems with a small ventral merophyte and narrow, elongated cortical cells, as well as scattered, unbranched rhizoids

Phylogenetic justification: *C. brittae* has been compared to extant members of both Jungermanniales and Porelliales on the basis of morphology [35, 36], and so we follow Heinrichs et al. [37] and treat it as a stem lineage member of Jungermanniidae.

Minimum age: 143.1

Maximum age: 515.5

Age justification: Material was collected from the Talyndzhan formation of the Bureja basin. The Talyndzhan formation was initially determined to be of Callovian or Callovian-Oxfordian age based on floral assemblage. This has been revised in light of palynological evidence and the top of the Talyndzhan is believed to correspond to the Jurassic-Cretaceous boundary [38, 39]. Thus, a minimum age based on the upper boundary of the Jurassic 143.1 [30], and so 144.9 Ma, is proposed.

C_13: Marchantiophyta: Jungermanniales – Porellales

Fossil taxon and specimen: *Diettertia montanensis* [no. 281 of the University of Montana Paleontological Collection (UMPC)] from the Missouri river, Montana, USA.

Phylogenetic justification: This genus has complicate-bilobed, subtransversely inserted leaves, massive stems with a small ventral merophyte and narrow, elongated cortical cells, as well as scattered, unbranched rhizoids

Phylogenetic justification: *Dietteria montanensis* has been assigned as a member of Jungermanniales stem lineage based on its overall morphology and anatomy, including complicate-bilobed, subtransversely inserted leaves, stems with a small merophyte with narrow elongated cortical cells [37, 40].

Minimum age: 112.8 Ma

Maximum age: 515.5 Ma

Age justification: Material was collected from the Kootenai Formation, Great Falls, Montana. Fossil assemblages, particularly molluscs, ostracods and charophyte algae indicate an Aptian age for the formation and an Albian age for the overlying Blackleaf formation [41, 42]. Thus, a minimum age is derived from the Aptian-Albian boundary, estimated at 113.10 ± 0.3 Ma, of 112.8 Ma.

C_14: Marchantiophyta: Porellineae – [Radulineae + Jubulineae]

Fossil taxon and specimen: *Gackstroemia cretacea* [AMNHBU ASJH-1, American Museum of Natural History, New York], from Myanmar [43].

Phylogenetic justification: Heinrichs et al. [43] placed *G. cretacea* within the Lepidolaenaceae based on a sterile gametophyte, including incubous foliation, complicate bilobed leaves, splitting to a dorsal lobe and a ventral *Frullania*-type water sac. Based on the presence of bifurcate underleaves, the fossil was assigned to the extant genus *Gackstroemia*. Lepidolaenaceae are placed within the Porellineae, and so *G. cretacea* can calibrate the divergence of *Porella* (Porellineae) from *Radula* (Radulineae) and *Frullania* (Jubulineae).

Minimum age: 98.17

Maximum age: 515.5 Ma

Age justification: The age of Burmese amber has long been contested, with estimates ranging between the Miocene and Cretaceous. Claims of an upper Albian age based on ammonite biostratigraphy [44] have not been verified [45]. However, a U-Pb age of 98.79 Ma \pm 0.62 Myr was obtained from zircons in volcanoclastics deposited contemporaneously with the amber [46]. Shi et al. (2012) and Yu et al. (2019) interpret this as a minimum age for the amber since it is bioeroded. Nevertheless, in the absence of other evidence, this U-Pb zircon age can be used to establish a minimum constraint of 98.17 Ma [44].

C_15: Marchantiophyta: Radulaceae – [Frullaniaceae – Lejeuneaceae]

Fossil taxon and specimen: *Radula cretacea* [PB22484, Nanjing Institute of Geology and Palaeontology, Chinese Academy of Sciences, Nanjing], from amber localities near the village of Tanai, Kachin State, Myanmar [47].

Phylogenetic justification: The specimen was assigned to the genus *Radula* (Radulaceae, Porellales) based on acute to acuminate leaf apices, gemmae exclusively produced from leaf-lobe marginal cells and female bracts present in two pairs [47].

Minimum age: 98.17

Maximum age: 515.5 Ma

Age justification: The age of Burmese amber has long been contested, with estimates ranging between the Miocene and Cretaceous. Claims of an upper Albian age based on ammonite biostratigraphy [44] have not been verified [45]. However, a U-Pb age of 98.79 Ma \pm 0.62 Myr was obtained from zircons in volcanoclastics deposited contemporaneously with the amber [46]. Shi et al. (2012) and Yu et al. (2019) interpret this as a minimum age for the amber since it is bioeroded. Nevertheless, in the absence of other evidence, this U-Pb zircon age can be used to establish a minimum constraint of 98.17 Ma [44].

C_16: Marchantiophyta: Frullaniaceae – Lejeuneaceae

Fossil taxon and specimen: *Frullania cretacea* [AMNH B-011, American Museum of Natural History, New York], from Myanmar [48]

Phylogenetic justification: *F. cretacea* was initially placed in the extant genus *Frullania* based on the reddish leaf colour, leaf cells with a single mammilla, lateral branching and helmet-shaped water sacs [49]. More recent findings support this placement, with the identification of a stylus shape similar to extant *Frullania*, bilobed underleaves and more elongate and toothed subgynoecial underleaves [48].

Minimum age: 98.17

Maximum age: 515.5 Ma

Age justification: The age of Burmese amber has long been contested, with estimates ranging between the Miocene and Cretaceous. Claims of an upper Albian age based on ammonite biostratigraphy [44] have not been verified [45]. However, a U-Pb age of 98.79 Ma \pm 0.62 Myr was obtained from zircons in volcanoclastics deposited contemporaneously with the amber [46]. Shi et al. (2012) and Yu et al. (2019) interpret this as a minimum age for the amber since it is bioeroded. Nevertheless, in the absence of other evidence, this U-Pb zircon age can be used to establish a minimum constraint of 98.17 Ma [44].

C_17: Bryophyta: Sphagnopsida – [Polytrichopsida – Bryopsida]

Fossil taxon and specimen: Sphagnales

Phylogenetic justification: Following Morris *et al.* [5]

Minimum age: 330.7 Ma

Maximum age: 515.5 Ma

Age justification: Following Morris *et al.* [5]

C_18: Bryophyta: Polytrichopsida – Bryopsida

Fossil taxon and specimen: *Palaeocampylopus buragoase*

Phylogenetic justification: Following Morris *et al.* [5]

Minimum age: 271.8 Ma

Maximum age: 515.5 Ma

Age justification: Following Morris *et al.* [5]

C_19: Bryophyta: Polytrichaceae – Tetrarhizaceae

Fossil Taxon and Specimen: *Meantoinia alophosioides* [Gemiferous gametophyte shoot in rock slab UAPC-ALTA P15393 B (slides B bot series a), University of Alberta Paleobotanical Collections (UAPC-ALTA), Edmonton, Alberta, Canada] from the Lower Cretaceous deposits at Apple Bay (Vancouver Island, British Columbia, Canada) [50].

Phylogenetic Justification: Bippus *et al.* [50] placed *M. alophosioides* within the extant family Polytrichaceae based on the presence of leaves with diagnostic anatomy, differentiated into sheathing base and free lamina, bearing photosynthetic lamellae, and a conducting strand present in the stem. A phylogenetic analysis of 91 morphological characters confirmed the position of *M. alophosioides* as a stem group member of Polytrichaceae [51].

Minimum age: 132 Ma

Maximum age: 515.5 Ma

Age justification: Sediments of the Apple Bay locality are regarded as Longarm Formation equivalent and originally proposed as Barremian based on the presence of monosulcate reticulate pollen grains [52]. Klymiuck and Stockey [53] argued that these grains are more likely associated with extinct gymnosperms, and instead propose a Valanginian age, supported by oxygen isotope analysis (D. Gröcke, personal communication, 2005, 2013). Thus, we propose the upper boundary of the Valanginian 132.6 ± 0.6 Ma and thus, 132 Ma as a minimum age constraint.

C_20: Bryophyta: *Atrichum* – *Polytrichum*

Fossil Taxon and Specimen: *Eopolytrichum antiquum* [PP44714, Department of Geology, Field Museum, Chicago] from Buffalo Creek member of the Gaillard Formation (Upper Cretaceous), Crawford County, Georgia, USA [54].

Phylogenetic Justification: Konopka *et al.* [54] determined that *Eopolytrichum* was a member of the extant family Polytrichaceae based on several features including the epiphragm, broadened columella apex, peristomial membrane structure, and echinulate spore sculpture, as well as the overall lamellae and leaf anatomy associated with Polytrichaceae. Based on a cladistic analysis of 99 characters Bippus *et al.* [51] recovered *Eopolytrichum* nested within the extant genus *Polytrichum*.

Minimum age: 82.9 Ma

Maximum age: 515.5 Ma

Age justification: *Eopolytrichum* was recovered from the sediments of the Allon quarry, assigned to the Buffalo Creek Member of the Gaillard Formation. Based on palynological analyses, a Campanian age was initially inferred [55]. However, this was later revised to a late Santonian age based on the correlation of the Allon locality palynological assemblage and that

of the Coastal Plain Province [56], which was assigned to Burnett's calcareous nannofossil zone CC17 [57], the base of which correlates with the Santonian/Campanian boundary. Thus, a minimum age constraint of 83.6 ± 0.7 is proposed.

C_21: Bryophyta: Funariidae – [Dicraniidae + Bryidae]

Fossil taxon and specimen: *Capimirinus riopretensis*

Phylogenetic justification: Following Morris *et al.* [5]

Minimum age: 268.3 Ma

Maximum age: 515.5 Ma

Age justification: Following Morris *et al.* [5]

C_22: Bryophyta: Dicraniidae – Bryiidae

Fossil Taxon and Specimen: *Tricarinella crassiphylla* [Gametophyte shoot in rock slab UAPC-ALTA P13311, University of Alberta Paleobotanical Collections, Edmonton, Alberta, Canada], from the Apple Bay locality, Longarm Formation equivalent, Quatsino Sound, northern Vancouver Island, British Columbia, Canada (latitude $50^{\circ}36'21''N$, $127^{\circ}39'25''W$; UTM 9U WG 951068) [58].

Phylogenetic justification: *Tricarinella crassiphylla* is distinguished predominantly based on 3 characters: tristichous phyllotaxis, bistratose leaf lamina and a homogeneous costa. This set of characters, along with the clear implication that *T. crassiphylla* was an acrocarpous species, leads to its placement within Grimmiales [58]. Grimmiales + Dicranales form part of the Dicraniidae, which is in turn sister to Bryiidae, and so was used to establish a minimum age constraint on the divergence of *Ceratodon* (Dicranales) and Bryiidae.

Minimum age: 133.3 Ma

Maximum age: 515.5 Ma

Age justification: Sediments of the Apple Bay locality are regarded as Longarm Formation equivalent and originally proposed as Barremian based on the presence of monosulcate reticulate pollen grains [52]. Klymiuck and Stockey [53] argued that these grains are more likely associated with extinct gymnosperms, and instead propose a Valanginian age, supported by oxygen isotope analysis (D. Gröcke, personal communication, 2005, 2013). Thus, we propose the upper boundary of the Valanginian 133.9 ± 0.6 Ma and thus, 133.3 Ma as a minimum age constraint.

C_23: Bryophyta: Hypnanae - Brynanae

Fossil taxon and specimen: *Krassiloviella limbelloides* [Gametophyte shoot in rock slabs UAPC-ALTA P17596 C bot and D, University of Alberta Paleobotanical Collections (UAPC-ALTA), Edmonton, Alberta, Canada] from the northern shore of Apple Bay, Quatsino Sound, on the west side of Vancouver Island, British Columbia, Canada (lat. $50^{\circ}36'21''N$, long. $127^{\circ}39'25''W$; UTM 9U WG 951068) [59].

Phylogenetic justification: *K. limbelloides* was described as a pleurocarpous moss based on the presence of pinnate branching, homocostate and pluricostate leaves [59]. This was further supported by features including a monopodial main stem, second-order branching, the lack of conducting strands in the stems, and thin laminal cell walls. This combination of features indicates membership within the superorder Hypnanae, and the extinct family Tricostaceae.

Minimum age: 132 Ma

Maximum age: 515.5 Ma

Age justification: Sediments of the Apple Bay locality are regarded as Longarm Formation equivalent and originally proposed as Barremian based on the presence of monosulcate reticulate pollen grains [52]. Klymiuck and Stockey [53] argued that these grains are more likely associated with extinct gymnosperms, and instead propose a Valanginian age, supported by oxygen isotope analysis (D. Gröcke, personal communication, 2005, 2013). Thus, we propose the upper boundary of the Valanginian 132.6 ± 0.6 Ma and thus, 132 Ma as a minimum age constraint.

C_24: Tracheophyta: Lycopodiophyta – Euphyllophyta

Fossil taxon and specimen: *Zosterophyllum* sp. and *Chelinospora?* sp. from the Ordovician successions of central Sweden.

Phylogenetic justification: The minimum age is derived following Morris *et al.* [5].

Minimum age: 420.7 Ma

Maximum age: 458.88 Ma

Age justification: The minimum age is derived following following Morris *et al.* [5]. The soft maximum is based on a probably trilete spore (*Chelinospora?* sp.) from the Ordovician successions of central Sweden [60]. Spores were sampled from the Borensult-1 drillcore, comprising a well-dated succession of the Middle to Upper Ordovician. The successions are dated based on conodont biostratigraphy [61] and occur below the Kinnekulle K-bentonite, itself dated using U/Pb zircons to 454.41 ± 0.17 Ma [62, 63], confirming a Sandbian age. The spore was sampled from 48.2m, corresponding to the *Baltoniodus variabilis* biozone, with a maximum age corresponding to the top of the Sandbian, 458.18 ± 0.7 Ma [61, 64].

C_25: Lycopodiophyta: Isoetales – Lycopodiales

Fossil taxon and specimen: *Leclercqia complexa*

Phylogenetic justification: Following Morris *et al.* [5]

Minimum age: 392.1 Ma

Maximum age: 451 Ma

Age justification: Following Morris *et al.* [5]

C_26: Lycopodiophyta: Isoetales – Selaginellaes

Fossil taxon and specimen: *Yuguangia ordinata* [Specimen BUP.H-y07, Department of Geology of Peking University, PRC], from the Haikou Formation of Yuguang village, Zhanyi District, Yunnan Province, southwestern China [65].

Phylogenetic justification: *Yuguangia* is placed within Lycophytina based on its overall morphology, including external and internal anatomy and reproductive structures. A cladistic analysis by Hao *et al.* [65] based on 19 morphological characters placed *Yuguangia* as more closely related to the extant order Isoetales than Selaginellales, and thus it allows a minimum age constraint on the divergence of *Isoetes* and *Selaginella*.

Minimum age: 377.7 Ma

Maximum age: 451 Ma

Age justification: A Givetian age for the Haikou Formation was determined based on the presence of a miospore assemblage of *Cirratriradites monogrammos*, *Archaeozonotriletes variabilis*, *Cymbosporites magnificus* and *Geminospora lemurata* [66]. Thus, we propose a minimum age constraint coinciding with the upper boundary of the Givetian (388.9 ± 1.2 Ma)

C_27: Lycopodiophyta: *Selaginella selaginoides* – [rhizophoric *Selaginella*]

Fossil Taxon and Specimen: *Selaginella resimus* [BMNH V.62303-V.62330, Palaeontology Department, Natural History Museum, London] from the Puddlebrook assemblage (Lower Carboniferous), Drybrook Sandstone of the Forest of Dean, Gloucestershire [67].

Phylogenetic Justification: Rowe [67] described small lycophyte leafy shoots with consistently narrow stems and terminal strobili, placing it within *Selaginella*. The leaves are isophyllous and decussately arranged, supporting a placement within the subgenus *Boreoselaginella*. Thus, *S. resimus* provides a minimum age for the divergence of *S. selaginoides* and the rhizophoric clade, following Westrand and Korall [68, 69].

Minimum age: 322.8 Ma

Maximum age: 451 Ma

Age Justification: The plant-bearing outcrop of the Puddlebrook assembly is part of the Drybrook Sandstone at Plump Hill. The correlation between Puddlebrook and Plump Hill is supported by macrofossil remains of the lycophyte *Lepidostrobophyllum fimbriatum* and the pteridosperm *Diplopteridium*, both commonly occurring at Puddlebrook [70, 71]. Plump Hill is characterised by the miospores *Perotriletes tessellatus*, *Schultzospora* and *Carbaneuletes circularis*, indicative of an Asbian/Holkerian age. An Upper Visean age for the Puddlebrook assembly is further supported by the presence *Tetrapterites visensis*, which is known from the Upper Visean of Carboniferous limestone of the Menai region of Wales, and the presence of *Cribosporites cribellatus* [72], which supports an Asbian level. Following Poty *et al.* [73], the Asbian correlates with the Chesterian and the Warnantian, which both share an upper boundary with the Serpukhovian (323.2 Ma \pm 0.4) and thus a minimum age of 322.8 Ma is proposed.

C_28: Lycopodiophyta: *Selaginella* subg. *Stachygynandrum* – *Selaginella kraussiana*

Fossil taxon and specimen: *Selaginella* sp. [GZG.BST.21966-2002, Geoscience Centre, University of Göttingen] from the mid-Cretaceous Kachin Burmese amber.

Phylogenetic justification: Based on the presence of bilateral strobili, Schmidt *et al.* [74] were able to confidently assign six fossil species to the subgenus *Stachygynandrum*.

Minimum age: 98.17 Ma

Maximum age: 451 Ma

Age justification: The age of Burmese amber has long been contested, with estimates ranging between the Miocene and Cretaceous. Claims of an upper Albian age based on ammonite biostratigraphy [44] have not been verified [45]. However, a U-Pb age of 98.79 Ma \pm 0.62 Myr was obtained from zircons in volcanoclastics deposited contemporaneously with the amber [46]. Shi *et al.* (2012) and Yu *et al.* (2019) interpret this as a minimum age for the amber since it is bioeroded. Nevertheless, in the absence of other evidence, this U-Pb zircon age can be used to establish a minimum constraint of 98.17 Ma [44].

C_29: Lycopodiophyta: Lycopodioideae – Lycopodielloideae

Fossil Taxon and Specimen: *Retitriletes austroclavatidites* [GIX-SHA08-5-2; R32/1, Forschungsstelle für Paläobotanik at the Institut für Geologie und Paläontologie, Westfälische Wilhelms-Universität Münster, Germany], from the Section Peak Formation

Phylogenetic Justification: *Retitriletes* is a spore genus which bears strong resemblance with the spores of extant *Lycopodium* on the basis of their isotetrahedral structure and distinctive reticulate ornamentation on the distal surface [75-77]. They are the first spores which can reliably be assigned to the Lycopodioideae, and thus we follow Bauret *et al.* [76] in providing a minimum age for the divergence of the Lycopodioideae and Lycopodielloideae.

Minimum age: 199.2 Ma

Maximum age: 451 Ma

Age justification: The earliest record of *Retitriletes* is recorded by Bomfleur *et al.* [78] in the upper part of the Section Peak at Priestly Glacier cliff. The assemblage is assigned to Price's APJ1 assemblage based on the presence of intrastriate *Classopollis ssp.* and the absence of *Ichyosporites crateris-punctatus* [79], advocating an early-middle Sinemurian age. Radiometric ages of detrital zircons from c. 10m below the sample indicated ages of 214 ± 3 Ma and from c. 15m above the sample of 198 ± 2 Ma [80] which is congruent with a Sinemurian assignment. Thus, a minimum age constraint of 199.5 ± 0.3 Ma was proposed [81].

C_30: Euphyllophyta

Fossil taxon and specimen: *Kenrickia bivena* [Axis in slab 557820-5, Smithsonian Institution, NMNH] from the Battery Point Formation (Early Devonian), Quebec, Canada

Phylogenetic justification: A cladistic analysis by Toledo *et al.* [82] of 50 morphological characters supported *Kenrickia* as sister to all other radiatopsids, a lineage within euphyllophytes and outgroup to remaining progymnosperms and seed plants.

Minimum age: 393.2 Ma

Maximum age: 451 Ma

Age justification: Material was collected from the south shore of Gaspé Bay, near Douglstown, QC, Canada. The fossils form part of an allochthonous assemblage preserved in fluvial to coastal deposits [83, 84]. Palynological assemblages indicate the deposits near Douglstown belong to the late Emsian [85] and so our minimum age is based on the boundary of the Emsian, 394.3 ± 1.1 Ma [86].

C_31: Monilophyta: *Equisetum* - Pteridophyta

Fossil taxon and specimen: *Ibyka amphikoma*

Phylogenetic justification: Following Morris *et al.* [5]

Minimum age: 384.7 Ma

Maximum age: 451 Ma

Age justification: Following Morris *et al.* [5]

C_32: Monilophyta: *Equisetum subg. Equisetum* – *Equisetum subg. Hippochaete*

Fossil taxon and specimen: *Equisetum fluviatoides* [US1-103] from 34 metres above the base of the Ravenscrag Formation at Ravenscrag Butte, 17 Km southwest of Eastend, Saskatchewan, Canada [87].

Phylogenetic justification: *Equisetum fluviatoides* is identified as belonging to the crown group *Equisetum* based on vegetative and reproductive similarities to the extant species *E. fluviatile* [87]. A phylogenetic analysis by Elgorriaga *et al.* [88] placed *E. fluviatoides* as sister to *E. fluviatile* or *E. diffusum* within the subgenus *Equisetum*. The present study analysed the same phenotype and molecular data in a Bayesian framework and was not able to resolve the position of *E. fluviatoides* within the subgenus *Equisetum* but in a more basal position as sister to clade containing *E. fluviatile*, *E. diffusum* and *E. arvense*. Following our results, *E. fluviatoides* was used to calibrate the divergence of subgenus *Equisetum* from subgenus *Hippochaete*, here representing the crown group of *Equisetum*.

Minimum age: 64.86 Ma

Maximum age: 451.0 Ma

Age justification: The minimum age constraint is based on the holotype of *Equisetum fluviatoides*, described from 34 metres above the base of the Ravenscrag Formation at Ravenscrag Butte, Saskatchewan, Canada by McIver and Basinger [87] who estimate an

early Palaeocene age on the basis that the Ravenscrag Formation conformably overlies the Ferris/No.1 Coal Zone which at this site is established to approximate the Cretaceous-Palaeocene Boundary based on palynostratigraphy, about 3 metres below the top of magnetostratigraphic zone 29r [89]. Thus, the horizon from which the holotype of *Equisetum fluviatoides* was derived falls within the 29n magnetozone, the minimum age of which can be constrained by the 29n-28r boundary, which is dated to 64.86 Ma [22].

C_33: Monilophyta: Marratiopsida – Psilotopsida

Fossil taxon and specimen: *Psaronius simplicaulis*

Phylogenetic justification: Following Morris *et al.* [5]

Minimum age: 318.71 Ma

Maximum age: 451.0 Ma

Age justification: Following Morris *et al.* [5]

C_34: Monilophyta: *Ophioglossum* – *Sceptridium*

Fossil taxon and specimen: *Botrychium wightonii* [S1333 in the Palaeobotanical Herbarium, Department of Botany, University of Alberta] from the Palaeocene Genesee locality, North Saskatchewan River, Canada.

Phylogenetic justification: Based on the morphology of sporophores and trophophores, Rothwell & Stockey [90] concluded that *B. wightonii* could be assigned to the extant genus *Botrychium* with high confidence.

Minimum age: 65.5

Maximum age: 451.0 Ma

Age justification: Material was collected from the Genesee locality, along the bank of the North Saskatchewan River 45 miles southwest of Edmonton, Canada. The megafloreal assemblage at the Genesee locality indicates a Tiffanian age [91]. More recently, Stockey *et al.* [92] indicated that the Genesee flora belonged to the upper Scollard Formation, pollen zone P1, which correlates with the *Wodehouseia fimbriata* biozone [93], indicating an earliest Palaeocene age. U-Pb dating of the Scollard Formation and the overlying Battle Formation indicates a lower bound of 65.5 Ma for the *Wodehouseia fimbriata* biozone [94].

C_35: Monilophyta: Gleicheniales – [Cyatheales – Polypodiales]

Fossil taxon and specimen: *Oligocarpia gothanii* [S138027–S138989, S147989 and KC 1983–1, Department of Palaeobotany of the Swedish Natural History Museum, Stockholm] from the Mid Permian of the Shihhotse valley near Ch'en-chia-yu near Taiyuan City in Shanxi Province, China [95].

Phylogenetic justification: *Oligocarpia* has been associated with gleicheniaceae ferns on the basis of both spore and leaf morphology [95, 96].

Minimum age: 268.3 Ma

Maximum age: 451.0 Ma

Age justification: *Oligocarpia gothanii* was collected from the beds of the south side of the Shihhotse valley near Ch'en-chia-yu near Taiyuan City in Shanxi Province, China [95]. These beds belong to the Lower Shihhotse Formation, which has been correlated with the Kungarian-Roadian stages of the Mid Permian [97, 98]. Thus, our minimum age is based on the upper boundary of the Roadian stage, estimated at 268.8 ± 0.5 Ma.

C_36: Monilophyta: Cyatheaceae – Thyrsopteridaceae

Fossil Taxon and Specimen: *Oguracaulis banksii* [Z2286, Tasmania Museum and Art Gallery, Hobart, Tasmania] from the Lune River Fossil Site (Early to Middle Jurassic), Ida Bay, Tasmania [99].

Phylogenetic Justification: Phylogenetic analysis based on characters derived from fossilized stems placed *Oguracaulis banksii* as sister to *Thyrsopteris* [100].

Minimum age: 178 Ma

Maximum age: 451 Ma

Age justification: The flora of the Lune River area, Tasmania was initially estimated as Tertiary based on the similarity of the assemblage to others across Tasmania [101]. However, a revision of the taxonomic affinities of the flora suggested a Late Jurassic to Early Cretaceous age [102]. A basalt outcrop upslope of the fossil localities was shown to be comagmatic with Jurassic dolerite [103] and volcanolithic sandstone containing the fossils was recently dated at 182 ± 4 Ma (U-Pb zircon) [104], indicating a Pliensbachian-Toarcian age. Thus, we propose a minimum age constraint of 178 Ma.

C_37: Monilophyta: Lindsaceae – Lonchitidaceae

Fossil Taxon and Specimen: Lindsaceae gen. sp. indet. [GZG.BST.21952 in the Geoscientific Collections of the Georg August University Göttingen, Germany] from the Burmese Amber, Myanmar.

Phylogenetic Justification: The fusion of multiple veins below elongate terminal sori allows unambiguous assignment to Lindsaceae [105].

Minimum age: 98.17

Maximum age: 451

Age Justification: The age of Burmese amber has long been contested, with estimates ranging between the Miocene and Cretaceous. Claims of an upper Albian age based on ammonite biostratigraphy [44] have not been verified [45]. However, a U-Pb age of $98.79 \text{ Ma} \pm 0.62 \text{ Myr}$ was obtained from zircons in volcanoclastics deposited contemporaneously with the amber [46]. Shi et al. (2012) and Yu et al. (2019) interpret this as a minimum age for the amber since it is bioeroded. Nevertheless, in the absence of other evidence, this U-Pb zircon age can be used to establish a minimum constraint of 98.17 Ma [44].

C_38: Monilophyta: Pteridiinae + Polypodiinae

Fossil Taxon and Specimen: *Pteris* sp. [Palaeontological Institute, Moscow, Catalogue no. 4799-10] from the mid-Cenomanian of the Nammoura locality, Mount-Liban, Lebanon [106].

Phylogenetic Justification: The specimen was assigned to the extant genus *Pteris* based on strong similarities in pinnule shape and coenosori position [106]. Schneider *et al.* supported its position as a pteridoid fern, but thought that no further assignment was possible and it was used as a minimum age for the divergence of Pteridaceae and Polypodiinae [107].

Minimum age: 100.3 Ma

Maximum age: 451 Ma

Age justification: The Nammoura lagerstätte is estimated as latest middle Cenomanian in age based on the foraminiferan biostratigraphy and the characteristic species *Rhapydionina laurinensis* [106]. The upper units of the bed are marked by the presence of orbitolinids, indicating a middle Cenomanian age, which are absent in the lower units. As the unit from which the specimen was collected is not apparent, we took the upper boundary of the Cenomanian as the minimum age constraint, estimated at 100.5 ± 0.1 , thus 100.3 Ma

C_39: Monilophyta: Eupolypods I + Eupolypods II

Fossil Taxon and Specimen: *Woodwardia* sp. from the Jose Creek Member, McRae Formation, New Mexico, USA [108].

Phylogenetic Justification: Leaf impressions demonstrating diagnostic traits including a toothed margin, sickle-shaped pinnules, veins that anastomose at the midvein to form costal areoles, veins outside the costal areoles that run to the margin, and the absence of freely ending veinlets indicate a membership to the extant genus *Woodwardia*, a member of the Eupolypods II clade (Aspleniineae) [108].

Minimum age: 72 Ma

Maximum age: 451

Age justification: Leaf impression were collected from the lower Jose Creek Member of the McRae Formation. U-Pb age estimates from multiple horizons within the McRae formation have confirmed a Late Campanian age, rather than the previously suggested Maastrichtian [109]. We use the upper boundary of the Campanian, 72.2 ± 0.2 Ma, thus 72 Ma, as a minimum age constraint.

Discussion: *Holttumpoteris burmensis*, a putative crown member of Eupolypods II was recently describe from Early Cretaceous amber deposits in Myanmar [110]. The authors decided that it was likely a member of the derived family Thelypteridaceae, but based on the lack of preservation of key characters could not confidently assign it to an extant family.

C_40: Spermatophyta: Acrogymnospermae – Angiospermae

Fossil taxon and specimen: *Cordaixylon iowensis*

Phylogenetic justification: Following Morris *et al.* [5]

Minimum age: 308.14 Ma

Maximum age: 365.63 Ma

Age justification: Following Morris *et al.* [5]

C_41: Acrogymnospermae: [Ginkgoopsida + Cycadopsida] – Pinopsida

Fossil taxon and specimen: *Cordaixylon iowensis*

Phylogenetic justification: Following Morris *et al.* [5]

Minimum age: 308.14 Ma

Maximum age: 365.63 Ma

Age justification: Following Morris *et al.* [5]

C_42: Acrogymnospermae: Ginkgoopsida – Cycadopsida

Fossil taxon and specimen: *Iratinia australis* [MCT1487-PB, Museu de Ciências da Terra, Rio de Janeiro, Brazil] from the Irati Formation of the Kungurian of the Paraná Basin, Brazil.

Phylogenetic justification: Though initially proposed as a lycopsid [111], based on multiple morpho-anatomical traits Spiekermann *et al.* [112] determined that *I. australis* belonged to the extant order Cycadales.

Minimum age: 274 Ma

Maximum age: 365.63 Ma

Age justification: Material was collected from the Irati Formation, the lowermost unit of the Passa Dois Group [ref]. The Assistência Member from the upper part of the Irati Formation was U-Pb zircon dated to 279.9 ± 4.8 to 280 ± 3 Ma [113], placing it within the Kungurian.

The Taquaral Member at the base of the formation was also placed in the Kungurian based on palynological evidence [114]. Thus, a Kungurian age is determined for the Irati Formation and a minimum age of 274.4 ± 0.5 Ma.

C_43: Pinopsida: Taxaceae - Cupressaceae

Fossil taxon and specimen: *Palaeotaxus rediviva* from the Rhaetian of southern Sweden.

Phylogenetic justification: *Palaeotaxus* forms the first credible record of the family Taxaceae and was placed within Taxaceae in a phylogenetic analysis of morphological characters [115, 116].

Minimum age: 201.2 Ma

Maximum age: 321.48 Ma

Age justification: Material was collected from the upper coal bed of the Skromberga Colliery in Scania, Sweden. The Skromberga locality has been assigned to the Bjuv member [117] and Pott & McLoughlin and McElwain et al indicate a Rhaetian age based on the correlation of the *Thaumatopteris-Lepidopteris* flora [118-120]. Thus, a minimum age is established based on the upper boundary of the Rhaetian 201.4 ± 0.2 Ma.

C_44: Pinopsida: *Cunninghamia* – *Juniperus*

Fossil taxon and specimen: *Chamaecypris corpulenta* from the Upper Cretaceous Comox Formation of Vancouver Island, British Columbia.

Phylogenetic justification: Based on morphological similarities of the seed cone structure and the cones on the axes, *C. corpulenta* is considered the oldest known representative of the Cupresseae [121, 122].

Minimum age: 83.2

Maximum age: 321.48 Ma

Age justification: Material was recovered from the carbonaceous shales from the No.2 coal seam within the Cumberland Member of the Comox Formation, Vancouver Island [123]. Marine fauna above and below the Comox formation indicates a Santonian age and thus a minimum of 83.7 ± 0.5 Ma.

C_45: Pinopsida: Pinales – Gnetales

Fossil taxon and specimen: *Eathiestrobus mackenziei*

[Specimen Pb483, Hunterian Museum, Glasgow, Scotland] from Eathie, Black Isle, Scotland.

Phylogenetic justification: Based on a combination of diagnostic seed, cone and scale morphologies, *Eathiestrobus* was considered a member of Pinaceae [124]. However, the wide bract is a feature that distinguishes *Eathiestrobus* from all extant genera.

Minimum age: 153.62 Ma

Maximum age: 321.48 Ma

Age justification: Coastal Jurassic exposures at Eathie are from the Kimmeridge Clay Formation [125]. Ammonite stratigraphy indicates belonging to the *Cymodoce* chronozone [126], yet dinoflagellate stratigraphy supports a placement in the *Baylei* zone (Lower Kimmeridgian). The top of the *Cymodoce* zone correlates with the magnetostratigraphic zone M24, the top of which is dated to 153.62 Ma [127].

C_46: Pinopsida: *Pinus* – *Cedrus*

Fossil taxon and specimen: *Pinus yorkshirensis* [BU4737 at the Lapworth Museum, University of Birmingham, United Kingdom] from the early Cretaceous Wealden of Yorkshire, United Kingdom.

Phylogenetic justification: A cladistic analysis of morphological characters resolved *P. yorkshirensis* as a crown-group member of *Pinus* [128].

Minimum age: 129 Ma

Maximum age: 321.48 Ma

Age justification: Material comes from the Speeton Clay Formation, north of Speeton, Yorkshire. Palynological samples collected from the same material include the dinoflagellate species *Callaiosphaeridium asymmetricum*, *Odontochitina operculata* and ?*Cribroperidinium confossum* [128] which constrains the sediment to Late Hauterivian to Early Barremian [129]. Based on the range bases of these taxa, Ryberg *et al.* conclude that the sediment was formed during the Hauterivian-Barremian transition (131-129 Ma) and thus, we propose a minimum age of 129 Ma.

C_47: Pinopsida: *Pinus* subg. *Pinus* – *Pinus* subg. *Strobus*

Fossil taxon and specimen: *Pinus triphylla* [E.C Jeffrey Collection, Harvard University, USA] from the Late Cretaceous Raritan Formation, Staten Island, USA.

Phylogenetic justification: *P. triphylla* has been considered a stem lineage member of the subgenus *Pinus* based on the synapomorphy of possessing two vascular bundles in the needles [130, 131].

Minimum age: 89.37 Ma

Maximum age: 321.48 Ma

Age justification: Material was collected from the Androvette Clay Pit, Staten Island, part of the Raritan Formation. Based on palynological evidence, a Turonian age is proposed for the formation [55, 132, 133]. Thus, we propose a minimum age based on the lower boundary of the Turonian of 89.75 ± 0.38 .

C_48: Pinopsida: *Pinus* subsection *Australes*

Fossil taxon and specimen: *Pinus storeyana* [University of California, Museum of Paleontology, holotype 7234] from the Coal Valley Formation, Celetom Quarry, Nevada, USA

Phylogenetic justification: Based on seed cone morphology, Saladin *et al.* placed *P. storeyana* within the *Attenuatae* subclade [130, 134].

Minimum age: 9 Ma

Maximum age: 321.48 Ma

Age justification: K-Ar dating of the Coal Valley Formation, Wassuk Group, Nevada suggests it was deposited between 13 and 9 Ma [135]. Thus, a minimum age of 9 Ma is proposed

C_49: Angiospermae

Fossil taxon and specimen: Tricolpate pollen grain

Phylogenetic justification: Following Morris *et al.* [5]

Minimum age: 125 Ma

Maximum age: 247Ma

Age justification: Following Morris *et al.* [5]

C_50: Angiospermae: Nymphaeales

Fossil taxon and specimen: Tricolpate pollen grain
Phylogenetic justification: Following Morris *et al.* [5]
Minimum age: 125 Ma
Maximum age: 247Ma
Age justification: Following Morris *et al.* [5]

C_51: Angiospermae: Austrobaileyales

Fossil taxon and specimen: Tricolpate pollen grain
Phylogenetic justification: Following Morris *et al.* [5]
Minimum age: 125 Ma
Maximum age: 247Ma
Age justification: Following Morris *et al.* [5]

C_52: Mesangiospermae

Fossil taxon and specimen: Tricolpate pollen grain
Phylogenetic justification: Following Morris *et al.* [5]
Minimum age: 125 Ma
Maximum age: 247Ma
Age justification: Following Morris *et al.* [5]

C_53: Magnoliids: [Magnoliales + Laurales] + Piperales

Fossil taxon and specimen: *Endressinia brasiliana*
Phylogenetic justification: Following Morris *et al.* [5]
Minimum age: 110.87 Ma
Maximum age: 247Ma
Age justification: Following Morris *et al.* [5]

C_54: Piperales: *Houttuynia* – *Saruma*

Fossil taxon and specimen: *Saururus tuckerae*
Phylogenetic justification: Following Morris *et al.* [5]
Minimum age: 44.3 Ma
Maximum age: 247Ma
Age justification: Following Morris *et al.* [5]

C_55: Eudicotyledonae

Fossil taxon and specimen: *Hyrkantha decussata*
Phylogenetic justification: Following Morris *et al.* [5]
Minimum age: 119.6 Ma
Maximum age: 128.63 Ma
Age justification: Following Morris *et al.* [5]

C_56: Eudicotyledonae: Vitales

Fossil taxon and specimen: *Paleoclusia chevalieri*
Phylogenetic justification: Following Barba-Montoya *et al.* [136]

Minimum age: 85.84 Ma
Maximum age: 128.63 Ma
Age justification: Following Barba-Montoya *et al.* [136]

C_57: Eudicotyledonae: Rosids minus Vitales

Fossil taxon and specimen: *Paleoclusia chevalieri*
Phylogenetic justification: Following Barba-Montoya *et al.* [136]
Minimum age: 85.8 Ma
Maximum age: 128.63 Ma
Age justification: Following Barba-Montoya *et al.* [136]

C_58: Eudicotyledonae: Ericales

Fossil taxon and specimen: *Paleoenkianthus sayrevillensis*
Phylogenetic justification: Following Barba-Montoya *et al.* [136]
Minimum age: 85.8 Ma
Maximum age: 128.63 Ma
Age justification: Following Barba-Montoya *et al.* [136]

C_59: Eudicotyledonae: Myrtales

Fossil taxon and specimen: *Esqueiria futabensis*
Phylogenetic justification: Following Barba-Montoya *et al.* [136]
Minimum age: 83.3 Ma
Maximum age: 128.63 Ma
Age justification: Following Barba-Montoya *et al.* [136]

C_60: Eudicotyledonae: Asteraceae

Fossil taxon and specimen: *Tubulifloridites antipodica*
Phylogenetic justification: Following Barba-Montoya *et al.* [136]
Minimum age: 41.5 Ma
Maximum age: 128.63 Ma
Age justification: Following Barba-Montoya *et al.* [136]

C_61: Eudicotyledonae: Salicaceae

Fossil taxon and specimen: *Pseudosalix handleyi*
Phylogenetic justification: Following Barba-Montoya *et al.* [136]
Minimum age: 48.57 Ma
Maximum age: 128.63 Ma
Age justification: Following Barba-Montoya *et al.* [136]

C_62: Eudicotyledonae: Solanales

Fossil taxon and specimen: *Solanites crassus*
Phylogenetic justification: Following Barba-Montoya *et al.* [136]
Minimum age: 37.3 Ma
Maximum age: 128.63 Ma

Age justification: Following Barba-Montoya *et al.* [136]

C_63: Monocotyledonae

Fossil taxon and specimen: *Liliacidites sp.*

Phylogenetic justification: Following Barba-Montoya *et al.* [136]

Minimum age: 113 Ma

Maximum age: 128.63 Ma

Age justification: Following Barba-Montoya *et al.* [136]

C_64: Monocotyledonae: Dioscoreales + Core Monocots

Fossil taxon and specimen: *Cratolirion bognerianum* [MB.Pb. 1997/1233 (repository: Museum für Naturkunde, Berlin)] from the Crato Formation (Early Cretaceous), Santana do Cariri, Brazil [137].

Phylogenetic justification: Phylogenetic analyses indicated a membership within the Pandanales + Dioscoreales lineage (Petrosaviidae) [137].

Minimum age: 119.5 Ma

Maximum age: 128.63 Ma

Age justification: The age of the [5, 138]Crato Formation has been much debated but a 123 Ma ± 3.5 Myro been obtained for the overlying Ipubi Formation using Re-Os isotope dating [139], providing for a minimum constraint of 119.5 Ma [138].

C_65: Monocotyledonae: Riponogaceae (*Colcichum – Smilax*) ** NEW **

Fossil taxon and specimen: *Ripogonum tasmanicum* from the Macquarie Harbour Formation (Early Eocene) from the central west coast of Tasmania [140]

Phylogenetic justification: Based on a phylogenetic analysis of morphological characters, *Ripogonum tasmanicum* was placed within the extant genus *Ripogonum* (Ripogonaceae, Liliales). Phylogenetic analyses identify Riponogaceae and Smilaceae being sister lineages respective to Colchicaceae. **Minimum age:** 50.5 Ma

Maximum age: 128.63 Ma

Age justification: Fossil leaves were collected from mudstone deposits at the Lowana Road site, the Macquarie Harbour Formation, Tasmania. Spores, pollen and dinoflagellates from the site were attributed to the Planktonic Foraminiferal Zone P8-9 [141] corresponding to a Mid-Ypresian age. The minimum age of the deposit is determined based on correlation of the Upper *Malvacipollis diversus* spore-pollen zone with zones P7 to P8 of Partridge [142] (52.5-50.5 Ma). Thus, a minimum age of 50.5 Ma is provided.

C_66: Monocotyledonae: Arecales – Poales

Fossil taxon and specimen: *Sabalites carolinensis*

Phylogenetic justification: Following Barba-Montoya *et al.* [136]

Minimum age: 83.41 Ma

Maximum age: 128.63 Ma

Age justification: Following Barba-Montoya *et al.* [136]

C_67: Poales: BOP-PACMAD

Fossil taxon and specimen: *Changii indicum* [Slide 13160, Q-14-3, Birbal Sahni I. Palaeobotany, Lucknow, India] recovered from Coprolites from red clays in Lameta

Formation, associated with the Deccan Volcanics (infra- and inter-trappan) at Pisdura (East and South sections), Central India. [143].

Phylogenetic justification: Based on phylogenetic analyses, Prasad *et al.* [143] determined *C. indicum* as a member of Erhartoideae (Poaceae). More recent analyses have questioned this placement, and have suggested that it could belong to a number of tribes. As such, we follow Christin *et al.* [144] in a more conservative placement as a stem group member of the BOP clade.

Minimum age: 66.261 Ma

Maximum age: 128.63 Ma

Age justification: Prasad *et al.* [61] interpreted the minimum age of *Changii indicum* as 65.0 Ma, while Iles *et al.* [145] adopted the radiometric and magnetostratigraphic dating of the Deccan Traps by Courtillot and Ren [146], supported by age constraints of associated dinosaur coprolites to infer a minimum constraint of 66.0 Ma, based on the dating of the Maastrichtian-Palaeocene Boundary; this would revise to 65.98 Ma following Ogg & Hinnov [138]. However, the dating of the Deccan Traps has been revised [147] and, given the fossil occurrences predate the volcanic sequences [61] we can establish a minimum constraint based on the age of the oldest phase of volcanism, dated to 66.288 Ma \pm 0.027 Myr, thus 66.261 Ma.

C_68: Poales: *Brachypodium* – *Oryza* ** NEW **

Fossil taxon and specimen: *Stipa florissanti* [Syntypes: 34750, 34751 (USNM)] from Florissant and Florissant Fossil Beds National Monument, CO, USA [148]

Phylogenetic justification: Based on morphology this fossil was assigned the extant grass lineage *Stipa*. Extant *Stipa* has a strictly Old World distribution [149], so the generic assignment has been questioned [150]. We follow Iles *et al.* [145] and their assignment to the stem group Stipeae, itself nested within Pooideae [151].

Minimum age: 33.705 Ma

Maximum age: 128.63 Ma

Age justification: The age interpretation of the Florissant Formation is usually based on Evanoff *et al.* [152] who used $^{40}\text{Ar}/^{39}\text{Ar}$ dating of sanidine crystals in detrital pumice, yielding an age of 34.07 Ma \pm 0.1 Myr. However, its detrital origin renders this a maximum constraint on the age of the fossil. Nevertheless, Florissant Formation falls fully within the C13r magnetozone [153] and the base of the succeeding C13n magnetozone is dated to 33.705 [22]

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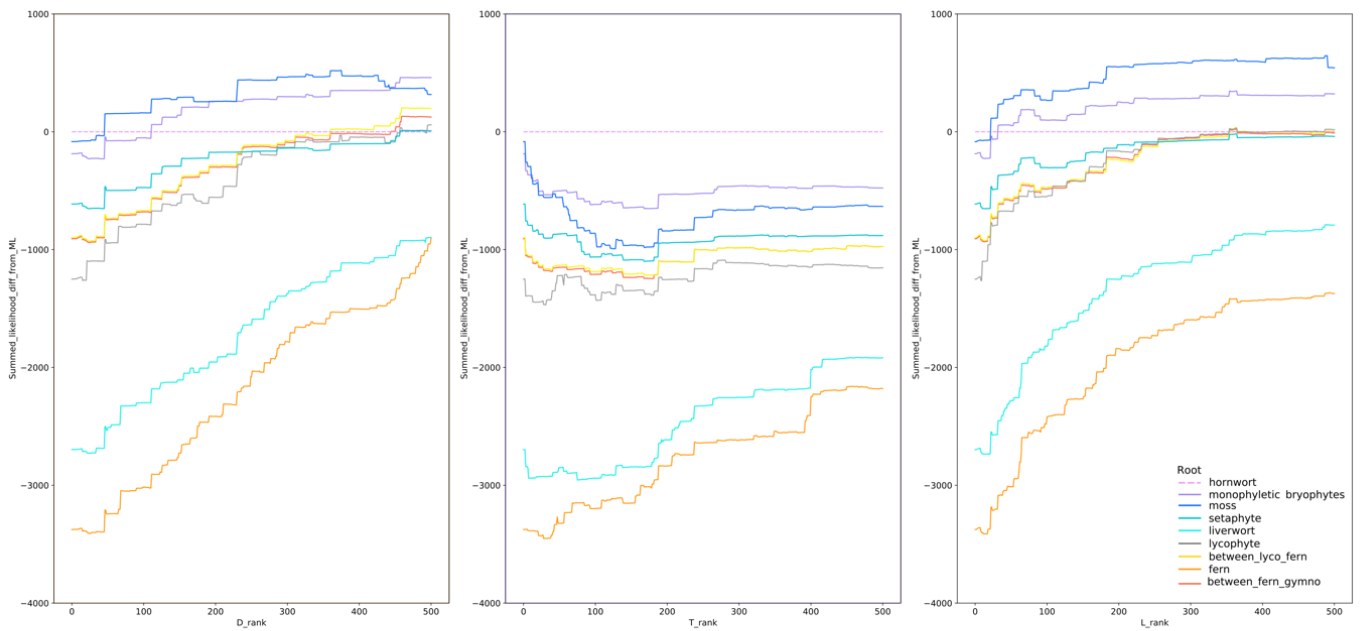
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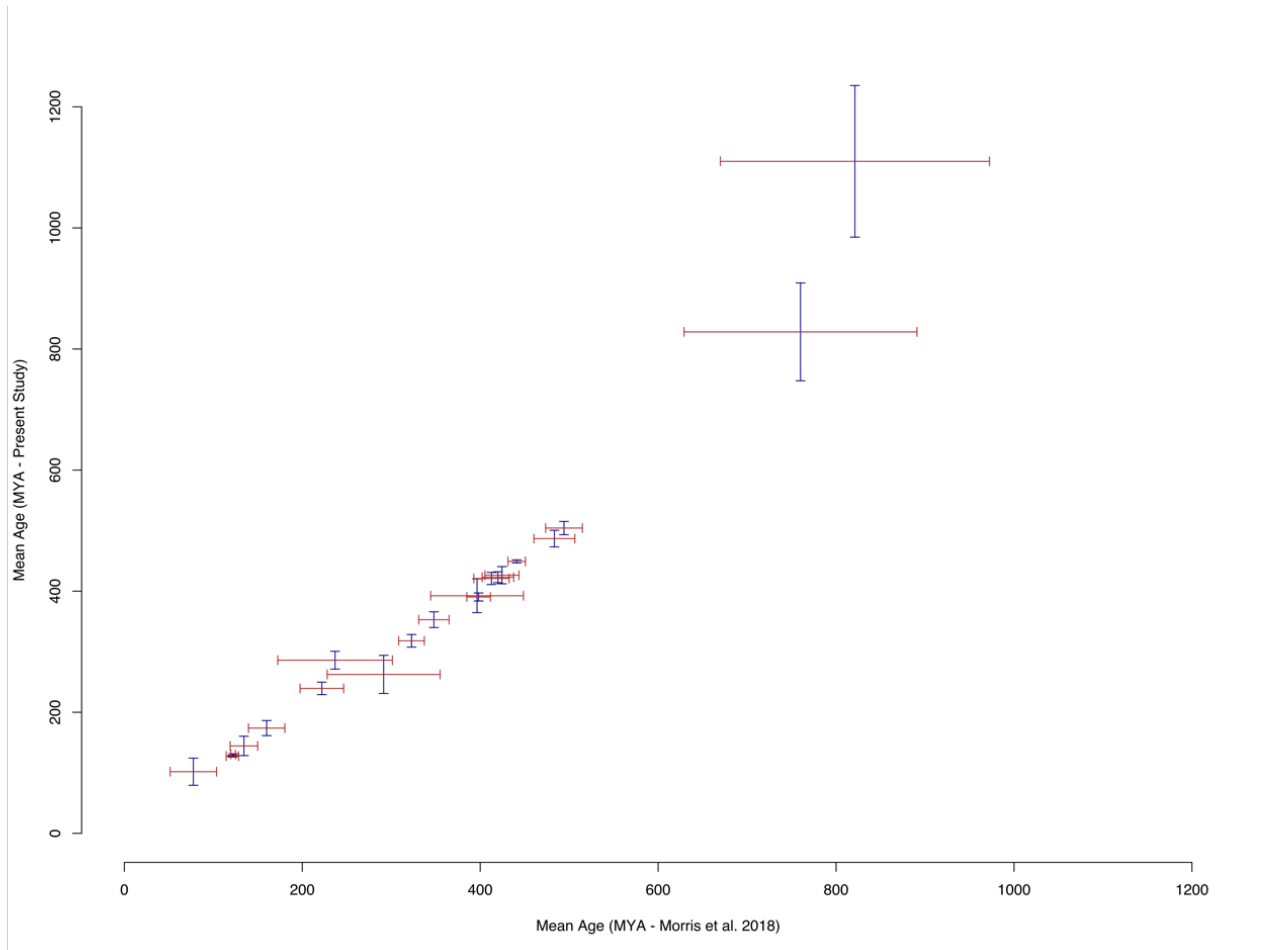
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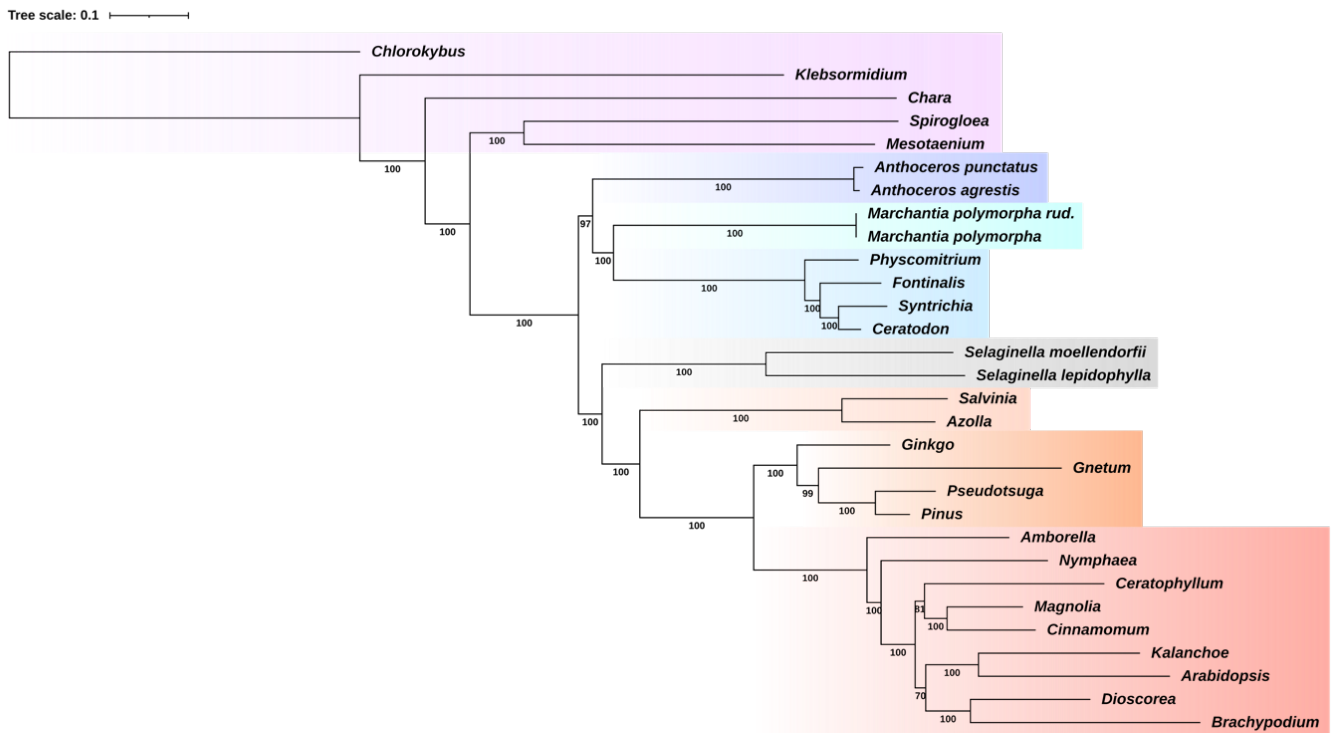
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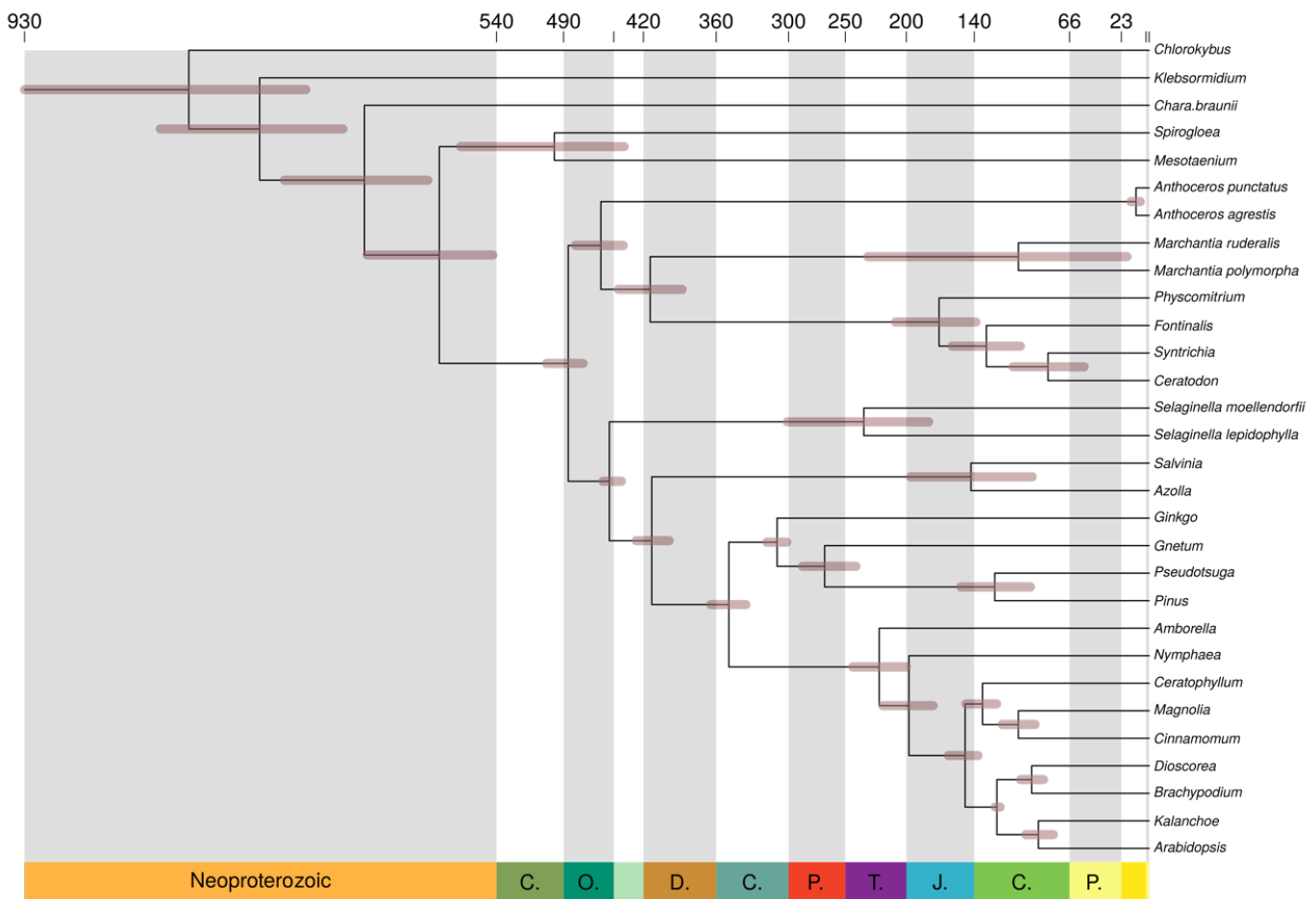
Supplementary Fig. 1 | Evaluating the root signal inferred by ALE. The impact of high duplication, transfer or loss rate families on inferred root order was investigated by iteratively removing top-ranked families and recalculating the summed likelihood on remaining families. While the roots of the credible set remained the top three by summed likelihood, their order varied as high D, T or L-rate families were removed. Removal of high-D and high-L families reduced support for the hornwort root (the ML root on all data), while removal of high-T families increased support for the hornwort root relative to the second- and third-ranked roots. As D, T and L rates were calculated based on the ML root, the preference of high-T families for alternative roots might reflect the difficulty of placing the sparsely-sampled hornworts accurately in individual gene trees.



Supplementary Fig. 2 | A comparison of divergence time estimates between the current study and the closest benchmark, that of Morris et al 2018. Points on each axis are centred on the mean age, with the 95% highest posterior density represented by the width of the bars. 95% HPDs were calculated from 2,000 post-burnin samples over 2,000,000 MCMC generations.

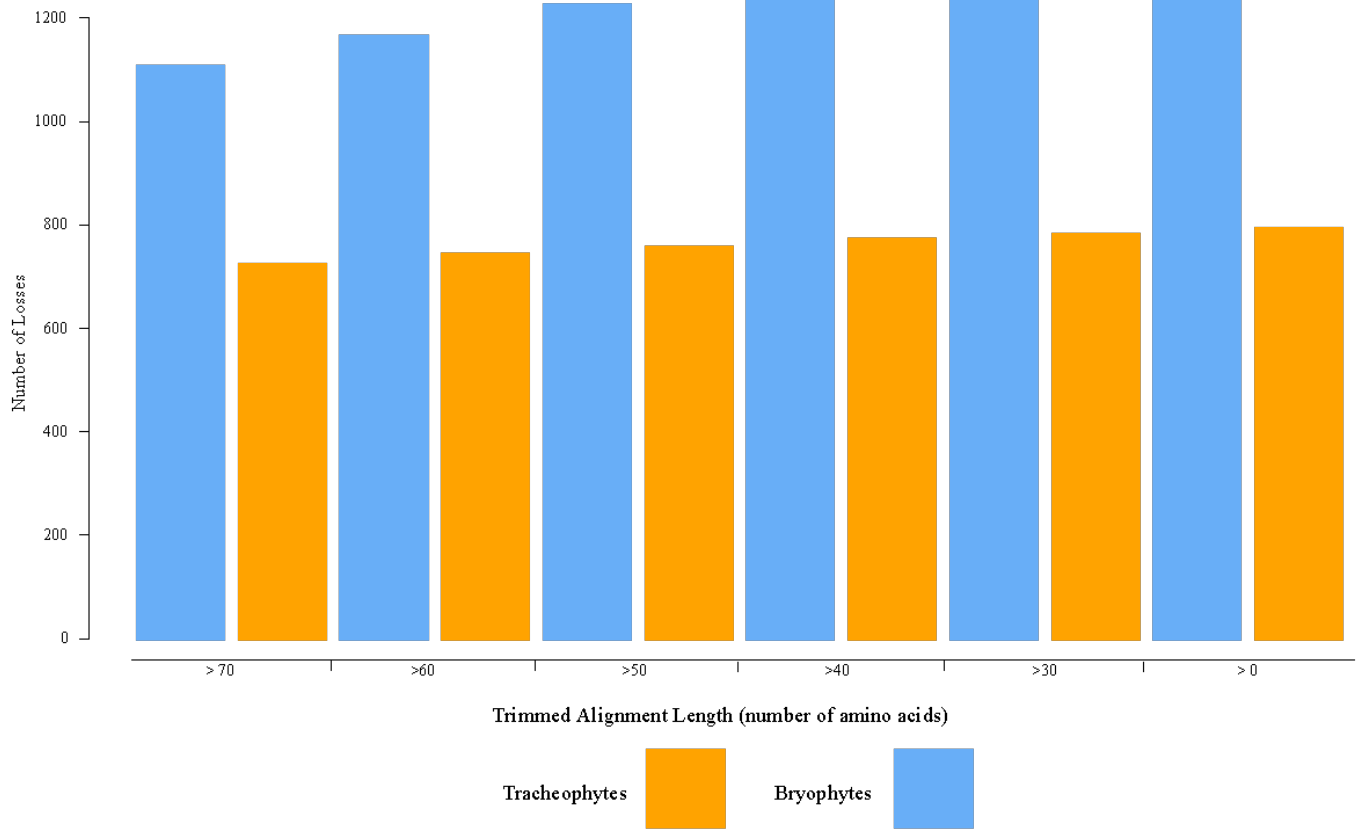


Supplementary Fig. 3 | Phylogenetic tree of embryophytes used for ancestral reconstruction. The maximum likelihood tree was inferred from an alignment of 30 species, 185 single copy orthologs and 71855 sites under the LG+C60+G4+F model in IQ-Tree (Nguyen *et al.*, 2015). Bootstrap support values are placed underneath the corresponding branches. The branch lengths are proportional to the average number of substitutions.

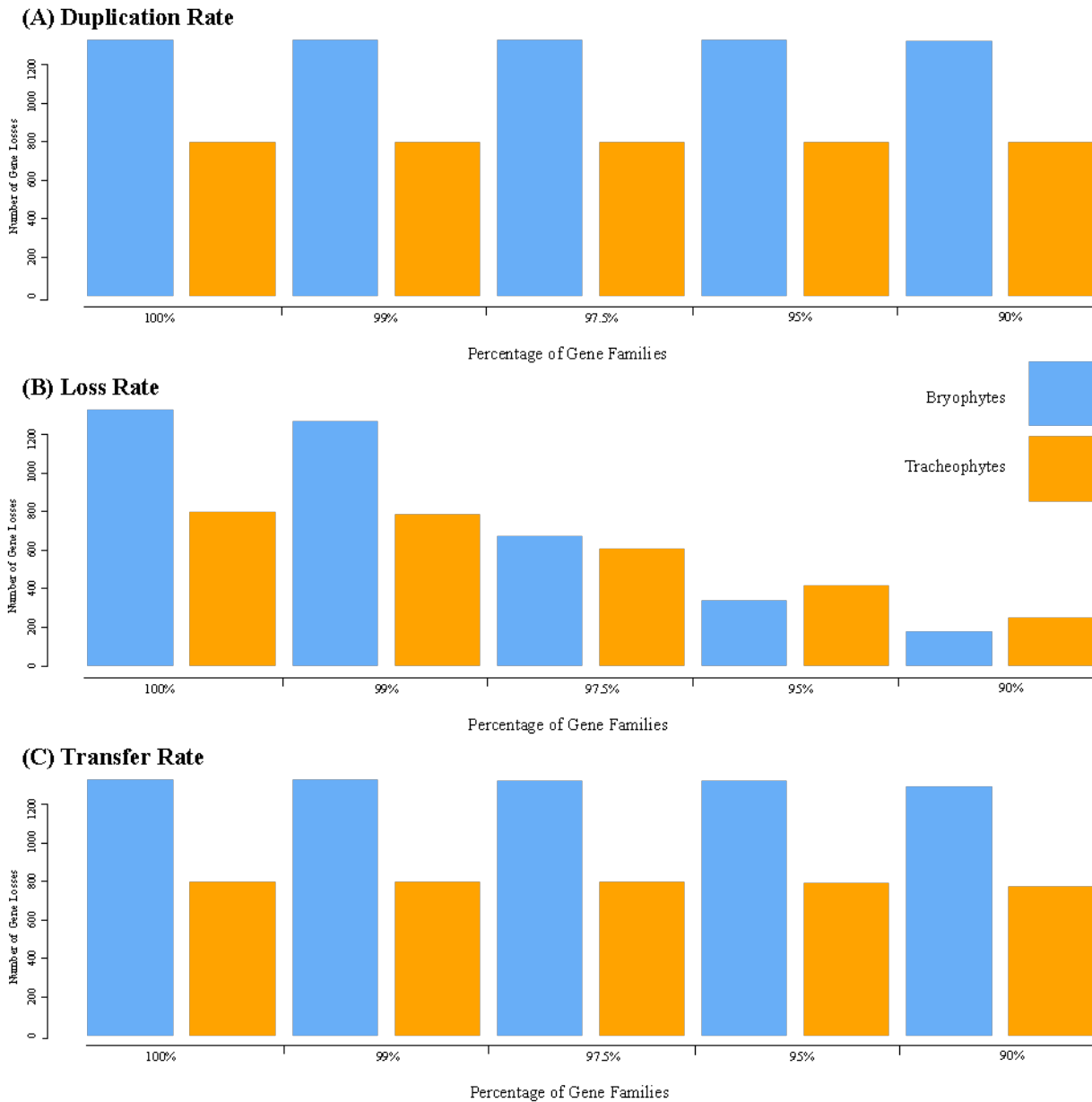


Supplementary Fig 4 | Time scale of embryophyte tree used for ancestral reconstruction.

The amino acid sequence alignment was combined with 18 fossil calibrations subsampled from the list of calibrations from the transcriptomic analysis. The normal approximation method was used in MCMCTree, with branch lengths estimated under the LG+G4 model. Calibrations were modelled as a uniform distribution between a minimum and soft maximum. 95% highest posterior density (HPD) width is shown as bars on nodes.

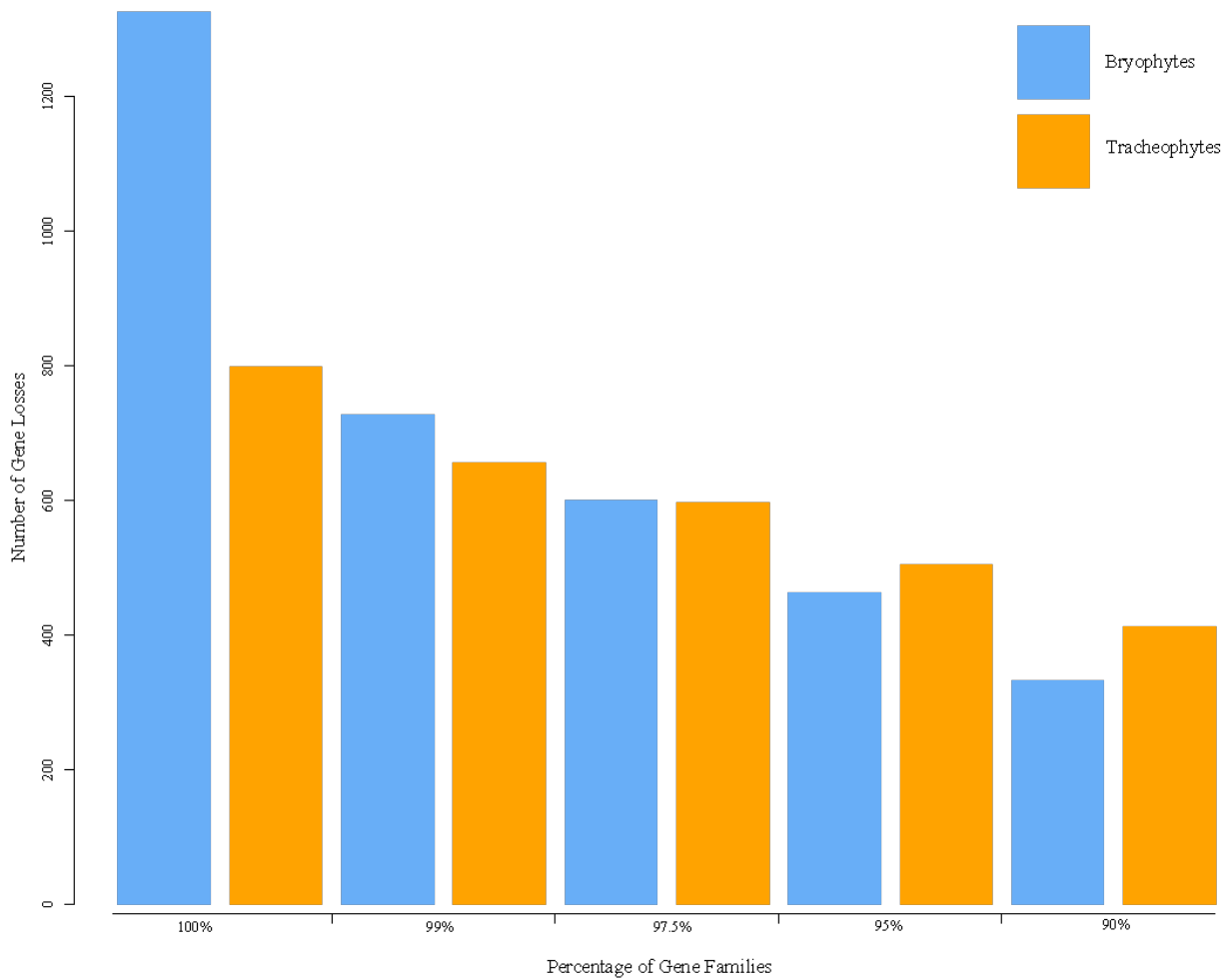


Supplementary Fig. 5 | Filtering gene families by alignment length. The inferred number of losses on the bryophyte (blue) and tracheophyte (orange) stem lineage after filtering the trimmed gene families by their alignment length - that is, discarding gene families with trimmed alignments shorter than a given threshold.

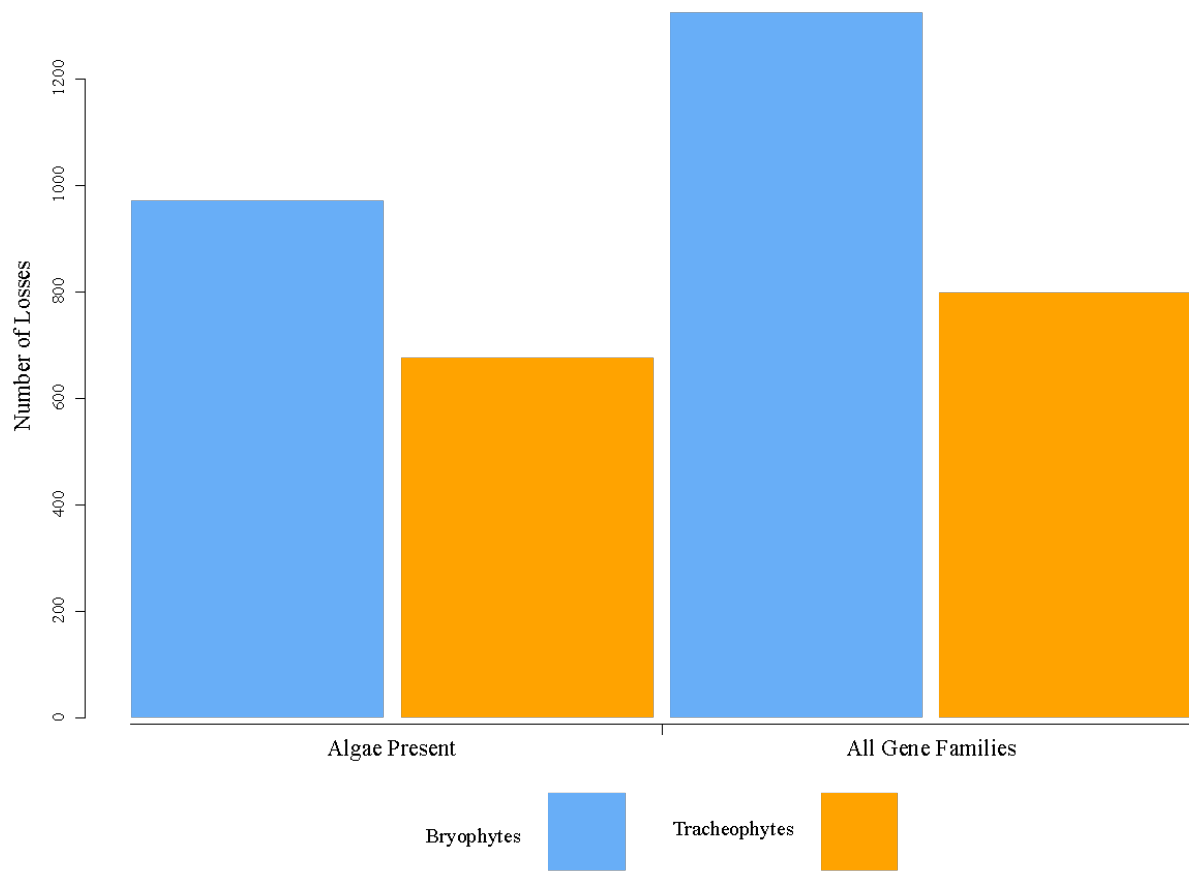


Supplementary Fig. 6 | Filtering gene families by rates of duplication, loss and transfer.

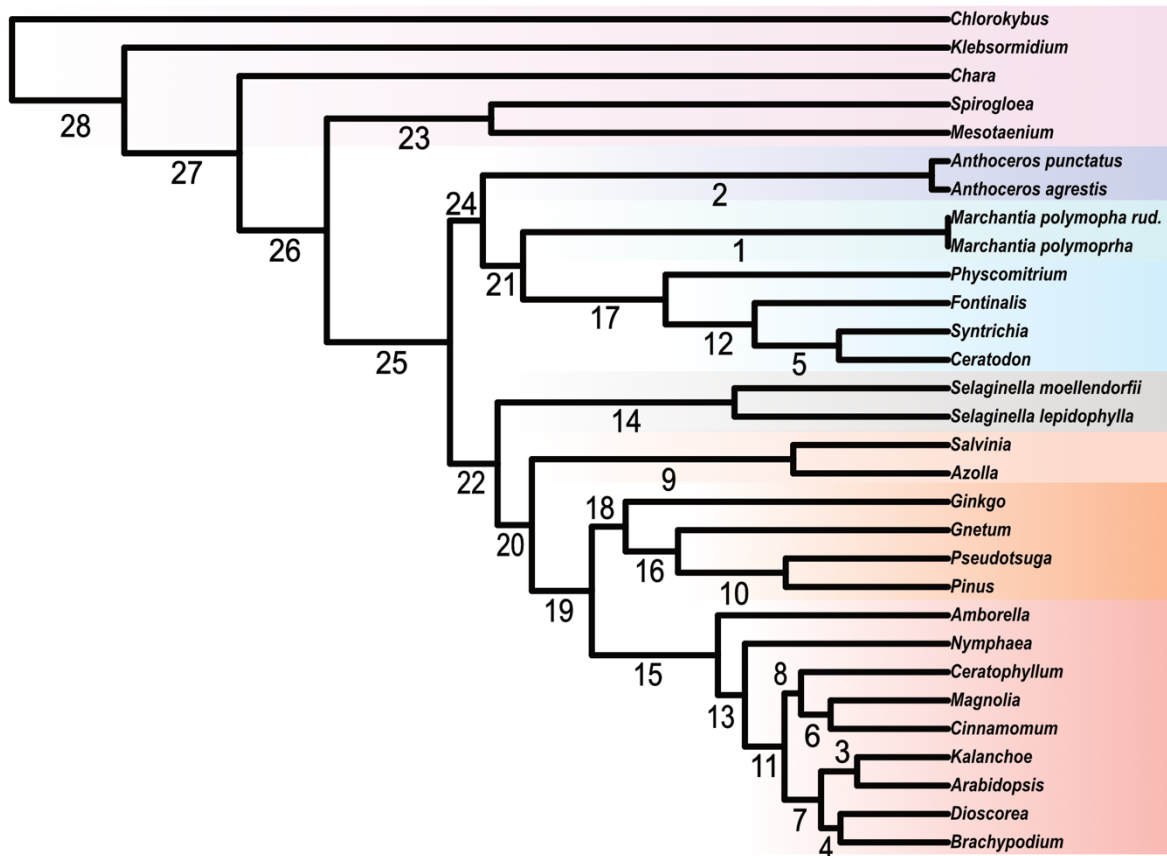
The inferred number of losses on the bryophyte (blue) and tracheophyte (orange) stem lineage after filtering the trimmed gene families by their (A) duplication rate (B) loss rate and (C) transfer rate. In each case, families were filtered to remove the top 1, 2.5, 5 and 10% of families. Filtering by duplication or transfer rate had no substantial impact on bryophyte or tracheophyte stem losses. As expected, filtering families by loss rate (B) reduced the number of inferred losses in the remaining families on both the bryophyte and tracheophyte stems. The reduction in losses was greater for the bryophyte stem, suggesting that the greatest number of losses are concentrated in a minority of gene families (see main text for discussion of some particularly notable losses).



Supplementary Fig. 7 | Filtering gene families by size (number of genes). The inferred number of losses on the bryophyte (blue) and tracheophyte (orange) stem lineage after filtering the trimmed gene families by their size (number of genes). Families were filtered to remove the top 1, 2.5, 5 and 10% of families.



Supplementary Fig. 8 | Filtering gene families by algal copies. The inferred number of losses on the bryophyte (blue) and tracheophyte (orange) stem lineage after filtering the trimmed gene families such that they must include at least 1 copy in 1 of the algal lineages (Algae Present).

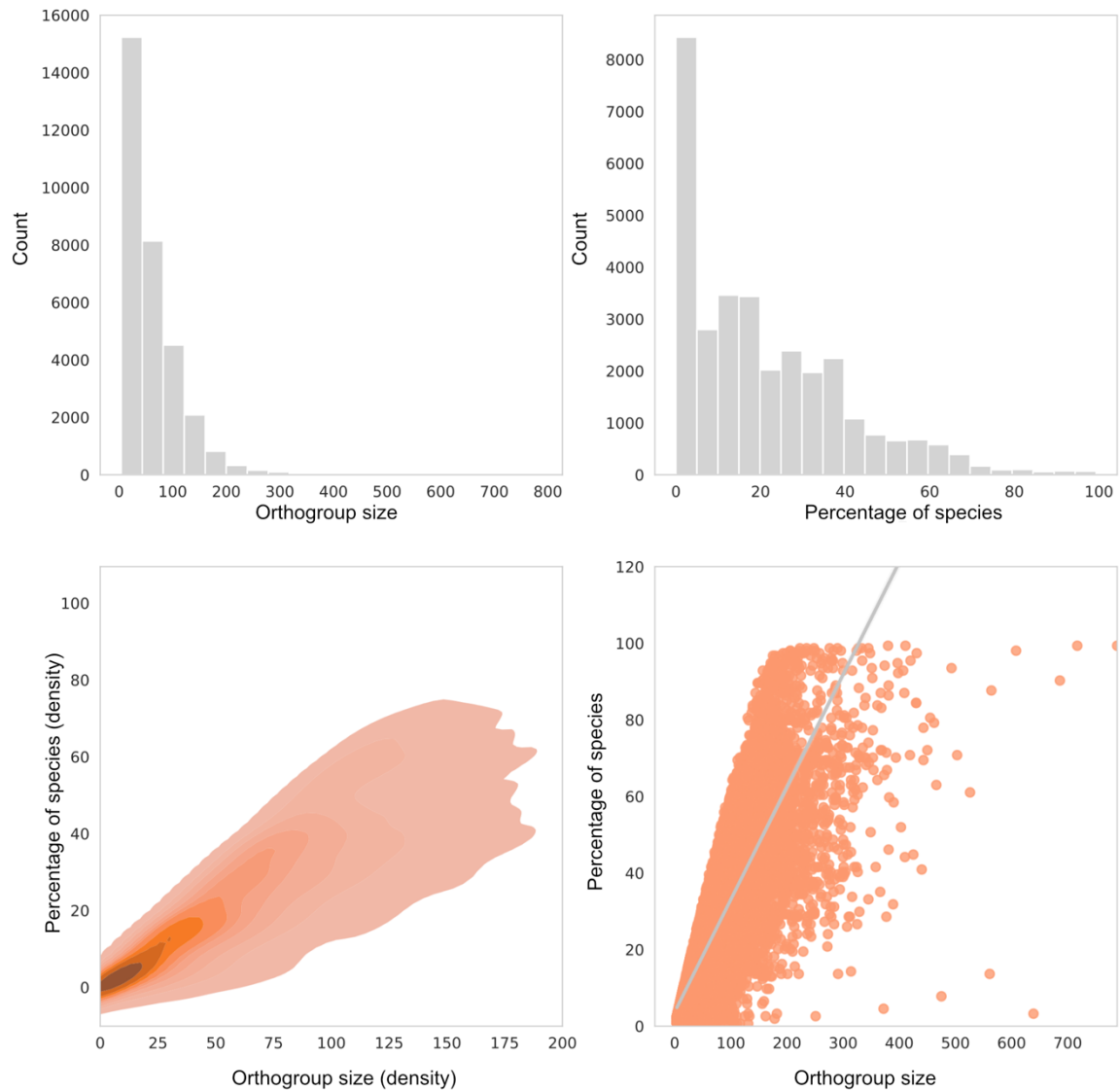


Supplementary Fig. 9 | Gain and loss of orthologous gene family clusters across the embryophyte tree (reference species tree). A maximum likelihood time-calibrated tree, with labels assigned to each branch. See Supplementary Fig. 10 for gains and losses of orthologous gene family clusters mapped to each branch of this reference tree.

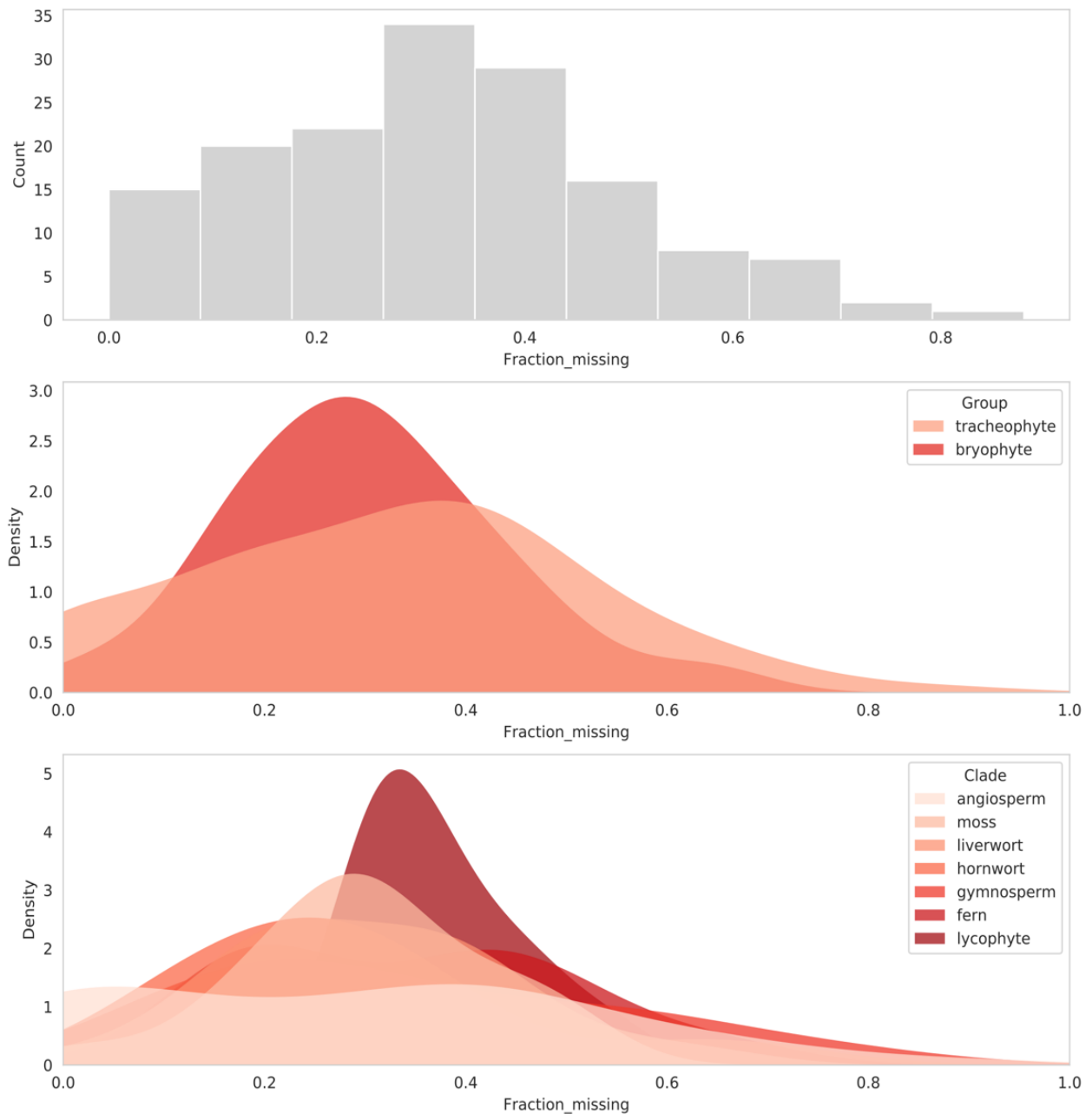
Orthologues Present Orthologues Lost Orthologues Gained

28	27	26	25	24	23	22	21	20	19	18	17	16	15	14	
ABI1	ABI1	ABI1	ABI1	ABI1	ABI1	ABI1	ABI1	ABI1	ABI1	ABI1	TTC1	ABI1	ABI1	ABI1	
DHARS	AUX1	AUX1	AUX1	AUX1	BLUS1	AUX1	AUX1	ALMT9	ALMT9	ALMT9	TPC1	ALMT9	ALMT9	AUX1	
EGL3	BLUS1	BLUS1	BHP	BHP	DHARS	BEARSKIN1	BHP	ARF7	ARF7	ARF7	TMM	ARF7	ARF7	BHP	
GL2	DHARS	DHARS	BLUS1	BLUS1	BLUS1	EGL3	BHP	AUX1	AUX1	AUX1	SPCH	AUX1	AUX1	BLUS1	
GL3	EGL3	EGL3	DHARS	DHARS	GL2	BLUS1	DHARS	BEARSKIN1	BEARSKIN1	BEARSKIN1	SLAC1	BEARSKIN1	BEARSKIN1	DHARS	
KAN1	GL2	GL2	EGL3	EGL3	GL3	BODENLOS	BHP	BHP	BHP	BHP	SCRM	BHP	BHP	EGL3	
KAN2	GL3	GL3	ERECTA	ERECTA	MDAR3	DHARS	ERECTA	BLUS1	BLUS1	BLUS1	SCR	BLUS1	BLUS1	EPFL9	
KAN4	KAN1	KAN1	FAMA	FAMA	OST1	EGL3	FAMA	BODENLOS	BODENLOS	BODENLOS	ROP9	BODENLOS	BODENLOS	ERECTA	
MDAR3	KAN2	KAN2	GL2	GL2	PDR3	EPFL9	CNCG15	CNCG15	CNCG15	CNCG15	RBR1	CNCG15	CNCG15	FAMA	
OST1	KAN4	KAN4	GL3	GL3	PHOT1	ERECTA	GL3	DHARS	DHARS	DHARS	PIN-FORMED	DHARS	DHARS	GL2	
PDR3	MDAR3	MDAR3	HT1	HT1	PHOT2	FAMA	HT1	EGL3	EGL3	EGL3	PHOT2	EGL3	EGL3	GL3	
PIN-FORMED	OST1	OST1	KAN1	KAN1	PIN-FORMED	GL2	KAN1	EPF1	EPF1	EPF1	PHOT1	EPF1	EPF1	GORK	
RBR1	PDR3	PDR3	KAN2	KAN2	RBR1	GL3	KAN2	EPFL9	EPFL9	EPFL9	PDR3	EPFL9	EPFL9	HT1	
ROP9	PHOT1	PHOT1	KAN4	KAN4	ROP9	GORK	KAN4	ERECTA	ERECTA	ERECTA	OST1	ERECTA	ERECTA	KAN1	
SLAC1	PHOT2	PHOT2	MDAR3	MDAR3	STM	HT1	MDAR3	FAMA	FAMA	FAMA	MUTE	FAMA	FAMA	KAN2	
STM	PIN-FORMED	PIN-FORMED	MUTE	MUTE	TMO5	KAN1	MUTE	GL2	GL2	GL2	MDAR3	GL2	GL2	KAN4	
TPC1	RBR1	RBR1	OST1	OST1	GL3	KAN2	OST1	GL3	GL3	GL3	KAN4	GL3	GL3	LYM1	
TTG1	ROP9	ROP9	PDR3	PDR3	YODA	KAN4	PDR3	GORK	GORK	GORK	KAN2	GORK	GORK	MDAR3	
	SLAC1	SCRM	PHOT1	PHOT1	LBD16	LBD16	PHOT1	HT1	HT1	HT1	KAN1	HT1	HT1	MUTE	
	STM	STM	PHOT2	PHOT2	Lonesome highway	PHOT2	KAN1	KAN1	KAN1	KAN1	KAN1	HT1	KAN1	KAN1	OST1
	TPC1	STM	PIN-FORMED	PIN-FORMED	LYM1	PIN-FORMED	KAN2	KAN2	KAN2	KAN2	GL3	KAN2	KAN2	PDR3	
	TTG1	TMO5	PYL8	PYL8	MDAR3	RBR1	KAN4	KAN4	KAN4	KAN4	GL2	KAN4	KAN4	PHOT1	
	TPC1	RBR1	RBR1	RBR1	MUTE	ROP9	LBD16	LBD16	LBD16	LBD16	FAMA	LBD16	LBD16	PHOT2	
	YODA	YODA	YODA	YODA	SCR	Lonesome highway	Lonesome highway	Lonesome highway	Lonesome highway	Lonesome highway	ERECTA	Lonesome highway	Lonesome highway	PIN-FORMED	
					SCR	PDR3	SCR	LYK5	LYK5	LYK5	EGL3	LYK5	LYK5	PYL8	
					SCR	PHOT1	SLAC1	LYM1	LYM1	LYM1	DHARS	LYM1	LYM1	RBR1	
					SLAC1	SLAC1	SPCH	MDAR3	MDAR3	MDAR3	BLUS1	MDAR3	MDAR3	ROP9	
					SPCH	SPCH	PIN-FORMED	MUTE	MUTE	MUTE	BHP	MUTE	MUTE	SCR	
					TMM	TMM	MYB60	MYB60	MYB60	MYB60	AUX1	MYB60	MYB60	SCRM	
					TMM	TMM	TPC1	OST1	OST1	OST1	AB1	OST1	OST1	SLAC1	
					TMO5	TPC1	RBR1	TTG1	PDR3	PDR3	PDR3	PDR3	PDR3	SPCH	
					TPC1	TTG1	YODA	PHOT1	PHOT1	PHOT1	PHOT2	PHOT1	PHOT1	STM	
					WER	WER	SCR	PIN-FORMED	PIN-FORMED	PIN-FORMED	PIN-FORMED	PIN-FORMED	PIN-FORMED	TMM	
					WOX4	WOX4	SLAC1	PRSL1	PRSL1	PRSL1	PRSL1	PRSL1	PRSL1	TMO5	
					YODA	YODA	SOMBRERO	PYL8	PYL8	PYL8	PYL8	PYL8	PYL8	WER	
							SPCH	RBR1	RBR1	RBR1	RBR1	RBR1	RBR1	YODA	
							STM	ROP9	ROP9	ROP9	ROP9	ROP9	ROP9		
							TMM	SCR	SCR	SCR	SCR	SCR	SCR		
							TMO5	SCRM	SCRM	SCRM	SCRM	SCRM	SCRM		
							TPC1	SLAC1	SLAC1	SLAC1	SLAC1	SLAC1	SLAC1		
							TTG1	SOMBRERO	SOMBRERO	SOMBRERO	SOMBRERO	SOMBRERO	SOMBRERO		
							WER	SPCH	SPCH	SPCH	SPCH	SPCH	SPCH		
							WOX4	STM	STM	STM	STM	STM	STM		
							YODA	TMM	TMM	TMM	TMM	TMM	TMM	BEARSKIN1	
								TMO5	TMO5	TMO5	TMO5	TMO5	TMO5	BODENLOS	
								TPC1	TPC1	TPC1	TPC1	TPC1	TPC1	Lonesome highway	
								TTG1	TTG1	TTG1	TTG1	TTG1	TTG1	PRSL1	
								WER	WER	WER	WER	WER	WER	SOMBRERO	
								WOX4	WOX4	WOX4	WOX4	WOX4	WOX4	TPC1	
								YODA	YODA	YODA	YODA	YODA	YODA	WOX4	

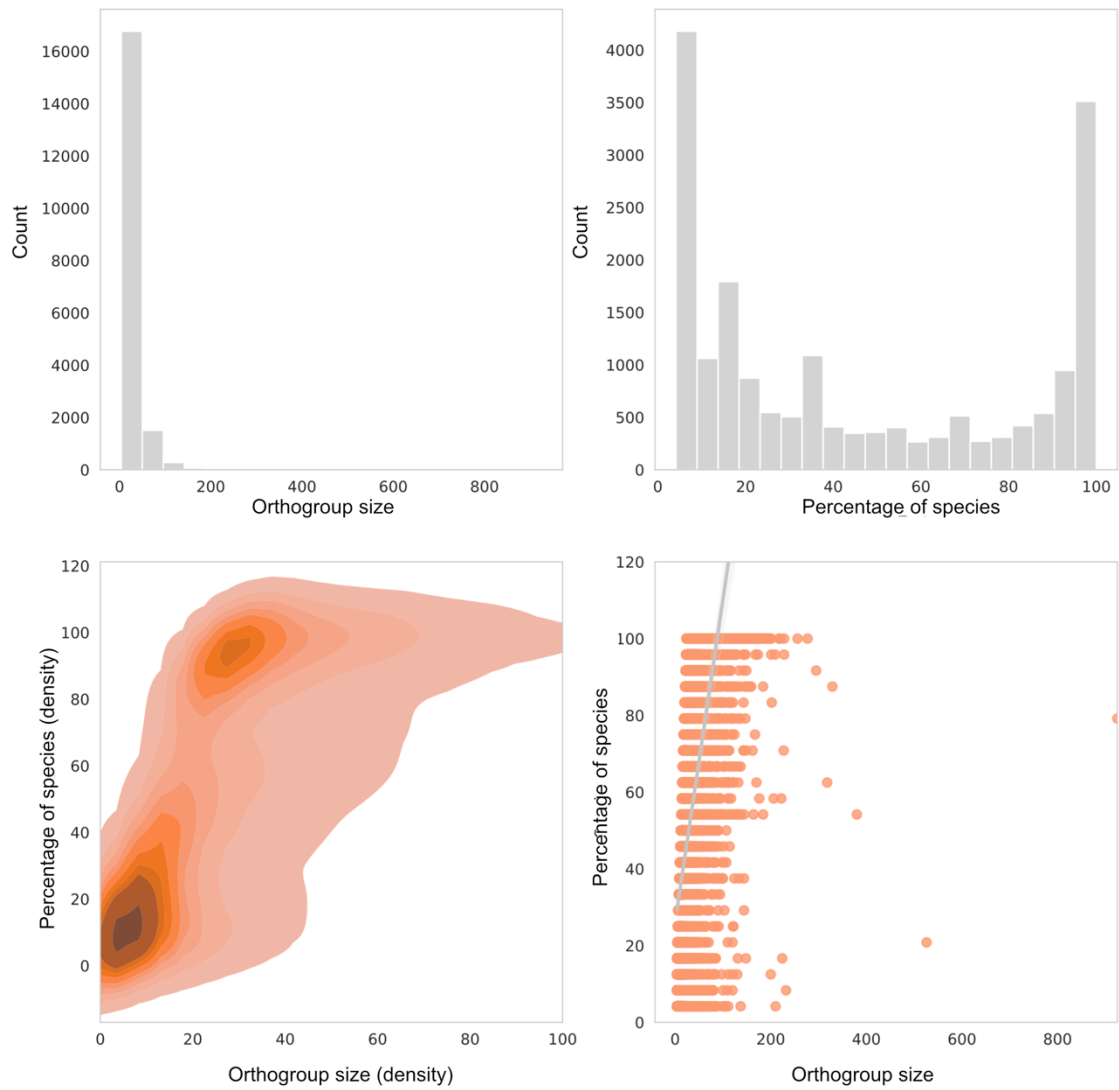
Supplementary Fig. 10 | Gain and loss of orthologous gene family clusters across the embryophyte tree. The labels correspond to a list of the key orthologous gene family clusters present at each branch. Fifty critical genes for Arabidopsis, stomata, vasculature and symbiosis related function/development were used to illustrate the gain and loss of orthologous gene family clusters across the embryophyte tree. Gene families are labelled according to their characterised representatives; for example, the “SLAC1” gene family is the family that contains Arabidopsis SLAC1. Gene family clusters that originated on the branch are coloured in green, whilst ones that are lost are coloured red and all genes present on the branch are in black.



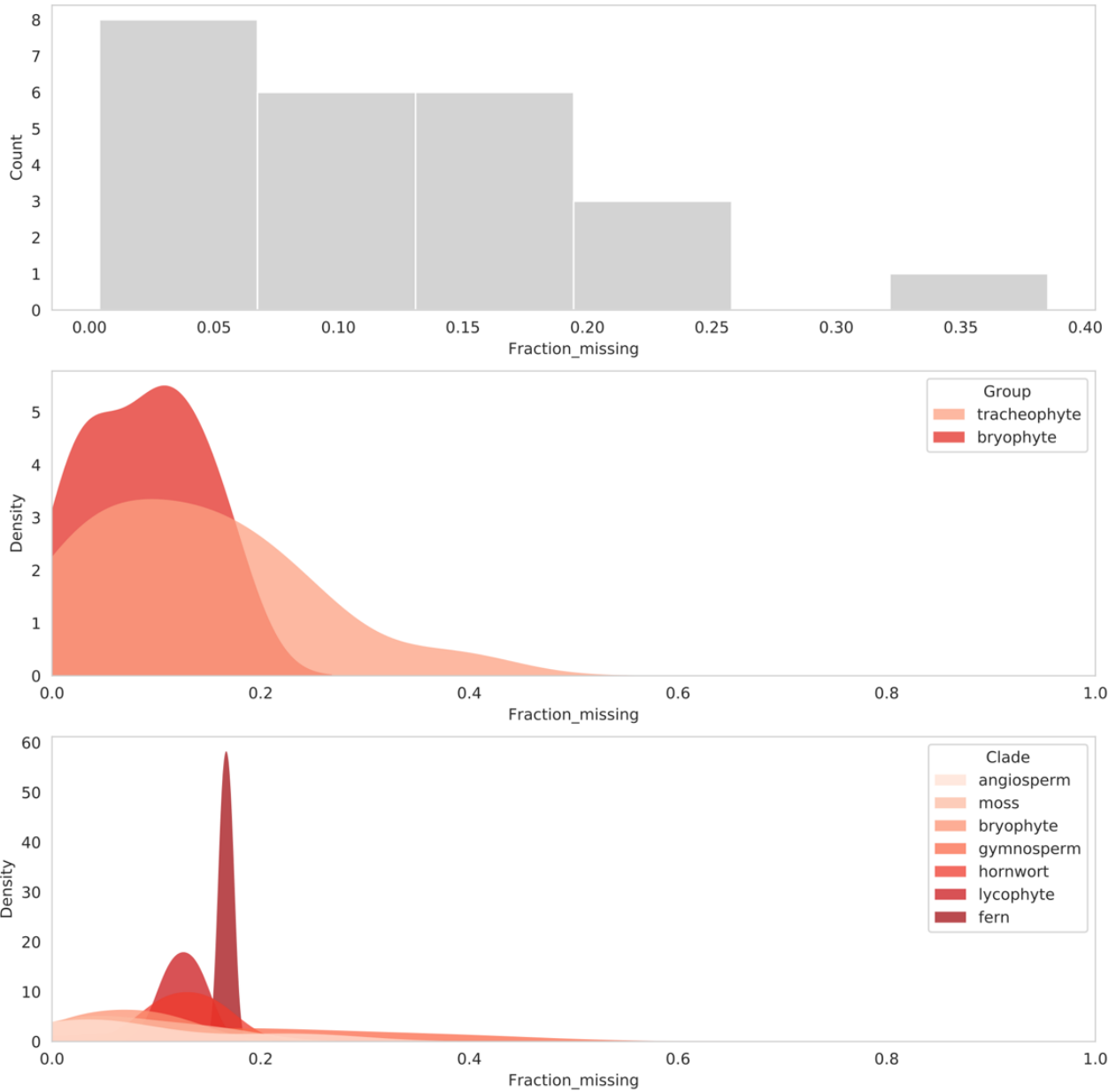
Supplementary Fig. 11 | Gene family clusters for genome and transcriptome dataset – 154 species. Top left, distribution of gene family size. Top right, distribution of gene family species representation. Bottom left, density of gene family species representation and size. Bottom right, regression of gene family species representation and size.



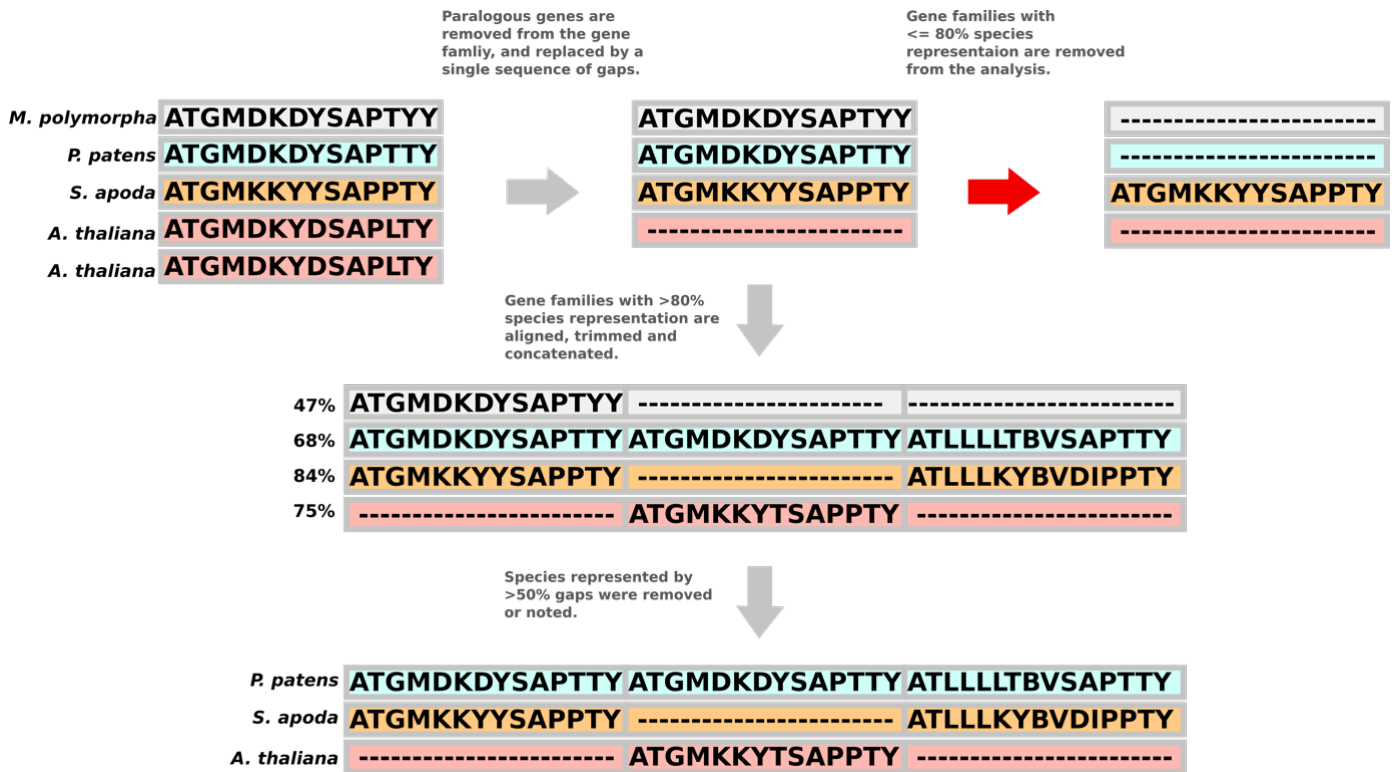
Supplementary Fig. 12 | BUSCO completion of genome and transcriptome dataset – 154 species. Top, histogram of BUSCO completeness. Middle, distribution of BUSCO completeness separated into the two main plant groups. Bryophytes, on average, have more complete BUSCOs than tracheophytes. Bottom, distribution of BUSCO completeness separated by lineage. Angiosperms had the most varied completeness values. As the dataset comprised of both genomes and transcriptomes, the underlying variation per lineage may be ascribed to the more incomplete transcriptomes.



Supplementary Fig. 13 | Gene family clusters for genome dataset - 24 species. Top left, distribution of gene family size. Top right, distribution of gene family species representation. Bottom left, density of gene family species representation and size. Bottom right, regression of gene family species representation and size.



Supplementary Fig. 14 | BUSCO completion of genome only dataset – 24 species. Top, histogram of BUSCO completeness. Middle, distribution of BUSCO completeness separated into the two main plant groups. Bryophytes, on average having more complete BUSCOs than tracheophytes. Bottom, distribution of genome completeness separated by lineage. BUSCO completeness is significantly higher for the genome only dataset.



Supplementary Fig. 15 | Graphical representation of orthology inference approach. Species represented by more than one gene copy are removed from the gene family alignment and replaced by a single sequence of gaps. Gene families with more than 80% of the original species present are retained. Gene families are then aligned, trimmed and concatenated together to form a super matrix. Species with less than 50% gaps in the alignment are removed or noted if within 10% error – this step prevents single species that are continually removed from gene families to be in the final alignment.