Science Advances

Supplementary Materials for

Internal tsunamigenesis and ocean mixing driven by glacier calving in Antarctica

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Fig. S1. Bathymetry and crevasse displacement rate for William Glacier. (a) Seabed bathymetry of Börgen Bay from multibeam echosounder data, and coastline/topography from

Landsat imagery, with crevasse displacement rate and location shown in panel b. (b) Location of the depth profile of the grounded ice front presented in panel c. (c) Depth profile of the grounded ice front.



Fig. S2. Börgen Bay oceanographic characteristics. Potential temperature versus salinity for the Börgen Bay CTD profiles, coloured by beam transmission. Contours of potential density anomaly are marked. High beam transmission denotes clearer waters; low beam transmission in deeper (denser) waters denotes waters with comparatively high concentration of sedimentary particles. Glacial meltwater lies along the steep diagonal dashed line defined by Gade (75), whilst subglacial discharge lies along the near-horizontal dashed line. The data points with low beam transmission lie along the Gade (75) line at depth; these are stations featuring strong glacial meltwater input, which has low beam transmission due to the presence of sedimentary particles in the water. This water rises at the front of William Glacier and spreads as a plume across Börgen Bay at approximately 50m depth. The mixing caused by the glacier calving resulted in the locus of points lying above both lines.



Fig. S3. Near-surface underway hydrographic data. Underway near-surface temperature along the ship track in Börgen Bay, for the periods before (upper panel) and after (lower panel) the calving event. The marked decrease in temperature at the time of the calving event is apparent, consistent with upward mixing of cooler waters from below.



Fig. S4. **Near-surface underway hydrographic data in temperature-salinity space**. The change in upper-layer temperature (Fig. S3) is concurrent with a shift toward higher salinity, indicating that it was caused by upwards mixing of more saline water below, as opposed to injection of cold (fresh) meltwater at the surface.



Fig. S5: Numerical simulations with 50% tapered bathymetry. As for Fig. 7, but with the bathymetry in the centre of the domain tapered by 50% towards a flat bottom. (a-e) Potential temperature at the time intervals given in the individual panels. The peak inflow for the imposed wave is $U_0 = 1.536$ m s⁻¹ and the central ridge has been tapered to 50% of its height from the bottom. At 3 hours the impulsive flow reaches its peak velocity and by 6 hours it is effectively zero. Note the nonlinear colour scale, which emphasises the temperatures where most of the mixing takes place. (f) Domain average kinetic energy for 4 different tapers as per the legend. (g) A series of 4-hour averages of squared buoyancy frequency, spatially averaged over the top 60 m.



Fig. S6: Numerical simulations with varying wave strengths and bathymetric tapers. 2D model potential temperature after 24 hours for a range of model parameters. (a,c,e,g,i) varies the strength of the incoming wave, as per the individual captions; mixing between 50 and 100 m is seen to be critically dependent upon this strength. For (b,d,f,h,j), the strength of the wave is held constant at 1.536 m s⁻¹ and the height of the central ridge is tapered towards a flat bottom by the percentage in the individual caption. Note the progressively weaker effect of the mixing as the bathymetric taper is increased.



Fig. S7. Histograms of dissipation due to non-tsunami processes. (a) Histogram of tidal conversion computed at 30 arc-second resolution. (b) Bed friction dissipation. (c) Wind / near-inertial shear-induced dissipation. (d) Surface cooling-induced dissipation. In each panel, mean values for the West Antarctic Peninsula shelf are stated.

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